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SULCAL: A MODEL of SULFUR CHEMISTRY in a PLUME

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SULCAL: A MODEL OF SULFUR CHEMISTRY IN A PLUME

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ABSTRACT

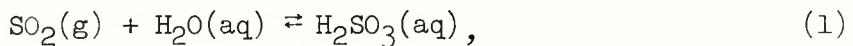
A computer program has been written that models the following features of the chemical behavior of sulfur emitted to the atmosphere from fossil-fuel burning power plants: (1) the rapid reactions of SO_2 with aerosol droplets to produce the dissolved sulfite species H_2SO_3 , HSO_3^- and SO_3^{2-} ; (2) the rapid reactions of SO_3 to produce the dissolved sulfate species HSO_4^- and SO_4^{2-} ; (3) the neutralization of the acid thus produced by atmospheric ammonia and (4) the eventual formation of particulate ammonium sulfate; (5) the slow oxidation of SO_2 to sulfate species by hydroxyl radical, and (6) the slow oxidation of species of SO_2 in aerosol droplets to sulfate species by dissolved ozone and oxygen. The model employs averaged concentrations based on the Gaussian plume and can calculate deposition rates for gaseous and particulate material as a function of such variables as distance from the source, wind speed, meteorological stability class, temperature, relative humidity, and the ambient concentrations of OH radical, ozone, and ammonia.

1. INTRODUCTION

Industrially produced sulfur enters the atmosphere in reactive forms, primarily SO_2 and SO_3 , that cannot be treated as inert substances while they remain in the atmosphere. This report describes a stand-alone model of atmospheric transport which includes the effects of (1) rapid chemical reactions that produce new species and (2) slow reactions by which SO_2 or its products from rapid reactions are oxidized to sulfates in various forms. As shall be seen, the most important consequence of these reactions is the production of an aerosol of liquid acidic droplets; hence, one object of the calculation is to estimate the amount and composition of the aerosol in a unit volume of the plume and its deposition rate. The model was developed for use as a sub-model in the Air Transport Model (ATM).¹

1.1 The Rapid Reactions of Sulfur

Both gaseous SO_2 and SO_3 can undergo rapid chemical reactions with constituents of the atmosphere. Gaseous SO_2 can dissolve in aerosol droplets producing the following products in solution:*

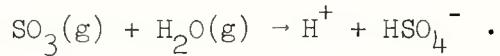


These reactions are reversible as well as rapid and as a result the concentration of the species involved are related to one another by

*The abbreviations (g), (aq), and (c) used in chemical reactions denote species in the gaseous, aqueous, or solid states, respectively.

equilibrium constants (i.e., K_1 , K_2 , and K_3 in Table I) that have known values² as a function of temperature. These, and the other equilibrium constants in Table I, include in their expressions a quotient of so-called activity coefficients (γ_i) which, as shall be seen, are functions of the composition of the solution involved. (The activity coefficients reflect the non-ideal behavior of the individual ions in the solution.) The tendency of SO_2 to form the aqueous species H_2SO_3^- , HSO_3^- , and SO_3^{2-} is weak and at the low concentrations of SO_2 in the atmosphere (usually <0.1 ppm) reaction occurs only if condensed water is already present.

In contrast, gaseous SO_3 reacts vigorously and completely with water, even in the gaseous form, to produce acid droplets



The bisulfate ion HSO_4^- undergoes rapid and reversible dissociation,



As a result, no trace of SO_3 gas is found in the atmosphere. If it is introduced, it is converted completely to acid droplets containing the ions H^+ , HSO_4^- and SO_4^{2-} , the concentrations of which are related to one another by K_4 (Table I).

The acid (H^+) produced by the solution of SO_2 and SO_3 in water can be neutralized by atmospheric ammonia in another set of rapid and reversible reactions, producing the aqueous ammonium ion,



and K_5 and K_6 apply. These reactions can lead to the formation of particles of ammonium sulfate, $(\text{NH}_4)_2\text{SO}_4$, if they proceed far enough and the relative humidity is less than ~83%. When both liquid and solid are

Table I Equilibrium Constants^a

Text Symbol ^b	Definition ^b	Log K(25°C)	Log K=A+B/T(K)		Program Name
			A	B	
K_1	$K_1 = \frac{[H_2SO_3]}{P_{SO_2} a_w}$	0.090	-4.4912	1365.9	EKS
K_2	$K_2 = \frac{[H^+][HSO_3^-]}{[H_2SO_3]} \cdot (g_{H^+} g_{HSO_3^-})$	-1.765	-4.815	909.4	EK1S
K_3	$K_3 = \frac{[H^+][SO_3^{2-}]}{[HSO_3^-]} \cdot \left(\frac{g_{H^+} g_{SO_3^{2-}}}{g_{HSO_3^-}} \right)$	-7.220	-8.8549	487.36	EK2S
K_4	$K_4 = \frac{[H^+][SO_4^{2-}]}{[HSO_4^-]} \left(\frac{g_{H^+} g_{SO_4^{2-}}}{g_{HSO_4^-}} \right)$	-1.994	-5.8348	1145.2	EK3S
K_5	$K_5 = \frac{[NH_3]}{P_{NH_3}}$	1.756	-4.233	1785.5	EKN
K_6	$K_6 = \frac{[NH_4^+]}{[NH_3][H^+]} \left(\frac{g_{NH_4^+}}{g_{H^+}} \right)$	9.251	0.1026	2727.5	EK1N

Table I (continued)

Text ^b Symbol	Definition ^b	Log K(25°C)	Log K=A+B/T(K)		Program Name
			A	B	
K_7	$= [NH_4^+]^2 [SO_4^{2-}] (g_{NH_4^+}^2 g_{SO_4^{2-}})$	-0.221 ^c	0.929 ^c	-343.1	EKSS
K_8	$= \frac{[H_2CO_3]}{P_{CO_2} a_w}$	-1.462	-5.0160	1059.74	EKC
K_9	$= \frac{[H^+][HCO_3^-]}{[H_2CO_3]} (g_H^+ g_{HCO_3^-})$	-6.368	-5.0300	-398.85	EK1C
K_{10}	$= \frac{[H^+][CO_3^{2-}]}{[HCO_3^-]} \left(\frac{g_H^+ g_{CO_3^{2-}}}{g_{HCO_3^-}} \right)$	-10.328	-7.7260	-775.85	EK2C
K_{13}	$= \frac{[O_3]}{P_{O_3}}$	-2.146 ^d	-8.488 ^d	1891 ^d	EKOZ
K_w	$= \frac{[H^+][OH^-]}{a_w}$	-14.00	-4.2134	-2916.54	EKW
$P_{H_2O}^o$	$= \frac{P_{H_2O}}{a_w}$	-1.505	6.2057	-2298.9	PWO

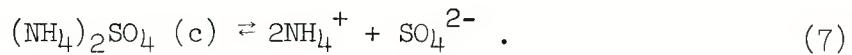
^aNumerical values are based on NBS Technical Note 270-3, et seq² unless otherwise indicated and should represent K values with adequate accuracy from 0 to 40°C.

^bSubscript on K refers to Eq. No. in text; concentrations of species in solution [i] are in molal units (mmol/g H₂O); gas pressures P_i are in atmospheres; activity coefficients of ions g_i are calculated as described in Appendix C.1.

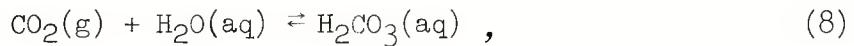
^cAdjusted to give the observed solubility of (NH₄)₂SO₄ in water.

^dBased on data of Rowson quoted in Seidell,¹⁷ Vol. II, p 1240.

present, the following equilibrium (and K_7) applies,



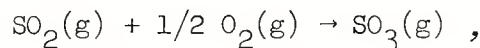
Another atmospheric gas that can affect the acidity (pH) of aerosol droplets is CO_2 .



All three reactions are rapid and reversible and the concentrations of the species involved are related by K_8 , K_9 , K_{10} .

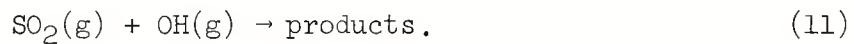
1.2 The Slow Oxidation of SO_2

The slower reactions to be considered are those by which SO_2 and some of the products of its rapid reactions (all of which are species of tetravalent sulfur) are oxidized to sulfate (hexavalent sulfur) species in the atmosphere. The direct oxidation of SO_2 by atmospheric oxygen



(followed instantly by reaction of SO_3 with H_2O) is slow and proceeds by a photochemical path. More rapid oxidation can occur by at least three other paths.

Recently Castleman et al³ and Davis and Klauber⁴ have concluded that the most rapid reaction involving a gaseous oxidant is that with the hydroxyl (OH) radical,



The rate expression and rate constant (k_{11}) are given in Table II. The concentration of OH radical, an intermediate in other reactions, is found by Wang et al⁵ to be in the range $<2 \cdot 10^{-13}$ to $2 \cdot 10^{-12}$ atm, and

Table II Rate Expressions for S(IV) Oxidation

Rate Expression	k^a	Log $k = A + B/T(K)$		Pro- gram Name
	(25°C)	A	B	
(1) Oxidation of $SO_2(g)$ by $OH(g)$, ^(b) $- \frac{dP}{dt} SO_2 = k_{11} P_{SO_2} P_{OH} .$	$k_{11} = 1.83 \cdot 10^7$	7.26	(0) ^e	RK4
(2) Oxidation of HSO_3^- by O_3 , ^(c) $- \frac{d[HSO_3^-]}{dt} = k_{12} [HSO_3^-] g_{X^-} [O_3] .$	$k_{12} = 9.8 \cdot 10^5$	14.61	-2570 ^e	RK3
(3) Oxidation of SO_3^{2-} by O_2 , ^(d) $- \frac{d[SO_3^{2-}]}{dt} = [K_{14}^{'} + K_{14}^{''} ([H^+] g_{H^+})^{1/2}] [SO_3^{2-}] g_{X^{2-}} .$	$k_{14}^{'} = 0.013$ $k_{14}^{''} = 59$	11.438 15.095	-3973 -3973	RK1 RK2

^aSubscript on k refers to the Eq. No. in text. Concentration and pressure units are as in Table I, footnote b. Time is in sec.

^bCastleman et al,³ Davis and Klauber.⁴

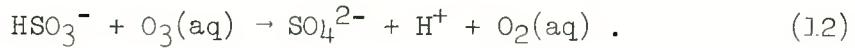
^cPenkett⁶ measured k_{12} at 9.6°C.

^dMcKay,⁷ Fuller and Crist.⁸

^eOur estimate.

hence the half-time for reaction 11 could be as short as a few hours.

Of the reactions that occur in solution, the oxidation of HSO_3^- by ozone (O_3) may be the fastest,⁶

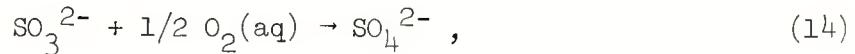


This reaction is preceded by



With an atmospheric concentration of ozone of 10^{-7} atm, we estimate from k_{12} and K_{13} (Tables I and II) a half life of 2 min. for HSO_3^- in solution.

Another reaction which is known to occur in solution involves the oxidation of SO_3^{2-} by dissolved oxygen.^{7,8} While the overall reaction is



the mechanism is apparently complex and is not known in detail. The reaction can be catalyzed by some metal ions and retarded by others.

In pure solutions the kinetics are described by the equations and constants in Table II.

1.3 Other Reactions

There are, of course, other reactions that could affect the behavior of sulfur in the atmosphere, including reversible equilibria such as the solution of NO_2 in aerosol droplets and slow reactions such as the dissolution of a flyash particle in an acid droplet or the catalytic oxidation of SO_2 at particle surfaces. There are numerous other constituents of a plume which could react with SO_2 or which certainly will affect the concentrations of ozone and hydroxyl radical. Such other reactions are not included in the present model, either because they are poorly

understood at present, or because they would greatly complicate the model and its validation.

2. THE MODEL

From the foregoing the essential features of the chemistry of sulfur in the atmosphere which might be included in a model are (1) the slow oxidation of SO_2 to produce acidic aerosol droplets in which (2) various rapidly attained and shifting chemical equilibria are established (Table I) including (3) progressive neutralization by ammonia, which--if the aerosol remains in the atmosphere long enough--can ultimately produce a relatively innocuous solid, $(\text{NH}_4)_2\text{SO}_4$. The computations performed by SULCAL to accomplish this may be summarized as follows (Fig. 1):*

1. M. SULCAL, the main program, first calculates the emission rate of plume components from the source.
2. S. SIG is called to obtain the effective volume of the plume for the calculation of component concentrations at a starting point a short distance (x) from the source.
3. M. SULCAL then calls S. KUTTA⁹ to begin the calculation of the composition of the plume at increasing values of x .
4. S. KUTTA calls S. RATE repeatedly to determine the rate of change of the concentration of the plume components with x , to be used in a Runge-Kutta integration of the component concentrations vs. x .
5. Each time S. RATE is entered it first calls S. EXPCT which determines by extrapolation or interpolation on x the starting values of several quantities to be refined in the calculation of the composition of the plume.

* Here and elsewhere M and S denote main and subroutine programs, respectively.

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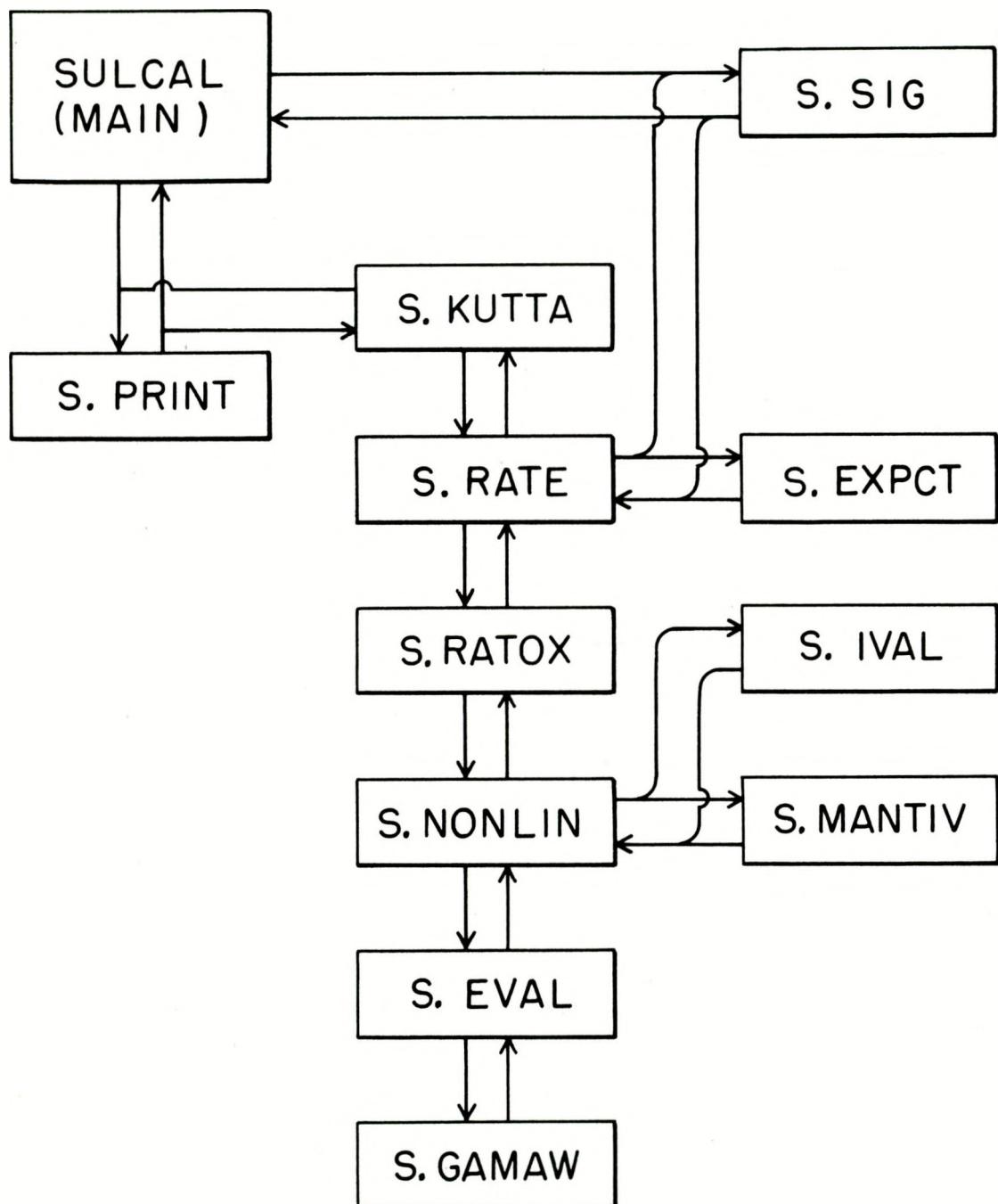


Fig. 1 Program SULCAL.

6. S. RATE next calls S. SIG to determine the effective volume of the plume at the current value of x and its rate of change with x.
7. S. RATE then calls S. RATOX to calculate the rate of S(IV) oxidation.
8. S. RATOX first calls S. NONLIN¹⁰ which employs the secant method to refine quantities needed to calculate the composition of the gas and aerosol phases of the plume.
9. S. NONLIN calls S. EVAL repeatedly to calculate these compositions to test for convergence of the quantities being refined.
10. S. EVAL calls S. GAMAW repeatedly to refine values of needed ionic activity coefficients and the activity of water, quantities which are themselves dependent on the composition of the droplets.
11. With the amount and composition of each phase in the plume returned to S. RATOX, the rate of S(IV) oxidation is calculated in the gas phase and in the droplets.
12. With the oxidation rates returned to S. RATE, the rate of change of the component concentrations are calculated, including also the effects of dilution and deposition; these are returned to S. KUTTA.
13. S. KUTTA returns to M. SULCAL integrated values of the component concentrations at specified values of x.
14. S. KUTTA and M. SULCAL call S. PRINT to put out results as the calculation proceeds.

The important features of the calculation are described in the following

sections. The program is listed in Appendix A.

2.1 The Plume Components

There are seven chemical constituents whose concentrations specify the composition of the plume in sufficient detail for the purposes of the model. They include S(IV), S(VI), CO₂, H₂O, NH₃, O₃, and OH. The first four come in part from the stack. It is assumed that S(IV) is emitted from the stack entirely as SO₂(g), and S(VI) entirely as droplets with the composition SO₃·nH₂O. The concentrations of all seven components in the surrounding air are specified as input data.

The rates (Q_i) at which SO₂, SO₃·nH₂O, CO₂ and H₂O--along with unreacted atmospheric gases--are emitted from the stack are calculated from input data (Table III), including the composition of the coal and the rate at which it is burned, the excess of air used, the fraction of sulfur converted to SO₃, and the moles of water (n) condensed by a mole of SO₃. The flow rate of gas (V_p°) at the mouth of the stack is calculated from the mass flow rate and the specified exit temperature.

The total concentration (C_i) of each of the seven components is to be calculated vs. x. A Gaussian plume is assumed for the estimation of dilution and deposition.

2.2 Effective Concentrations and Plume Volume

The simplest conditions under which to study chemical reactions would involve well-mixed solutions of uniform concentration and temperature. Though the conditions in a plume are far from this ideal, it is possible to define mean concentrations and temperature, and effective rate and equilibrium constants under the assumption of rapid mixing such that the net rate of oxidation to sulfur (VI) is the same

Table III Input Data

	Units	Typical Value	Text Symbol	Program Name
Carbon in coal	wt %	82.57 ^a		WPCTC
Hydrogen in coal	"	6.93		WPCTH
Sulfur in coal	"	1.0		WPCTS
Rate of coal combustion	kg/sec	100		QCOAL
Fraction of S converted to SO ₃	--	0.02		FCTS6
Fractional excess of air in combustion	--	0.25		EXAIR
Moles H ₂ O condensed per moles SO ₃	--	5	n	FCON
Temperature of plume at mouth of stack	°C	177		TCP
Ambient temperature	°C	21.11		TCA
Ambient concentration of S(IV)	mmol/m ³	0.004	A ₁	A(1)
Ambient concentration of S(VI)	"	0	A ₂	A(2)
Ambient concentration of O ₃	"	0.0032	A ₃	A(3)
Ambient concentration of OH	"	2·10 ⁻⁸	A ₄	A(4)
Ambient concentration of CO ₂	"	12.25	A ₅	A(5)
Ambient concentration of NH ₃	"	0.0006	A ₆	A(6)
Ambient concentration of H ₂ O	"	930	A ₇	A(7)
Effective height of stack	m	400	h	H
Meteorological stability class	--	1 to 6 ^b		NSC
Roughness length of terrain	--	1 to 6 ^b		NRL
Wind speed	m/sec	(a)	u _x	UX
Deposition velocity of gas	m/sec	0.01	v _d ^g	VDG
Deposition velocity of particle	m/sec	0.01	v _d ^p	VDP
Initial distance from source	m	300		XNEXT
Initial trial value of Δx	m	0.1		DELX

	Units	Typical Value	Text Symbol	Program Name
Fractional error of integrated quantities	--	$3 \cdot 10^{-6}$		ERRR
No. of distance intervals	--	--		NX
Size of distance interval	m	--		XINCR (NX)

^aThis value is higher than is typical of power plant fuels.

^bThe numbers 1 to 6 denote stability classes A to F and roughness lengths of 1, 4, 10, 40, 100, and 400 m, respectively. The range of wind speeds depends on the stability class (Table IV).

for the idealized uniform case as for a real plume or for the model Gaussian distribution of concentrations. This can be achieved for reactions involving concentrations through second order in the Gaussian distributed concentrations by the artifice of an effective volume related to the Gaussian dispersion coefficients. Higher order reaction rates can be preserved if desired through appropriate scaling of the rate constants. The justification is given in Appendix B. We use the Gaussian distribution here not because we think it fits the plume distribution better than other forms, but rather because these are existing measurements of plume volume which have been expressed in terms of the Gaussian dispersion.

In terms of the effective plume volume V_p , the concentration (C_i) of each plume component is given by

$$C_i = Q_i/V_p + A_i(1 - V_p^*/V_p), \quad (15)$$

where Q_i is the emission rate of component i (mmol/sec), A_i its ambient concentration (mmol/m³), and V_p^* the initial plume volume flow rate (m³/sec). Thus the effective plume volume is a volume flow rate (m³/sec) at a given x , or, a volume (m³) of a length of plume u_x where u_x is the wind velocity (m/sec).

The effective plume volume V_p is related to the dispersion coefficients σ_y and σ_z (m) of the Gaussian distribution by the expression (see Appendix B)

$$V_p = 4\pi \sigma_y \sigma_z u_x^\beta \text{ (m}^3/\text{sec)} , \quad (16)$$

where β is a correction factor which accounts for the change in the concentration distribution due to the interaction with (or "reflection"

from) the ground. The factor β is given by the expression (Appendix B)

$$\beta = \left[1 + \frac{1}{\sqrt{2}} \exp \left(-h^2 / \sigma_z^2 \right) \right]^{-1}, \quad (17)$$

where h is the effective stack height (m). Note that when $\sigma_z \ll h$ then $\beta = 1$ and when $\sigma_z \gg h$ then $\beta = .59$. The value of β changes smoothly as the ratio h/σ_z passes through unity.

The dispersion coefficients used in the effective reaction volume determination are instantaneous or short-time averaged coefficients (σ_s), and are smaller than the long-time averaged dispersion coefficients (σ_ℓ) described by Hosker¹¹ (Table IV) which are used to calculate ground level concentrations and depositions. Both coefficients are a function of the meteorological stability class, the distance x from the stack, and a roughness length characteristic of the terrain. To these dispersion coefficients we have added a small correction σ°

$$\sigma^\circ = \sqrt{V_p^\circ / 4\pi u_x} , \quad (18)$$

which keeps the plume volume finite at the source, preventing infinite concentrations and reaction rates.

The rate of change of average component concentration with distance due to surface deposition is given by

$$\frac{dC_i}{dx} = - C_i V_d \gamma / u_x , \quad (19)$$

where V_d is the conventionally defined deposition velocity and γ is

$$\gamma = (\sqrt{2/\pi} \sigma_z) \exp \left(-h^2 / 2\sigma_z^2 \right) . \quad (20)$$

γ and V_p are calculated for each specified value of x in S. SIG; σ_z is the conventional long-time averaged dispersion coefficient.

The short-time dispersion coefficient σ_s can be related to the long-

Table IV Calculation of Dispersion Coefficients^a

$$\sigma_y(x) = C_3 x / \sqrt{1 + 0.001x} + \sigma^o$$

$$\sigma_z(x) = \ln(c_1 x^{d_1}) \cdot \frac{a_1 x^{b_1}}{1 + a_2 x^{b_2}} + \sigma^o$$

$$(\sigma^o = \sqrt{V_p^o / (4\pi u_x)})$$

Stability Class	u_x (m/sec)	a_1	b_1	a_2	b_2	c_3
A (1)	1 - 2.5	0.112	1.06	$5.38 \cdot 10^{-4}$	0.815	0.22
B (2)	1.5 - 5	0.130	0.950	$6.52 \cdot 10^{-4}$	0.750	0.16
C (3)	2 - >6	0.112	0.920	$9.05 \cdot 10^{-4}$	0.718	0.11
D (4)	5 - >10	0.098	0.889	$1.35 \cdot 10^{-3}$	0.688	0.08
E (5)	2 - 5	0.0609	0.895	$1.96 \cdot 10^{-3}$	0.684	0.06
F (6)	2 - 3	0.0638	0.783	$1.36 \cdot 10^{-3}$	0.672	0.04

Roughness Length (cm)	c_1	d_1
1 (1) ^b	1.56	0.0480
4 (2)	2.02	0.0269
10 (3)	2.718	0
40 (4)	5.16	-0.060
100 (5)	7.37	-0.0957
400 (6)	11.7	-0.128

^a From Hosker¹¹; these are long-time averaged (σ_y) values. A factor has been omitted from the expression for σ_z since it differs little from unity (< 10%) for the distances of interest (< 30 km).

^b Numbers are values of NSC and NRL supplied as input data (Table III).

term coefficients σ_ℓ by

$$\sigma_s = f \sigma_\ell . \quad (21)$$

The form of f is given in Appendix B.

In equation 15, additivity of gas volumes is assumed. By the assumption of the ideal gas law, the mean absolute temperature of the plume is then given as

$$T = \frac{T_a}{1 - (1 - T_a/T_p) (V_p^\circ/V_p)} , \quad (22)$$

where T_p and T_a are respectively the absolute temperature of the plume at the mouth of the stack and of the ambient air.

2.3 The Rate of Change of Component Concentrations

The seven derivatives dc_i/dx are evaluated at a given x in S. RATE each time it is called by S. KUTTA. There are three possible contributions to each derivative which arise from (1) dilution, (2) deposition, and (3) oxidation,

$$\frac{dc(i)}{dx} = R_1(i) + R_2(i) + R_3(i) . \quad (23)$$

The terms due to dilution are of the form

$$R_1(i) = - [c(i) - A(i)] \frac{d \ln V_p}{dx} , \quad (24)$$

obtained by differentiation of eq 15.

For the terms arising from deposition two deposition velocities are specified V_d^g for gases and V_d^p for aerosol particles. For the components $S(VI)$ and H_2O only particulate deposition is involved

$$R_2(S(VI)) = - c(S(VI)) \cdot V_d^p \gamma / u_x , \quad (25)$$

$$R_2(H_2O) = - (g_w/0.018) \cdot V_d^p \gamma / u_x . \quad (26)$$

(Here g_w is the amount of condensed water in g/m^3 .) For the components S(IV) and NH_3 the expressions are of the form

$$R_2(i) = - \left\{ [P(i)/RT] V_d^g + [C(i) - P(i)/RT] V_d^p \right\} \gamma/u_x . \quad (27)$$

$P(i)$ is the pressure of the gaseous species. For the other components (CO_2 , O_3 , and OH) the deposition rates are assumed to be zero.

The rates of oxidation of S(IV) , calculated in and returned from S. RATOX are used in the following R_3 terms:

$$R_3(\text{S(IV)}) = - R_T/u_x , \quad (28)$$

$$R_3(\text{S(VI)}) = R_T/u_x , \quad (29)$$

$$R_3(\text{O}_3) = -R_{\text{OZ}}/u_x , \quad (30)$$

where R_T is the total oxidation rate and R_{OZ} is the oxidation rate from reaction of HSO_3^- with ozone in the droplets (both in $\text{mmol m}^{-3} \text{ sec}^{-1}$).

The depletion of OH by reaction with SO_2 is neglected since its concentration is controlled by other reactions not included in the model.

2.4 The Composition of the Plume

The remainder of the calculation is concerned mainly with determining for a given set of component concentrations the partial pressures of gases, the amount and composition of the aerosol droplets, and the amount--if any--of particulate ammonium sulfate in order that the oxidation rates of the various species of S(IV) (eqs 11, 12, and 14) can be calculated.

Given the concentrations of the components S(IV) , S(VI) , NH_3 , CO_2 , and H_2O , the following material balance equations may be written:

$$c(s(IV)) = P_{SO_2}/RT + ([H_2SO_3] + [HSO_3^-] + [SO_3^{2-}])g_w , \quad (31)$$

$$c(s(VI)) = ([HSO_4^-] + [SO_4^{2-}])g_w + c_s , \quad (32)$$

$$c(CO_2) = P_{CO_2}/RT + ([H_2CO_3] + [HCO_3^-] + [CO_3^{2-}])g_w , \quad (33)$$

$$c(NH_3) = P_{NH_3}/RT + ([NH_3] + [NH_4^+])g_w + 2c_s , \quad (34)$$

$$c(H_2O) = P_{H_2O}/RT + g_w/0.018 . \quad (35)$$

(The pressures P_i are in atm, all solution concentrations $[i]$ are in mmol/gH₂O--i.e., molal units--and c_s is the amount of $(NH_4)_2SO_4$ solid in mmol/m³.) Upon appropriate substitution of equilibrium constants (Table I) in these equations, the unknown quantities which remain are $[H^+]$, g_w , c_s , and the ionic activity coefficients.

Only four activity coefficients are distinguished (see Appendix C.1),

$$g_{H^+}, g_{NH_4^+}, g_{X^-}, g_{X^{2-}} ,$$

the activity coefficients of all singly charged anions being assumed to be the same and similarly for doubly charged anions. The activity coefficients, and the osmotic coefficient (ϕ) needed for the calculation of the activity of water, are rather complicated but not too sensitive functions of the corresponding ion concentrations $[H^+]$, $[NH_4^+]$, $[X^-]$, and $[X^{2-}]$ (Appendix C.1). They are refined each time S. EVAL is entered with trial values of $[H^+]$, g_w , and c_s by repeated cycling between S. EVAL and S. GAMAW.

The quantities $[H^+]$, g_w , and c_s (if >0) are refined by cycling between S. NONLIN and S. EVAL until the following conditions are met:

(1) Charge Balance

$$[\text{H}^+] + [\text{NH}_4^+] = [\text{X}^-] + 2[\text{X}^{2-}] \quad (36)$$

(2) Water Activity Balance

$$\frac{P_{\text{H}_2\text{O}}}{P_{\text{H}_2\text{O}}^{\circ}} = a_w \quad (37)$$

Here $P_{\text{H}_2\text{O}}^{\circ}$ is the saturation partial pressure of water in the atmosphere and a_w is calculated from the composition of the solution (i.e., from ϕ). This condition specifies that the vapor pressure of the droplets as defined by their composition and the temperature will equal the partial pressure of water in the atmosphere.

(3) Ammonium Sulfate Saturation (only if $C_s > 0$)

$$K_7 = [\text{NH}_4^+]^2 [\text{SO}_4^{2-}] g_{\text{NH}_4^+}^2 g_{\text{X}^{2-}}^2 \quad (38)$$

The residual deviations from these conditions are defined to equal or to approximate the fractional discrepancy in the concentration of ions in the droplets

$$y_1 = \left([\text{H}^+] + [\text{NH}_4^+] - [\text{X}^-] - 2[\text{X}^{2-}] \right) / \sum [i] \quad (39)$$

$$y_2 = \left(\frac{P_{\text{H}_2\text{O}}}{P_{\text{H}_2\text{O}}^{\circ}} - 1/a_w \right) / (0.018 \cdot \sum [i]) \quad (40)$$

$$y_3 = \left(\frac{[\text{NH}_4^+]^2 [\text{SO}_4^{2-}] g_{\text{NH}_4^+}^2 g_{\text{X}^{2-}}^2}{K_7} \right)^{1/3} - 1 \quad (41)$$

This appears to give an appropriate weight to each.

Initial estimates of g_w , $[\text{H}^+]$, $[\text{NH}_4^+]$, $[\text{X}^-]$ and $[\text{SO}_4^{2-}]$ are made the first time S. RATOX is entered as described in Appendix C.2. The initial estimate of C_s is made if and when the droplets become supersaturated with ammonium sulfate ($y_3 > 0$), otherwise $C_s = 0$. Subsequent estimates of these quantities are made in S. EXPCT before entry into

S. RATOX (Fig. 1).

The quantities $[H^+]$ and g_w are refined in S. NONLIN in logarithmic form since both can adopt widely ranging values and have either a first power or higher effect on the concentration of species in solution.

The quantity g_w can go to zero if the concentration ratio $[NH_4^+]/[SO_4^{2-}]$ reaches 2 and the solution is saturated with $(NH_4)_2SO_4$. If significant amounts of acid remain unreacted, however, the drops cannot dry out.* The value of g_w is monitored and if it becomes lower than 10^{-8} g/m^3 the computation of solution composition is discontinued and only the oxidation of SO_2 by OH (Eq. 11) is calculated.

The quantity C_s differs from $[H^+]$ and g_w in its effect on the calculations. When it is small it has a small effect on the composition of the droplets, but when it approaches half the amount of the NH_3 component present the difference $C(NH_3) - 2C_s$ becomes important in determining the amount of the droplets. Because of these considerations, the variable C_s is refined in S. NONLIN in the form

$$x_3 = \log[C(NH_3)/(C(NH_3) - 2C_s)] . \quad (42)$$

*While there is evidence that under these conditions an acid salt $[(NH_4)_2SO_4]_3 \cdot H_2SO_4$ can form, for simplicity we consider only the formation of $(NH_4)_2SO_4$.

3. TEST CALCULATIONS

A series of calculations, summarized in Table V, were carried out to determine the sensitivity of the various output quantities to what are thought to be typical ranges of the input quantities. The underscored values are reference values. Some of the output quantities for the reference case are plotted versus the distance from the source in Fig. 2. The other output quantities in the table result by varying each input quantity from its reference value while holding all other inputs at the reference values. For example, entries in the first row were generated using 2.5 wt. % sulfur in coal, stability class C, wind speed 2 m/sec, OH concentration $2 \cdot 10^{-8}$ mmol/m³, NH₃ concentration $6 \cdot 10^{-4}$ mmol/m³, and temperature 21.11°C.

In all calculations ERRR, the error limit parameter in S. KUTTA, was set at $3 \cdot 10^{-6}$. In test calculations this produced output values differing by < 0.2% from the limiting values obtained as ERRR was further decreased. In the calculation of the reference case: S. KUTTA called S. RATE 1345 times in performing the integration from X = 0.3 to 25 km; S. NONLIN called S. EVAL an average of 5.7 times per step of the integration; each time S. EVAL was called, it called S. GAMAW an average of 6.1 times. The computation time for the reference case on the IBM-360-91 computer was 16 sec.

Under the conditions tested, the aerosol droplets contain mostly sulfuric acid with a relatively small amount of ammonium sulfate, the latter being limited by and equivalent to the supply of ambient ammonia. The concentration of the droplets is strongly dependent on the ambient

Table V Sensitivity of Values Calculated by SULCAL to Variation of Input Values

Input Values ^{a,b}		Output Values at 25 km From Stack ^c					
		SO ₂ (10 ⁻² mmol/m ³)	Sulfate (10 ⁻³ mmol/m ³)	SO ₂ Ox. Rate (%/hr)	Liq Water (10 ⁻³ g/m ³)	H ₂ SO ₄ in Droplets (mmol/g H ₂ O)	(NH ₄) ₂ SO ₄ in Droplets (mmol/g H ₂ O)
Sulphur in Coal (Wt %)	2.5	14.3	19.3	3.2	7.7	2.5	0.039
	1.0	6.0	7.9	3.2	3.1	2.4	0.096
	0.5	3.2	4.0	3.2	1.6	2.4	0.189
Meteorological Stability Class	A	4.8	6.3	3.2	2.6	2.4	0.12
	C	6.0	7.9	3.2	3.1	2.4	0.096
	F	14.8	19.7	3.2	6.5	3.0	0.046
Wind Speed (m/sec)	1	9.8	25.3	3.2	9.0	2.8	0.033
	2	6.0	7.9	3.2	3.1	2.4	0.096
	4	3.5	2.5	3.2	1.0	2.2	0.29
	8	2.0	0.84	3.2	0.32	1.7	0.93
Ambient OH (10 ⁻⁸ mmol/m ³)	10	3.9	28.2	16.0	11.2	2.5	0.027
	2	6.0	7.9	3.2	3.1	2.4	0.096
	0.5	6.4	3.0	0.8	1.2	2.3	0.26
	0.1	6.6	1.6	0.16	0.60	2.1	0.50
Ambient NH ₃ (10 ⁻⁴ mmol/m ³)	20	6.0	7.9	3.2	3.0	2.3	0.33
	6	6.0	7.9	3.2	3.1	2.4	0.096
Ambient Temp. (°C)	18.33	6.0	7.9	3.2	740	0.010	0.00040
	19.44	6.0	7.9	3.2	14	0.54	0.021
	21.11	6.0	7.9	3.2	3.1	2.4	0.096

Table V (continued)

Input Values ^{a,b}	Output Values at 25 km From Stack ^c					
	SO ₂ (10 ⁻² mmol/m ³)	Sulfate (10 ⁻³ mmol/m ³)	SO ₂ Ox. Rate (%/hr)	Liq Water (10 ⁻³ g/m ³)	H ₂ SO ₄ in Droplets (mmol/g H ₂ O)	(NH ₄) ₂ SO ₄ in Droplets (mmol/g H ₂ O)
Ambient Temp. (°C)	23.89	6.0	7.9	3.2	1.6	4.7
	26.67	6.0	7.9	3.2	1.2	6.5
	29.44	6.0	7.9	3.2	0.95	7.9

^aUnderscored numbers are the reference values. Each alternate value was taken while all other input values at the reference values.

^bOther input values were: Wt % carbon in coal = 82.6; Wt % Hydrogen in coal 6.9%; Combustion rate = 100 Kg coal/sec; Effective stack height = 400 m; Deposition velocity = 0.01 m/sec; ambient water concentration = 930 mmol/m³.

^cConcentrations are averages across the plume (See Appendix B).

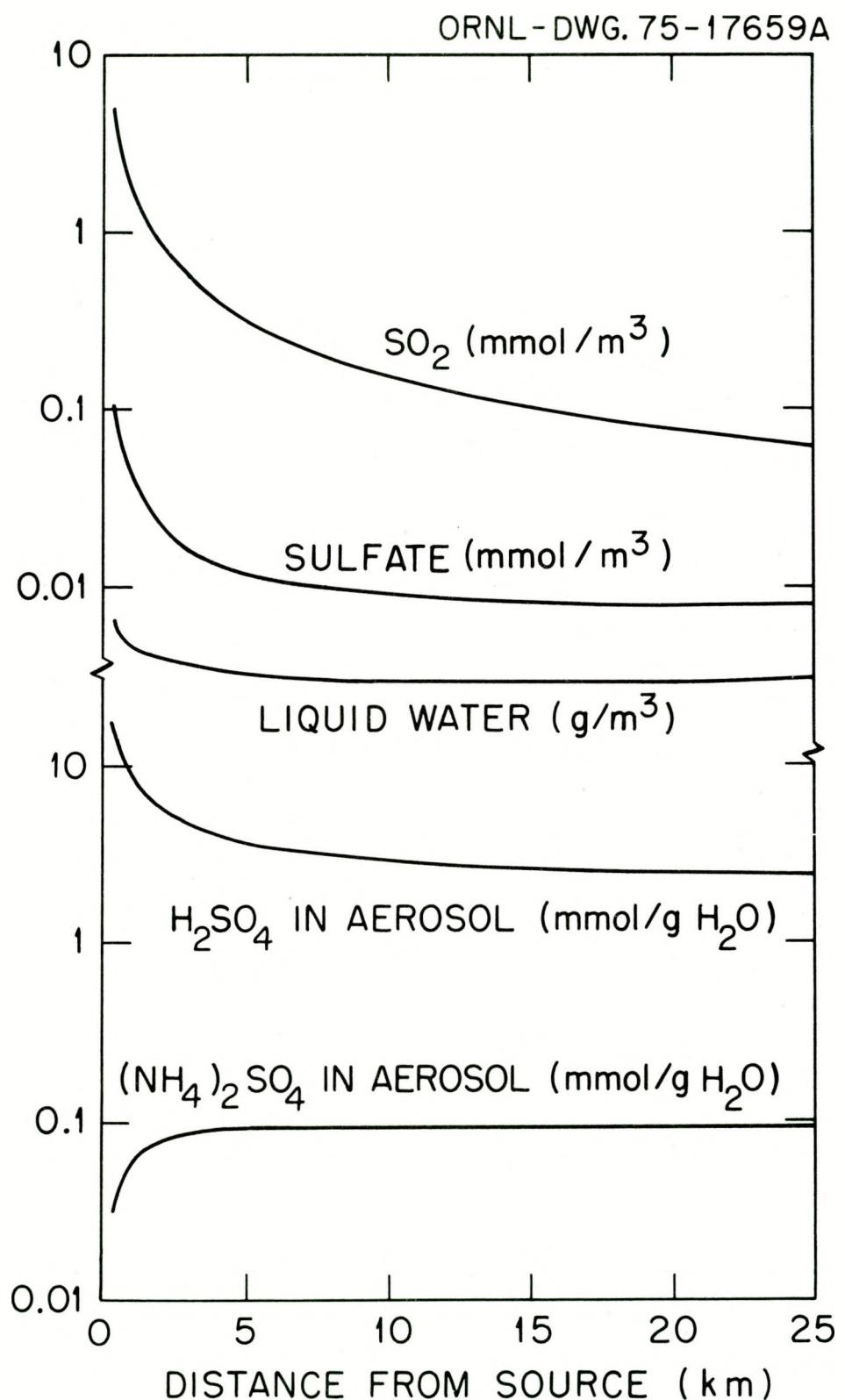


Fig. 2 Quantities calculated by SULCAL for the reference case.

temperature because of the strong inverse effect of the temperature on the relative humidity (the total concentration of water in the ambient air was not varied). The amounts of SO_2 , sulfuric acid and condensed water in the plume change directly as the sulfur emission rate, which here is varied by changing the sulfur content in the coal. Increased wind speed strongly decreases the concentrations of SO_2 , sulfuric acid and liquid water.

Under all the conditions tested in Table V, only the OH radical is important in the oxidation of SO_2 and the percentage of SO_2 oxidized per hour depends almost entirely on the ambient concentration of the OH radical. This concentration is controlled at the assumed values by unspecified chemical reactions. The conditions under which the oxidation of SO_2 species in the droplets (reactions 12 and 14) becomes significant may be determined as follows: The ratio of the rate of oxidation of HSO_3^- in solution by ozone (R_{OZ}) to the rate of oxidation of gaseous SO_2 by OH radical (R_{OH}) is given approximately by

$$\frac{R_{\text{OZ}}}{R_{\text{OH}}} = (K_1 K_2 K_{13} k_{12}/k_{11}) \cdot RT \cdot \left(\frac{g_w P_{\text{O}_3}}{[\text{H}^+] P_{\text{OH}}} \right). \quad (43)$$

The ratio of the rates in solution of SO_3^{2-} oxidation by oxygen R_{O} to HSO_3^- oxidation by ozone is given by

$$\frac{R_{\text{O}}}{R_{\text{OZ}}} = \frac{K_3}{K_{13}} \left(\frac{k_{14}' + k_{14}'' [\text{H}^+]}{k_{12}} \right) \left/ [\text{H}^+] P_{\text{O}_3} \right. \quad (44)$$

(Both these expressions are approximations because activity coefficients have been neglected.) With P_{O_3} and P_{OH} at their reference values ($8 \cdot 10^{-8}$ and $5 \cdot 10^{-13}$ atm, respectively) these ratios of rates at 25°C

become

$$\frac{R_{Oz}}{R_{OH}} \approx 3 \cdot 10^{-5} \frac{g_w}{[H^+]} \quad (45)$$

and

$$\frac{R_O}{R_{OH}} \approx \frac{1 \cdot 10^{-6}}{[H^+]} + \frac{5 \cdot 10^{-3}}{[H^+]^{1/2}} \quad (46)$$

Thus the oxidation of HSO_3^- by ozone should become significant when the ratio $g_w/[H^+]$ exceeds $\sim 10^4$. The rate of SO_3^{2-} oxidation by dissolved oxygen should then become significant when the pH exceeds ~ 4 . Since g_w and $[H^+]$ change in opposite directions and almost logarithmically with the relative humidity, the rate of oxidation of SO_2 species in aerosol droplets increase rapidly in importance as the relative humidity increases (Fig. 3). This rate is also favored by conditions that produce a more complete neutralization of the acid droplets. In general, neutralization is favored by a lower sulfur emission rate and a lower OH radical concentration, both tending to reduce the rate of production of the acid SO_3 . It is favored as well by a high ambient concentration of ammonia and a high wind speed, both tending to deliver more ammonia to the acid droplets.

We emphasize that, in addition to the approximations in the present model arising from the use of averaged concentrations in the plume, the oxidation mechanisms are oversimplified and incomplete. The concentration of OH radical, which may often determine the rate of the most rapid oxidation reaction, is controlled by other chemical reactions not included in the model. The species of SO_2 in solution oxidized by ozone may include others in addition to (or perhaps instead of) HSO_3^- .

ORNL-DWG. 75-17658A

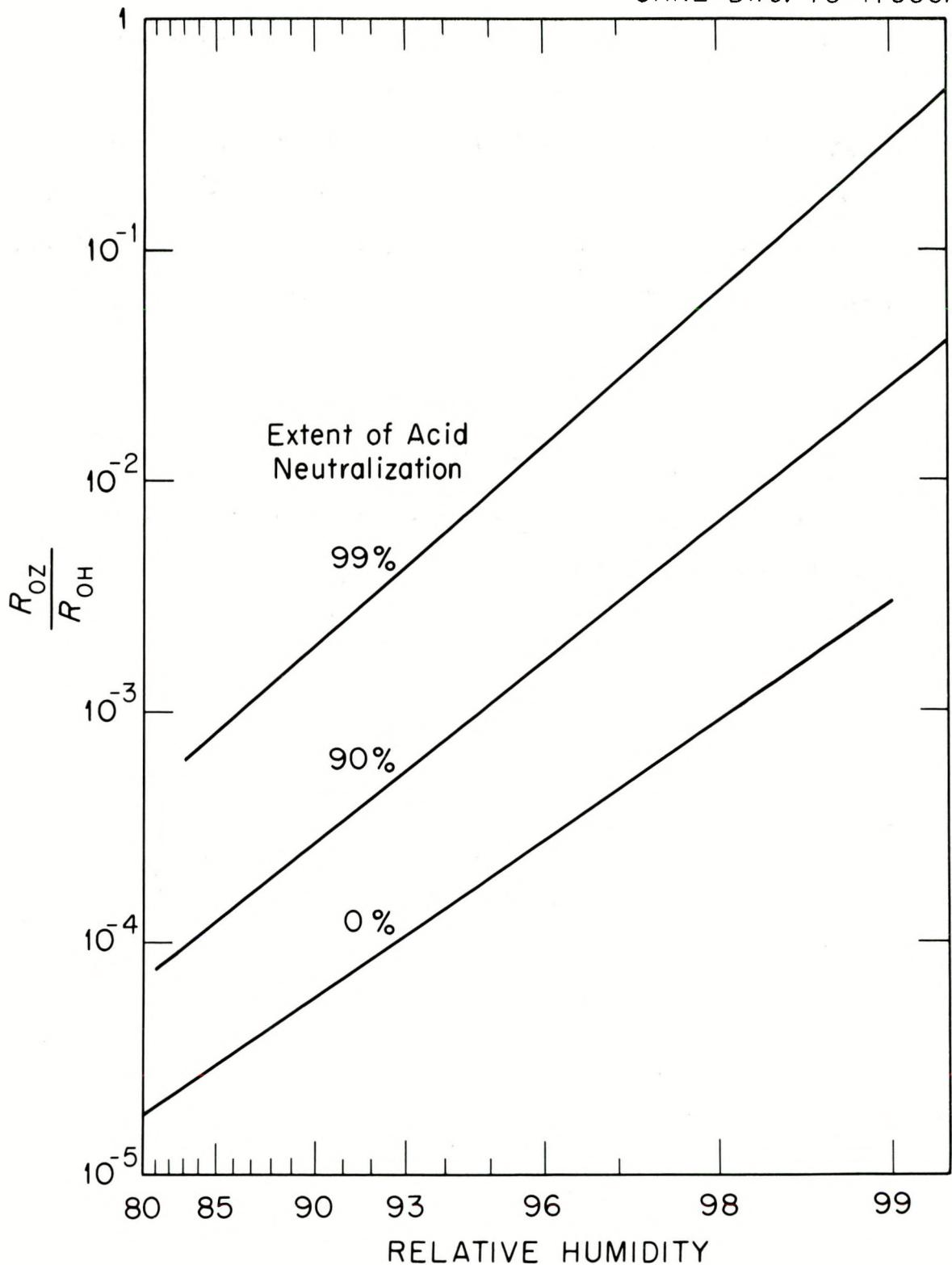


Fig. 3 The effect of the relative humidity and the extent of acid neutralization on the ratio of the rates of SO_2 oxidation by ozone (R_{OZ}) and by hydroxyl radical (R_{OH}) at $25^\circ C$.

The oxidation of SO_3^{2-} by oxygen can be catalyzed or poisoned by certain metal ions in solution. Finally, possible catalytic oxidation of SO_2 at the surface of solid particles and droplets has not been included. It does not seem worthwhile to include all such effects at present, nor may it be so even when sufficient knowledge has been gained. As more information does become available, however, it will be relatively easy to modify the rate expressions in the S. RATOX and to add others as may be desired.

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APPENDIX A

Program SULCAL

1. Partial List of Symbols

<u>In Text</u>	<u>Definition</u>	<u>In Program</u>
a_w	Activity of water	AWI
B^o_{JK} , B^1_{JK}	Parameters for calculating interaction coefficients between cation J and anion K $B^o_{JK} = B(1, J, K)$ $B^1_{JK} = B(2, J, K)$	$B(I, J, K)$
β	Correction factors (Eq. 17) for reflection of plume from ground, for short-time (BETAS) and time-averaged (BETAL) plumes	BETAS, BETAL
C_i or $C(i)$	Total concentration of component; (mmol/m ³)	QC(I)
C_s	Concentration of $(\text{NH}_4)_2\text{SO}_4$ particles in plume (mmol/m ³)	CSS
ERRR	Error limit employed in S. KUTTA	ERRR
f	Factor relating short-time and long-time sigmas (Eq. 22, App. B)	FLS
g_i	Activity coefficient of an ion; $G(1,1) = g_{\text{H}^+}$, $G(1,2) = g_{\text{NH}_4^+}$, $G(2,1) = g_{\text{x}^-}$, $G(2,2) = g_{\text{x}^{2-}}$	$G(I, J)$
g_w	Concentration of condensed water in plume (g H ₂ O/m ³)	CGW
γ	Factor (Eq. 20) which corrects effective plume concentration to ground level concentrations	GAMMA
P_j	Pressure of a gaseous constituent (atm)	$P(J)$
P_w^o	Vapor pressure of pure water (atm)	PWO
ϕ	Osmotic coefficient of aerosol droplets (Eq. 4, App. C)	P
Q_i	Emission rate of component; (mmol/sec)	$Q(I)$
R	The gas constant (m ³ mmol ⁻¹ °K ⁻¹)	R

<u>In Text</u>	<u>Definition</u>	<u>In Program</u>
$\frac{1}{RT}$	Factor for conversion of pressure (atm) to concentration (mmol/m ³)	CONST
$\sigma_x^*, \sigma_y^*, \sigma_z^*$	Time-averaged dispersion coefficients for a point source	SGPY, SGPZ
σ^o	Correction which makes σ_y and σ_z finite at $x = 0$ (Eq. 18)	SGO
$\sigma_{\ell,y}, \sigma_{\ell,z}$	Time-averaged dispersion coefficient (m)	SGLY, SGLZ
$\sigma_{s,y}, \sigma_{s,z}$	Short-time dispersion coefficient (m)	SGSY, SGSZ
T	Absolute temperature of plume (°K)	T
u_x	Wind velocity (m/sec)	UX
v_d^g, v_d^p	Disposition velocity for gases and particles (m/sec)	VDG, VDP
v_p^*	Effective volume (m ³ /sec) (the effective volume of a length of plume equal to u_x , the wind velocity in m/sec). VRG is calculated using short-time dispersion coefficients and VLG is calculated with time-averaged dispersion coefficients	VRG, VLG
v_p^o	Gas flow rate at mouth of stack (m ³ /sec)	VRGO
x	Distance (m) from source in wind direction	X

2. Program Listing

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C PROGRAM FOR CALCULATION OF CHEMICAL FORMS OF SULFUR IN AND THEIR          BLCK
C DEPOSITION FROM A PLUME                                         BLCK
C BLOCK DATA                                         BLCK
C IMPLICIT REAL*8 (A-H,O-Z)                                         BLCK
C R = GAS CONSTANT (M**3/ATM/K,MMOL)                                         BLCK
C PI = 3.1416, PLN = LN(10), FMW = GMS H2O/MMOL                                         BLCK
C B(I,J,K) = INTERACTION COEF'S. OF CATION I WITH ANION J, TWO          BLCK
C PARAMETERS FOR EACH IJ PAIR                                         BLCK
C Z = ION CHARGE (FIRST SUBSCRIPT IS 1 FOR CATIONNS AND 2 FOR ANIONS)          BLCK
COMMON/BLK/R,PI,PLN,FMW,B(4,4,2),Z(2,4)                                         BLCK
      DATA B(1,1,1),B(1,1,2),B(1,2,1),B(1,2,2),B(2,1,1),          BLCK
      B(2,1,2),B(2,2,1),B(2,2,2)                                         BLCK
      8/0.2722D+0,3.0803D+0,0.0502D+0,-4.0054D+0,0.0D+0,C.319D+0,          BLCK
      6-0.04873D+0,-1.18580D+0/                                         BLCK
      DATA PI,PLN,R,FMW/3.1415927D+0,2.302585D+0,82.0597D-9,18.0153D-3/          BLCK
      8,Z(1,1),Z(1,2),Z(2,1),Z(2,2)/1.0D+0,1.0D+0,-1.0D+0,-2.0D+0/          BLCK
      END                                         BLCK

C
C MAIN PROGRAM                                         MAIN
C IMPLICIT REAL*8 (A-H,O-Z)                                         MAIN
C EXTERNAL RATE                                         MAIN
COMMON/BLK/R,PI,PLN,FMW,B(4,4,2),Z(2,4)                                         MAIN
C NOX = TOTAL NUMBER OF TIMES S. RATE IS ENTERED FOR A GIVEN CASE          MAIN
COMMON/CASE/NOX                                         MAIN
COMMON/RTE/NSC,NRL,VRGO,TP,H,UX,A(7),TA,DELX,VDG,VDP                                         MAIN
COMMON/SIGMA/SG0,SGPY,SGPZ,SGLY,SGLZ,SGSY,SGSZ,BETAL,BETAS          MAIN
COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,RTO,NCAL,NITN          MAIN
COMMON/EQK/EKN,EK1N,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW          MAIN
COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,          MAIN
  &PN,PC,PS                                         MAIN
COMMON/ACT/DHA,G(2,4),P,PW0,CONST                                         MAIN
COMMON/GIVN/WPCTC,WPCTH,WPCTS,QCOAL,FCTS6,EXAIR,FCON,TCP,TCA,          MAIN
  &XIAST,XNEXT,ERRR                                         MAIN
COMMON/VAR/QC(7)                                         MAIN
  DIMENSION Q(7),FDMY(7),XINCR(50)                                         MAIN
C WPCTI=WT. % OF ELEMENT I IN COAL                                         MAIN
C QCOAL=COMBUSTION RATE IN KG/SEC                                         MAIN
C FCTS6=FRACT OF S CONVERTED TO S(VI),EXAIR=FRACTION OF EXCESS AIR USED          MAIN
C FCON=NO.OF MOLES OF WATER CONDENSED/MOLE SO3, TCP=TEMP OF EXIT GAS (C)          MAIN
C H=APPARENT STACK HEIGHT IN METERS                                         MAIN
  1 READ(20,100) WPCTC,WPCTH,QCOAL,FCTS6,EXAIR,FCON,TCE,H          MAIN
C STABILITY CLASS,ROUGHNESS LENGTH,WIND SPEED (M/SEC)                                         MAIN
C AND AMBIENT CONCENTRATIONS (MMOL/M**3) AND TEMPERATURE (C)          MAIN
  2 READ(20,100) NSC,NRL,UX,VDG,VDP,(A(I),I=1,7),TCA,WPCTS          MAIN
C CALCULATE MOLES OF GASES PER KG OF COAL BURNED          MAIN
  RMLC=WPCTC/1.200115                                         MAIN
  RMLW=0.5*WPCTH/0.100797                                         MAIN
  RMLS4=(WPCTS/3.2064)*(1.0-FCTS6)                                         MAIN
  RMLS6=(WPCTS/3.2064)*FCTS6                                         MAIN
C CALCULATE MOLES OF OXYGEN CONSUMED PER KG OF COAL BURNED          MAIN
  CMLO2=RMLC+0.5*RMLW+RMLS4+1.5*RMLS6                                         MAIN
C CALCULATE TOTAL MOLES OF EFFLUENT GAS PER KG OF COAL EURNED          MAIN
  RMLGAS=RMLC+RMLW+RMLS4+((79.05/20.95)*(1.0+EXAIR)+EXAIR)*CMLO2          MAIN
  &-FCON*RMLS6                                         MAIN
  TP=TCP+273.15                                         MAIN
  TA=TCA+273.15                                         MAIN
C CALCULATE RATE OF EMISSION OF GAS (M**3/SEC)                                         MAIN
  VRGO=RMLGAS*QCOAL*(82.0597D-6)*TP                                         MAIN

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C CALCULATE EMISSION RATES OF COMPONENTS IN MMOLS/SEC          MAIN
BNFCT=1000.0*QCOAL                                         MAIN
Q(1)=RMLS4*BNFCT                                         MAIN
Q(2)=RMLS6*BNFCT                                         MAIN
Q(3)=0.0                                         MAIN
Q(4)=0.0                                         MAIN
Q(5)=RMLC*BNFCT                                         MAIN
Q(6)=0.0                                         MAIN
Q(7)=RMLW*BNFCT                                         MAIN
NOX=1                                         MAIN

C READ FIRST VALUE OF X (M), DELTA X AND ERROR LIMIT          MAIN
READ(20,100) XNEXT,DELX,ERRR                           MAIN
C READ NO. OF STEPS IN X AND THE SIZE OF EACH             MAIN
READ(20,100) NX,(XINCR(I),I=1,NX)                      MAIN
NCAL=0                                         MAIN
NITN=0                                         MAIN

C CALCULATE CONCENTRATIONS AT STARTING VALUE OF X (MMOLS/M**3) MAIN
CALL SIG(XNEXT,VRG,PEX)                                 MAIN
DO 10 I=1,7                                         MAIN
QC(I)=Q(I)/VRG+A(I)*(1.0-VRGO/VRG)                  MAIN
10 CONTINUE                                         MAIN
CALL PRINT(1)                                         MAIN

C START CALCULATION OF QC(I) VS. X                      MAIN
DO 20 I=1,NX                                         MAIN
XLAST=XNEXT                                         MAIN
XNEXT=XNEXT+XINCR(I)                                MAIN
CALL KUTTA(XLAST,XNEXT,QC,7,DELX,ERRR,40,RATE)      MAIN
20 CONTINUF                                         MAIN

C CALL RATE FOR FINAL CALCULATION OF PLUME COMPOSITION      MAIN
CALL RATE(XNEXT,OC,FDMY,1)                           MAIN
XLAST=XNEXT                                         MAIN
CALL PRINT(2)                                         MAIN
GO TO 2                                         MAIN

100 FORMAT(7G10.0)                                     MAIN
END                                         MAIN

C
C SUBROUTINE TO CALCULATE EFFECTIVE VOLUME OF PLUME (VRG)      SIG
C AND THE DEPOSITION FACTOR (PEXD)                                SIG
C CALCULATION OF SIGMA(Y) AND SIGMA(Z) ACCORDING TO HOSKER      SIG
SUBROUTINE SIG(X,VRG,PEXD)                                SIG
IMPLICIT REAL*8(A-H,O-Z)                                SIG
DIMENSION GC3(6),GA1(6),GB1(6),GA2(6),GB2(6),GC1(6),GD1(6),FCTL(6) SIG
COMMON/NOX/
COMMON/RTE/NSC,NRL,VRGO,TP,H,UX,A(7),TA,DELX,VDG,VDP SIG
COMMON/SIGMA/SG0,SGPY,SGPZ,SGLY,SGLZ,SGSY,SGSZ,BETAL,BETAS SIG
DATA GC3/0.22D+0,0.16D+0,0.11D+0,0.08D+0,0.06D+0, SIG
8.04D+0/                                                SIG
DATA GA1/0.112D+0,0.130D+0,0.112D+0,0.098D+0, SIG
8.0609D+0,0.0638D+0/                                 SIG
DATA GB1/1.06D+0,0.950D+0,0.920D+0,0.889D+0,0.895D+0, SIG
8.0783D+0/                                              SIG
DATA GA2/0.000538D+0,0.000652D+0,0.000905D+0, SIG
8.00135D+0,0.00196D+0,0.00136D+0/                   SIG
DATA GB2/0.815D+0,0.750D+0,0.718D+0,0.688E+0,0.684D+0, SIG
8.0672D+0/                                              SIG
DATA GC1/1.56D+0,2.02D+0,2.71828D+0,5.16D+0,7.37D+0, SIG
8.117D+0/                                              SIG
DATA GD1/0.048CD+0,0.0269D+0,0.0D+0,-0.06E+0,-0.0957D+0 SIG

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E,-0.128D+0/ SIG
  DATA FCTL/5.0D+2,2.5D+2,1.0D+2,0.5D+2,0.2D+2,0.1D+2/ SIG
C UPON FIRST ENTRY ASSIGN VALUES TO PARAMETERS THAT DEPEND ON STABILITY SIG
C   CLASS, ROUGHNESS LENGTH, AND WIND SPEED SIG
  IF(NOX.GT.1) GO TO 1 SIG
  PI=3.1415927 SIG
  C3=GC 3(NSC) SIG
  A1=GA 1(NSC) SIG
  B1=GB 1(NSC) SIG
  A2=GA 2(NSC) SIG
  B2=GB 2(NSC) SIG
  C1=GC 1(NRL) SIG
  D1=GD 1(NRL) SIG
  R1=FCTL(NSC) SIG
  FLS=DSQRT(X/(X+50.0*RI)) SIG
  SGO=DSQRT(VRG0/(4.0*PI*UX)) SIG
C CALCULATE POINT-SOURCE TIME-AVERAGED SIGMA(Y) SIG
  1 DTY=1.0+0.001*X SIG
  SGPY=C3*X/DSQRT(DTY) SIG
  F=DLOG(C1*X**D1) SIG
  G=A1*X**B1/(1.0+A2*X**B2) SIG
C CALCULATE POINT-SOURCE TIME-AVERAGED SIGMA(Z) SIG
  SGPZ=F*G SIG
C CALCULATE TIME-AVERAGED SIGMA(Y) AND SIGMA(Z) FOR REAL SOURCE SIG
  SGLY=SGPY+SGO SIG
  SGLZ=SGPZ+SGO SIG
C CALCULATE SHORT-TIME SIGMA(Y) AND SIGMA(Z) FOR REAL SOURCE SIG
  SGSY=FLS*SGPY+SGO SIG
  SGSZ=FLS*SGPZ+SGO SIG
C EVALUATE CORRECTION FACTOR BETA SIG
  EXFS=-(H/SGSZ)**2 SIG
  IF(EXFS.LT.-25.0) EXFS=-25.0 SIG
  EXFL=-(H/SGLZ)**2 SIG
  IF(EXFL.LT.-25.0) EXFL=-25.0 SIG
  BETAL=1.0/(1.0+0.70711*DEXP(EXFL)) SIG
  BETAS=1.0/(1.0+0.70711*DEXP(EXFS)) SIG
C CALCULATE EFFECTIVE REACTION VOLUME VRG (M**3) SIG
C   AND EFFECTIVE DEPOSITION VOLUME VLG (M**3) SIG
  VRG=4.0*PI*SGSY*SGSZ*UX*BETAS SIG
  VLG=4.0*PI*SGLY*SGLZ*UX*BETAL SIG
C CALCULATE FACTOR RELATED TO DRY DEPOSITION SIG
  GAMMA=DSQRT(2.0/PI)/SGLZ*DEXP(EXPL/2.0) SIG
  FFD=(VRG/VLG)*GAMMA SIG
  RETURN SIG
  END SIG

C
C SUBROUTINE CALLED BY KUTTA TO CALCULATE RATE OF CHANGE OF COMPONENT RATE
C   CONCENTRATIONS YI(I) WITH DISTANCE X (DYDX) RATE
C IN = INTEGER WHICH (WHEN 1) SIGNALS CALCULATION OF DYEX AT X RATE
  SUBROUTINE RATE(X,YI,DYDX,IN) RATE
  IMPLICIT REAL*8(A-H,O-Z) RATE
  DIMENSION YI(7),DYDX(7) RATE
  COMMON/BLK/R,PI,PLN,FMW,B(4,4,2),Z(2,4) RATE
  COMMON/CASE/NOX RATE
  COMMON/RTE/NSC,NRL,VRG0,TP,H,UX,A(7),TA,DELX,VDG,VDP RATE
  COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,RTO,NCAL,NITN RATE
  COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S, RATE
  EPN,PC,PS RATE

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COMMON/ACT/DHA,G(2,4),P,PW0,CONST          RATE
COMMON/VAR/Y(7)                           RATE
DO 10 I=1,7                           RATE
Y(I)=YT(I)                           RATE
10 CONTINUE                           RATE
C HAVE DROPLETS DRIED UP?          RATE
IF(CGW.LE.0.0) GO TO 2          RATE
C ESTIMATE STARTING VALUPS OF CGW,PH,CSS, AND ACTIVITY COEFFICIENTS      RATE
CALL EXPCT(NOX,IN,X,CGW,PH,CSS,C(1,2),C(2,1),C(2,2))          RATE
C EVALUATE DLN(V)/D(X) NUMERICALLY          RATE
2 X1=X+DELX/2.0          RATE
X2=X-DFLX/2.0          RATE
CALL SIG(X,VRG,FEXD)          RATE
CALL SIG(X1,VRG1,FEX)          RATE
CALL SIG(X2,VRG2,FEX)          RATE
DLVDX=2.0*(VRG1-VRG2)/(DELX*(VRG1+VRG2))          RATE
C CALCULATE TEMPERATURE IN PLUME (K)          RATE
T=TA/(1.0-(1.0-TA/TP)*VRG0/VRG)          RATE
CONST=1.0/(R*T)          RATE
C CALCULATE RATE OF OXIDATION OF S(IV)          RATE
CALL RATOX          RATE
NOX=NOX+1          RATE
C CALCULATE DYDX(I)          RATE
DYDX(1)=-(RTT+(PS*CONST*(VDG-VDP)+Y(1)*VDP)*FEXD)/UX          RATE
DYDX(2)=(RTT-Y(2)*VDP*FEXD)/UX          RATE
DYDX(3)=-RTOZ/UX          RATE
DYDX(4)=0.0          RATE
DYDX(5)=0.0          RATE
DYDX(6)=-(PN*CCNST*(VDG-VDP)+Y(6)*VDP)*FEXD/UX          RATE
DYDX(7)=-CGW/FMW*VDP*FEXD/UX          RATE
DO 1 I=1,7          RATE
DYDX(I)=DYDX(I)-(Y(I)-A(I))*DLVDX          RATE
1 CONTINUE          RATE
RETURN          RATE
END          RATE
C
C SUBROUTINE TO CALCULATE RATE OF S(IV) OXIDATION          RTOX
SUBROUTINE RATOX          RTOX
IMPLICIT REAL*8(A-H,O-Z)          RTOX
EXTERNAL IVAL          RTOX
EXTERNAL EVAL          RTOX
DIMENSION X(10)          RTOX
COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)          RTOX
COMMON/CASE/NOX          RTOX
COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,PTO,NCAL,NITN          RTOX
COMMON/FQK/FKN,EK1N,EK2C,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW          RTOX
COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,          RTOX
SPN,PC,PS          RTOX
COMMON/ACT/DHA,G(2,4),P,PW0,CONST          RTOX
COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW          RTOX
C CONSTANTS FOR THE CALCULATION OF RATE AND EQUILIBRIUM CONSTANTS          RTOX
C AS A FUNCTION OF THE TEMPERATURE          RTDX
DATA EKNO,EKNT,EKCO,EKCT,EKSO,EKST,EK1NO,EK1NT,          RTOX
EK1SO,EK1ST,EK2SO,EK2ST,EK1C0,EK1CT,EK2C0,EK2CT,EK3SO,          RTOX
EK3ST,EKSS0,EKSST,EKWO,EKWT,EKOZO,EKOZT,RK10,RK1T,RK20,RK2T,          RTOX
ERK30,RK3T,RK40,RK4T,EPW0,EPWT          RTOX
8/-4.233D+0,1.7855D+0,-5.016D+0,1.05974D+0,-4.4912D+0,1.3659D+0,          RTOX
80.1026D+0,2.7275D+0,-4.815D+0,0.9094D+0,-8.8549D+0,0.48736D+0,          RTOX

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8-5.030D+0,-0.39885D+0,-7.7260D+0,-0.77585D+0,-5.8348D+0,1.1452D+0, RTOX
80.929D+0,-0.3431D+0,-4.2134D+0,-2.91654D+0,-8.488D+0,1.891D+0, RTOX
811.4378D+0,-3.9725D+0,15.0947D+0,-3.9725D+0,14.61D+0,-2.57D+0, RTOX
82.65D+0,0.0D+0,6.2057D+0,-2.2989D+0/ RTOX

C ALL CONCENTRATIONS ARE IN MMOLES/METER**3 OR MMOLS/G RTOX
C CALCULATE TEMPERATURE DEPENDENT PARAMETERS RTOX
  FT=1000.0/T RTOX
  DHA=(4.91D-1+8.0D-4*(T-273.15))*FLN RTOX
  EKN=DEXP((EKNO+EKN*T*FT)*FLN) RTOX
  EKC=DEXP((EKCO+EKCT*T*FT)*FLN) RTOX
  FKS=DEXP((EKS0+EKS*T*FT)*FLN) RTCX
  EK1N=DEXP((EK1NO+EK1NT*T*FT)*FLN) RTOX
  EK1S=DEXP((EK1SO+EK1ST*T*FT)*FLN) RTOX
  EK2S=DEXP((EK2SO+EK2ST*T*FT)*FLN) RTCX
  EK1C=DEXP((EK1CO+EK1CT*T*FT)*FLN) RTOX
  EK2C=DEXP((EK2CO+EK2CT*T*FT)*FLN) RTOX
  EK3S=DEXP((EK3SO+EK3ST*T*FT)*FLN) RTCX
  EKSS=DEXP((EKSS0+EKSST*T*FT)*FLN) RTOX
  FKW=DEXP((EKWO+EKWT*T*FT)*FLN) RTOX
  EKOZ=DEXP((EKOZ0+EKOZT*T*FT)*FLN) RTCX
  RK1=DEXP((RK10+RK1T*T*FT)*FLN) RTOX
  RK2=DEXP((RK20+RK2T*T*FT)*FLN) RTOX
  RK3=DEXP((RK30+RK3T*T*FT)*FLN) RTCX
  RK4=DEXP((RK40+RK4T*T*FT)*FLN) RTOX
  PWO=DEXP((EPWO+EPWT*T*FT)*FLN) RTOX

C TFST TO SEE IF CGW AND PH ARE GIVEN RTOX
  IF(NOX.GT.1) GO TO 1 RTOX
  CSS=0.0 RTOX

C ESTIMATE STARTING VALUES OF CGW, ION CON'NS AND GAMAS RTOX
  PNI=3.0 RTOX
  TB=(PWO*CONST-CW+PNI*C6)/2.0 RTOX
  TC=-PNI*C6*CW RTOX
  CGW=PMW*(-TB+DSQRT(TB**2-TC)) RTOX
  X(1)=DLOG10(CGW) RTOX
  G(1,1)=1.0 RTOX
  G(2,1)=1.0 RTOX
  G(2,2)=1.0 RTCX
  CALL NONLIN(X,1,ICONV,IVAL) RTOX
  CGW=DEXP(X(1)*FLN) RTOX
  PH=-DLOG10(C(1,1)) RTOX

C BEGIN REFINEMENT RTOX
C HAVE DROPLETS DRIED OUT? RTOX
  1 IF(CGW.LE.0.0) GO TO 10 RTOX
    X(1)=DLOG10(CGW) RTOX
    X(2)=PH RTOX

C ARE DROPLETS UNSATURATED? RTOX
  IF(CSS.LE.0.0) GO TO 2 RTOX

C IS ENOUGH ACID PRESENT TO PREVENT DRYING OUT OF DROPLETS? RTOX
  IF(CN/C6.LT.2.0) GO TO 4 RTOX
  10 CGW=0.0 RTOX
    CSS=C6 RTOX
    RTO=0.0 RTOX
    RTOZ=0.0 RTOX
    PS=C4/CONST RTOX
    PN=(CN-2.0*CSS)/CONST RTOX
    PC=CC/CONST RTOX
    GO TO 11 RTCX
  2 CSS=0.0 RTOX

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      CALL NONLIN(X,2,ICONV,EVAL)          RTOX
C IS SOLUTION SUPERSATURATED WITH (NH4)2SO4?    RTCX
      GFS=G(1,2)**2*G(2,2)          RTOX
      DSS=CM1N**2*CM4S*GFS-FKSS      RTOX
      IF(DSS.LT.0.0) GO TO 6          RTCX
C IS ENOUGH ACID PRESENT TO PREVENT DRYING OUT OF DROPLETS?    RTOX
      IF((CN/C6).LT.2.0.AND.CGW.GT.1.0D-8) GO TO 3      RTOX
      GO TO 10                         RTCX
C ESTIMATE AMOUNT OF SOLID (NH4)2SO4      RTOX
      3 CSS=DSS*CGW/((CM1N+4.0*CM4S)*CM1N*GFS)      RTOX
      IF((CN-2.0*CSS).GT.0.0) GO TO 4      RTCX
      CSS=0.4*CN                         RTOX
      4 X(3)=DLOG10(CN/(CN-2.0*CSS))      RTOX
      CALL NONLIN(X,3,ICONV,EVAL)          RTOX
C HAS SOLUTION BECOME UNSATURATED WITH (NH4)2SO4?    RTOX
      IF(X(3).LE.0.0) GO TO 2          RTOX
C HAVE DROPLETS DRIED UP?          RTOX
      IF(CGW.LE.1.0D-8) GO TO 10      RTOX
      CSS=CN/2.0*(1.0-DEXP(-X(3)*FLN))      RTOX
      5 CGW=DEXP(X(1)*FLN)          RTOX
      PH=X(2)                         RTOX
      CMH=DEXP(-PH*FLN)          RTOX
C BEGIN CALCULATION OF RATE OF OXIDATION OF S(IV)      RTOX
C RATE OF OXIDATION OF SO32- BY OXYGEN IN SOLUTION (MMOL/M**3/SEC)    RTCX
      RQ1=RK1*G(2,2)          RTOX
      RQ2=RK2*G(1,1)**0.5*G(2,2)      RTOX
      RTO=(RQ1+RQ2*CMH**0.5)*CM2S*CGW      RTCX
C RATE OF OXIDATION OF HS03- BY OZONE IN SOLUTION (MMOL/M**3/SEC)    RTOX
      RQ3=RK3*G(2,1)          RTOX
      CM0Z=EK0Z*COZ/(CONST+EK0Z*CGW)      RTOX
      RTOZ=RQ3*CM0Z*CM1S*CGW      RTOX
C RATE OF OXIDATION OF SO2 BY OH RADICAL IN GAS PHASE (MMOL/M**3/SEC)    RTOX
      11 RTOH=RK4*COH*PS*CONST      RTOX
C TOTAL OXIDATION RATE OF S(IV) (MMOL/METER**3/SEC)      RTOX
      RTT=RTOH+RTOZ+RTO          RTOX
      RETURN                         RTOX
      END                           RTCX
C
C SUBROUTINE TO REFINE CGW, PH, CSS BY SECANT METHOD      NLIN
      SUBROUTINE NONLIN(XIN,N,ICONV,EVAL)      NLIN
C PROGRAM AUTHORS R. E. FUNDERLIC AND J. RINZEL,      NLIN
C COMPUTING TECHNOLOGY CENTER, UNION CARBIDE CORP., NUCLEAR DIV.,      NLIN
C OAK RIDGE, TENN.      NLIN
C
C CTC ORD PROGRAM NO. 9066      NLIN
      IMPLICIT REAL*8 (A-H,O-Z)      NLIN
      LOGICAL WRT      NLIN
      COMMON/HONON/WRT      NLIN
      DIMENSION XIN(N),X(20,21),XBAR(20,2),A(21,21),AINV(21,21),P(21),      NLIN
      1B(21),XNORM(21),BNORM(20)      NLIN
      WRT=.FALSE.      NLIN
      ICONV=0      NLIN
      NP1=N+1      NLIN
      DO 110 J=1,NP1      NLIN
      DO 120 I=1,N      NLIN
      IF(J.NE.1) GO TO 130      NLIN
      X(I,J)=XIN(I)      NLIN
      GO TO 120      NLIN

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130 X(I,J)=XTN(I) NLIN
  IF (J-1.EQ.I) X(I,J)=0.005+X(I,J)
120 CONTINUE NLIN
110 CONTINUE NLIN
  DO 1 J=1,NP1 NLIN
  B(J)=1. NLIN
  A(NP1,J)=1. NLIN
1 CALL EVAL(X(1,J),A(1,J)) NLIN
  DO 2 I=1,NP1 NLIN
  DO 2 J=1,NP1 NLIN
  AINV(I,J)=0.0D0 NLIN
  IF (I.EQ.J) AINV(I,J)=1.0D0 NLIN
2 CONTINUE NLIN
  DO 4 J=1,NP1 NLIN
  SUM=0. NLIN
  DO 3 I=1,N NLIN
  SUM=SUM+DABS(A(I,J)) NLIN
4 XNORM(J)=SUM NLIN
  CALL MATINV(A,AINV,NP1) NLIN
  DO 5 I=1,N NLIN
5 XBAR(I,1)=X(I,NP1) NLIN
  ITER=40 NLIN
  DO 15 K=1,ITER NLIN
  XBNUM=0. NLIN
  XBDEN=0. NLIN
  DO 7 I=1,N NLIN
  XB=0. NLIN
  DO 6 J=1,NP1 NLIN
6 XB=XB + AINV(J,NP1)*X(I,J) NLIN
  XBAR(I,2)=XB NLIN
  XBNUM=XBNUM + (XBAR(I,2)-XBAR(I,1))**2 NLIN
7 XBDEN=XBDEN + XBAR(I,2)**2 NLIN
  ICONV=ICONV+1 NLIN
  IF ((XBNUM/XBDEN).LT.1.D-14) GO TO 16 NLIN
  DO 8 I=1,N NLIN
8 XBAR(I,1)=XBAR(I,2) NLIN
  BIGNOR=0. NLIN
  DO 9 J=1,NP1 NLIN
  BNORM(J)=XNORM(J)*DABS(B(J)) NLIN
  IF (BNORM(J).LE.BIGNOR) GO TO 9 NLIN
  BIGNOR=BNORM(J) NLIN
  JCOL=J NLIN
9 CONTINUE NLIN
  CALL EVAL(XBAR,P) NLIN
  P(NP1)=1. NLIN
  XNORM(JCOL)=0. NLIN
  DO 10 I=1,N NLIN
  X(I,JCOL)=XBAR(I,1) NLIN
10 XNORM(JCOL)=XNORM(JCOL)+DABS(P(I)) NLIN
  DO 11 I=1,NP1 NLIN
  B(I)=0. NLIN
  DO 11 J=1,NP1 NLIN
11 B(I)=B(I)+AINV(I,J)*P(J) NLIN
  DO 13 I=1,NP1 NLIN
  IF (I.EQ.JCOL) GO TO 13 NLIN
  DO 12 J=1,NP1 NLIN
12 AINV(I,J)=AINV(I,J)-AINV(JCOL,J)*(B(I)/B(JCOL)) NLIN
13 CONTINUE NLIN

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      DO 14 I=1,NP1          NLIN
14 AINV(JCOL,I)=AINV(JCOL,I)/B(JCOL)          NLIN
15 CONTINUE          NLIN
      WRITE(20,101)          NLIN
101 FORMAT('1*****THE FOLLOWING CASE DID NOT CONVERGE*****')          NLIN
16 WRT=.TRUE.          NLIN
      CALL EVAL(XBAR,P)          NLIN
      DO 17 I=1,N          NLIN
17 XIN(I)=XBAR(I,2)          NLIN
      RETURN          NLIN
      END          NLIN

      SUBROUTINE MATINV(A,B,III)          MINV
      IMPLICIT REAL*8 (A-H,C-Z)          MINV
      DIMENSION A(21,21), B(21,21)          MINV
      KK=III          MINV
      NV=III          MINV
      KKM=KK-1          MINV
      DO9I=1,KKM          MINV
      S=0.0          MINV
      DO1J=I,KK          MINV
      R=DABS(A(J,I))          MINV
      IF(R.LT.S) GO TO 1          MINV
      S=R          MINV
      L=J          MINV
1 CONTINUE          MINV
      IF (L.EQ.I) GO TO 5          MINV
      IF(L-I) 2,5,2          MINV
2 DO3J=I,KK          MINV
      S=A(I,J)          MINV
      A(I,J)=A(L,J)          MINV
3 A(L,J)=S          MINV
      DO4J=1,NV          MINV
      S=B(I,J)          MINV
      B(I,J)=B(L,J)          MINV
4 B(L,J)=S          MINV
5 IF(A(I,I).EQ.0.) GO TO 9          MINV
      IPO=I+1          MINV
      DO8J=IPO,KK          MINV
      IF(A(J,I).EQ.0.) GO TO 8          MINV
      S=A(J,I)/A(I,I)          MINV
      A(J,I)=0.0          MINV
      DO6K=IPO,KK          MINV
6 A(J,K)=A(J,K)-A(I,K)*S          MINV
      DO7K=1,NV          MINV
7 B(J,K)=B(J,K)-B(I,K)*S          MINV
8 CONTINUE          MINV
9 CONTINUE          MINV
      KMO=KK-1          MINV
      DO11K=1,NV          MINV
      B(KK,K)=B(KK,K)/A(KK,KK)          MINV
      DO11I=1,KMO          MINV
      N=KK-I          MINV
      DO10J=N,KMO          MINV
10 B(N,K)=B(N,K)-A(N,J+1)*B(J+1,K)          MINV
11 B(N,K)=B(N,K)/A(N,N)          MINV
      RETURN          MINV
      END          MINV

```

```

C
C SUBROUTINE CALLED BY S NONLIN TO CALCULATE ION CONCENTRATIONS C(I) ,          EVAL
C ACTIVITY COEFFICIENTS G(I) , AND THE RESIDUAL IN THE WATER                      EVAL
C ACTIVITY BALANCE Y(1) FOR THE CURRENT VALUE OF CGW AS GIVEN BY X(1)          EVAL
C SUBROUTINE IVAL(X,Y)                                                       EVAL
C IMPLICIT REAL*8(A-H,O-Z)                                              EVAL
C DIMENSION X(10),Y(21)                                              EVAL
C COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)                                EVAL
C COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTCH,RTOZ,RTO,NCAL,NITN          EVAL
C COMMON/EQK/EKN,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW          EVAL
C COMMON/ACT/DHA,G(2,4),P,PWO,CONST                                         EVAL
C COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW                                         EVAL
C CGW=DEXP(X(1)*FLN)                                              EVAL
C GHS=G(1,1)*G(2,2)/G(2,1)                                              EVAL
C QHS=EK3S/GHS                                                       EVAL
C BEGIN REFINEMENT OF ACTIVITY COEFFICIENTS AND ION CONCENTRATIONS          EVAL
C FOR CURRENT VALUE OF X(1)                                              EVAL
DO 10 I=1,100
  BT=(QHS+(CN-C6)/CGW)/2.0                                              EVAL
  CT=QHS*(CN-2.0*C6)/CGW                                              EVAL
  C(1,1)=-BT+DSQRT(BT**2-CT)                                              EVAL
  C(1,2)=CN/CGW                                                       EVAL
  TERM=C6/(CGW*(C(1,1)+QHS))                                              EVAL
  C(2,1)=C(1,1)*TERM                                              EVAL
  C(2,2)=QHS*TERM                                              EVAL
  QL=QHS                                                       EVAL
  CALL GAMAW(DHA,2,2,B,C,Z,G,P,AWI)                                         EVAL
  GHS=G(1,1)*G(2,2)/G(2,1)                                              EVAL
  QHS=EK3S/GHS                                                       EVAL
  IF(((QL-QHS)/QHS)**2.LT.1.0D-12) GO TO 20                           EVAL
10 CONTINUE
20 RAW=(CW-CGW/FLN)/(PWO*CONST)                                              EVAL
C CALCULATE RESIDUAL IN WATER ACTIVITY BALANCE                           EVAL
  Y(1)=(AWI-RAW)*FMW*(C(1,1)+C(1,2)+C(2,1)+C(2,2))          EVAL
  RETURN
END

C
C SUBROUTINE, CALLED BY NONLIN, CALCULATES ION CONCENTRATIONS ,          EVAL
C ACTIVITY COEFFICIENTS AND RESIDUALS FOR CURRENT VALUES OF          EVAL
C CGW,PH AND CSS AS GIVEN BY X(1),X(2), AND X(3)                      EVAL
C SUBROUTINE EVAL(X,Y)                                              EVAL
C IMPLICIT REAL*8 (A-H,O-Z)                                              EVAL
C DIMENSION X(10),Y(21)                                              EVAL
C COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)                                EVAL
C COMMON/CASE/NOX                                              EVAL
C COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTCH,RTOZ,RTO,NCAL,NITN          EVAL
C COMMON/EQK/EKN,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW          EVAL
C COMMON/CON/CMH,CNN,CM1N,CNC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,    EVAL
&PN,PC,PS                                              EVAL
C COMMON/ACT/DHA,G(2,4),P,PWO,CONST                                         EVAL
C COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW                                         EVAL
  IF(CSS.EQ.0.0) GO TO 10                                              EVAL
  CSS=CN/2.0*(1.0-DEXP(-X(3)*FLN))                                              EVAL
10 CMH=DEXP(-X(2)*FLN)                                              EVAL
  CGW=DEXP(X(1)*FLN)                                              EVAL
  C(1,1)=CMH                                              EVAL
C BEGIN REFINEMENT OF ACTIVITY COEFFICIENTS          EVAL
DO 1 J=1,100

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CALL GAMAW(DHA,2,2,B,C,Z,G,P,AWI)          EVAL
GQ1=G(2,1)*G(1,1)                          EVAL
GQ2=G(2,1)/(G(2,2)*G(1,1))                EVAL
EQ1N=EK1N*G(1,1)/G(1,2)                    EVAL
EQC=EKC*AWI                                EVAL
EQ1C=EK1C/GQ1                                EVAL
EQ2C=EK2C*GQ2                                EVAL
EQS=EKs*AWI                                EVAL
EQ1S=EK1S/GQ1                                EVAL
EQ2S=EK2S*GQ2                                EVAL
EQ3S=EK3S*GQ2                                EVAL
EQSS=EKSS/(G(1,2)**2*G(2,2))                EVAL
EQW=EKW*AWI/GQ1                                EVAL
CMHSQ=CMH*CMH                                EVAL
PN=(CN-2.0*CSS)/(CONST+EKN*(1.0+EQ1N*CMH)*CGW)  EVAL
CMN=EKN*PN                                EVAL
CM1N=EQ1N*CMN*CMH                            EVAL
PC=CC/(CONST+EQC*(1.0+EQ1C/CMH+EQ1C*EQ2C/CMHSQ)*CGW)  EVAL
CMC=EQC*PC                                EVAL
CM1C=EQ1C*CMC/CMH                            EVAL
CM2C=EQ2C*CM1C/CMH                            EVAL
PS=C4/(CONST+EQS*(1.0+EQ1S/CMH+EQ2S*EQ1S/CMHSQ)*CGW)  EVAL
CMS=EQS*PS                                EVAL
CM1S=EQ1S*CMS/CMH                            EVAL
CM2S=EQ2S*CM1S/CMH                            EVAL
CMOH=EQW/CMH                                EVAL
CM3S=(C6-CSS)/((1.0+EQ3S/CMH)*CGW)          EVAL
CM4S=EQ3S*CM3S/CMH                            EVAL
CMA1=CM1C+CM1S+CM3S+CMOH                    EVAL
CMA2=CM2C+CM2S+CM4S                            EVAL
DLSQI=((C(1,2)-CM1N)/CM1N)**2+((C(2,1)-CMA1)/CMA1)**2  EVAL
S+((C(2,2)-CMA2)/CMA2)**2                    EVAL
C(1,2)=CM1N                                EVAL
C(2,1)=CMA1                                EVAL
C(2,2)=CMA2                                EVAL
IF(DLSQI.LT.1.0D-12) GO TO 2                EVAL
1 CONTINUE                                EVAL
C CALCULATE RESIDUALS IN CHARGE BALANCE Y(1), WATER ACTIVITY BALANCE
C Y(2), AND (NH4) 2SC4 SOLUBILITY BALANCE Y(3) (IF APPLICABLE)  EVAL
C Y(2), AND (NH4) 2SC4 SOLUBILITY BALANCE Y(3) (IF APPLICABLE)  EVAL
2 SUMI=C(1,1)+C(1,2)+C(2,1)+C(2,2)          EVAL
Y(1)=(C(1,1)+C(1,2)-C(2,1)-2.0*C(2,2))/SUMI  EVAL
SUMN=CMN+CMS+CPC                            EVAL
AWI=DEXP(-FMW*SUMI*P)                        EVAL
CRAW=1.0/AWI+FMW*SUMN                        EVAL
ARAW=PWO*CONST/(CW-CGW/FMW)                  EVAL
Y(2)=(ARAW-CRAW)/(FMW*SUMI)                  EVAL
IF(CSS.EQ.0.0) GO TO 20                      EVAL
AQSS=CM1N**2*CM4S                            EVAL
Y(3)=(AQSS/EQSS)**(1.0/3.0)-1.0            EVAL
20 NCAL=NCAL+1                                EVAL
NITN=NITN+J                                EVAL
RETURN                                     EVAL
END                                         EVAL
C
C SUBROUTINE TO CALCULATE ACTIVITY COEFFICIENTS AND ACTIVITY OF WATER  GMAW
C DHA = DEBEYE-HUCKEL CONSTANT, NC = NO. OF CATIONS  GMAW
C NA = NO. OF ANIONS, B = INTERACTION COEFFICIENTS,  GMAW
C C = ION CONCENTRATIONS (MMOL/G H2O), Z = ION CHARGE  GMAW

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C P = OSMOTIC COEFFICIENT, AWT = WATER ACTIVITY CALCULATED FROM          GMAW
C ION CONCENTRATIONS          GMAW
C SUBROUTINE GAMAW(DHA,NC,NA,B,C,Z,G,P,AWI)          GMAW
C IMPLICIT REAL*8(A-H,O-Z)          GMAW
C DIMENSION B(4,4,2),BG(4,4),BP(4,4),C(2,4),Z(2,4),G(2,4),ZQ(2,4)          GMAW
C SP=0          GMAW
C SI=0          GMAW
C DO 10 K=1,2          GMAW
C IF (K.EQ.1) GO TO 1          GMAW
C N=NA          GMAW
C GO TO 2          GMAW
1 N=NC          GMAW
2 DO 10 T=1,N          GMAW
C ZQ(K,T)=Z(K,T)*Z(K,T)          GMAW
C SP=SP+C(K,T)          GMAW
C SI=SI+ZQ(K,T)*C(K,T)          GMAW
10 CONTINUE          GMAW
C CI=0.5*SI          GMAW
C IF (CI.GT.0.0) GO TO 5          GMAW
C CALL PRINT(3)          GMAW
C CALL EXIT          GMAW
5 SCI=DSQRT(CI)          GMAW
C TI=1.0+SCI          GMAW
C FG=-DHA*SCI/TI          GMAW
C FP=-2.0*DHA*(TI-1.0/TI-2.0*DLOG(TI))          GMAW
C FPI=DEXP(-2.0*SCI)          GMAW
C FGI=(1.0-FPI*(1.0+2.0*SCI-2.0*CI))/(4.0*CI)          GMAW
C SBM=0          GMAW
C DO 20 I=1,NC          GMAW
C DO 20 J=1,NA          GMAW
C BG(I,J)=B(I,J,1)+FGI*B(I,J,2)          GMAW
C BP(I,J)=B(I,J,1)+FPI*B(I,J,2)          GMAW
C SBM=SBM+BP(I,J)*C(1,I)*C(2,J)          GMAW
20 CONTINUE          GMAW
C DO 30 I=1,NC          GMAW
C SG=0          GMAW
C DO 40 J=1,NA          GMAW
C SG=SG+BG(I,J)*C(2,J)          GMAW
40 CONTINUE          GMAW
C G(1,I)=DEXP(ZQ(1,I)*FG+SG)          GMAW
C 30 CONTINUE          GMAW
C DO 50 J=1,NA          GMAW
C SG=0          GMAW
C DO 60 I=1,NC          GMAW
C SG=SG+BG(I,J)*C(1,I)          GMAW
60 CONTINUE          GMAW
C G(2,J)=DEXP(ZQ(2,J)*FG+SG)          GMAW
C 50 CONTINUE          GMAW
C P=1.0+(FP+SBM)/SP          GMAW
C AWI=DEXP(-0.0180153*SP*P)          GMAW
C RETURN          GMAW
C END          GMAW

C SUBROUTINE TO CALCULATE STARTING VALUES OF ARGUMENTS          XPCT
C NOX = NO OF CALCULATION CALLED BY KUTTA          XPCT
C IN = SIGNAL (IF 1) THAT THIS IS FIRST CALCULATION OF EACH INCREMENT          XPCT
C X = DISTANCE FROM SOURCE, Y = ARRAY OF ARGUMENTS          XPCT
C SUBROUTINE EXPCT(NOX,IN,X,Y)          XPCT

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IMPLICIT REAL*8(A-H,O-Z)          XPCT
DIMENSION Y(6),YL(6),YLL(6),DYDX(6)  XPCT
IF(NOX.EQ.1) GO TO 211             XPCT
IF(NOX.EQ.2) GO TO 121             XPCT
IF(IN.EQ.0.AND.IX.EQ.1.AND.ID.EQ.1) GO TO 111  XPCT
IF(IN.EQ.1.AND.ID.EQ.1) GO TO 311  XPCT
IF(IN.EQ.0.AND.IX.EQ.3) GO TO 132  XPCT
IF(IN.EQ.0.AND.IX.EQ.1.AND.ID.EQ.2) GO TO 112  XPCT
IF(IN.EQ.1.AND.ID.EQ.2) GO TO 312  XPCT
C IX=1 DO NOT SAVE X, IX=2 SAVE X, IX=3 SAVF X AND LAST X  XPCT
C IY=1 DO NOT SAVE Y, IY=2 SAVE Y, IY=3 SAVE Y AND LAST Y AND  XPCT
C CALCULATE VALUES OF DYDX  XPCT
C ID=1 DO NOT CHANGE Y, ID=2 INTERPOLATE OR EXTRAPOLATE Y  XPCT
C FIRST ENTRY, SAVE X  XPCT
211 IX=2  XPCT
    IY=1  XPCT
    ID=1  XPCT
    GO TO 2  XPCT
C SECOND ENTRY, SAVE Y VALUES FROM PREVIOUS CALCULATION  XPCT
121 IX=1  XPCT
    IY=2  XPCT
    ID=1  XPCT
    GO TO 5  XPCT
C NOX=3,4,OR 5, RETURN CURRENT VALUES OF Y  XPCT
111 IX=1  XPCT
    IY=1  XPCT
    ID=1  XPCT
    RETURN  XPCT
C FIRST CALL FROM KUTTA IN SECOND INCREMENT, SAVE X AND XL  XPCT
311 IX=3  XPCT
    IY=1  XPCT
    ID=1  XPCT
    GO TO 3  XPCT
C SECOND CALL FROM KUTTA IN AN INCREMENT AFTER FIRST, SAVE Y AND YL,  XPCT
C CALCULATE DYDX AND RETURN NEW VALUES OF Y  XPCT
132 IX=1  XPCT
    IY=3  XPCT
    ID=2  XPCT
    GO TO 6  XPCT
C 3RD, 4TH, OR 5TH CALL FROM KUTTA IN AN INCREMENT AFTER FIRST,  XPCT
C RETURN NEW Y VALUES  XPCT
112 IX=1  XPCT
    IY=1  XPCT
    ID=2  XPCT
    GO TO 8  XPCT
C FIRST CALL FROM KUTTA IN AN INCREMENT AFTER SECOND,  XPCT
C SAVE X AND XL, AND RETURN NEW Y VALUES  XPCT
312 IX=3  XPCT
    IY=1  XPCT
    ID=2  XPCT
    3 XLL=XL  XPCT
    2 XL=X  XPCT
    IF(ID.EQ.2) GO TO 9  XPCT
    RETURN  XPCT
5 DO 10 I=1,6  XPCT
10 YL(I)=Y(I)  XPCT
    RETURN  XPCT
6 DO 20 I=1,6  XPCT

```

```

YLL(T)=YL(I)          XPC1
YL(I)=Y(I)            XPC1
20 DYDX(I)=(YL(I)-YLL(I))/(XL-XLL) XPC1
8 DO 30 I=1,6          XPC1
30 Y(I)=YL(I)+DYDX(I)*(X-XL) XPC1
RETURN                XPC1
9 DO 40 I=1,6          XPC1
40 Y(I)=YL(I)+DYDX(I)*(XL-XLL) XPC1
RETURN                XPC1
END                  XPC1

C
C SUBROUTINE TO DETERMINE COMPONENT CONC'N BY RUNGA-KUTTA INTEGRATION KUTA
C SUBROUTINE KUTTA(XL,XU,Y,NE,DEL,ACCURC,IMAX,EQUA) KUTA
C PRCGPAM AUTHOR F. D. HAMMERLING, KUTA
C COMPUTING TECHNOLOGY CENTER, UNION CARBIDE CORP., NUCLEAR DIV., KUTA
C OAK RIDGE, TN. KUTA
CC MODIFIED AS INDICATED FOR THIS PROGRAM KUTA
    EXTERNAL EQUA KUTA
    DIMENSION Y(10),YI(10),YN(10),K1(10),K2(10),K3(10), K4(10),K5(10), KUTA
    > F(10),E(10),F1(10) KUTA
    DOUBLE PRECISION Y,F,E,F1,XN,XL,DEL,XU,H,TEST,ACCURC, YN,YI,K1, KUTA
    > K2,K3,K4,K5 KUTA
    LOGICAL QUIT,FIRST KUTA
    ITTER=0 KUTA
    N=NE KUTA
    XN=XL KUTA
    H=DEL KUTA
    FIRST=.TRUE. KUTA
    QUIT=.FALSE. KUTA
    DO 10 I=1,N KUTA
10   YN(I)=Y(I) KUTA
20 IF (XN+H.LT.XU) GO TO 30 KUTA
    DEL=H KUTA
    H=XU-XN KUTA
    IF (FIRST) DEL=H KUTA
    QUIT=.TRUE. KUTA
CC THE INTEGER (IN = 1 OR 0) HAS BEEN ADDED TO ARGUEMENT LIST KUTA
CC OF S. EQUA TO SIGNAL FIRST CALCULATION OF EACH INCREMENT KUTA
30 CALL EQUA(XN,YN,F1,1) KUTA
CC PRINT OUT RESULTS OF THIS FIRST CALCULATION AT THE BEGINNING OF KUTA
CC EACH INTERVAL IN X SPECIFIED IN THE MAIN PROGRAM KUTA
    IF (FIRST) CALL PRINT(2) KUTA
40 DO 50 I=1,N KUTA
    K1(I)=H*F1(I)/3. KUTA
50   YI(I)=YN(I)+K1(I) KUTA
    CALL EQUA(XN+H/3.,YI,F,0) KUTA
    DO 60 I=1,N KUTA
    K2(I)=H*F(I)/3. KUTA
60   YI(I)=YN(I)+K1(I)/2.+K2(I)/2. KUTA
    CALL EQUA(XN+H/3.,YI,F,0) KUTA
    DO 70 I=1,N KUTA
    K3(I)=H*F(I)/3. KUTA
70   YI(I)=YN(I)+3.*K1(I)/8.+9.*K3(I)/8. KUTA
    CALL EQUA(XN+H/2.,YI,F,0) KUTA
    DO 80 I=1,N KUTA
    K4(I)=H*F(I)/3. KUTA
80   YI(I)=YN(I)+3.*K1(I)/2.-9.*K3(I)/2.+6.*K4(I) KUTA
    CALL EQUA(XN+H,YI,F,0) KUTA

```

```

TEST=0.0
DO 90 I=1,N
  K5(I)=H*F(I)/3.
  E(I)=(K1(I)-9.*K3(I)/2.+4.*K4(I)-K5(I)/2.)/5.
  TEST=DMAX1(TEST,DABS(E(I)))
90  CONTINUE
  IF (TEST.LT.ACURC) GO TO 100
  ITTER=ITTER+1
  IF (ITTER.GE.IMAX) GO TO 140
  H=H/2.
  QUIT=.FALSE.
  FIRST=.FALSE.
  GO TO 40
100 DO 110 I=1,N
110  YN(I)=YN(I)+(K1(I)+4.*K4(I)+K5(I))/2.
  XN=XN+H
  FIRST=.FALSE.
  IF (TEST.GE.ACURC/32.) GO TO 120
  H=2.*H
  ITTER=ITTER-1
120 IF (.NOT.QUIT) GO TO 20
  DO 130 I=1,N
130  Y(I)=YN(I)
  GO TO 150
140 WRITE(21,1000)
150 RETURN
1000 FORMAT(1X,'IMAX EXCEEDED')
END

C
C SUBROUTINE FOR PRINTING OUTPUT ON LINE PRINTER
SUBROUTINE PRINT(N)
IMPLTCIT REAL*8(A-H,O-Z)
COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
COMMON/CASE/NOX
COMMON/RTE/NSC,NRL,VRGO,TP,H,UX,A(7),TA,DELX,VDG,VDP
COMMON/SIGMA/SGO,SGPY,SGPZ,SGLY,SGLZ,SGSY,SGSZ,BETAL,BETAS
COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,RTO,NCAL,NITN
COMMON/EQK/EKN,EK1N,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW
COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,
&PN,PC,PS
COMMON/ACT/DHA,G(2,4),P,PW0,CONST
COMMON/GIVN/WPCTC,WPCTH,WPCTS,QCCAL,FCTS6,EXAIR,FCON,TCP,TCA,
&XLAST,XNEXT,ERRR
COMMON/VAR/C4,C6,COZ,CCH,CC,CN,CW
  GO TO (1,2),N
1  WRITE(21,100) WPCTC,WPCTH,WPCTS,QCCAL,FCTS6,EXAIR,FCON,H,
  &TCF,NSC,NRL,UX,VDG,VDP,TCA
  WRITE(21,125) (A(I),I=1,7)
  WRITE(21,150) ERRR,SGO
  RETURN
2  WRITE(21,300) XLAST,DELX,SGPY,SGLY,SGSY,SGPZ,SGLZ,SGSZ,
  &CGW,PH,CSS,T,NCX,NCAL,NITN,
  &RTT,RTOH,RTOZ,RTO,G(1,1),G(1,2),G(2,1),G(2,2)
  WRITE(21,325) C4,C6,COZ,CCH,CC,CN,CW
  WRITE(21,340) PN,PC,PS
  IF(CGW.GT.0.0) GO TO 21
  WRITE(21,348)
21 WRITE(21,350) CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S

```

```

100 FORMAT(1H1,'SU IFUR CHEMISTRY IN A GAUSSIAN PLUME'///)
  81X,' WPCTC  WPCTH  WPCTS  QCCAL  FCTS6  EXAIR',
  8'   FCN      H      TCP NSC NRL      UX      VDG',
  8'   VDP      TCA'/
  81X,F7.2,8F8.2,2I4,4F8.2/)
125 FORMAT(
  81X,'      AC4      AC6      A0Z      A0H      ACC      ACN',
  8'      ACW'/1X,1P7E9.2///)
150 FORMAT(1X,'  ERRR      SGO  '/1X,1PE8.1,0PF8.2/)
300 FORMAT(
  821X,'      X      DELX      SGPY      SGLY      SGSY      SGPZ',
  8'      SGLZ      SGSZ'/21X,F7.1,7F8.2,//
  821X,'      CGW      PH      CSS      T      NOX      NCL      NITN',
  821X,1PE9.2,0P2F8.4,F8.2,2I6,I7//
  821X,'      RTT      RTOH      RTOZ      RTO      GH',
  8'      GN      GHS      GS'/21X,1P4E9.2,0P4F8.4/)
325 FORMAT(
  811X,'      C4      C6      COZ      CCH      CC      CN',
  8'      CW'/11X,1P7E9.2/)
340 FORMAT(
  811X,'      PN      PC      PS'/11X,1P3E8.1/)
348 FORMAT(1X,'DROPLETS HAVE DRIED UP, LAST SCLN. COME. WAS: '/')
350 FORMAT(
  81X,'      CMN      CM1N      CMC      CM1C      CM2C',
  8'      CMS      CM1S      CM2S      CM3S      CM4S'/
  81X,1P10E12.4//)
      RETURN
      END

```

3. Sample Input and Output

The following input values, used in the reference case (Table V), were read by the indicated READ statements in M. SULCAL:

```

[READ(20, 100) WPCTC,WPCTH,QCOAL,FCTS6,EXAIR,FCON,TCP,H]
 82.57      6.93     100.0      0.02      0.25      5.0     177.0
 400.0

[READ(20, 100) NSC,NRL,UX,VDG,VDP,(A(I), I=1, 7),TCA,WPCTS]
 3          5        2.0        0.01      0.01      0.004      0.0
 0.0032    2.0E-8     12.26     6.0E-4     929.7     21.11     1.0

[READ(20, 100) XNEXT,DELX,ERRR]
 300.0      0.1  0.000003

[READ(20, 100) NX,(XINCR(I), I=1,NX)]
 4      700.0     4000.0     5000.0    15000.0

[100 FORMAT (7G 10.0)]

```

The output produced is shown on the next two pages. The first three rows of values, mostly input, are independent of x. In row 2 are the ambient concentrations of the seven components S(IV), S(VI), O₃, OH, CO₂, NH₃, and H₂O. Rows 4 through 9 are repeated for each specified value of x. Row 4 includes the current value of the increment in x (DELX) in S. KUTTA. In row 5, NOX, NCL and NITN are, respectively, the accumulated entries into S. RATE, S. EVAL and S. GAMAW. Row 6 includes the oxidation rates: total, by OH, by O₃, and by oxygen; and the activity coefficients of the aqueous ions H⁺, NH₄⁺, X⁻, and X²⁻. In row 7 are the effective concentrations of components in the plume. Row 8 includes the partial pressures of NH₃, CO₂ and SO₂ gases. Finally in row 9 are the concentrations of the aqueous species NH₃, NH₄⁺, H₂CO₃, HCO₃⁻, CO₃²⁻, H₂SO₃, HSO₃⁻, SO₃²⁻, HSO₄⁻ and SO₄²⁻.

SULFUR CHEMISTRY IN A GAUSSIAN PLUME

WPCTC	WPCTH	WPCTS	QCOAL	FCFS6	EXAIR	FCON	H	TCP	NSC	NRL	UX	VDG	VDP	TCA
82.57	6.93	1.00	100.00	0.02	0.25	5.00	400.00	177.00	3	5	2.00	0.01	0.01	21.11

AC4	AC6	AOZ	AOH	ACC	ACN	ACW
4.00D-03	0.0	3.20D-03	2.00D-08	1.23D 01	6.00D-04	9.30D 02

FRRR	SG0
3.0D-06	8.84

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
300.0	0.10	28.94	37.78	15.73	29.31	38.15	15.82

CGW	PH	CSS	T	NOX	NCI	NITN
6.39D-03	-1.2005	0.0	330.21	2	5	24

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
3.00D-05	3.00D-05	5.27D-16	1.32D-18	49.3771	0.3923	59.5242	0.0157

C4	C6	COZ	COH	CC	CN	CW
4.89D 00	9.98D-02	2.19D-03	1.37D-08	1.11D 03	4.11D-04	1.19D 03

PN	PC	PS
9.4D-15	3.0D-02	1.3D-04

CMN	CM1N	CMC	CM1C	CM2C	CMS	CM1S	CM2S	CM3S	CM4S
1.3990D-13	6.4383D-02	3.6043D-05	1.0670D-15	4.3346D-25	1.0746D-05	2.0027D-12	4.0459D-19	1.5294D 01	3.1787D-01

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
1000.0	12.80	77.41	86.25	27.26	75.91	84.75	26.90

CGW	PH	CSS	T	NOX	NCI	NITN
4.47D-03	-0.9569	0.0	305.47	595	2973	16466

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
1.32D-05	1.32D-05	1.41D-14	2.28D-17	6.0477	0.4091	9.8348	0.0165

C4	C6	COZ	COH	CC	CN	CW
1.65D 00	3.79D-02	2.86D-03	1.79D-08	3.82D 02	5.36D-04	1.02D 03

PN	PC	PS
2.0D-14	9.6D-03	4.1D-05

CMN	CM1N	CMC	CM1C	CM2C	CMS	CM1S	CM2S	CM3S	CM4S
8.3464D-13	1.2014D-01	1.4468D-04	1.2400D-13	7.3285D-23	2.1002D-05	5.6631D-10	3.3968D-16	7.7830D 00	6.9631D-01

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
5000.0	102.40	223.19	232.03	61.94	236.07	244.92	65.00

CGW	PH	CSS	T	NOX	NCI	NITN
3.20D-03	-0.6334	0.0	296.23	1160	5805	35040

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
2.60D-06	2.60D-06	3.84D-14	9.02D-17	1.4106	0.4276	2.7777	0.0196

C4	C6	COZ	COH	CC	CN	CW
----	----	-----	-----	----	----	----

2.97D-01 1.20D-02 3.14D-03 1.96D-08 7.93D 01 5.88D-04 3.45D 02

PN PC PS
1.0D-13 1.9D-03 7.2D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S
6.3393D-12 1.8356D-01 5.7870D-05 1.4439D-12 1.5260D-21 7.8317D-06 8.3611D-09 1.2068D-14 2.9836D 00 7.4973D-01

X DELX SGPY SGLY SGSY SGPZ SGLZ SGSZ
10000.0 204.80 329.80 338.65 87.30 355.37 364.22 93.39

CGW PH CSS T NOX NCL NITN
3.04D-03 -0.5419 0.0 295.23 1340 6705 40674

RTT RTOH RTOZ RTO GH GN GHS GS
1.28D-06 1.28D-06 2.97D-14 7.95D-17 1.1468 0.4362 2.2439 0.0211

C4 C6 COZ COH CC CN CW
1.45D-01 9.11D-03 3.17D-03 1.98D-08 4.54D 01 5.93D-04 9.37D 02

PN PC PS
1.5D-13 1.1D-03 3.5D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S
9.7110D-12 1.9500D-01 3.5642D-05 1.6543D-12 1.9528D-21 4.1551D-06 8.5414D-09 1.4235D-14 2.3121D 00 6.8261D-01

X DELX SGPY SGLY SGSY SGPZ SGLZ SGSZ
25000.0 819.20 534.71 543.56 136.05 552.64 561.49 140.32

CGW PH CSS T NOX NCL NITN
3.10D-03 -0.4734 0.0 294.67 1540 7707 46631

RTT RTOH RTOZ RTO GH GN GHS GS
5.30D-07 5.30D-07 1.73D-14 5.09D-17 1.0142 0.4435 1.9694 0.0224

C4 C6 COZ COH CC CN CW
5.95D-02 7.87D-03 3.19D-03 1.99D-08 2.63D 01 5.94D-04 9.33D 02

PN PC PS
1.8D-13 6.4D-04 1.4D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S
1.2339D-11 1.9167D-01 2.1560D-05 1.5005D-12 1.9100D-21 1.7838D-06 5.6055D-09 1.0265D-14 1.9130D 00 6.2662D-01



APPENDIX B

EFFECTIVE VOLUME OF FREE PLUME

The reactions in a plume are complicated by the situation that the reaction rates vary across and down the plume because of variations in concentration, temperature and mixing. The detailed treatment of these varying conditions is beyond the scope of this investigation. We want to use an idealized model for this study and yet relate the result to the conditions which are obtained in the open environment. The dispersion of effluents in the atmosphere has been successfully modeled by a Gaussian distribution in which the variance or dispersion σ follows empirically determined relations that are functions of time, distance from source, and the prevailing meteorological and topographical conditions. We will show that when modeling reaction rates it is possible to use a simple uniform model which can yield the same reaction rates as a well mixed Gaussian distribution.

In the following treatment, let q_i be the total amount of reactant i , in a unit length of the plume in the x direction, let C_i be the density or concentration of reactant i , and let R be the total rate of production of the reaction product. Consider a binary reaction

$$\frac{\text{Rate}}{\text{Volume}} = k C_1^{\alpha_1} C_2^{\alpha_2}, \quad (1)$$

then

$$R = k \int_V C_1^{\alpha_1} C_2^{\alpha_2} dv. \quad (2)$$

Let

$$C_i = \frac{q_i}{2\pi\sigma_y\sigma_z} e^{-y^2/2\sigma_y^2} e^{-z^2/2\sigma_z^2}. \quad (3)$$

Then the rate of reaction in a volume of unit length in the x-direction and infinite in extent in the y and z direction is

$$\begin{aligned}
 R &= \int_{-\infty}^{\infty} dy \int_{-\infty}^{\infty} dz \ k \ C_1^{\alpha_1} C_2^{\alpha_2} \\
 &= \frac{k q_1^{\alpha_1} q_2^{\alpha_2}}{2\pi\sigma_y \sigma_z} \int_{-\infty}^{\infty} dy \int_{-\infty}^{\infty} dz e^{-(\alpha_1+\alpha_2)y^2/2\sigma_y^2} e^{-(\alpha_1+\alpha_2)z^2/2\sigma_z^2} \\
 &= \frac{k q_1^{\alpha_1} q_2^{\alpha_2}}{(2\pi\sigma_y \sigma_z)^{\alpha_1+\alpha_2}} (2\pi\sigma_y \sigma_z) \cdot \frac{1}{\alpha_1 + \alpha_2} \\
 &\equiv \frac{k q_1^{\alpha_1} q_2^{\alpha_2}}{(4\pi\sigma_y \sigma_z)^{\alpha_1+\alpha_2}} (4\pi\sigma_y \sigma_z) \left(\frac{2^{\alpha_1+\alpha_2-1}}{\alpha_1 + \alpha_2} \right), \quad (4)
 \end{aligned}$$

where we have assumed a uniform or effective temperature. Now suppose the same amounts of reactants were spread uniformly through a volume V such that

$$\bar{C}_i = q_i/V. \quad (5)$$

Then

$$\bar{R} = k \int_V \bar{C}_1^{\alpha_1} \bar{C}_2^{\alpha_2} dv = \frac{k q_1^{\alpha_1} q_2^{\alpha_2}}{V^{\alpha_1+\alpha_2}} \cdot V. \quad (6)$$

Now if we choose as a control volume $V = 4\pi\sigma_y \sigma_z$ then

$$\bar{R} = \frac{\alpha_1 + \alpha_2}{2^{\alpha_1+\alpha_2-1}} R = F R. \quad (7)$$

The choice $V = 4\pi\sigma_y \sigma_z$ yields reaction rates which are correct through second order. Note this includes the case $\alpha_1 = 0$ with $\alpha_2 = 1$ or 2.

Other cases can be handled by scaling the rate constants by the relation

$$\bar{k} = k \cdot \frac{2^{\alpha_1 + \alpha_2 - 1}}{\alpha_1 + \alpha_2} , \quad (8)$$

where \bar{k} is the effective rate constant. Similar results may be obtained for equilibria, and they may be extended to reactions and equilibria of the form $k C_1^{\alpha_1} C_2^{\alpha_2} \dots C_n^{\alpha_n}$ by replacing $(\alpha_1 + \alpha_2)$ in the equations above by $(\alpha_1 + \alpha_2 + \dots + \alpha_n)$.

1. Mean Concentration When Plume Touches the Ground

The foregoing discussion assumes a free plume and neglects interactions with the surface. An effect of the ground is to limit the downward spread of the plume and to restrict the dilution of the effluents near the land surface. This has been successfully modeled by assuming an "image plume", i.e., a "reflection" of the plume at the ground. The effect of the "image plume" is included in the present treatment as follows:

Consider the total reaction rate

$$R = k \int_V C_1 C_2 dv , \quad (9)$$

where the concentrations C_i have the distribution

$$C_i(x, y, z) = \frac{q_i}{2\pi\sigma_y\sigma_z} e^{-y^2/2\sigma_y^2} y \left[e^{-(z-h)^2/2\sigma_z^2} + e^{-(z+h)^2/2\sigma_z^2} \right] , \quad (10)$$

where q_i is the total amount of material available to react in a unit length of plume. Then

$$R = \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \int_{-\infty}^{\infty} e^{-2y^2/2\sigma_y^2} dy \int_{-\infty}^{\infty} \left(e^{-(z-h)^2/2\sigma_z^2} + e^{-(z+h)^2/2\sigma_z^2} \right)^2 dz$$

$$= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \cdot \sqrt{2\pi} \sigma_y \int_0^{\infty} dz \left[\left(e^{-2(z-h)^2/2\sigma_z^2} \right. \right.$$

$$\left. \left. + e^{-(z+h)^2/2\sigma_z^2} \right) + 2e^{-\frac{(z-h)^2}{2\sigma_z^2} - \frac{(z+h)^2}{2\sigma_z^2}} \right]$$

$$= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \left[\frac{2\pi\sigma_y \sigma_z}{2} + \sqrt{2\pi} \sigma_y 2 \int_0^{\infty} e^{-(z^2 - 2hz + h^2 + z^2 + 2hz + h^2)/2\sigma_z^2} dz \right]$$

$$= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \left[\frac{2\pi\sigma_y \sigma_z}{2} + \sqrt{2\pi} \sigma_y e^{-h^2/\sigma_z^2} 2 \int_0^{\infty} e^{-2z^2/2\sigma_z^2} dz \right]$$

$$= \frac{kq_1 q_2}{(4\pi\sigma_y \sigma_z)^2} \cdot (4\pi\sigma_y \sigma_z) \left[1 + \frac{e^{-h^2/\sigma_z^2}}{\sqrt{2}} \right] = k \left(\frac{q_1}{\beta V} \right) \left(\frac{q_2}{\beta V} \right) (\beta V) \beta \left[1 + \frac{e^{-h^2/\sigma_z^2}}{\sqrt{2}} \right],$$

(11)

where $V = 4\pi\sigma_y\sigma_z$. To give the same reaction rate assuming $\bar{T} = Q/V$ we must choose

$$\beta = \frac{1}{1 + \exp(-h^2/\sigma^2)/\sqrt{2}} \quad . \quad (12)$$

Note for point source, $\beta(x=0) = 1$, $\beta(x=\sigma) = \frac{\sqrt{2}}{1+\sqrt{2}} = .586$.

2. Plume Depletion Due to Washout

According to the washout concept, the depletion from the box of the chemical model is expressed in terms of the washout coefficient Λ as,

$$\Delta m = \Lambda m \Delta t, \quad (13)$$

where m is the amount of (aerosol or water soluble) material in the box.

But in terms of the volume V of the box, $C = m/V$, hence

$$\left(\frac{dC}{dt}\right)_{\text{washout}} = -\Lambda C. \quad (14)$$

Washout occurs under conditions of higher than normal ambient humidity.

Washout is not included in the present model.

3. Plume Depletion Due to Dry Deposition

The ground level concentration according to the Gaussian plume model is given by

$$C(x, y, 0) = \frac{Q(x)}{2\pi\sigma_y\sigma_z u} e^{-y^2/2\sigma_y^2} e^{-h^2/2\sigma_z^2} \quad . \quad (15)$$

In terms of the concentration \bar{C} of the chemical box model

$$\bar{C}(x) = Q(x)/4\pi\sigma_y^2\sigma_z^2 u_x , \quad (16)$$

or

$$C(x, y, 0) = 4\bar{C} e^{-y^2/2\sigma_y^2} e^{-h^2/2\sigma_z^2} . \quad (17)$$

The dry deposition in time Δt at a distance x is

$$\begin{aligned} \Delta m &= \int_{-\infty}^{\infty} dy \int_{x-1/2}^{x+1/2} dx C(x, y, 0) V_d \Delta t = 4\bar{C} V_d \Delta t \sqrt{2\pi} \sigma_y u_x e^{-h^2/2\sigma_z^2} \\ &= - \left(\frac{d\bar{C}}{dt} \right)_{dep} (4\pi\sigma_y^2\sigma_z^2 u_x) \Delta t , \end{aligned} \quad (18)$$

where V_d is the deposition "velocity".

Hence

$$\left(\frac{dC}{dt} \right)_{dep} = - \bar{C} 2V_d \left(\frac{e^{-h^2/2\sigma_z^2}}{\sqrt{2\pi} \sigma_z} \right) . \quad (19)$$

4. Choice of Sigma

The behavior of a natural plume is far more complicated than the idealized case we have considered so far. Over a period of time convection cells move, the plume may wander, fan, fumigate, etc., so that the volume swept out is considerably larger than that which obtains in any instant. It is the larger volume which is described by the dispersion coefficients of Pasquill-Gifford and Briggs-Smith, which are

defined for a point source. For the shorter distances at least, the instantaneous plume will be considerably smaller. At very short distances a "concentration catastrophe" arises if the plume dimensions are allowed to shrink to zero. Two artifices are used to avoid this crisis. One is the concept of a "virtual (point) source" behind the stack, i.e. the distances along the plume are measured from a point behind the stack and can not go to zero. The second artifice (which we use) involves adding a constant such that $4\pi\sigma_y\sigma_{zux}$ reduces to the initial plume volume flow rate V_p° at $x = 0$. At small distances the additive term σ° dominates while at large distances it can be neglected.

In defining the various dispersion coefficients below we use an asterisk to denote a point source, a subscript ℓ for long-time (larger) σ 's and a subscript s for short-time (smaller) σ 's. We assume that large scale effects such as deposition are governed by these coefficients given by

$$\sigma_\ell = \sigma_\ell^* + \sigma^\circ . \quad (20)$$

The instantaneous dispersion coefficients used in determining the size of the "box" we define by

$$\sigma_s = f\sigma_\ell^* + \sigma^\circ , \quad (21)$$

where f is a function of the distance x , $f < 1$ and f approaches unity at large distances. The exact form of f is not crucial and we have elected to use a form due to Scriven,¹² which we write in the form

$$f = \sqrt{\frac{x}{x+50L}} , \quad (22)$$

where L is given in the table below.

<u>Stability Class</u>	<u>L (meters)</u>
A	500
B	250*
C	100
D	50*
E	20*
F	10

*Interpolated values.

APPENDIX C

ACTIVITY COEFFICIENTS, THE WATER ACTIVITY, AND INITIAL VALUES

1. The Estimation of Ion Activity Coefficients and the Osmotic Coefficient

For an ideal solution, the activity coefficients of ions are unity and the activity of water is given by its mole fraction,

$$a_w = \frac{n_w}{n_w + \sum_i n_i + \sum_j n_j} = \frac{55.51}{55.51 + \sum_i [i] + \sum_j [j]} \quad (1)$$

(here i and j refer to cations and anions, respectively). Such behavior is approached as solutions become more and more dilute. The aerosol droplets in the atmosphere often are quite concentrated, however, and the deviations from ideal behavior can be considerable (Fig. 4). Hence it is necessary to estimate the values of ion activity coefficients for use in the equilibrium constant expressions (Table I) and in the calculation of the activity of water. The following is based on the treatments of Pitzer and Brewer¹³ and Pitzer¹⁴.

The activity coefficients of ions g_i and g_j in a solution and the activity of water (a_w) are related by the Gibbs-Duhem equation,

$$\sum_i [i] d \ln([i] g_i) + \sum_j [j] d \ln([j] g_j) = -55.51 d \ln a_w. \quad (2)$$

The activity of water is usually expressed in terms of the osmotic coefficient ϕ ,

$$\ln a_w = -0.018(\sum_i [i] + \sum_j [j])\phi. \quad (3)$$

The relationship between the osmotic coefficient and the activity coefficient is obtained by combining Eqs. 2 and 3 and integrating,

$$\phi = 1 + \frac{\sum_i [i] d \ln g_i + \sum_j [j] d \ln g_j}{\sum_i [i] + \sum_j [j]}. \quad (4)$$

The integrations are from 0 concentration to the individual ion concentrations in the solution for which ϕ is desired.

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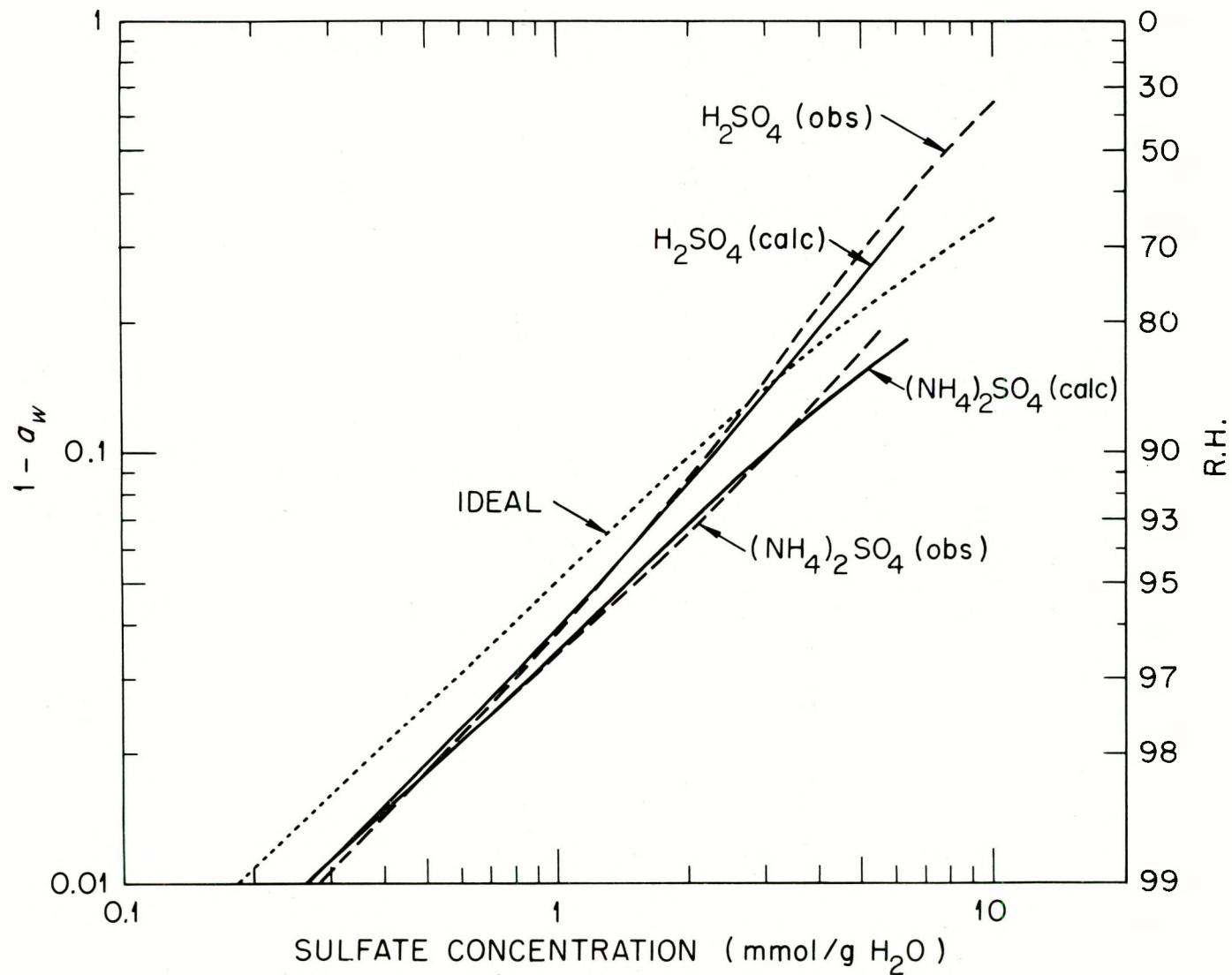


Fig. 4 Variation of the concentration of aerosol droplets with the activity of water a_w (the relative humidity/100) at 25°C.

The following expressions are assumed for the individual ionic activity coefficients;

$$\ln g_i = \frac{-z_i^2 S I^{1/2}}{1 + I^{1/2}} + \sum_j B_{ij}^g [j] , \quad (5a)$$

$$\ln g_j = \frac{-z_j^2 S I^{1/2}}{1 + I^{1/2}} + \sum_i B_{ij}^g [i] , \quad (5b)$$

wherein z_i and z_j are the ion charges, I is the ionic strength

$$I = \frac{1}{2} (\sum_i z_i^2 [i] + \sum_j z_j^2 [j]) , \quad (6)$$

and S is the Debye-Hückel constant, a theoretical quantity which depends only on the density and dielectric constant of the solvent and on the temperature. It is given with adequate accuracy for present purposes by the approximate expression

$$S = 1.1305 + 0.01842 t(^{\circ}C) \quad (7)$$

The coefficients B_{ij}^g , one for each pairwise combination of cations and anions is assumed to be a function only of I .

$$B_{ij}^g (I) = B_{ij}^{\circ} + B_{ij}^1 F(I) \quad (8)$$

$$F(I) = [1 - (1 + 2I^{1/2} - 2I) \cdot \exp(-2I^{1/2})]/4I . \quad (8a)$$

B_{ij}° and B_{ij}^1 are adjustable constants.

The above expressions for g_i and g_j (eqs 5-8) when introduced into eq 4 give the following expression for ϕ

$$\phi = 1.0 + \frac{-2SG(I) + \sum_i \sum_j B_{ij}^g [i][j]}{\sum_i [i] + \sum_j [j]} , \quad (9)$$

$$G(I) = (1 + I^{1/2}) - \frac{1}{(1 + I^{1/2})} - 2\ln(1 + I^{1/2}) , \quad (9a)$$

$$B_{ij}^{\phi} = B_{ij}^{\circ} + B_{ij}^l \exp(-2I^{1/2}) \quad (9b)$$

Since four kinds of ions are distinguished, two cations (H^+ and NH_4^+) and two anions (x^- and x^{2-}), there are four pairwise combinations of cations with anions and 8 adjustable parameters. Six of the following values were assigned by fitting the appropriate expressions above to the known osmotic and activity coefficients of pure H_2SO_4 and $(NH_4)_2SO_4$ solutions given by Robinson and Stokes¹⁵ (Fig. 4), and the observed dissociation of HSO_4^- ion in pure sulfuric acid solutions (Young et al¹⁶).

i	j	B_{ij}°	B_{ij}^l
H^+	x^-	0.272	3.080
H^+	x^{2-}	0.050	-4.005
NH_4^+	x^-	0.0*	0.319*
NH_4^+	x^{2-}	-0.049	-1.186

The remaining two (with asterisks) were adjusted to give approximately the observed solubility of $(NH_4)_2SO_4$ in H_2SO_4 solutions (Seidell,¹⁷ Vol. II p. 756), Fig. 5.

2. Initial Estimate of Quantities Needed for the Calculation of the Plume Composition

The initial estimates of the amount of condensed water g_w and of the ion concentrations $[H^+]$, $[NH_4^+]$, $[x^2]$ and $[x^{2-}]$ are made by assuming that the only ions present in solution are H^+ , NH_4^+ , HSO_4^- , and SO_4^{2-} and that all the NH_3 is present as NH_4^+ . The following expressions then apply;

$$[NH_4^+] = C(NH_3)/g_w, \quad (10)$$

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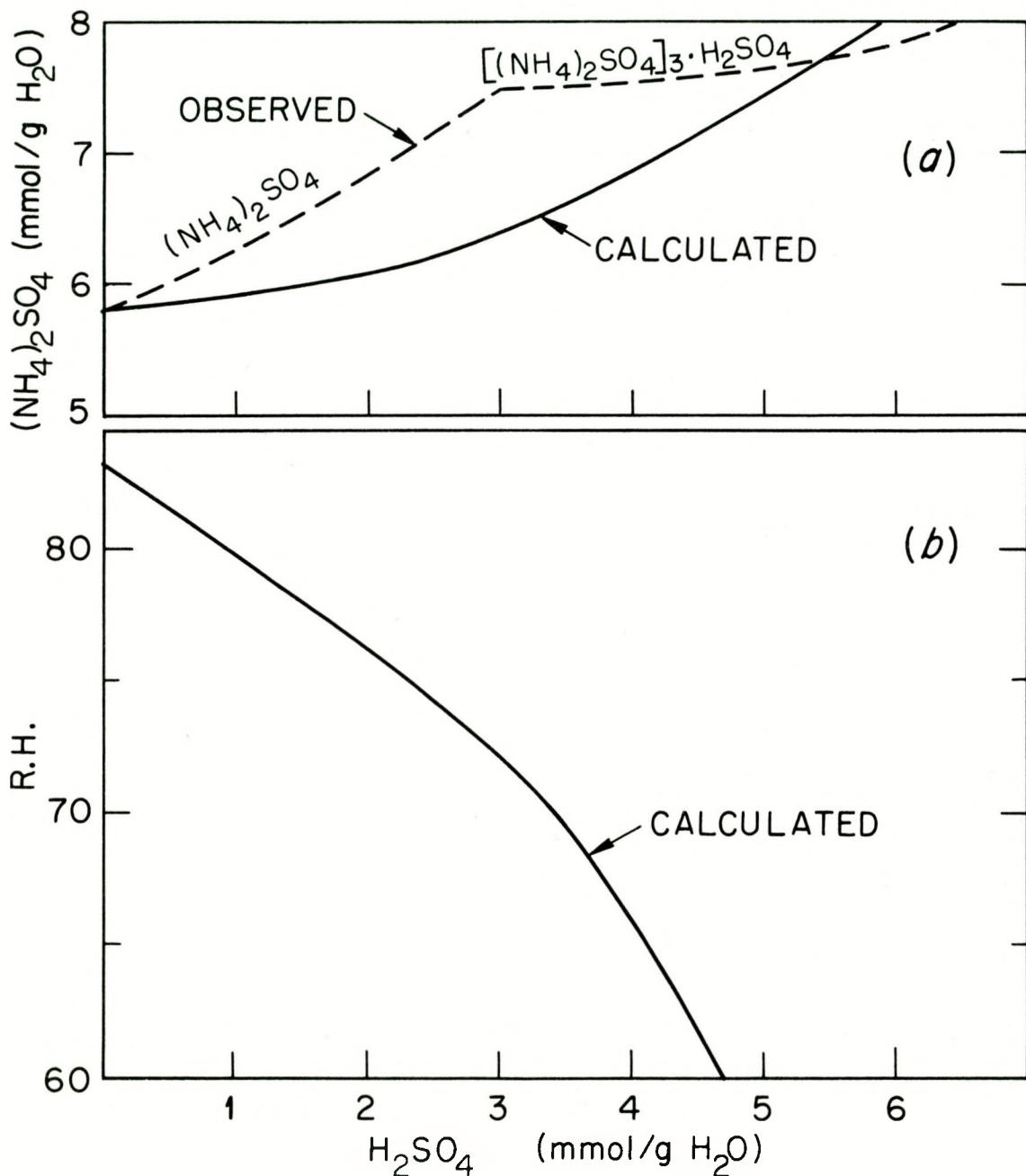


Fig. 5 The sulfuric acid - ammonium sulfate - water system at 25°C.

$$[x^-] = [HSO_4^-] = \frac{[H^+]}{[H^+] + Q_4} \cdot \frac{c(S(VI))}{g_w} , \quad (11)$$

$$[x^{2-}] = [SO_4^{2-}] = \frac{Q_4}{[H^+] + Q_4} \cdot \frac{c(S(VI))}{g_w} , \quad (12)$$

where Q_4 is the concentration quotient from K_4 (Table I),

$$Q_4 = \frac{[H^+][SO_4^{2-}]}{[HSO_4^-]} = K_4 \cdot \frac{g_{x^-}}{g_{H^+}g_{x^{2-}}} . \quad (13)$$

Introducing Eqs. 10 to 13 into the charge balance equation

$$[H^+] + [NH_4^+] = [HSO_4^-] + 2[SO_4^{2-}] , \quad (14)$$

a quadratic in $[H^+]$ is obtained with the solution

$$[H^+] = -B + \left(B^2 - C \right)^{1/2} , \quad (15)$$

wherein

$$B = \left[Q_4 + [c(NH_3) - c(S(VI))] / g_w \right] / 2 , \quad (15a)$$

$$C = Q_4 [c(NH_3) - 2c(S(VI))] / g_w . \quad (15b)$$

The only unknowns are g_w and the ion activity coefficients.

The calculation begins by solving the following equation for g_w ,

$$[c(H_2O) - g_w / 0.018] \frac{RT}{P_w^\circ} = \frac{g_w}{g_w + 3 \cdot c(S(VI)) \cdot 0.018} . \quad (16)$$

Each side is an expression of the activity of water, the right side being an approximation of the mole fraction of water in solution assuming three moles of ions are produced by a mole of S(VI). The solution for g_w is

$$g_w = \left[-B' + (B'^2 - C')^{1/2} \right] \cdot 0.018 , \quad (17)$$

$$B' = [P_w^\circ / RT - c(H_2O) + 3c(S(VI))] / 2 , \quad (17a)$$

$$C' = -3c(S(VI)) \cdot c(H_2O) . \quad (17b)$$

The activity coefficients are set initially at unity. The ion concentrations thus obtained from Eqs. 10-15 are used to estimate activity coefficients in S. GAMAW, and the calculation is repeated by cycling between S. EVAL AND S. GAMAW until the activity coefficients become constant (only a few interactions are needed). Then the water activity given by S. GAMAW is compared with that given by the amount of gaseous water,

$$a_w = [C(H_2O) - g_w/0.018]RT/P_w^\circ \quad (18)$$

and g_w is adjusted by repeated cycling between S. NONLIN and S. EVAL until a good estimate of g_w (and ion concentrations and activity coefficients) is obtained.

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