

**SULCAL:**  
**A MODEL of SULFUR CHEMISTRY in a PLUME**

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Program

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## TABLE OF CONTENTS

	PAGE
Abstract . . . . .	1
1. Introduction . . . . .	2
1.1 The Rapid Reactions of Sulfur . . . . .	2
1.2 The Slow Oxidation of $\text{SO}_2$ . . . . .	7
1.3 Other Reactions . . . . .	9
2. The Model . . . . .	11
2.1 The Plume Components . . . . .	14
2.2 Effective Concentrations and Plume Volume . . . . .	14
2.3 The Rate of Change of Component Concentration . . . . .	20
2.4 The Composition of the Plume . . . . .	21
3. Test Calculations . . . . .	25
References . . . . .	33
Appendices	
Appendix A Program SULCAL . . . . .	35
1. Partial List of Symbols . . . . .	36
2. Program Listing . . . . .	38
3. Sample Input and Output . . . . .	53
Appendix B Effective Volume of Free Plume . . . . .	57
1. Mean Concentration When Plume Touches the Ground . .	60
2. Plume Depletion Due to Washout . . . . .	62
3. Plume Depletion Due to Dry Deposition . . . . .	62
4. Choice of Sigma . . . . .	63

	PAGE
Appendix C Activity Coefficients, The Water Activity and Initial Values . . . . .	67
1. The Estimation of Ion Activity Coefficients and the Osmotic Coefficient . . . . .	68
2. Initial Estimate of Quantities Needed for the Calculation of the Plume Composition . . . . .	71

## LIST OF FIGURES

	PAGE
1. Program SULCAL .....	12
2. Quantities Calculated by SULCAL for the Reference Case ....	28
3. The Effect of the Relative Humidity and the Extent of Acid Neutralization on the Ratio of the Rates of $\text{SO}_2$ Oxidation by Ozone ( $R_{\text{O}_3}$ ) and by Hydroxyl Radical ( $R_{\text{OH}}$ ) at $25^\circ\text{C}$ .....	31
4. Variation of the Concentration of Aerosol Droplets with the Activity of Water $a_w$ (the Relative Humidity/100) at $25^\circ\text{C}$ ...	69
5. The Sulfuric Acid - Ammonium Sulfate - Water System at $25^\circ\text{C}$ .....	72





## LIST OF TABLES

	PAGE
Table I    Equilibrium Constants .....	4
Table II   Rate Expressions for S(IV) Oxidation .....	8
Table III   Input Data .....	15
Table IV   Calculation of Dispersion Coefficients .....	19
Table V   Sensitivity of Values Calculated by SULCAL to Variation of Input Values .....	26

## ABSTRACT

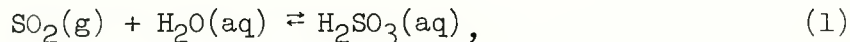
A computer program has been written that models the following features of the chemical behavior of sulfur emitted to the atmosphere from fossil-fuel burning power plants: (1) the rapid reactions of  $\text{SO}_2$  with aerosol droplets to produce the dissolved sulfite species  $\text{H}_2\text{SO}_3$ ,  $\text{HSO}_3^-$  and  $\text{SO}_3^{2-}$ ; (2) the rapid reactions of  $\text{SO}_3$  to produce the dissolved sulfate species  $\text{HSO}_4^-$  and  $\text{SO}_4^{2-}$ ; (3) the neutralization of the acid thus produced by atmospheric ammonia and (4) the eventual formation of particulate ammonium sulfate; (5) the slow oxidation of  $\text{SO}_2$  to sulfate species by hydroxyl radical, and (6) the slow oxidation of species of  $\text{SO}_2$  in aerosol droplets to sulfate species by dissolved ozone and oxygen. The model employs averaged concentrations based on the Gaussian plume and can calculate deposition rates for gaseous and particulate material as a function of such variables as distance from the source, wind speed, meteorological stability class, temperature, relative humidity, and the ambient concentrations of OH radical, ozone, and ammonia.

## 1. INTRODUCTION

Industrially produced sulfur enters the atmosphere in reactive forms, primarily  $\text{SO}_2$  and  $\text{SO}_3$ , that cannot be treated as inert substances while they remain in the atmosphere. This report describes a stand-alone model of atmospheric transport which includes the effects of (1) rapid chemical reactions that produce new species and (2) slow reactions by which  $\text{SO}_2$  or its products from rapid reactions are oxidized to sulfates in various forms. As shall be seen, the most important consequence of these reactions is the production of an aerosol of liquid acidic droplets; hence, one object of the calculation is to estimate the amount and composition of the aerosol in a unit volume of the plume and its deposition rate. The model was developed for use as a sub-model in the Air Transport Model (ATM).<sup>1</sup>

1.1 The Rapid Reactions of Sulfur

Both gaseous  $\text{SO}_2$  and  $\text{SO}_3$  can undergo rapid chemical reactions with constituents of the atmosphere. Gaseous  $\text{SO}_2$  can dissolve in aerosol droplets producing the following products in solution:\*



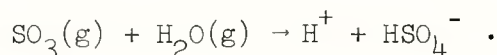
These reactions are reversible as well as rapid and as a result the concentration of the species involved are related to one another by

---

\*The abbreviations (g), (aq), and (c) used in chemical reactions denote species in the gaseous, aqueous, or solid states, respectively.

equilibrium constants (i.e.,  $K_1$ ,  $K_2$ , and  $K_3$  in Table I) that have known values<sup>2</sup> as a function of temperature. These, and the other equilibrium constants in Table I, include in their expressions a quotient of so-called activity coefficients ( $g_i$ ) which, as shall be seen, are functions of the composition of the solution involved. (The activity coefficients reflect the non-ideal behavior of the individual ions in the solution.) The tendency of  $\text{SO}_2$  to form the aqueous species  $\text{H}_2\text{SO}_3$ ,  $\text{HSO}_3^-$ , and  $\text{SO}_3^{2-}$  is weak and at the low concentrations of  $\text{SO}_2$  in the atmosphere (usually  $<0.1$  ppm) reaction occurs only if condensed water is already present.

In contrast, gaseous  $\text{SO}_3$  reacts vigorously and completely with water, even in the gaseous form, to produce acid droplets

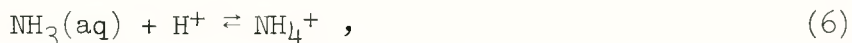


The bisulfate ion  $\text{HSO}_4^-$  undergoes rapid and reversible dissociation,



As a result, no trace of  $\text{SO}_3$  gas is found in the atmosphere. If it is introduced, it is converted completely to acid droplets containing the ions  $\text{H}^+$ ,  $\text{HSO}_4^-$  and  $\text{SO}_4^{2-}$ , the concentrations of which are related to one another by  $K_4$  (Table I).

The acid ( $\text{H}^+$ ) produced by the solution of  $\text{SO}_2$  and  $\text{SO}_3$  in water can be neutralized by atmospheric ammonia in another set of rapid and reversible reactions, producing the aqueous ammonium ion,



and  $K_5$  and  $K_6$  apply. These reactions can lead to the formation of particles of ammonium sulfate,  $(\text{NH}_4)_2\text{SO}_4$ , if they proceed far enough and the relative humidity is less than  $\sim 83\%$ . When both liquid and solid are

Table I Equilibrium Constants<sup>a</sup>

Text Symbol <sup>b</sup>	Definition <sup>b</sup>	Log K(25°C)	Log K=A+B/T(K)		Program Name
			A	B	
K <sub>1</sub>	$= \frac{[\text{H}_2\text{SO}_3]}{P_{\text{SO}_2} a_w}$	0.090	-4.4912	1365.9	EKS
K <sub>2</sub>	$= \frac{[\text{H}^+][\text{HSO}_3^-]}{[\text{H}_2\text{SO}_3]} \cdot \left( \frac{g_{\text{H}^+} g_{\text{HSO}_3^-}}{g_{\text{H}_2\text{SO}_3}} \right)$	-1.765	-4.815	909.4	EK1S
K <sub>3</sub>	$= \frac{[\text{H}^+][\text{SO}_3^{2-}]}{[\text{HSO}_3^-]} \cdot \left( \frac{g_{\text{H}^+} g_{\text{SO}_3^{2-}}}{g_{\text{HSO}_3^-}} \right)$	-7.220	-8.8549	487.36	EK2S
K <sub>4</sub>	$= \frac{[\text{H}^+][\text{SO}_4^{2-}]}{[\text{HSO}_4^-]} \cdot \left( \frac{g_{\text{H}^+} g_{\text{SO}_4^{2-}}}{g_{\text{HSO}_4^-}} \right)$	-1.994	-5.8348	1145.2	EK3S
K <sub>5</sub>	$= \frac{[\text{NH}_3]}{P_{\text{NH}_3}}$	1.756	-4.233	1785.5	EKN
K <sub>6</sub>	$= \frac{[\text{NH}_4^+]}{[\text{NH}_3][\text{H}^+]} \cdot \left( \frac{g_{\text{NH}_4^+}}{g_{\text{H}^+}} \right)$	9.251	0.1026	2727.5	EK1N

Table I (continued)

Text Symbol <sup>b</sup>	Definition <sup>b</sup>	Log K(25°C)	Log K=A+B/T(K)		Program Name
			A	B	
$K_7$	$= \frac{[\text{NH}_4^+]^2 [\text{SO}_4^{2-}]}{(\text{g}_{\text{NH}_4^+}^2 \text{g}_{\text{SO}_4^{2-}})}$	-0.221 <sup>c</sup>	0.929 <sup>c</sup>	-343.1	EKSS
$K_8$	$= \frac{[\text{H}_2\text{CO}_3]}{P_{\text{CO}_2} a_w}$	-1.462	-5.0160	1059.74	EKC
$K_9$	$= \frac{[\text{H}^+][\text{HCO}_3^-]}{[\text{H}_2\text{CO}_3]} (\text{g}_{\text{H}^+} \text{g}_{\text{HCO}_3^-})$	-6.368	-5.0300	-398.85	EK1C
$K_{10}$	$= \frac{[\text{H}^+][\text{CO}_3^{2-}]}{[\text{HCO}_3^-]} (\frac{\text{g}_{\text{H}^+} \text{g}_{\text{CO}_3^{2-}}}{\text{g}_{\text{HCO}_3^-}})$	-10.328	-7.7260	-775.85	EK2C
$K_{13}$	$= \frac{[\text{O}_3]}{P_{\text{O}_3}}$	-2.146 <sup>d</sup>	-8.488 <sup>d</sup>	1891 <sup>d</sup>	EKOZ
$K_w$	$= \frac{[\text{H}^+][\text{OH}^-]}{a_w}$	-14.00	-4.2134	-2916.54	EKW
$P_{\text{H}_2\text{O}}^\circ$	$= \frac{P_{\text{H}_2\text{O}}}{a_w}$	-1.505	6.2057	-2298.9	PWO

<sup>a</sup>Numerical values are based on NBS Technical Note 270-3, et seq<sup>2</sup> unless otherwise indicated and should represent K values with adequate accuracy from 0 to 40°C.

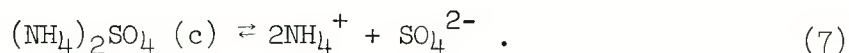
<sup>b</sup>Subscript on K refers to Eq. No. in text; concentrations of species in solution [i] are in molal units (mmol/g H<sub>2</sub>O); gas pressures P<sub>i</sub> are in atmospheres; activity coefficients of ions g<sub>i</sub> are calculated as described in Appendix C.1.

<sup>c</sup>Adjusted to give the observed solubility of (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> in water.

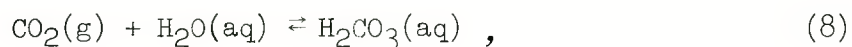
<sup>d</sup>Based on data of Rowson quoted in Seidell,<sup>17</sup> Vol. II, p 1240.



present, the following equilibrium (and  $K_7$ ) applies,



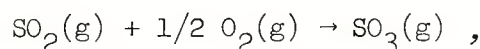
Another atmospheric gas that can affect the acidity (pH) of aerosol droplets is  $\text{CO}_2$ .



All three reactions are rapid and reversible and the concentrations of the species involved are related by  $K_8$ ,  $K_9$ ,  $K_{10}$ .

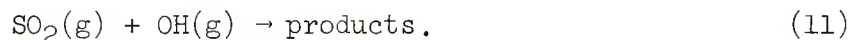
## 1.2 The Slow Oxidation of $\text{SO}_2$

The slower reactions to be considered are those by which  $\text{SO}_2$  and some of the products of its rapid reactions (all of which are species of tetravalent sulfur) are oxidized to sulfate (hexavalent sulfur) species in the atmosphere. The direct oxidation of  $\text{SO}_2$  by atmospheric oxygen



(followed instantly by reaction of  $\text{SO}_3$  with  $\text{H}_2\text{O}$ ) is slow and proceeds by a photochemical path. More rapid oxidation can occur by at least three other paths.

Recently Castleman et al<sup>3</sup> and Davis and Klauber<sup>4</sup> have concluded that the most rapid reaction involving a gaseous oxidant is that with the hydroxyl (OH) radical,



The rate expression and rate constant ( $k_{11}$ ) are given in Table II. The concentration of OH radical, an intermediate in other reactions, is found by Wang et al<sup>5</sup> to be in the range  $<2 \cdot 10^{-13}$  to  $2 \cdot 10^{-12}$  atm, and

Table II Rate Expressions for S(IV) Oxidation

Rate Expression	$k^a$	Log $k=A+B/T(K)$		Program Name
	(25°C)	A	B	
(1) Oxidation of $SO_2(g)$ by $OH(g)$ , <sup>(b)</sup> $-\frac{dP_{SO_2}}{dt} = k_{11} P_{SO_2} P_{OH}$	$k_{11} = 1.83 \cdot 10^7$	7.26	(0) <sup>e</sup>	RK4
(2) Oxidation of $HSO_3^-$ by $O_3$ , <sup>(c)</sup> $-\frac{d[HSO_3^-]}{dt} = k_{12} [HSO_3^-] g_X [O_3]$	$k_{12} = 9.8 \cdot 10^5$	14.61	-2570 <sup>e</sup>	RK3
(3) Oxidation of $SO_3^{2-}$ by $O_2$ , <sup>(d)</sup> $-\frac{d[SO_3^{2-}]}{dt} = [K_{14}' + K_{14}''([H^+] g_{H^+})^{1/2}] [SO_3^{2-}] g_{X2}$	$k_{14}' = 0.013$	11.438	-3973	RK1
	$k_{14}'' = 59$	15.095	-3973	RK2

∞

<sup>a</sup>Subscript on k refers to the Eq. No. in text. Concentration and pressure units are as in Table I, footnote b. Time is in sec.

<sup>b</sup>Castleman et al,<sup>3</sup> Davis and Klauber.<sup>4</sup>

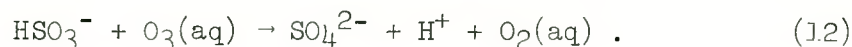
<sup>c</sup>Penkett<sup>6</sup> measured  $k_{12}$  at 9.6°C.

<sup>d</sup>McKay,<sup>7</sup> Fuller and Crist.<sup>8</sup>

<sup>e</sup>Our estimate.

hence the half-time for reaction 11 could be as short as a few hours.

Of the reactions that occur in solution, the oxidation of  $\text{HSO}_3^-$  by ozone ( $\text{O}_3$ ) may be the fastest,<sup>6</sup>

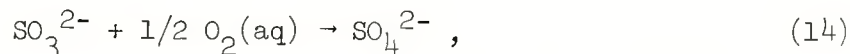


This reaction is preceeded by



With an atmospheric concentration of ozone of  $10^{-7}$  atm, we estimate from  $k_{12}$  and  $K_{13}$  (Tables I and II) a half life of 2 min. for  $\text{HSO}_3^-$  in solution.

Another reaction which is known to occur in solution involves the oxidation of  $\text{SO}_3^{2-}$  by dissolved oxygen.<sup>7,8</sup> While the overall reaction is



the mechanism is apparently complex and is not known in detail. The reaction can be catalyzed by some metal ions and retarded by others. In pure solutions the kinetics are described by the equations and constants in Table II.

### 1.3 Other Reactions

There are, of course, other reactions that could affect the behavior of sulfur in the atmosphere, including reversible equilibria such as the solution of  $\text{NO}_2$  in aerosol droplets and slow reactions such as the dissolution of a flyash particle in an acid droplet or the catalytic oxidation of  $\text{SO}_2$  at particle surfaces. There are numerous other constituents of a plume which could react with  $\text{SO}_2$  or which certainly will affect the concentrations of ozone and hydroxyl radical. Such other reactions are not included in the present model, either because they are poorly

understood at present, or because they would greatly complicate the model and its validation.

## 2. THE MODEL

From the foregoing the essential features of the chemistry of sulfur in the atmosphere which might be included in a model are (1) the slow oxidation of  $\text{SO}_2$  to produce acidic aerosol droplets in which (2) various rapidly attained and shifting chemical equilibria are established (Table I) including (3) progressive neutralization by ammonia, which--if the aerosol remains in the atmosphere long enough--can ultimately produce a relatively innocuous solid,  $(\text{NH}_4)_2\text{SO}_4$ . The computations performed by SULCAL to accomplish this may be summarized as follows (Fig. 1):\*

1. M. SULCAL, the main program, first calculates the emission rate of plume components from the source.
2. S. SIG is called to obtain the effective volume of the plume for the calculation of component concentrations at a starting point a short distance (x) from the source.
3. M. SULCAL then calls S. KUTTA<sup>9</sup> to begin the calculation of the composition of the plume at increasing values of x.
4. S. KUTTA calls S. RATE repeatedly to determine the rate of change of the concentration of the plume components with x, to be used in a Runge-Kutta integration of the component concentrations vs. x.
5. Each time S. RATE is entered it first calls S. EXPCT which determines by extrapolation or interpolation on x the starting values of several quantities to be refined in the calculation of the composition of the plume.

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\* Here and elsewhere M and S denote main and subroutine programs, respectively.

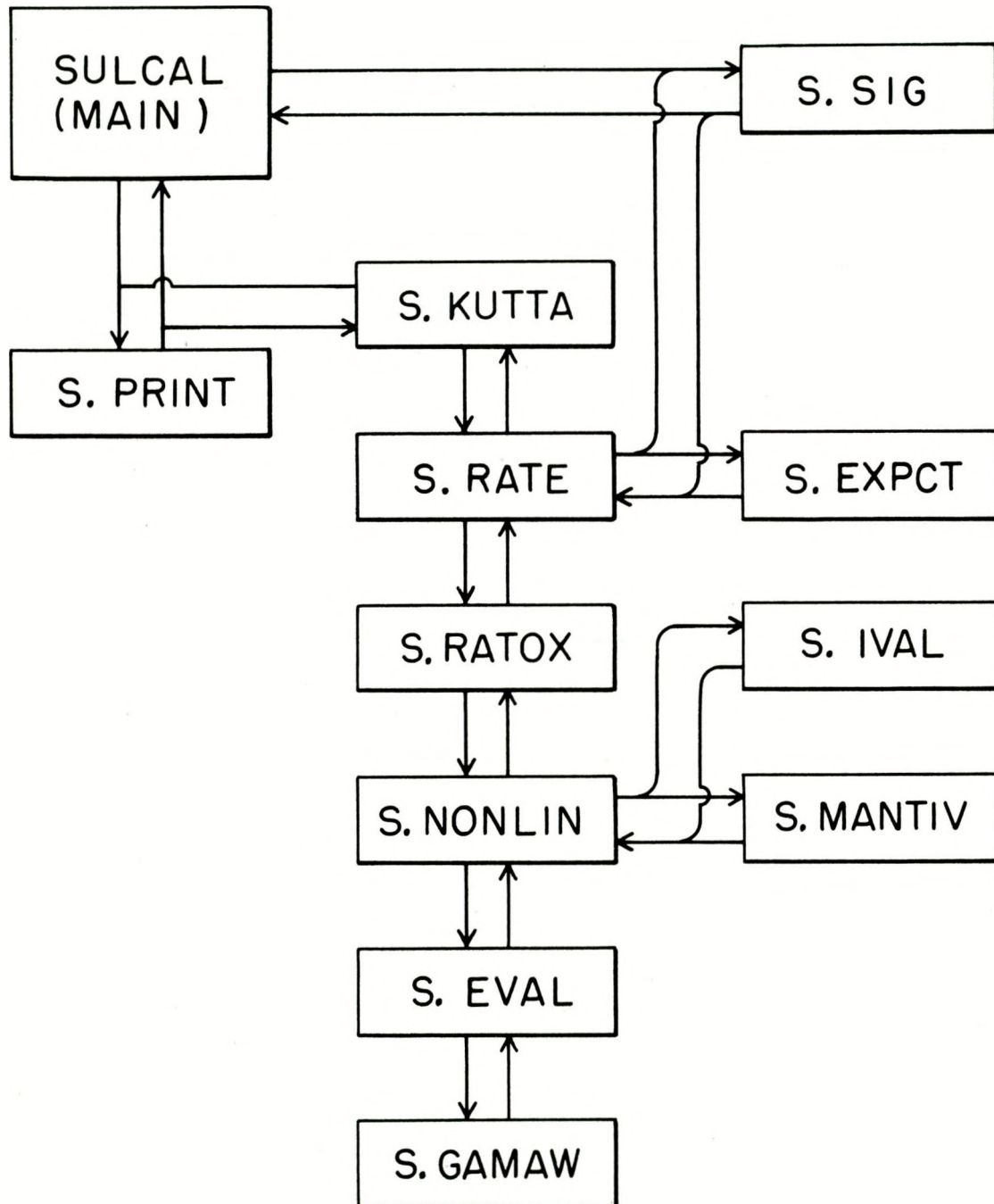


Fig. 1 Program SULCAL.

6. S. RATE next calls S. SIG to determine the effective volume of the plume at the current value of  $x$  and its rate of change with  $x$ .
7. S. RATE then calls S. RATOX to calculate the rate of S(IV) oxidation.
8. S. RATOX first calls S. NONLIN<sup>10</sup> which employs the secant method to refine quantities needed to calculate the composition of the gas and aerosol phases of the plume.
9. S. NONLIN calls S. EVAL repeatedly to calculate these compositions to test for convergence of the quantities being refined.
10. S. EVAL calls S. GAMAW repeatedly to refine values of needed ionic activity coefficients and the activity of water, quantities which are themselves dependent on the composition of the droplets.
11. With the amount and composition of each phase in the plume returned to S. RATOX, the rate of S(IV) oxidation is calculated in the gas phase and in the droplets.
12. With the oxidation rates returned to S. RATE, the rate of change of the component concentrations are calculated, including also the effects of dilution and deposition; these are returned to S. KUTTA.
13. S. KUTTA returns to M. SULCAL integrated values of the component concentrations at specified values of  $x$ .
14. S. KUTTA and M. SULCAL call S. PRINT to put out results as the calculation proceeds.

The important features of the calculation are described in the following

sections. The program is listed in Appendix A.

## 2.1 The Plume Components

There are seven chemical constituents whose concentrations specify the composition of the plume in sufficient detail for the purposes of the model. They include S(IV), S(VI),  $\text{CO}_2$ ,  $\text{H}_2\text{O}$ ,  $\text{NH}_3$ ,  $\text{O}_3$ , and OH. The first four come in part from the stack. It is assumed that S(IV) is emitted from the stack entirely as  $\text{SO}_2(\text{g})$ , and S(VI) entirely as droplets with the composition  $\text{SO}_3 \cdot \text{nH}_2\text{O}$ . The concentrations of all seven components in the surrounding air are specified as input data.

The rates ( $Q_i$ ) at which  $\text{SO}_2$ ,  $\text{SO}_3 \cdot \text{nH}_2\text{O}$ ,  $\text{CO}_2$  and  $\text{H}_2\text{O}$ --along with unreacted atmospheric gases--are emitted from the stack are calculated from input data (Table III), including the composition of the coal and the rate at which it is burned, the excess of air used, the fraction of sulfur converted to  $\text{SO}_3$ , and the moles of water ( $n$ ) condensed by a mole of  $\text{SO}_3$ . The flow rate of gas ( $V_p^\circ$ ) at the mouth of the stack is calculated from the mass flow rate and the specified exit temperature.

The total concentration ( $C_i$ ) of each of the seven components is to be calculated vs.  $x$ . A Gaussian plume is assumed for the estimation of dilution and deposition.

## 2.2 Effective Concentrations and Plume Volume

The simplest conditions under which to study chemical reactions would involve well-mixed solutions of uniform concentration and temperature. Though the conditions in a plume are far from this ideal, it is possible to define mean concentrations and temperature, and effective rate and equilibrium constants under the assumption of rapid mixing such that the net rate of oxidation to sulfur (VI) is the same



Table III Input Data

	Units	Typical Value	Text Symbol	Program Name
Carbon in coal	wt %	82.57 <sup>a</sup>		WPCTC
Hydrogen in coal	"	6.93		WPCTH
Sulfur in coal	"	1.0		WPCTS
Rate of coal combustion	kg/sec	100		QCOAL
Fraction of S converted to SO <sub>3</sub>	--	0.02		FCTS6
Fractional excess of air in combustion	--	0.25		EXAIR
Moles H <sub>2</sub> O condensed per moles SO <sub>3</sub>	--	5	n	FCON
Temperature of plume at mouth of stack	°C	177		TCP
Ambient temperature	°C	21.11		TCA
Ambient concentration of S(IV)	mmol/m <sup>3</sup>	0.004	A <sub>1</sub>	A(1)
Ambient concentration of S(VI)	"	0	A <sub>2</sub>	A(2)
Ambient concentration of O <sub>3</sub>	"	0.0032	A <sub>3</sub>	A(3)
Ambient concentration of OH	"	2·10 <sup>-8</sup>	A <sub>4</sub>	A(4)
Ambient concentration of CO <sub>2</sub>	"	12.25	A <sub>5</sub>	A(5)
Ambient concentration of NH <sub>3</sub>	"	0.0006	A <sub>6</sub>	A(6)
Ambient concentration of H <sub>2</sub> O	"	930	A <sub>7</sub>	A(7)
Effective height of stack	m	400	h	H
Meteorological stability class	--	1 to 6 <sup>b</sup>		NSC
Roughness length of terrain	--	1 to 6 <sup>b</sup>		NRL
Wind speed	m/sec	(a)	u <sub>x</sub>	UX
Deposition velocity of gas	m/sec	0.01	v <sub>d</sub> <sup>g</sup>	VDG
Deposition velocity of particle	m/sec	0.01	v <sub>d</sub> <sup>p</sup>	VDP
Initial distance from source	m	300		XNEXT
Initial trial value of Δx	m	0.1		DELX

	Units	Typical Value	Text Symbol	Program Name
Fractional error of integrated quantities	--	$3 \cdot 10^{-6}$		ERRR
No. of distance intervals	--	--		NX
Size of distance interval	m	--		XINCR (NX)

<sup>a</sup>This value is higher than is typical of power plant fuels.

<sup>b</sup>The numbers 1 to 6 denote stability classes A to F and roughness lengths of 1, 4, 10, 40, 100, and 400 m, respectively. The range of wind speeds depends on the stability class (Table IV).

for the idealized uniform case as for a real plume or for the model Gaussian distribution of concentrations. This can be achieved for reactions involving concentrations through second order in the Gaussian distributed concentrations by the artifice of an effective volume related to the Gaussian dispersion coefficients. Higher order reaction rates can be preserved if desired through appropriate scaling of the rate constants. The justification is given in Appendix B. We use the Gaussian distribution here not because we think it fits the plume distribution better than other forms, but rather because these are existing measurements of plume volume which have been expressed in terms of the Gaussian dispersion.

In terms of the effective plume volume  $V_p$ , the concentration ( $C_i$ ) of each plume component is given by

$$C_i = Q_i/V_p + A_i(1 - V_p^\circ/V_p), \quad (15)$$

where  $Q_i$  is the emission rate of component  $i$  (mmol/sec),  $A_i$  its ambient concentration (mmol/m<sup>3</sup>), and  $V_p^\circ$  the initial plume volume flow rate (m<sup>3</sup>/sec). Thus the effective plume volume is a volume flow rate (m<sup>3</sup>/sec) at a given  $x$ , or, a volume (m<sup>3</sup>) of a length of plume  $u_x$  where  $u_x$  is the wind velocity (m/sec).

The effective plume volume  $V_p$  is related to the dispersion coefficients  $\sigma_y$  and  $\sigma_z$  (m) of the Gaussian distribution by the expression (see Appendix B)

$$V_p = 4\pi \sigma_y \sigma_z u_x \beta \text{ (m}^3\text{/sec) ,} \quad (16)$$

where  $\beta$  is a correction factor which accounts for the change in the concentration distribution due to the interaction with (or "reflection"

from) the ground. The factor  $\beta$  is given by the expression (Appendix B)

$$\beta = \left[ 1 + \frac{1}{\sqrt{2}} \exp(-h^2/\sigma_z^2) \right]^{-1}, \quad (17)$$

where  $h$  is the effective stack height (m). Note that when  $\sigma_z \ll h$  then  $\beta = 1$  and when  $\sigma_z \gg h$  then  $\beta = .59$ . The value of  $\beta$  changes smoothly as the ratio  $h/\sigma_z$  passes through unity.

The dispersion coefficients used in the effective reaction volume determination are instantaneous or short-time averaged coefficients ( $\sigma_s$ ), and are smaller than the long-time averaged dispersion coefficients ( $\sigma_\ell$ ) described by Hosker<sup>11</sup> (Table IV) which are used to calculate ground level concentrations and depositions. Both coefficients are a function of the meteorological stability class, the distance  $x$  from the stack, and a roughness length characteristic of the terrain. To these dispersion coefficients we have added a small correction  $\sigma^\circ$

$$\sigma^\circ = \sqrt{V_p^\circ / 4\pi u_x}, \quad (18)$$

which keeps the plume volume finite at the source, preventing infinite concentrations and reaction rates.

The rate of change of average component concentration with distance due to surface deposition is given by

$$\frac{dC_i}{dx} = -C_i V_d \gamma / u_x, \quad (19)$$

where  $V_d$  is the conventionally defined deposition velocity and  $\gamma$  is

$$\gamma = (\sqrt{2/\pi}/\sigma_z) \exp(-h^2/2\sigma_z^2). \quad (20)$$

$\gamma$  and  $V_p$  are calculated for each specified value of  $x$  in S. SIG;  $\sigma_z$  is the conventional long-time averaged dispersion coefficient.

The short-time dispersion coefficient  $\sigma_s$  can be related to the long-

Table IV Calculation of Dispersion Coefficients<sup>a</sup>

$$\sigma_y(x) = c_3 x / \sqrt{1 + 0.001x} + \sigma^o$$

$$\sigma_z(x) = \ln(c_1 x^{d_1}) \cdot \frac{a_1 x^{b_1}}{1 + a_2 x^{b_2}} + \sigma^o$$

$$(\sigma^o = \sqrt{V_p^o / (4\pi u_x)})$$

Stability Class	$u_x$ (m/sec)	$a_1$	$b_1$	$a_2$	$b_2$	$c_3$
<sup>b</sup>						
A (1)	1 - 2.5	0.112	1.06	$5.38 \cdot 10^{-4}$	0.815	0.22
B (2)	1.5 - 5	0.130	0.950	$6.52 \cdot 10^{-4}$	0.750	0.16
C (3)	2 - >6	0.112	0.920	$9.05 \cdot 10^{-4}$	0.718	0.11
D (4)	5 - >10	0.098	0.889	$1.35 \cdot 10^{-3}$	0.688	0.08
E (5)	2 - 5	0.0609	0.895	$1.96 \cdot 10^{-3}$	0.684	0.06
F (6)	2 - 3	0.0638	0.783	$1.36 \cdot 10^{-3}$	0.672	0.04

Roughness Length (cm)		$c_1$	$d_1$
1	(1) <sup>b</sup>	1.56	0.0480
4	(2)	2.02	0.0269
10	(3)	2.718	0
40	(4)	5.16	-0.060
100	(5)	7.37	-0.0957
400	(6)	11.7	-0.128

<sup>a</sup>From Hosker<sup>11</sup>; these are long-time averaged ( $\sigma_\theta$ ) values. A factor has been omitted from the expression for  $\sigma_z$  since it differs little from unity (< 10%) for the distances of interest (< 30 km).

<sup>b</sup>Numbers are values of NSC and NRL supplied as input data (Table III).

term coefficients  $\sigma_\ell$  by

$$\sigma_s = f \sigma_\ell . \quad (21)$$

The form of  $f$  is given in Appendix B.

In equation 15, additivity of gas volumes is assumed. By the assumption of the ideal gas law, the mean absolute temperature of the plume is then given as

$$T = \frac{T_a}{1 - (1 - T_a/T_p) (V_p^\circ/V_p)} , \quad (22)$$

where  $T_p$  and  $T_a$  are respectively the absolute temperature of the plume at the mouth of the stack and of the ambient air.

### 2.3 The Rate of Change of Component Concentrations

The seven derivatives  $dC_i/dx$  are evaluated at a given  $x$  in S. RATE each time it is called by S. KUTTA. There are three possible contributions to each derivative which arise from (1) dilution, (2) deposition, and (3) oxidation,

$$\frac{d C(i)}{dx} = R_1(i) + R_2(i) + R_3(i) . \quad (23)$$

The terms due to dilution are of the form

$$R_1(i) = - [C(i) - A(i)] \frac{d \ln V_p}{dx} , \quad (24)$$

obtained by differentiation of eq 15.

For the terms arising from deposition two deposition velocities are specified  $V_d^g$  for gases and  $V_d^p$  for aerosol particles. For the components S(VI) and  $H_2O$  only particulate deposition is involved

$$R_2(S(VI)) = - C(S(VI)) \cdot V_d^p \gamma/u_x , \quad (25)$$

$$R_2(H_2O) = - (g_w/0.018) \cdot V_d^p \gamma/u_x . \quad (26)$$

(Here  $g_w$  is the amount of condensed water in  $g/m^3$ .) For the components S(IV) and  $NH_3$  the expressions are of the form

$$R_2(i) = - \left\{ [P(i)/RT] V_d^g + [C(i) - P(i)/RT] V_d^p \right\} \gamma/u_x . \quad (27)$$

$P(i)$  is the pressure of the gaseous species. For the other components ( $CO_2$ ,  $O_3$ , and  $OH$ ) the deposition rates are assumed to be zero.

The rates of oxidation of S(IV), calculated in and returned from S. RATOX are used in the following  $R_3$  terms:

$$R_3(S(IV)) = - R_T/u_x , \quad (28)$$

$$R_3(S(VI)) = R_T/u_x , \quad (29)$$

$$R_3(O_3) = -R_{OZ}/u_x , \quad (30)$$

where  $R_T$  is the total oxidation rate and  $R_{OZ}$  is the oxidation rate from reaction of  $HSO_3^-$  with ozone in the droplets (both in  $mmol\ m^{-3}\ sec^{-1}$ ). The depletion of  $OH$  by reaction with  $SO_2$  is neglected since its concentration is controlled by other reactions not included in the model.

#### 2.4 The Composition of the Plume

The remainder of the calculation is concerned mainly with determining for a given set of component concentrations the partial pressures of gases, the amount and composition of the aerosol droplets, and the amount--if any--of particulate ammonium sulfate in order that the oxidation rates of the various species of S(IV) (eqs 11, 12, and 14) can be calculated.

Given the concentrations of the components S(IV), S(VI),  $NH_3$ ,  $CO_2$ , and  $H_2O$ , the following material balance equations may be written:

$$C(S(IV)) = P_{SO_2}/RT + ([H_2SO_3] + [HSO_3^-] + [SO_3^{2-}])g_w , \quad (31)$$

$$C(S(VI)) = ([HSO_4^-] + [SO_4^{2-}])g_w + C_s , \quad (32)$$

$$C(CO_2) = P_{CO_2}/RT + ([H_2CO_3] + [HCO_3^-] + [CO_3^{2-}])g_w , \quad (33)$$

$$C(NH_3) = P_{NH_3}/RT + ([NH_3] + [NH_4^+])g_w + 2C_s , \quad (34)$$

$$C(H_2O) = P_{H_2O}/RT + g_w/0.018 . \quad (35)$$

(The pressures  $P_i$  are in atm, all solution concentrations  $[i]$  are in mmol/gH<sub>2</sub>O--i.e., molal units--and  $C_s$  is the amount of (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> solid in mmol/m<sup>3</sup>.) Upon appropriate substitution of equilibrium constants (Table I) in these equations, the unknown quantities which remain are  $[H^+]$ ,  $g_w$ ,  $C_s$ , and the ionic activity coefficients.

Only four activity coefficients are distinguished (see Appendix C.1),

$$g_{H^+}, g_{NH_4^+}, g_{X^-}, g_{X^{2-}} ,$$

the activity coefficients of all singly charged anions being assumed to be the same and similarly for doubly charged anions. The activity coefficients, and the osmotic coefficient ( $\phi$ ) needed for the calculation of the activity of water, are rather complicated but not too sensitive functions of the corresponding ion concentrations  $[H^+]$ ,  $[NH_4^+]$ ,  $[X^-]$ , and  $[X^{2-}]$  (Appendix C.1). They are refined each time S. EVAL is entered with trial values of  $[H^+]$ ,  $g_w$ , and  $C_s$  by repeated cycling between S. EVAL and S. GAMAW.

The quantities  $[H^+]$ ,  $g_w$ , and  $C_s$  (if >0) are refined by cycling between S. NONLIN and S. EVAL until the following conditions are met:



(1) Charge Balance

$$[H^+] + [NH_4^+] = [X^-] + 2[X^{2-}] \quad (36)$$

(2) Water Activity Balance

$$P_{H_2O}/P_{H_2O}^\circ = a_w \quad (37)$$

Here  $P_{H_2O}^\circ$  is the saturation partial pressure of water in the atmosphere and  $a_w$  is calculated from the composition of the solution (i.e., from  $\phi$ ). This condition specifies that the vapor pressure of the droplets as defined by their composition and the temperature will equal the partial pressure of water in the atmosphere.

(3) Ammonium Sulfate Saturation (only if  $C_s > 0$ )

$$K_7 = [NH_4^+]^2 [SO_4^{2-}] g_{NH_4^+}^2 g_{X^{2-}} \quad (38)$$

The residual deviations from these conditions are defined to equal or to approximate the fractional discrepancy in the concentration of ions in the droplets

$$y_1 = ([H^+] + [NH_4^+] - [X^-] - 2[X^{2-}]) / \Sigma[i] \quad (39)$$

$$y_2 = (P_{H_2O}^\circ / P_{H_2O} - 1/a_w) / (0.018 \cdot \Sigma[i]) \quad (40)$$

$$y_3 = \left( \frac{[NH_4^+]^2 [SO_4^{2-}] g_{NH_4^+}^2 g_{X^{2-}}}{K_7} \right)^{1/3} - 1 \quad (41)$$

This appears to give an appropriate weight to each.

Initial estimates of  $g_w$ ,  $[H^+]$ ,  $[NH_4^+]$ ,  $[X^-]$  and  $[SO_4^{2-}]$  are made the first time S. RATOX is entered as described in Appendix C.2. The initial estimate of  $C_s$  is made if and when the droplets become super-saturated with ammonium sulfate ( $y_3 > 0$ ), otherwise  $C_s = 0$ . Subsequent estimates of these quantities are made in S. EXPCT before entry into

S. RATOX (Fig. 1).

The quantities  $[H^+]$  and  $g_w$  are refined in S. NONLIN in logarithmic form since both can adopt widely ranging values and have either a first power or higher effect on the concentration of species in solution.

The quantity  $g_w$  can go to zero if the concentration ratio  $[NH_4^+]/[SO_4^{2-}]$  reaches 2 and the solution is saturated with  $(NH_4)_2SO_4$ . If significant amounts of acid remain unreacted, however, the drops cannot dry out.\*

The value of  $g_w$  is monitored and if it becomes lower than  $10^{-8} \text{ g/m}^3$  the computation of solution composition is discontinued and only the oxidation of  $SO_2$  by OH (Eq. 11) is calculated.

The quantity  $C_s$  differs from  $[H^+]$  and  $g_w$  in its effect on the calculations. When it is small it has a small effect on the composition of the droplets, but when it approaches half the amount of the  $NH_3$  component present the difference  $C(NH_3) - 2C_s$  becomes important in determining the amount of the droplets. Because of these considerations, the variable  $C_s$  is refined in S. NONLIN in the form

$$x_3 = \log[C(NH_3)/(C(NH_3) - 2C_s)] . \quad (42)$$

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\*While there is evidence that under these conditions an acid salt  $[(NH_4)_2SO_4]_3 \cdot H_2SO_4$  can form, for simplicity we consider only the formation of  $(NH_4)_2SO_4$ .

### 3. TEST CALCULATIONS

A series of calculations, summarized in Table V, were carried out to determine the sensitivity of the various output quantities to what are thought to be typical ranges of the input quantities. The underscored values are reference values. Some of the output quantities for the reference case are plotted versus the distance from the source in Fig. 2. The other output quantities in the table result by varying each input quantity from its reference value while holding all other inputs at the reference values. For example, entries in the first row were generated using 2.5 wt. % sulfur in coal, stability class C, wind speed 2 m/sec, OH concentration  $2 \cdot 10^{-8}$  mmol/m<sup>3</sup>, NH<sub>3</sub> concentration  $6 \cdot 10^{-4}$  mmol/m<sup>3</sup>, and temperature 21.11°C.

In all calculations ERRR, the error limit parameter in S. KUTTA, was set at  $3 \cdot 10^{-6}$ . In test calculations this produced output values differing by < 0.2% from the limiting values obtained as ERRR was further decreased. In the calculation of the reference case: S. KUTTA called S. RATE 1345 times in performing the integration from X = 0.3 to 25 km; S. NONLIN called S. EVAL an average of 5.7 times per step of the integration; each time S. EVAL was called, it called S. GAMAW an average of 6.1 times. The computation time for the reference case on the IBM-360-91 computer was 16 sec.

Under the conditions tested, the aerosol droplets contain mostly sulfuric acid with a relatively small amount of ammonium sulfate, the latter being limited by and equivalent to the supply of ambient ammonia. The concentration of the droplets is strongly dependent on the ambient

Table V Sensitivity of Values Calculated by SULCAL to Variation of Input Values

Input Values <sup>a,b</sup>		Output Values at 25 km From Stack <sup>c</sup>					
		SO <sub>2</sub> (10 <sup>-2</sup> mmol/m <sup>3</sup> )	Sulfate (10 <sup>-3</sup> mmol/m <sup>3</sup> )	SO <sub>2</sub> Ox. Rate (%/hr)	Liq Water (10 <sup>-3</sup> g/m <sup>3</sup> )	H <sub>2</sub> SO <sub>4</sub> in Droplets (mmol/g H <sub>2</sub> O)	(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub> in Droplets (mmol/g H <sub>2</sub> O)
Sulphur in Coal (Wt %)	2.5	14.3	19.3	3.2	7.7	2.5	0.039
	1.0	6.0	7.9	3.2	3.1	2.4	0.096
	0.5	3.2	4.0	3.2	1.6	2.4	0.189
Meteorological Stability Class	A	4.8	6.3	3.2	2.6	2.4	0.12
	C	6.0	7.9	3.2	3.1	2.4	0.096
	F	14.8	19.7	3.2	6.5	3.0	0.046
Wind Speed (m/sec)	1	9.8	25.3	3.2	9.0	2.8	0.033
	2	6.0	7.9	3.2	3.1	2.4	0.096
	4	3.5	2.5	3.2	1.0	2.2	0.29
	8	2.0	0.84	3.2	0.32	1.7	0.93
Ambient OH (10 <sup>-8</sup> mmol/m <sup>3</sup> )	10	3.9	28.2	16.0	11.2	2.5	0.027
	2	6.0	7.9	3.2	3.1	2.4	0.096
	0.5	6.4	3.0	0.8	1.2	2.3	0.26
	0.1	6.6	1.6	0.16	0.60	2.1	0.50
Ambient NH <sub>3</sub> (10 <sup>-4</sup> mmol/m <sup>3</sup> )	20	6.0	7.9	3.2	3.0	2.3	0.33
	6	6.0	7.9	3.2	3.1	2.4	0.096
Ambient Temp. (°C)	18.33	6.0	7.9	3.2	740	0.010	0.00040
	19.44	6.0	7.9	3.2	14	0.54	0.021
	21.11	6.0	7.9	3.2	3.1	2.4	0.096

Table V (continued)

Input Values <sup>a,b</sup>		Output Values at 25 km From Stack <sup>c</sup>					
		SO <sub>2</sub> (10 <sup>-2</sup> mmol/m <sup>3</sup> )	Sulfate (10 <sup>-3</sup> mmol/m <sup>3</sup> )	SO <sub>2</sub> Ox. Rate (%/hr)	Liq Water (10 <sup>-3</sup> g/m <sup>3</sup> )	H <sub>2</sub> SO <sub>4</sub> in Droplets (mmol/g H <sub>2</sub> O)	(NH <sub>4</sub> ) SO <sub>4</sub> in Droplets (mmol/g H <sub>2</sub> O)
Ambient Temp. (°C)	23.89	6.0	7.9	3.2	1.6	4.7	0.18
	26.67	6.0	7.9	3.2	1.2	6.5	0.25
	29.44	6.0	7.9	3.2	0.95	7.9	0.31

<sup>a</sup>Underscored numbers are the reference values. Each alternate value was taken while all other input values at the reference values.

<sup>b</sup>Other input values were: Wt % carbon in coal = 82.6; Wt % Hydrogen in coal 6.9%; Combustion rate = 100 Kg coal/sec; Effective stack height = 400 m; Deposition velocity = 0.01 m/sec; ambient water concentration = 930 mmol/m<sup>3</sup>.

<sup>c</sup>Concentrations are averages across the plume (See Appendix B).

ORNL - DWG. 75-17659A

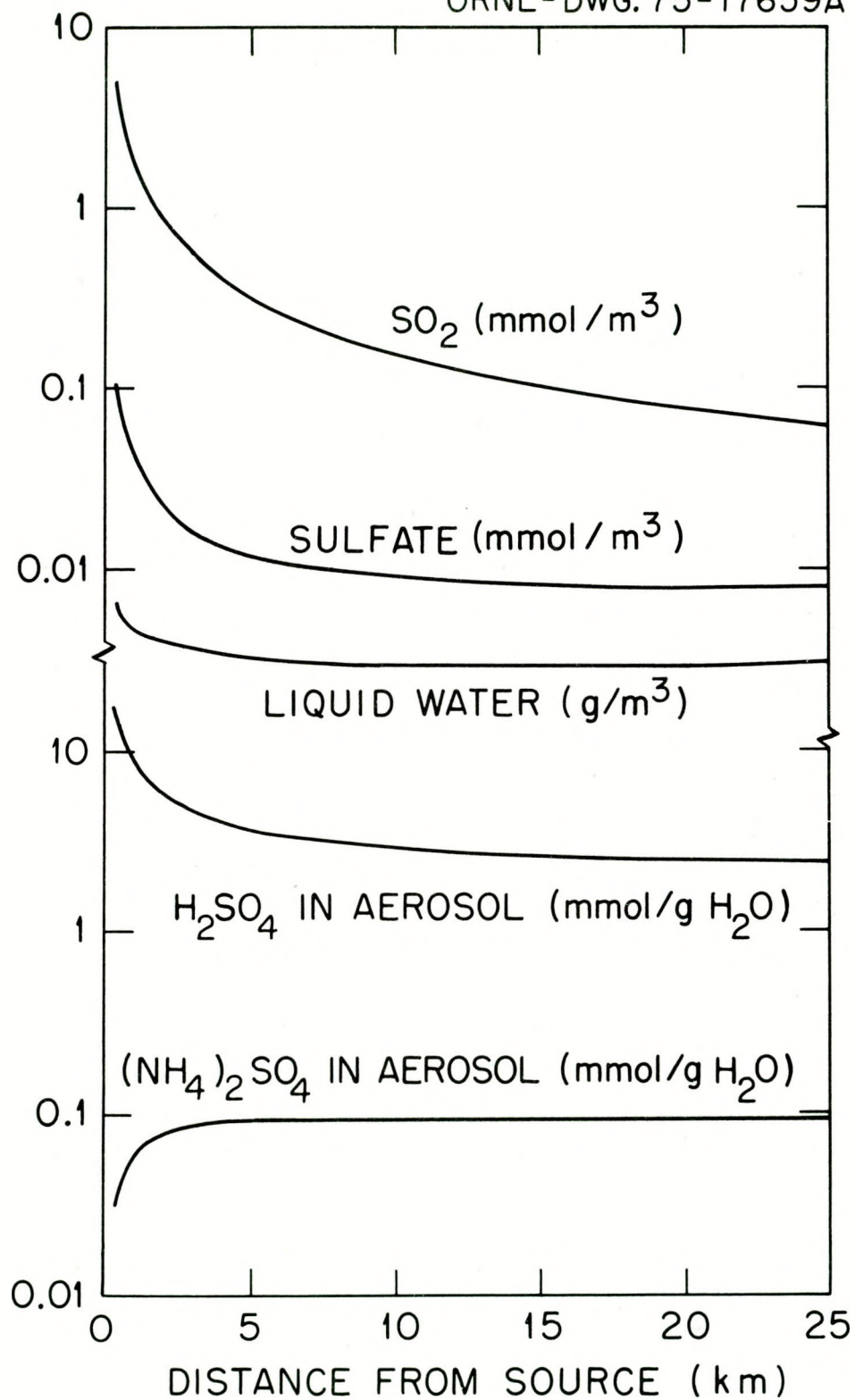


Fig. 2 Quantities calculated by SULCAL for the reference case.

temperature because of the strong inverse effect of the temperature on the relative humidity (the total concentration of water in the ambient air was not varied). The amounts of  $\text{SO}_2$ , sulfuric acid and condensed water in the plume change directly as the sulfur emission rate, which here is varied by changing the sulfur content in the coal. Increased wind speed strongly decreases the concentrations of  $\text{SO}_2$ , sulfuric acid and liquid water.

Under all the conditions tested in Table V, only the OH radical is important in the oxidation of  $\text{SO}_2$  and the percentage of  $\text{SO}_2$  oxidized per hour depends almost entirely on the ambient concentration of the OH radical. This concentration is controlled at the assumed values by unspecified chemical reactions. The conditions under which the oxidation of  $\text{SO}_2$  species in the droplets (reactions 12 and 14) becomes significant may be determined as follows: The ratio of the rate of oxidation of  $\text{HSO}_3^-$  in solution by ozone ( $R_{\text{Oz}}$ ) to the rate of oxidation of gaseous  $\text{SO}_2$  by OH radical ( $R_{\text{OH}}$ ) is given approximately by

$$\frac{R_{\text{Oz}}}{R_{\text{OH}}} = (K_1 K_2 K_{13} k_{12}/k_{11}) \cdot RT \cdot \left( \frac{g_w P_{\text{O}_3}}{[\text{H}^+] P_{\text{OH}}} \right). \quad (43)$$

The ratio of the rates in solution of  $\text{SO}_3^{2-}$  oxidation by oxygen  $R_0$  to  $\text{HSO}_3^-$  oxidation by ozone is given by

$$\frac{R_0}{R_{\text{Oz}}} = \frac{K_3}{K_{13}} \left( \frac{k_{14}' + k_{14}'' [\text{H}^+]}{k_{12}} \right) \bigg/ [\text{H}^+] P_{\text{O}_3}. \quad (44)$$

(Both these expressions are approximations because activity coefficients have been neglected.) With  $P_{\text{O}_3}$  and  $P_{\text{OH}}$  at their reference values ( $8 \cdot 10^{-8}$  and  $5 \cdot 10^{-13}$  atm, respectively) these ratios of rates at  $25^\circ\text{C}$



become

$$\frac{R_{Oz}}{R_{OH}} \approx 3 \cdot 10^{-5} \frac{g_w}{[H^+]} \quad (45)$$

and

$$\frac{R_O}{R_{OH}} \approx \frac{1 \cdot 10^{-6}}{[H^+]} + \frac{5 \cdot 10^{-3}}{[H^+]} I/2 \quad (46)$$

Thus the oxidation of  $HSO_3^-$  by ozone should become significant when the ratio  $g_w/[H^+]$  exceeds  $\sim 10^4$ . The rate of  $SO_3^{2-}$  oxidation by dissolved oxygen should then become significant when the pH exceeds  $\sim 4$ . Since  $g_w$  and  $[H^+]$  change in opposite directions and almost logarithmically with the relative humidity, the rate of oxidation of  $SO_2$  species in aerosol droplets increase rapidly in importance as the relative humidity increases (Fig. 3). This rate is also favored by conditions that produce a more complete neutralization of the acid droplets. In general, neutralization is favored by a lower sulfur emission rate and a lower OH radical concentration, both tending to reduce the rate of production of the acid  $SO_3$ . It is favored as well by a high ambient concentration of ammonia and a high wind speed, both tending to deliver more ammonia to the acid droplets.

We emphasize that, in addition to the approximations in the present model arising from the use of averaged concentrations in the plume, the oxidation mechanisms are oversimplified and incomplete. The concentration of OH radical, which may often determine the rate of the most rapid oxidation reaction, is controlled by other chemical reactions not included in the model. The species of  $SO_2$  in solution oxidized by ozone may include others in addition to (or perhaps instead of)  $HSO_3^-$ .



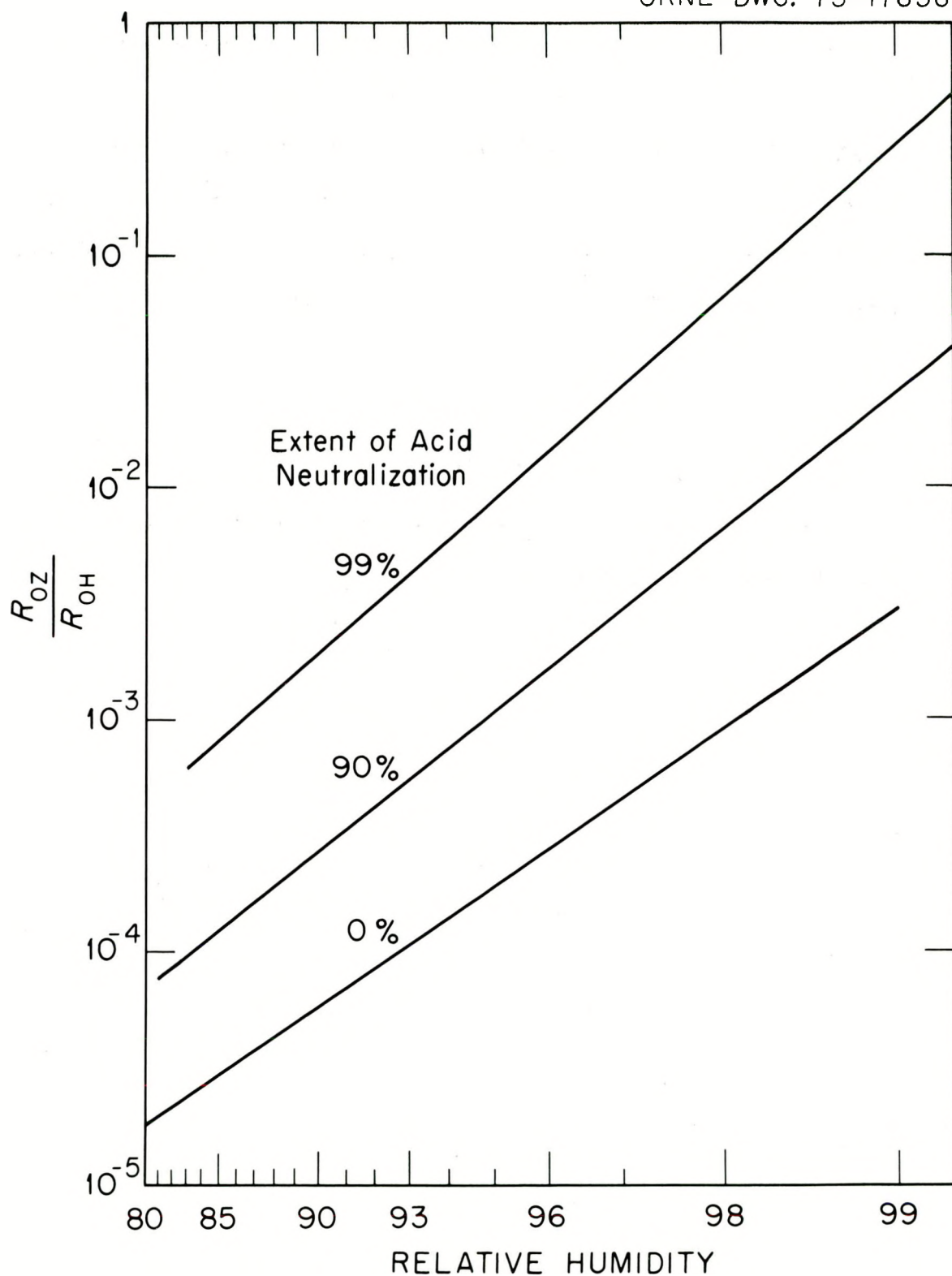


Fig. 3 The effect of the relative humidity and the extent of acid neutralization on the ratio of the rates of  $\text{SO}_2$  oxidation by ozone ( $R_{OZ}$ ) and by hydroxyl radical ( $R_{OH}$ ) at  $25^\circ\text{C}$ .

The oxidation of  $\text{SO}_3^{2-}$  by oxygen can be catalyzed or poisoned by certain metal ions in solution. Finally, possible catalytic oxidation of  $\text{SO}_2$  at the surface of solid particles and droplets has not been included. It does not seem worthwhile to include all such effects at present, nor may it be so even when sufficient knowledge has been gained. As more information does become available, however, it will be relatively easy to modify the rate expressions in the S. RATOX and to add others as may be desired.

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APPENDIX A

Program SULCAL

1. Partial List of Symbols

<u>In Text</u>	<u>Definition</u>	<u>In Program</u>
$a_w$	Activity of water	AWI
$B^{\circ}_{JK}, B^1_{JK}$	Parameters for calculating interaction coefficients between cation J and anion K $B^{\circ}_{JK} = B(1,J,K)$ $B^1_{JK} = B(2,J,K)$	B(I,J,K)
$\beta$	Correction factors (Eq. 17) for reflection of plume from ground, for short-time (BETAS) and time-averaged (BETAL) plumes	BETAS, BETAL
$C_i$ or $C(i)$	Total concentration of component; (mmol/m <sup>3</sup> )	QC(I)
$C_s$	Concentration of (NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub> particles in plume (mmol/m <sup>3</sup> )	CSS
ERRR	Error limit employed in S. KUTTA	ERRR
f	Factor relating short-time and long-time sigmas (Eq. 22, App. B)	FLS
$g_i$	Activity coefficient of an ion; $G(1,1) = g_{H^+}$ , $G(1,2) = g_{NH_4^+}$ , $G(2,1) = g_{X^-}$ , $G(2,2) = g_{X^{2-}}$	G(I,J)
$g_w$	Concentration of condensed water in plume (g H <sub>2</sub> O/m <sup>3</sup> )	CGW
$\gamma$	Factor (Eq. 20) which corrects effective plume concentration to ground level concentrations	GAMMA
$P_j$	Pressure of a gaseous constituent (atm)	P(J)
$P_w^{\circ}$	Vapor pressure of pure water (atm)	PWO
$\phi$	Osmotic coefficient of aerosol droplets (Eq. 4, App. C)	P
$Q_i$	Emission rate of component; (mmol/sec)	Q(I)
R	The gas constant (m <sup>3</sup> mmol <sup>-1</sup> °K <sup>-1</sup> )	R

<u>In Text</u>	<u>Definition</u>	<u>In Program</u>
$\frac{1}{RT}$	Factor for conversion of pressure (atm) to concentration (mmol/m <sup>3</sup> )	CONST
$\sigma_{\ell,y}^*, \sigma_{\ell,z}^*$	Time-averaged dispersion coefficients for a point source	SGPY, SGPZ
$\sigma^o$	Correction which makes $\sigma_y$ and $\sigma_z$ finite at $x = 0$ (Eq. 18)	SGO
$\sigma_{\ell,y}, \sigma_{\ell,z}$	Time-averaged dispersion coefficient (m)	SGLY, SGLZ
$\sigma_{s,y}, \sigma_{s,z}$	Short-time dispersion coefficient (m)	SGSY, SGSZ
T	Absolute temperature of plume (°K)	T
$u_x$	Wind velocity (m/sec)	UX
$V_d^g, V_d^p$	Disposition velocity for gases and particles (m/sec)	VDG, VDP
$V_p$	Effective volume (m <sup>3</sup> /sec) (the effective volume of a length of plume equal to $u_x$ , the wind velocity in m/sec). VRG is calculated using short-time dispersion coefficients and VLG is calculated with time-averaged dispersion coefficients	VRG, VLG
$V_p^g$	Gas flow rate at mouth of stack (m <sup>3</sup> /sec)	VRGO
x	Distance (m) from source in wind direction	X

## 2. Program Listing

```

C PROGRAM FOR CALCULATION OF CHEMICAL FORMS OF SULFUR IN AND THEIR
C DEPOSITION FROM A PLUME
      BLOCK DATA
      IMPLICIT REAL*8 (A-H,O-Z)
C R = GAS CONSTANT (M**3/ATM/K,MMOL)
C PI = 3.1416, FLN = LN(10), FMW = GMS H2O/MMOL
C B(I,J,K) = INTERACTION COEF'S. OF CATION I WITH ANION J, TWO
C PARAMETERS FOR EACH IJ PAIR
C Z = ION CHARGE (FIRST SUBSCRIPT IS 1 FOR CATIONS AND 2 FOR ANIONS)
      COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
      DATA B(1,1,1),B(1,1,2),B(1,2,1),B(1,2,2),B(2,1,1),
     & B(2,1,2),B(2,2,1),B(2,2,2)
     & /0.2722D+0,3.0803D+0,0.0502D+0,-4.0054D+0,0.0D+0,C.319D+0,
     & -0.04873D+0,-1.18580D+0/
      DATA PI,FLN,R,FMW/3.1415927D+0,2.302585D+0,82.0597D-9,18.0153D-3/
     & ,Z(1,1),Z(1,2),Z(2,1),Z(2,2)/1.0D+0,1.0D+0,-1.0D+0,-2.0D+0/
      END
C
C MAIN PROGRAM
      IMPLICIT REAL*8 (A-H,O-Z)
      EXTERNAL RATE
      COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
C NOX = TOTAL NUMBER OF TIMES S. RATE IS ENTERED FOR A GIVEN CASE
      COMMON/CASE/NOX
      COMMON/RTE/NSC,NRL,VRGO,TP,H,UX,A(7),TA,DELX,VDG,VDP
      COMMON/SIGMA/SGO,SGPY,SGPZ,SGLY,SGLZ,SGSY,SGSZ,BETAL,BETAS
      COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,RTO,NCAL,NITN
      COMMON/EQK/EKN,EK1N,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW
      COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,
     & PN,PC,PS
      COMMON/ACT/DHA,G(2,4),P,PW0,CONST
      COMMON/GIVN/WPCTC,WPCTH,WPCTS,QCOAL,FCTS6,EXAIR,FCON,TCP,TCA,
     & XLAST,XNEXT,ERRR
      COMMON/VAR/QC(7)
      DIMENSION Q(7),PDY(7),XINCR(50)
C WPCTI=WT. % OF ELEMENT I IN COAL
C QCOAL=COMBUSTION RATE IN KG/SEC
C FCTS6=FRACT OF S CONVERTED TO S(VI),EXAIR=FRACTION OF EXCESS AIR USED
C FCON=NO.OF MOLES OF WATER CONDENSED/MOLE SO3, TCP=TEMP OF EXIT GAS (C)
C H=APPARENT STACK HEIGHT IN METERS
      1 READ(20,100) WPCTC,WPCTH,QCOAL,FCTS6,EXAIR,FCON,TCP,H
C STABILITY CLASS,ROUGHNESS LENGTH,WIND SPEED (M/SEC)
C AND AMBIENT CONCENTRATIONS (MMOL/M**3) AND TEMPERATURE (C)
      2 READ(20,100) NSC,NRL,UX,VDG,VDP,(A(I),I=1,7),TCA,WPCTS
C CALCULATE MOLES OF GASES PER KG CF COAL BURNED
      RMLC=WPCTC/1.200115
      RMLW=0.5*WPCTH/0.100797
      RMLS4=(WPCTS/3.2064)*(1.0-FCTS6)
      RMLS6=(WPCTS/3.2064)*FCTS6
C CALCULATE MOLES OF OXYGEN CONSUMED PER KG OF COAL BURNED
      CMLO2=RMLC+0.5*RMLW+RMLS4+1.5*RMLS6
C CALCULATE TOTAL MOLES OF EFFLUENT GAS PER KG CF COAL EURNED
      RMLGAS=RMLC+RMLW+RMLS4+((79.05/20.95)*(1.0+EXAIR)+EXAIR)*CMLO2
     & -FCON*RMLS6
      TP=TCP+273.15
      TA=TCA+273.15
C CALCULATE RATE OF EMISSION OF GAS (M**3/SEC)
      VRGO=RMLGAS*QCOAL*(82.0597D-6)*TP

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C CALCULATE EMISSION RATES OF COMPONENTS IN MMOLS/SEC
  BNFACT=1000.0*QCOAL
  Q(1)=RMLS4*BNFACT
  Q(2)=RMLS6*BNFACT
  Q(3)=0.0
  Q(4)=0.0
  Q(5)=RMLC*BNFACT
  Q(6)=0.0
  Q(7)=RMIW*BNFACT
  NOX=1
C READ FIRST VALUE OF X (*), DELTA X AND ERROR LIMIT
  READ(20,100) XNEXT,DELX,ERRR
C READ NO. OF STEPS IN X AND THE SIZE OF EACH
  READ(20,100) NX,(XINCR(I),I=1,NX)
  NCAL=0
  NITN=0
C CALCULATE CONCENTRATIONS AT STARTING VALUE OF X (MMOLS/M**3)
  CALL SIG(XNEXT,VRG,PEX)
  DO 10 I=1,7
    QC(I)=Q(I)/VRG+A(I)*(1.0-VRG0/VRG)
  10 CONTINUE
  CALL PRINT(1)
C START CALCULATION OF QC(I) VS. X
  DO 20 I=1,NX
    XLAST=XNEXT
    XNEXT=XNEXT+XINCR(I)
    CALL KUTTA(XLAST,XNEXT,QC,7,DELX,ERRR,40,FATE)
  20 CONTINUE
C CALL RATE FOR FINAL CALCULATION OF PLUME COMPOSITION
  CALL RATE(XNEXT,QC,FDMY,1)
  XLAST=XNEXT
  CALL PRINT(2)
  GO TO 2
100 FORMAT(7G10.0)
END

C
C SUBROUTINE TO CALCULATE EFFECTIVE VOLUME OF PLUME (VRG)
C AND THE DEPOSITION FACTOR (PEXD)
C CALCULATION OF SIGMA(Y) AND SIGMA(Z) ACCORDING TO HOSKER
  SUBROUTINE SIG(X,VRG,PEXD)
    IMPLICIT REAL*8 (A-H,O-Z)
    DIMENSION GC3(6),GA1(6),GB1(6),GA2(6),GB2(6),GC1(6),GD1(6),FCTL(6)
    COMMON/CASE/NOX
    COMMON/RTE/NSC,NRL,VRG0,TP,H,UX,A(7),TA,DELX,VDG,VDP
    COMMON/SIGMA/SG0,SGPY,SGPZ,SGLY,SGLZ,SGSY,SGSZ,BETAL,BETAS
    DATA GC3/0.22D+0,0.16D+0,0.11D+0,0.08D+0,0.06D+0,
      80.04D+0/
    DATA GA1/0.112D+0,0.130D+0,0.112D+0,0.098D+0,
      80.0609D+0,0.0638D+0/
    DATA GB1/1.06D+0,0.950D+0,0.920D+0,0.889D+0,0.895D+0,
      80.783D+0/
    DATA GA2/0.000538D+0,0.000652D+0,0.000905D+0,
      80.00135D+0,0.00196D+0,0.00136D+0/
    DATA GB2/0.815D+0,0.750D+0,0.718D+0,0.688D+0,0.684D+0,
      80.672D+0/
    DATA GC1/1.56D+0,2.02D+0,2.71828D+0,5.16D+0,7.37D+0,
      811.7D+0/
    DATA GD1/0.048CD+0,0.0269D+0,0.0D+0,-0.06D+0,-0.0957D+0

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E, -0.128D+0/	SIG
DATA PCTL/5.0D+2, 2.5D+2, 1.0D+2, 0.5D+2, 0.2D+2, 0.1D+2/	SIG
C UPON FIRST ENTRY ASSIGN VALUES TO PARAMETERS THAT DEPEND ON STABILITY	SIG
C CLASS, ROUGHNESS LENGTH, AND WIND SPEED	SIG
IF (NOX.GT.1) GO TO 1	SIG
PI=3.1415927	SIG
C3=GC3 (NSC)	SIG
A1=GA1 (NSC)	SIG
B1=GB1 (NSC)	SIG
A2=GA2 (NSC)	SIG
B2=GB2 (NSC)	SIG
C1=GC1 (NRL)	SIG
D1=GD1 (NRL)	SIG
RL=FCTL (NSC)	SIG
PLS=DSQRT (X/(X+50.0*RL))	SIG
SG0=DSQRT (VRG0/(4.0*PI*UX))	SIG
C CALCULATE POINT-SOURCE TIME-AVERAGED SIGMA (Y)	SIG
1 DTY=1.0+0.001*X	SIG
SGPY=C3*X/DSQRT (DTY)	SIG
F=DLOG (C1*X**D1)	SIG
G=A1*X**B1/(1.0+A2*X**B2)	SIG
C CALCULATE POINT-SOURCE TIME-AVERAGED SIGMA (Z)	SIG
SGPZ=F*G	SIG
C CALCULATE TIME-AVERAGED SIGMA (Y) AND SIGMA (Z) FOR REAL SOURCE	SIG
SGLY=SGPY+SG0	SIG
SGLZ=SGPZ+SG0	SIG
C CALCULATE SHORT-TIME SIGMA (Y) AND SIGMA (Z) FOR REAL SOURCE	SIG
SGSY=PLS*SGPY+SG0	SIG
SGSZ=PLS*SGPZ+SG0	SIG
C EVALUATE CORRECTION FACTOR BETA	SIG
EXFS=-(H/SGSZ)**2	SIG
IF (EXFS.LT.-25.0) EXFS=-25.0	SIG
EXFL=-(H/SGLZ)**2	SIG
IF (EXFL.LT.-25.0) EXFL=-25.0	SIG
BETAL=1.0/(1.0+0.70711*DEXP (EXFL))	SIG
BETAS=1.0/(1.0+0.70711*DEXP (EXFS))	SIG
C CALCULATE EFFECTIVE REACTION VOLUME VRG (M**3)	SIG
C AND EFFECTIVE DEPOSITION VOLUME VLG (M**3)	SIG
VRG=4.0*PI*SGSY*SGSZ*UX*BETAS	SIG
VLG=4.0*PI*SGLY*SGLZ*UX*BETAL	SIG
C CALCULATE FACTOR RELATED TO DRY DEPOSITION	SIG
GAMA=DSQRT (2.0/PI) /SGLZ*DEXP (EXFL/2.0)	SIG
FXD=(VRG/VLG)*GAMA	SIG
RETURN	SIG
END	SIG
C	
C SUBROUTINE CALLED BY KUTTA TO CALCULATE RATE OF CHANGE OF COMPONENT	RATE
C CONCENTRATIONS YI(I) WITH DISTANCE X (DYDX)	RATE
C IN = INTEGER WHICH (WHEN 1) SIGNALS CALCULATION OF DYDX AT X	RATE
SUBROUTINE RATE (X, YI, DYDX, IN)	RATE
IMPLICIT REAL*8 (A-H, O-Z)	RATE
DIMENSION YI (7), DYDX (7)	RATE
COMMON/BLK/R, PI, PLN, FMW, B (4, 4, 2), Z (2, 4)	RATE
COMMON/CASE/NOX	RATE
COMMON/RTE/NSC, NRL, VRG0, TP, H, UX, A (7), TA, DELX, VDG, VDP	RATE
COMMON/ARG/CGW, PH, CSS, C (2, 4), T, RTT, RTOH, RTOZ, RTO, NCAL, NITN	RATE
COMMON/CON/CMH, CMN, CM1N, CHC, CM1C, CM2C, CHS, CM1S, CM2S, CM3S, CM4S,	RATE
EPN, PC, PS	RATE

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COMMON/ACT/DHA,G(2,4),P,PW0,CONST
COMMON/VAR/Y(7)
DO 10 I=1,7
  Y(I)=YI(I)
10 CONTINUE
C HAVE DROPLETS DRIED UP?
  IF(CGW.LE.0.0) GO TO 2
C ESTIMATE STARTING VALUFS OF CGW,PH,CSS, AND ACTIVITY COEFFICIENTS
  CALL EXPCT(NOX,IN,X,CGW,PH,CSS,C(1,2),C(2,1),C(2,2))
C EVALUATE DLN(V)/D(X) NUMERICALLY
  X1=X+DELX/2.0
  X2=X-DELX/2.0
  CALL SIG(X,VRG,FEXD)
  CALL SIG(X1,VRG1,FEX)
  CALL SIG(X2,VRG2,FEX)
  DLVDX=2.0*(VRG1-VRG2)/(DELX*(VRG1+VRG2))
C CALCULATE TEMPERATURE IN PLUME (K)
  T=TA/(1.0-(1.0-TA/TP)*VRG0/VRG)
  CONST=1.0/(R*T)
C CALCULATE RATE OF OXIDATION OF S(IV)
  CALL RATOX
  NOX=NOX+1
C CALCULATE DYDX(I)
  DYDX(1)=-(RTT+(PS*CONST*(VDG-VDP)+Y(1)*VDP)*FEXD)/UX
  DYDX(2)=(RTT-Y(2)*VDP*FEXD)/UX
  DYDX(3)=-RTOZ/UX
  DYDX(4)=0.0
  DYDX(5)=0.0
  DYDX(6)=-(PN*CONST*(VDG-VDP)+Y(6)*VDP)*FEXD/UX
  DYDX(7)=-CGW/FMW*VDP*FEXD/UX
  DO 1 I=1,7
    DYDX(I)=DYDX(I)-(Y(I)-A(I))*DLVDX
  1 CONTINUE
  RETURN
  END
C
C SUBROUTINE TO CALCULATE RATE OF S(IV) OXIDATION
  SUBROUTINE RATOX
    IMPLICIT REAL*8(A-H,O-Z)
    EXTERNAL IVAL
    EXTERNAL EVAL
    DIMENSION X(10)
    COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
    COMMON/CASE/NOX
    COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTOH,RTOZ,PTO,NCAL,NITN
    COMMON/EQK/FKN,EK1N,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW
    COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,
    &PN,PC,PS
    COMMON/ACT/DHA,G(2,4),P,PW0,CONST
    COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW
C CONSTANTS FOR THE CALCULATION OF RATE AND EQUILIBRIUM CONSTANTS
C AS A FUNCTION OF THE TEMPERATURE
    DATA EKNO,EKNT,EKCO,EKCT,EKS0,EKST,EK1N0,EK1NT,
    &EK1S0,EK1ST,EK2S0,EK2ST,EK1C0,EK1CT,EK2C0,EK2CT,EK3S0,
    &EK3ST,EKSS0,EKSSS,EKW0,EKWT,EKOZ0,EKOZT,RK10,RK1T,RK20,RK2T,
    &RK30,RK3T,RK40,RK4T,EPW0,EPWT
    E/-4.233D+0,1.7855D+0,-5.016D+0,1.05974D+0,-4.4912D+0,1.3659D+0,
    &0.1026D+0,2.7275D+0,-4.815D+0,0.9094D+0,-8.8549D+0,0.48736D+0,

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8-5.030D+0,-0.39885D+0,-7.7260D+0,-5.77585D+0,-5.8348D+0,1.1452D+0,	RTOX
80.929D+0,-0.3431D+0,-4.2134D+0,-2.91654D+0,-8.4888D+0,1.891D+0,	RTOX
811.4378D+0,-3.9725D+0,15.0947D+0,-3.9725D+0,14.61D+0,-2.57D+0,	RTOX
82.65D+0,0.0D+0,6.2057D+0,-2.2989D+0/	RTOX
C ALL CONCENTRATIONS ARE IN MMOL/L/METER**3 OR MMOL/G	RTCX
C CALCULATE TEMPERATURE DEPENDENT PARAMETERS	RTOX
PT=1000.0/T	RTOX
DHA=(4.91D-1+8.0D-4*(T-273.15))*PLN	RTOX
EKN=DEXP((EKN0+EKN*PT)*PLN)	RTOX
EKC=DEXP((EKC0+EKC*PT)*PLN)	RTOX
EKS=DEXP((EKS0+EKS*PT)*PLN)	RTCX
EK1N=DEXP((EK1N0+EK1N*PT)*PLN)	RTOX
EK1S=DEXP((EK1S0+EK1S*PT)*PLN)	RTOX
EK2S=DEXP((EK2S0+EK2S*PT)*PLN)	RTCX
EK1C=DEXP((EK1C0+EK1C*PT)*PLN)	RTOX
EK2C=DEXP((EK2C0+EK2C*PT)*PLN)	RTOX
EK3S=DEXP((EK3S0+EK3S*PT)*PLN)	RTCX
EKSS=DEXP((EKSS0+EKSS*PT)*PLN)	RTOX
FKW=DEXP((FKW0+FKW*PT)*PLN)	RTOX
EKOZ=DEXP((EKOZ0+EKOZ*PT)*PLN)	RTCX
RK1=DEXP((RK10+RK1*PT)*PLN)	RTOX
RK2=DEXP((RK20+RK2*PT)*PLN)	RTOX
RK3=DEXP((RK30+RK3*PT)*PLN)	RTCX
RK4=DEXP((RK40+RK4*PT)*PLN)	RTOX
PW0=DEXP((EPW0+EPW*PT)*PLN)	RTOX
C TEST TO SEE IF CGW AND PH ARE GIVEN	RTOX
IF(NOX.GT.1) GO TO 1	RTOX
CSS=0.0	RTOX
C ESTIMATE STARTING VALUES OF CGW, ION CON'NS AND GAMAS	RTOX
PNI=3.0	RTOX
TB=(PW0*CONST-CW+PNI*C6)/2.0	RTOX
TC=-PNI*C6*CW	RTOX
CGW=PMW*(-TB+DSQRT(TB**2-TC))	RTOX
X(1)=DLOG10(CGW)	RTOX
G(1,1)=1.0	RTOX
G(2,1)=1.0	RTOX
G(2,2)=1.0	RTCX
CALL NONLIN(X,1,ICONV,IVAL)	RTOX
CGW=DEXP(X(1)*PLN)	RTOX
PH=-DLOG10(C(1,1))	RTOX
C BEGIN REFINEMENT	RTOX
C HAVE DROPLETS DRIED OUT?	RTOX
1 IF(CGW.LE.0.0) GO TO 10	RTOX
X(1)=DLOG10(CGW)	RTOX
X(2)=PH	RTOX
C ARE DROPLETS UNSATURATED?	RTOX
IF(CSS.LE.0.0) GO TO 2	RTOX
C IS ENOUGH ACID PRESENT TO PREVENT DRYING OUT OF DROPLETS?	RTOX
IF(CN/C6.LT.2.0) GO TO 4	RTOX
10 CGW=0.0	RTOX
CSS=C6	RTOX
RTO=0.0	RTOX
RTOZ=0.0	RTOX
PS=C4/CONST	RTOX
PN=(CN-2.0*CSS)/CONST	RTOX
PC=CC/CONST	RTOX
GO TO 11	RTCX
2 CSS=0.0	RTOX

```

CALL NONLIN(X,2,ICONV,EVAL)
C IS SOLUTION SUPERSATURATED WITH (NH4)2SO4?
  GFS=G(1,2)**2*G(2,2)
  DSS=CM1N**2*CM4S*GFS-FKSS
  IF(DSS.LT.0.0) GO TO 6
C IS PNOUGH ACID PRESENT TO PREVENT DRYING OUT OF DROPLETS?
  IF((CN/C6).LT.2.0.AND.CGW.GT.1.0D-8) GO TO 3
  GO TO 10
C ESTIMATE AMOUNT OF SOLID (NH4)2SO4
  3 CSS=DSS*CGW/((CM1N+4.0*CM4S)*CM1N*GFS)
  IF((CN-2.0*CSS).GT.0.0) GO TO 4
  CSS=0.4*CN
  4 X(3)=DLOG10(CN/(CN-2.0*CSS))
  CALL NONLIN(X,3,ICONV,EVAL)
C HAS SOLUTION BECOME UNSATURATED WITH (NH4)2SO4?
  IF(X(3).LE.0.0) GO TO 2
C HAVE DROPLETS DRIED UP?
  IF(CGW.LE.1.0D-8) GO TO 10
  CSS=CN/2.0*(1.0-DEXP(-X(3)*FLN))
  6 CGW=DEXP(X(1)*FLN)
  PH=X(2)
  CMH=DEXP(-PH*FIN)
C BEGIN CALCULATION OF RATE OF OXIDATION OF S(IV)
C RATE OF OXIDATION OF SO32- BY OXYGEN IN SOLUTION (MMOL/M**3/SEC)
  RQ1=RK1*G(2,2)
  RQ2=RK2*G(1,1)**0.5*G(2,2)
  RTO=(RQ1+RQ2*CMH**0.5)*CM2S*CGW
C RATE OF OXIDATION OF HSO3- BY OZONE IN SOLUTION (MMOL/M**3/SEC)
  RQ3=RK3*G(2,1)
  CMOZ=EKOZ*COZ/(CONST+EKOZ*CGW)
  RTOZ=RQ3*CMOZ*CM1S*CGW
C RATE OF OXIDATION OF SO2 BY OH RADICAL IN GAS PHASE (MMOL/M**3/SEC)
  11 RTOH=RK4*COH*PS*CONST
C TOTAL OXIDATION RATE OF S(IV) (MMOL/METER**3/SEC)
  RTT=RTOH+RTOZ+RTO
  RETURN
END

C
C SUBROUTINE TO REFINE CGW, PH, CSS BY SECANT METHOD
  SUBROUTINE NONLIN(XIN,N,ICONV,EVAL)
C PROGRAM AUTHORS R. E. FUNDERLIC AND J. RINZEL,
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C
C CTC ORD PROGRAM NO. 9066
  IMPLICIT REAL*8 (A-H,O-Z)
  LOGICAL WRT
  COMMON/HONON/WRT
  DIMENSION XIN(N),X(20,21),XBAR(20,2),A(21,21),AINV(21,21),P(21),
  1B(21),XNORM(21),BNORM(20)
  WRT=.FALSE.
  ICONV=0
  NP1=N+1
  DO 110 J=1,NP1
  DO 120 I=1,N
  IF(J.NE.1) GO TO 130
  X(I,J)=XIN(I)
  GO TO 120

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130	X(I,J)=XIN(I)	NLIN
	IF(J-1.EQ.I) X(I,J)=0.005+X(I,J)	NLIN
120	CONTINUE	NLIN
110	CONTINUE	NLIN
	DO 1 J=1, NP1	NLIN
	B(J)=1.	NLIN
	A(NP1,J)=1.	NLIN
1	CALL EVAL(X(1,J),A(1,J))	NLIN
	DO 2 I=1, NP1	NLIN
	DO 2 J=1, NP1	NLIN
	AINV(I,J)=0.000	NLIN
	IF(I.EQ.J) AINV(I,J)=1.000	NLIN
2	CONTINUE	NLIN
	DO 4 J=1, NP1	NLIN
	SUM=0.	NLIN
	DO 3 I=1, N	NLIN
3	SUM=SUM+DABS(A(I,J))	NLIN
4	XNORM(J)=SUM	NLIN
	CALL MATINV(A,AINV,NP1)	NLIN
	DO 5 I=1, N	NLIN
5	XBAR(I,1)=X(I,NP1)	NLIN
	ITER=40	NLIN
	DO 15 K=1, ITER	NLIN
	XBNUM=0.	NLIN
	XB DEN=0.	NLIN
	DO 7 I=1, N	NLIN
	XB=0.	NLIN
	DO 6 J=1, NP1	NLIN
6	XB=XB + AINV(J,NP1)*X(I,J)	NLIN
	XBAR(I,2)=XB	NLIN
	XB NUM=XB NUM + (XBAR(I,2)-XBAR(I,1))**2	NLIN
7	XB DEN=XB DEN + XBAR(I,2)**2	NLIN
	ICONV=ICONV+1	NLIN
	IF((XB NUM/XB DEN).LT.1.D-14) GO TO 16	NLIN
	DO 8 I=1, N	NLIN
8	XBAR(I,1)=XBAR(I,2)	NLIN
	BIGNOR=0.	NLIN
	DO 9 J=1, NP1	NLIN
	BNORM(J)=XNORM(J)*DABS(B(J))	NLIN
	IF(BNORM(J).LE.BIGNOR) GO TO 9	NLIN
	BIGNOR=BNORM(J)	NLIN
	JCOL = J	NLIN
9	CONTINUE	NLIN
	CALL EVAL(XBAR,P)	NLIN
	P(NP1)=1.	NLIN
	XNORM(JCOL)=0.	NLIN
	DO 10 I=1, N	NLIN
	X(I,JCOL)=XBAR(I,1)	NLIN
10	XNORM(JCOL)=XNORM(JCOL)+DABS(P(I))	NLIN
	DO 11 I=1, NP1	NLIN
	B(I)=0.	NLIN
	DO 11 J=1, NP1	NLIN
11	B(I)=B(I)+AINV(I,J)*P(J)	NLIN
	DO 13 I=1, NP1	NLIN
	IF(I.EQ.JCOL) GO TO 13	NLIN
	DO 12 J=1, NP1	NLIN
12	AINV(I,J)=AINV(I,J)-AINV(JCOL,J)*(B(I)/B(JCOL))	NLIN
13	CONTINUE	NLIN



DO 14 I=1,NP1	NLIN
14 AINV(JCOL,I)=AINV(JCOL,I)/B(JCOL)	NLIN
15 CONTINUE	NLIN
WRITE(20,101)	NLIN
101 FORMAT('1*****THE FOLLOWING CASE DID NOT CONVERGE*****')	NLIN
16 WRT=.TRUE.	NLIN
CALL EVAL(XBAR,P)	NLIN
DO 17 I=1,N	NLIN
17 XIN(I)=XBAR(I,2)	NLIN
RETURN	NLIN
END	NLIN
SUBROUTINE MATINV(A,B,III)	MINV
IMPLICIT REAL*8 (A-H,C-Z)	MINV
DIMENSION A(21,21), B(21,21)	MINV
KK=III	MINV
NV=III	MINV
KKM=KK-1	MINV
DO9I=1,KKM	MINV
S=0.0	MINV
DO1J=I,KK	MINV
R=DABS(A(J,I))	MINV
IF(R.LT.S) GO TO 1	MINV
S=R	MINV
L=J	MINV
1 CONTINUE	MINV
IF (L.EQ.I) GO TO 5	MINV
IF(L-I) 2,5,2	MINV
2 DO3J=I,KK	MINV
S=A(I,J)	MINV
A(I,J)=A(L,J)	MINV
3 A(L,J)=S	MINV
DO4J=1,NV	MINV
S=B(I,J)	MINV
B(I,J)=B(L,J)	MINV
4 B(L,J)=S	MINV
5 IF(A(I,I).EQ.0.) GO TO 9	MINV
IPO=I+1	MINV
DO8J=IPO,KK	MINV
IF(A(J,I).EQ.0.) GO TO 8	MINV
S=A(J,I)/A(I,I)	MINV
A(J,I)=0.0	MINV
DO6K=IPO,KK	MINV
6 A(J,K)=A(J,K)-A(I,K)*S	MINV
DO7K=1,NV	MINV
7 B(J,K)=B(J,K)-B(I,K)*S	MINV
8 CONTINUE	MINV
9 CONTINUE	MINV
KMO=KK-1	MINV
DO11K=1,NV	MINV
E(KK,K)=B(KK,K)/A(KK,KK)	MINV
DO11I=1,KMO	MINV
N=KK-I	MINV
DO10J=N,KMO	MINV
10 B(N,K)=B(N,K)-A(N,J+1)*B(J+1,K)	MINV
11 B(N,K)=B(N,K)/A(N,N)	MINV
RETURN	MINV
END	MINV

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C
C SUBROUTINE CALLED BY S NONLIN TO CALCULATE ION CONCENTRATIONS C(I),
C ACTIVITY COEFFICIENTS G(I), AND THE RESIDUAL IN THE WATER
C ACTIVITY BALANCE Y(1) FOR THE CURRENT VALUE OF CGW AS GIVEN BY X(1)
SUBROUTINE IVAL(X,Y)
IMPLICIT REAL*8 (A-H,O-Z)
DIMENSION X(10),Y(21)
COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTCH,RTOZ,RTO,NCAL,NITN
COMMON/EQK/EKN,EK1N,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW
COMMON/ACT/DHA,G(2,4),P,PW0,CONST
COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW
CGW=DEXP(X(1)*FLN)
GHS=G(1,1)*G(2,2)/G(2,1)
QHS=EK3S/GHS
C BEGIN REFINEMENT OF ACTIVITY COEFFICIENTS AND ION CONCENTRATIONS
C FOR CURRENT VALUE OF X(1)
DO 10 I=1,100
BT=(QHS+(CN-C6)/CGW)/2.0
CT=QHS*(CN-2.0*C6)/CGW
C(1,1)=-BT+DSQRT(BT**2-CT)
C(1,2)=CN/CGW
TERM=C6/(CGW*(C(1,1)+QHS))
C(2,1)=C(1,1)*TERM
C(2,2)=QHS*TERM
QL=QHS
CALL GAMAW(DHA,2,2,B,C,Z,G,P,AWI)
GHS=G(1,1)*G(2,2)/G(2,1)
QHS=EK3S/GHS
IF(((QL-QHS)/QHS)**2.1E-12) GO TO 20
10 CONTINUE
20 RAW=(CW-CGW/PMW)/(PW0*CONST)
C CALCULATE RESIDUAL IN WATER ACTIVITY BALANCE
Y(1)=(AWI-RAW)*PMW*(C(1,1)+C(1,2)+C(2,1)+C(2,2))
RETURN
END
C
C SUBROUTINE CALLED BY NONLIN, CALCULATES ION CONCENTRATIONS,
C ACTIVITY COEFFICIENTS AND RESIDUALS FOR CURRENT VALUES OF
C CGW,PH AND CSS AS GIVEN BY X(1),X(2), AND X(3)
SUBROUTINE EVAL(X,Y)
IMPLICIT REAL*8 (A-H,O-Z)
DIMENSION X(10),Y(21)
COMMON/BLK/R,PI,FLN,FMW,B(4,4,2),Z(2,4)
COMMON/CASE/NOX
COMMON/ARG/CGW,PH,CSS,C(2,4),T,RTT,RTCH,RTOZ,RTO,NCAL,NITN
COMMON/EQK/EKN,EK1N,EKC,EK1C,EK2C,EKS,EK1S,EK2S,EK3S,EKSS,EKW
COMMON/CON/CMH,CMN,CM1N,CMC,CM1C,CM2C,CMS,CM1S,CM2S,CM3S,CM4S,
&PN,PC,PS
COMMON/ACT/DHA,G(2,4),P,PW0,CONST
COMMON/VAR/C4,C6,COZ,COH,CC,CN,CW
IF(CSS.EQ.0.0) GO TO 10
CSS=CN/2.0*(1.0-DEXP(-X(3)*FLN))
10 CMH=DEXP(-X(2)*FLN)
CGW=DEXP(X(1)*FLN)
C(1,1)=CMH
C BEGIN REFINEMENT OF ACTIVITY COEFFICIENTS
DO 1 J=1,100

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CALL GAMAW(DHA,2,2,B,C,Z,G,P,AWI)
GQ1=G(2,1)*G(1,1)
GQ2=G(2,1)/(G(2,2)*G(1,1))
EQ1N=EK1N*G(1,1)/G(1,2)
EQC=EKC*AWI
EQ1C=EK1C/GQ1
EQ2C=EK2C*GQ2
EQS=EKS*AWI
EQ1S=EK1S/GQ1
EQ2S=EK2S*GQ2
EQ3S=EK3S*GQ2
EQSS=EKSS/(G(1,2)**2*G(2,2))
EQW=EKW*AWI/GQ1
CMHSQ=CMH*CMH
PN=(CN-2.0*CSS)/(CONST+EKN*(1.0+EQ1N*CMH)*CGW)
CMN=EKN*PN
CM1N=EQ1N*CMN*CMH
PC=CC/(CONST+EQC*(1.0+EQ1C/CMH+EQ1C*EQ2C/CMHSQ)*CGW)
CMC=EQC*PC
CM1C=EQ1C*CMC/CMH
CM2C=EQ2C*CM1C/CMH
PS=C4/(CONST+EQS*(1.0+EQ1S/CMH+EQ2S*EQ1S/CMHSQ)*CGW)
CMS=EQS*PS
CM1S=EQ1S*CMS/CMH
CM2S=EQ2S*CM1S/CMH
CMOH=EQW/CMH
CM3S=(C6-CSS)/((1.0+EQ3S/CMH)*CGW)
CM4S=EQ3S*CM3S/CMH
CMA1=CM1C+CM1S+CM3S+CMOH
CMA2=CM2C+CM2S+CM4S
DLSQI=((C(1,2)-CM1N)/CM1N)**2+((C(2,1)-CMA1)/CMA1)**2
S+((C(2,2)-CMA2)/CMA2)**2
C(1,2)=CM1N
C(2,1)=CMA1
C(2,2)=CMA2
IF(DLSQI.LT.1.0D-12) GO TO 2
1 CONTINUE
C CALCULATE RESIDUALS IN CHARGE BALANCE Y(1), WATER ACTIVITY BALANCE
C Y(2), AND (NH4)2SO4 SOLUBILITY BALANCE Y(3) (IF APPLICABLE)
2 SUMI=C(1,1)+C(1,2)+C(2,1)+C(2,2)
Y(1)=(C(1,1)+C(1,2)-C(2,1)-2.0*C(2,2))/SUMI
SUMN=CMN+CMS+CMC
AWI=DEXP(-FMW*SUMI*P)
CRAW=1.0/AWI+FMW*SUMN
ARAW=PHO*CONST/(CW-CGW/FMW)
Y(2)=(ARAW-CRAW)/(FMW*SUMI)
IF(CSS.EQ.0.0) GO TO 20
AQSS=CM1N**2*CM4S
Y(3)=(AQSS/EQSS)**(1.0/3.0)-1.0
20 NCAL=NCAL+1
NITN=NITN+J
RETURN
END
C
C SUBROUTINE TO CALCULATE ACTIVITY COEFFICIENTS AND ACTIVITY OF WATER
C DHA = DEBYE-HUCKEL CONSTANT, NC = NO. OF CATIONS
C NA = NO. OF ANIONS, B = INTERACTION COEFFICIENTS,
C C = ION CONCENTRATIONS (MMOL/G H2O), Z = ION CHARGE

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IMPLICIT REAL*8(A-H,O-Z)	XPCT
DIMENSION Y(6),YL(6),YLL(6),DYDX(6)	XPCT
IF(NOX.EQ.1) GO TO 211	XPCT
IF(NOX.EQ.2) GO TO 121	XPCT
IF(IN.EQ.0.AND.IX.EQ.1.AND.ID.EQ.1) GO TO 111	XPCT
IF(IN.EQ.1.AND.ID.EQ.1) GO TO 311	XPCT
IF(IN.EQ.0.AND.IX.EQ.3) GO TO 132	XPCT
IF(IN.EQ.0.AND.IX.EQ.1.AND.ID.EQ.2) GO TO 112	XPCT
IF(IN.EQ.1.AND.ID.EQ.2) GO TO 312	XPCT
C IX=1 DO NOT SAVE X, IX=2 SAVE X, IX=3 SAVE X AND LAST X	XPCT
C IY=1 DO NOT SAVE Y, IY=2 SAVE Y, IY=3 SAVE Y AND LAST Y AND	XPCT
C CALCULATE VALUES OF DYDX	XPCT
C ID=1 DO NOT CHANGE Y, ID=2 INTERPOLATE OR EXTRAPOLATE Y	XPCT
C FIRST ENTRY, SAVE X	XPCT
211 IX=2	XPCT
IY=1	XPCT
ID=1	XPCT
GO TO 2	XPCT
C SECOND ENTRY, SAVE Y VALUES FROM PREVIOUS CALCULATION	XPCT
121 IX=1	XPCT
IY=2	XPCT
ID=1	XPCT
GO TO 5	XPCT
C NOX=3,4,OR 5, RETURN CURRENT VALUES OF Y	XPCT
111 IX=1	XPCT
IY=1	XPCT
ID=1	XPCT
RETURN	XPCT
C FIRST CALL FROM KUTTA IN SECOND INCREMENT, SAVE X AND XL	XPCT
311 IX=3	XPCT
IY=1	XPCT
ID=1	XPCT
GO TO 3	XPCT
C SECOND CALL FROM KUTTA IN AN INCREMENT AFTER FIRST, SAVE Y AND YL,	XPCT
C CALCULATE DYDX AND RETURN NEW VALUES OF Y	XPCT
132 IX=1	XPCT
IY=3	XPCT
ID=2	XPCT
GO TO 6	XPCT
C 3RD, 4TH, OR 5TH CALL FROM KUTTA IN AN INCREMENT AFTER FIRST,	XPCT
C RETURN NEW Y VALUES	XPCT
112 IX=1	XPCT
IY=1	XPCT
ID=2	XPCT
GO TO 8	XPCT
C FIRST CALL FROM KUTTA IN AN INCREMENT AFTER SECOND,	XPCT
C SAVE X AND XL, AND RETURN NEW Y VALUES	XPCT
312 IX=3	XPCT
IY=1	XPCT
ID=2	XPCT
3 XLL=XL	XPCT
2 XL=X	XPCT
IF(ID.EQ.2) GO TO 9	XPCT
RETURN	XPCT
5 DO 10 I=1,6	XPCT
10 YL(I)=Y(I)	XPCT
RETURN	XPCT
6 DO 20 I=1,6	XPCT

[illegible]

[illegible][illegible]

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100 FORMAT(1H1,'SU IFUR CHEMISTRY IN A GAUSSIAN PLUME'////
      &1X,' WPCTC WPCTH WPCTS QCCAL FCTS6 EXAIR',
      &' FCON H TCP NSC NRL UX VDG',
      &' VDP TCA'/
      &1X,F7.2,8F8.2,2I4,4F8.2/)
125 FORMAT(
      &1X,' AC4 AC6 AOZ AOH ACC ACN',
      &' ACW'/1X,1P7E9.2///)
150 FORMAT(1X,' ERRR SG0 '/1X,1PE8.1,0PP8.2/)
300 FORMAT(
      &21X,' X DELX SGPY SGLY SGSY SGPZ',
      &' SGLZ SGSZ'/21X,F7.1,7F8.2,///
      &21X,' CGW PH CSS T NOX NCL NITN'/
      &21X,1PE9.2,0P2F8.4,F8.2,2I6,I7//
      &21X,' RTT RTOH RTOZ RTO GH',
      &' GN GHS GS'/21X,1P4E9.2,0P4F8.4/)
325 FORMAT(
      &11X,' C4 C6 COZ CCH CC CN',
      &' CW'/11X,1P7E9.2/)
340 FORMAT(
      &11X,' PN PC PS'/11X,1P3E8.1/)
348 FORMAT(1X,'DROPLETS HAVE DRIED UP, LAST SCLN. COMP. WAS: '/')
350 POPMAT(
      &1X,' CMN CM1N CMC CM1C CM2C',
      &' CMS CM1S CM2S CM3S CM4S'/
      &1X,1P10E12.4//)
      RETURN
      END

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### 3. Sample Input and Output

The following input values, used in the reference case (Table V), were read by the indicated READ statements in M. SULCAL:

```
[READ(20, 100)  WPCTC,WPCTH,QCOAL,FCTS6,EXAIR,FCON,TCP,H]
82.57      6.93      100.0      0.02      0.25      5.0      177.0
400.0

[READ(20, 100)  NSC,NRL,UX,VDG,VDP,(A(I), I=1, 7),TCA,WPCTS]
      3      5      2.0      0.01      0.01      0.004      0.0
0.0032  2.0E-8  12.26  6.0E-4  929.7  21.11  1.0

[READ(20, 100)  XNEXT,DELX,ERRR]
300.0      0.1  0.000003

[READ(20, 100)  NX,(XINCR(I), I=1,NX)]
      4      700.0  4000.0  5000.0  15000.0

[100 FORMAT (7G 10.0)]
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The output produced is shown on the next two pages. The first three rows of values, mostly input, are independent of  $x$ . In row 2 are the ambient concentrations of the seven components  $S(IV)$ ,  $S(VI)$ ,  $O_3$ ,  $OH$ ,  $CO_2$ ,  $NH_3$ , and  $H_2O$ . Rows 4 through 9 are repeated for each specified value of  $x$ . Row 4 includes the current value of the increment in  $x$  (DELX) in S. KUTTA. In row 5, NOX, NCL and NITN are, respectively, the accumulated entries into S. RATE, S. EVAL and S. GAMAW. Row 6 includes the oxidation rates: total, by  $OH$ , by  $O_3$ , and by oxygen; and the activity coefficients of the aqueous ions  $H^+$ ,  $NH_4^+$ ,  $X^-$ , and  $X^{2-}$ . In row 7 are the effective concentrations of components in the plume. Row 8 includes the partial pressures of  $NH_3$ ,  $CO_2$  and  $SO_2$  gases. Finally in row 9 are the concentrations of the aqueous species  $NH_3$ ,  $NH_4^+$ ,  $H_2CO_3$ ,  $HCO_3^-$ ,  $CO_3^{2-}$ ,  $H_2SO_3$ ,  $HSO_3^-$ ,  $SO_3^{2-}$ ,  $HSO_4^-$  and  $SO_4^{2-}$ .



SULFUR CHEMISTRY IN A GAUSSIAN PLUME

WPCTC	WPCTH	WPCTS	QCOAL	FCTSG	EXAIR	PCON	H	PCP	NSC	NRL	UX	VDG	VDP	TCA
82.57	6.93	1.00	100.00	0.02	0.25	5.00	400.00	177.00	3	5	2.00	0.01	0.01	21.11

AC4	AC6	AOZ	AOH	ACC	ACN	ACW
4.00D-03	0.0	3.20D-03	2.00D-08	1.23D 01	6.00D-04	9.30D 02

FRRR	SGO
3.0D-06	8.84

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
300.0	0.10	28.94	37.78	15.73	29.31	38.15	15.82

CGW	PH	CSS	T	NOX	NCL	NITN
6.39D-03	-1.2005	0.0	330.21	2	5	24

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
3.00D-05	3.00D-05	5.27D-16	1.32D-18	49.3771	0.3923	59.5242	0.0157

C4	C6	COZ	COH	CC	CN	CW
4.89D 00	9.98D-02	2.19D-03	1.37D-08	1.11D 03	4.11D-04	1.19D 03

PN	PC	PS
9.4D-15	3.0D-02	1.3D-04

CMN	CM1N	CMC	CM1C	CM2C	CM5	CM1S	CM2S	CM3S	CM4S
1.3990D-13	6.4383D-02	3.6043D-05	1.0670D-15	4.3346D-25	1.0746D-05	2.0027D-12	4.0459D-19	1.5294D 01	3.1787D-01

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
1000.0	12.80	77.41	86.25	27.26	75.91	84.75	26.90

CGW	PH	CSS	T	NOX	NCL	NITN
4.47D-03	-0.9559	0.0	305.47	595	2973	16466

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
1.32D-05	1.32D-05	1.41D-14	2.28D-17	6.0477	0.4091	9.8348	0.0165

C4	C6	COZ	COH	CC	CN	CW
1.65D 00	3.79D-02	2.86D-03	1.79D-08	3.82D 02	5.36D-04	1.02D 03

PN	PC	PS
2.0D-14	9.6D-03	4.1D-05

CMN	CM1N	CMC	CM1C	CM2C	CM5	CM1S	CM2S	CM3S	CM4S
8.3464D-13	1.2014D-01	1.4469D-04	1.2400D-13	7.3285D-23	2.1002D-05	5.6631D-10	3.3968D-16	7.7839D 00	6.9631D-01

X	DELX	SGPY	SGLY	SGSY	SGPZ	SGLZ	SGSZ
5000.0	102.40	223.19	232.03	61.94	236.07	244.92	65.00

CGW	PH	CSS	T	NOX	NCL	NITN
3.20D-03	-0.6334	0.0	296.23	1160	5805	35040

RTT	RTOH	RTOZ	RTO	GH	GN	GHS	GS
2.60D-06	2.60D-06	3.84D-14	9.02D-17	1.4106	0.4276	2.7777	0.0196

C4	C6	COZ	COH	CC	CN	CW
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2.97D-01 1.20D-02 3.14D-03 1.96D-08 7.93D 01 5.89D-04 3.45D 02

PN PC PS  
1.0D-13 1.9D-03 7.2D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S  
6.3393D-12 1.8356D-01 5.7870D-05 1.4439D-12 1.5260D-21 7.8317D-06 8.3611D-09 1.2068D-14 2.9836D 00 7.4973D-01

X DELX SGPY SGLY SGSY SGPZ SGLZ SGSZ  
10000.0 204.80 329.80 338.65 87.30 355.37 364.22 93.39

CGW PH CSS T NOX NCL NITN  
3.04D-03 -0.5419 0.0 295.23 1340 6705 40674

RTT RTOH RTOZ RTO GH GN GHS GS  
1.28D-06 1.28D-06 2.97D-14 7.95D-17 1.1468 0.4362 2.2439 0.0211

C4 C6 COZ COH CC CN CW  
1.45D-01 9.11D-03 3.17D-03 1.98D-08 4.54D 01 5.93D-04 9.37D 02

PN PC PS  
1.5D-13 1.1D-03 3.5D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S  
9.7110D-12 1.9500D-01 3.5642D-05 1.6543D-12 1.9528D-21 4.1551D-06 8.5414D-09 1.4235D-14 2.3121D 00 6.8261D-01

X DELX SGPY SGLY SGSY SGPZ SGLZ SGSZ  
25000.0 819.20 534.71 543.56 136.05 552.64 561.49 140.32

CGW PH CSS T NOX NCL NITN  
3.10D-03 -0.4734 0.0 294.67 1540 7707 46631

RTT RTOH RTOZ RTO GH GN GHS GS  
5.30D-07 5.30D-07 1.73D-14 5.09D-17 1.0142 0.4435 1.9694 0.0224

C4 C6 COZ COH CC CN CW  
5.95D-02 7.87D-03 3.19D-03 1.99D-08 2.63D 01 5.94D-04 9.33D 02

PN PC PS  
1.8D-13 6.4D-04 1.4D-06

CMN CM1N CMC CM1C CM2C CMS CM1S CM2S CM3S CM4S  
1.2339D-11 1.9167D-01 2.1560D-05 1.5005D-12 1.9100D-21 1.7838D-06 5.6055D-09 1.0265D-14 1.9130D 00 6.2662D-01



APPENDIX B

EFFECTIVE VOLUME OF FREE PLUME

The reactions in a plume are complicated by the situation that the reaction rates vary across and down the plume because of variations in concentration, temperature and mixing. The detailed treatment of these varying conditions is beyond the scope of this investigation. We want to use an idealized model for this study and yet relate the result to the conditions which are obtained in the open environment. The dispersion of effluents in the atmosphere has been successfully modeled by a Gaussian distribution in which the variance or dispersion  $\sigma$  follows empirically determined relations that are functions of time, distance from source, and the prevailing meteorological and topographical conditions. We will show that when modeling reaction rates it is possible to use a simple uniform model which can yield the same reaction rates as a well mixed Gaussian distribution.

In the following treatment, let  $q_i$  be the total amount of reactant  $i$ , in a unit length of the plume in the  $x$  direction, let  $C_i$  be the density or concentration of reactant  $i$ , and let  $R$  be the total rate of production of the reaction product. Consider a binary reaction

$$\frac{\text{Rate}}{\text{Volume}} = k C_1^{\alpha_1} C_2^{\alpha_2}, \quad (1)$$

then

$$R = k \int_V C_1^{\alpha_1} C_2^{\alpha_2} dv. \quad (2)$$

Let

$$C_i = \frac{q_i}{2\pi\sigma_y\sigma_z} e^{-y^2/2\sigma_y^2} e^{-z^2/2\sigma_z^2}. \quad (3)$$

Then the rate of reaction in a volume of unit length in the x-direction and infinite in extent in the y and z direction is

$$\begin{aligned}
 R &= \int_{-\infty}^{\infty} dy \int_{-\infty}^{\infty} dz \, k \, C_1^{\alpha_1} C_2^{\alpha_2} \\
 &= \frac{k \, q_1^{\alpha_1} q_2^{\alpha_2}}{2\pi\sigma_y\sigma_z} \int_{-\infty}^{\infty} dy \int_{-\infty}^{\infty} dz \, e^{-(\alpha_1+\alpha_2)y^2/2\sigma_y^2} e^{-(\alpha_1+\alpha_2)z^2/2\sigma_z^2} \\
 &= \frac{k \, q_1^{\alpha_1} q_2^{\alpha_2}}{(2\pi\sigma_y\sigma_z)^{\alpha_1+\alpha_2}} (2\pi\sigma_y\sigma_z) \cdot \frac{1}{\alpha_1+\alpha_2} \\
 &\equiv \frac{k \, q_1^{\alpha_1} q_2^{\alpha_2}}{(4\pi\sigma_y\sigma_z)^{\alpha_1+\alpha_2-1}} (4\pi\sigma_y\sigma_z) \left( \frac{2^{\alpha_1+\alpha_2-1}}{\alpha_1+\alpha_2} \right), \quad (4)
 \end{aligned}$$

where we have assumed a uniform or effective temperature. Now suppose the same amounts of reactants were spread uniformly through a volume  $V$  such that

$$\bar{C}_i = q_i/V. \quad (5)$$

Then

$$\bar{R} = k \int_V \bar{C}_1^{\alpha_1} \bar{C}_2^{\alpha_2} dv = \frac{k q_1^{\alpha_1} q_2^{\alpha_2}}{V^{\alpha_1+\alpha_2}} \cdot V. \quad (6)$$

Now if we choose as a control volume  $V = 4\pi\sigma_y\sigma_z$  then

$$\bar{R} = \frac{\alpha_1+\alpha_2}{2^{\alpha_1+\alpha_2-1}} R = F R. \quad (7)$$

The choice  $V = 4\pi\sigma_y\sigma_z$  yields reaction rates which are correct through second order. Note this includes the case  $\alpha_1 = 0$  with  $\alpha_2 = 1$  or 2.

Other cases can be handled by scaling the rate constants by the relation

$$\bar{k} = k \cdot \frac{2^{\alpha_1 + \alpha_2 - 1}}{\alpha_1 + \alpha_2} , \quad (8)$$

where  $\bar{k}$  is the effective rate constant. Similar results may be obtained for equilibria, and they may be extended to reactions and equilibria of the form  $k C_1^{\alpha_1} C_2^{\alpha_2} \dots C_n^{\alpha_n}$  by replacing  $(\alpha_1 + \alpha_2)$  in the equations above by  $(\alpha_1 + \alpha_2 + \dots + \alpha_n)$ .

#### 1. Mean Concentration When Plume Touches the Ground

The forgoing discussion assumes a free plume and neglects interactions with the surface. An effect of the ground is to limit the downward spread of the plume and to restrict the dilution of the effluents near the land surface. This has been successfully modeled by assuming an "image plume", i.e., a "reflection" of the plume at the ground. The effect of the "image plume" is included in the present treatment as follows:

Consider the total reaction rate

$$R = k \int_V C_1 C_2 \, dv , \quad (9)$$

where the concentrations  $C_i$  have the distribution

$$C_i(x, y, z) = \frac{q_i}{2\pi\sigma_y\sigma_z} e^{-y^2/2\sigma_y^2} \left[ e^{-(z-h)^2/2\sigma_z^2} + e^{-(z+h)^2/2\sigma_z^2} \right] , \quad (10)$$

where  $q_1$  is the total amount of material available to react in a unit length of plume. Then

$$\begin{aligned}
 R &= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \int_{-\infty}^{\infty} e^{-2y^2/2\sigma_y^2} dy \int_{-\infty}^{\infty} \left( e^{-(z-h)^2/2\sigma_z^2} + e^{-(z+h)^2/2\sigma_z^2} \right)^2 dz \\
 &= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \cdot \sqrt{2\pi} \sigma_y \int_0^{\infty} dz \left[ e^{-2(z-h)^2/2\sigma_z^2} \right. \\
 &\quad \left. + e^{-(z+h)^2/2\sigma_z^2} + 2e^{-\frac{(z-h)^2}{2\sigma_z^2} - \frac{(z+h)^2}{2\sigma_z^2}} \right] \\
 &= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \left[ \frac{2\pi\sigma_y \sigma_z}{2} + \sqrt{2\pi} \sigma_y \int_0^{\infty} e^{-(z^2-2hz+h^2+z^2+2hz+h^2)/2\sigma_z^2} dz \right] \\
 &= \frac{kq_1 q_2}{(2\pi\sigma_y \sigma_z)^2} \left[ \frac{2\pi\sigma_y \sigma_z}{2} + \sqrt{2\pi} \sigma_y e^{-h^2/\sigma_z^2} \int_0^{\infty} e^{-2z^2/2\sigma_z^2} dz \right] \\
 &= \frac{kq_1 q_2}{(4\pi\sigma_y \sigma_z)^2} \cdot (4\pi\sigma_y \sigma_z) \left[ 1 + \frac{e^{-h^2/\sigma_z^2}}{\sqrt{2}} \right] = k \left( \frac{q_1}{\beta V} \right) \left( \frac{q_2}{\beta V} \right) (\beta V) \beta \left[ 1 + \frac{e^{-h^2/\sigma_z^2}}{\sqrt{2}} \right],
 \end{aligned}$$

where  $V = 4\pi\sigma_y\sigma_z$ . To give the same reaction rate assuming  $\tau = Q/V$  we must choose

$$\beta = \frac{1}{1 + \exp(-h^2/\sigma_z^2)/\sqrt{2}} \quad (12)$$

Note for point source,  $\beta(x=0) = 1$ ,  $\beta(x=\infty) = \frac{\sqrt{2}}{1+\sqrt{2}} = .586$ .

## 2. Plume Depletion Due to Washout

According to the washout concept, the depletion from the box of the chemical model is expressed in terms of the washout coefficient  $\Lambda$  as,

$$\Delta m = \Lambda m \Delta t, \quad (13)$$

where  $m$  is the amount of (aerosol or water soluble) material in the box.

But in terms of the volume  $V$  of the box,  $C = m/V$ , hence

$$\left(\frac{dC}{dt}\right)_{\text{washout}} = -\Lambda C. \quad (14)$$

Washout occurs under conditions of higher than normal ambient humidity.

Washout is not included in the present model.

## 3. Plume Depletion Due to Dry Deposition

The ground level concentration according to the Gaussian plume model is given by

$$C(x,y,0) = \frac{Q(x)}{2\pi\sigma_y\sigma_z u x} e^{-y^2/2\sigma_y^2} e^{-h^2/2\sigma_z^2}. \quad (15)$$



In terms of the concentration  $\bar{C}$  of the chemical box model

$$\bar{C}(x) = Q(x)/4\pi\sigma_y'\sigma_z'u_x, \quad (16)$$

or

$$C(x,y,0) = 4\bar{C} e^{-y^2/2\sigma_y'^2} e^{-h^2/2\sigma_z'^2}. \quad (17)$$

The dry deposition in time  $\Delta t$  at a distance  $x$  is

$$\begin{aligned} \Delta m &= \int_{-\infty}^{\infty} dy \int_{x-1/2}^{x+1/2} dx C(x,y,0) V_d \Delta t = 4\bar{C} V_d \Delta t \sqrt{2\pi} \sigma_y u_x e^{-h^2/2\sigma_z'^2} \\ &= - \left( \frac{d\bar{C}}{dt} \right)_{\text{dep}} (4\pi\sigma_y'\sigma_z'u_x) \Delta t, \end{aligned} \quad (18)$$

where  $V_d$  is the deposition "velocity".

Hence

$$\left( \frac{dC}{dt} \right)_{\text{dep}} = - \bar{C} 2V_d \left( \frac{e^{-h^2/2\sigma_z'^2}}{\sqrt{2\pi} \sigma_z} \right). \quad (19)$$

#### 4. Choice of Sigma

The behavior of a natural plume is far more complicated than the idealized case we have considered so far. Over a period of time convection cells move, the plume may wander, fan, fumigate, etc., so that the volume swept out is considerably larger than that which obtains in any instant. It is the larger volume which is described by the dispersion coefficients of Pasquill-Gifford and Briggs-Smith, which are

defined for a point source. For the shorter distances at least, the instantaneous plume will be considerably smaller. At very short distances a "concentration catastrophe" arises if the plume dimensions are allowed to shrink to zero. Two artifices are used to avoid this crisis. One is the concept of a "virtual (point) source" behind the stack, i.e. the distances along the plume are measured from a point behind the stack and can not go to zero. The second artifice (which we use) involves adding a constant such that  $4\pi\sigma_y\sigma_z u_x$  reduces to the initial plume volume flow rate  $V_p^\circ$  at  $x = 0$ . At small distances the additive term  $\sigma^\circ$  dominates while at large distances it can be neglected.

In defining the various dispersion coefficients below we use an asterisk to denote a point source, a subscript  $\ell$  for long-time (larger)  $\sigma$ 's and a subscript  $s$  for short-time (smaller)  $\sigma$ 's. We assume that large scale effects such as deposition are governed by these coefficients given by

$$\sigma_\ell = \sigma_\ell^* + \sigma^\circ . \quad (20)$$

The instantaneous dispersion coefficients used in determining the size of the "box" we define by

$$\sigma_s = f\sigma_\ell^* + \sigma^\circ , \quad (21)$$

where  $f$  is a function of the distance  $x$ ,  $f < 1$  and  $f$  approaches unity at large distances. The exact form of  $f$  is not crucial and we have elected to use a form due to Scriven,<sup>12</sup> which we write in the form

$$f = \frac{x}{\sqrt{x+50L}} , \quad (22)$$

where L is given in the table below.

<u>Stability Class</u>	<u>L (meters)</u>
A	500
B	250*
C	100
D	50*
E	20*
F	10

---

\*Interpolated values.



APPENDIX C

ACTIVITY COEFFICIENTS, THE WATER ACTIVITY, AND INITIAL VALUES

### 1. The Estimation of Ion Activity Coefficients and the Osmotic Coefficient

For an ideal solution, the activity coefficients of ions are unity and the activity of water is given by its mole fraction,

$$a_w = \frac{n_w}{n_w + \sum_i n_i + \sum_j n_j} = \frac{55.51}{55.51 + \sum_i [i] + \sum_j [j]} \quad (1)$$

(here i and j refer to cations and anions, respectively). Such behavior is approached as solutions become more and more dilute. The aerosol droplets in the atmosphere often are quite concentrated, however, and the deviations from ideal behavior can be considerable (Fig.4). Hence it is necessary to estimate the values of ion activity coefficients for use in the equilibrium constant expressions (Table I) and in the calculation of the activity of water. The following is based on the treatments of Pitzer and Brewer<sup>13</sup> and Pitzer.<sup>14</sup>

The activity coefficients of ions  $g_i$  and  $g_j$  in a solution and the activity of water ( $a_w$ ) are related by the Gibbs-Duhem equation,

$$\sum_i [i] d \ln([i]g_i) + \sum_j [j] d \ln([j]g_j) = -55.51 d \ln a_w. \quad (2)$$

The activity of water is usually expressed in terms of the osmotic coefficient  $\phi$ ,

$$\ln a_w = -0.018(\sum_i [i] + \sum_j [j])\phi. \quad (3)$$

The relationship between the osmotic coefficient and the activity coefficient is obtained by combining Eqs. 2 and 3 and integrating,

$$\phi = 1 + \frac{\sum_i [i] d \ln g_i + \sum_j [j] d \ln g_j}{\sum_i [i] + \sum_j [j]} \quad (4)$$

The integrations are from 0 concentration to the individual ion concentrations in the solution for which  $\phi$  is desired.

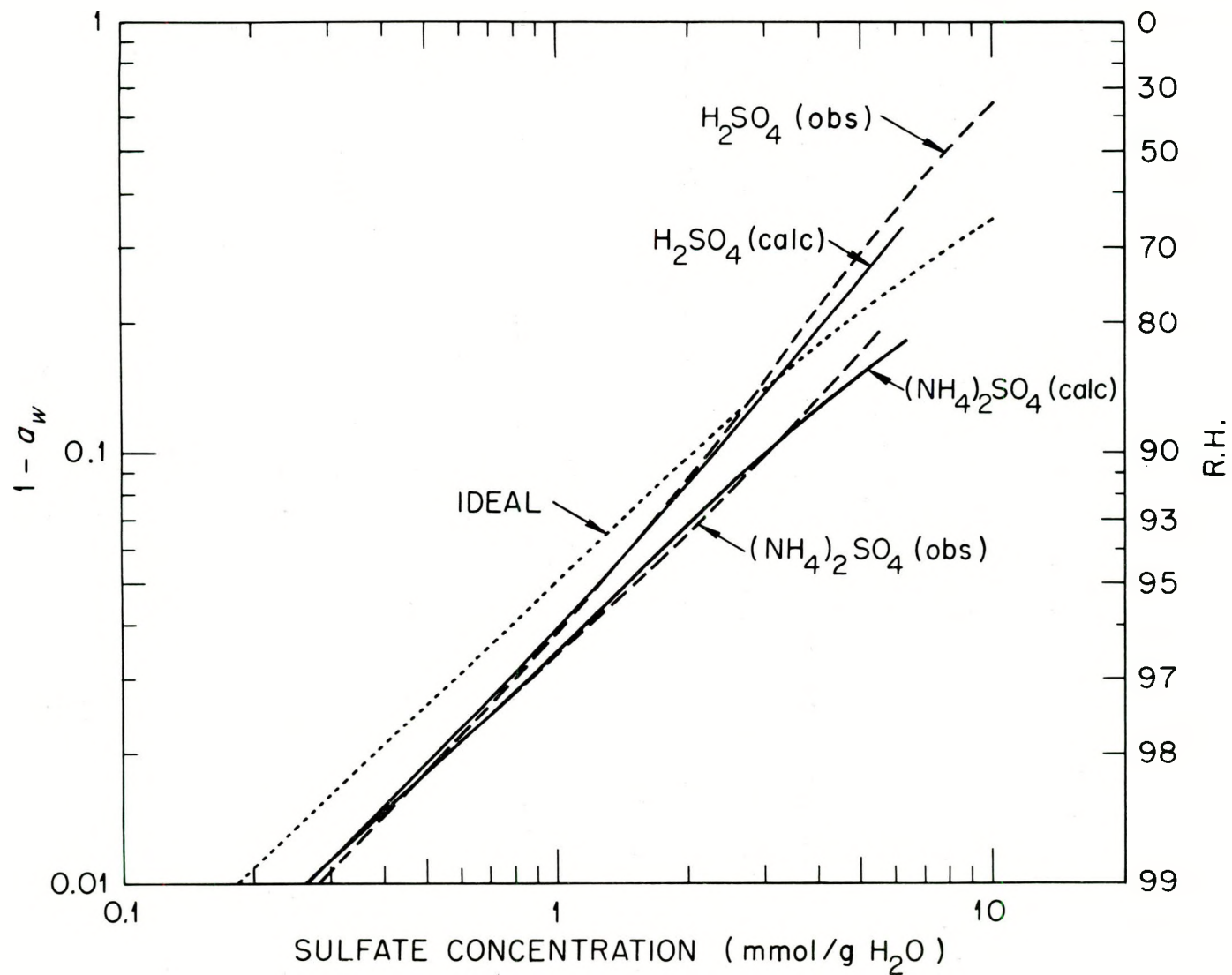


Fig. 4 Variation of the concentration of aerosol droplets with the activity of water  $a_w$  (the relative humidity/100) at 25°C.

The following expressions are assumed for the individual ionic activity coefficients;

$$\ln g_i = \frac{-Z_i^2 S I^{1/2}}{1 + I^{1/2}} + \sum_j B_{ij} g_{[j]} , \quad (5a)$$

$$\ln g_j = \frac{-Z_j^2 S I^{1/2}}{1 + I^{1/2}} + \sum_i B_{ij} g_{[i]} , \quad (5b)$$

wherein  $Z_i$  and  $Z_j$  are the ion charges,  $I$  is the ionic strength

$$I = \frac{1}{2} (\sum_i Z_i^2 [i] + \sum_j Z_j^2 [j]) , \quad (6)$$

and  $S$  is the Debye-Hückel constant, a theoretical quantity which depends only on the density and dielectric constant of the solvent and on the temperature. It is given with adequate accuracy for present purposes by the approximate expression

$$S = 1.1305 + 0.01842 t(^{\circ}\text{C}) \quad (7)$$

The coefficients  $B_{ij}^g$ , one for each pairwise combination of cations and anions is assumed to be a function only of  $I$ .

$$B_{ij}^g(I) = B_{ij}^{\circ} + B_{ij}^1 F(I) \quad (8)$$

$$F(I) = [1 - (1 + 2I^{1/2} - 2I) \cdot \exp(-2I^{1/2})] / 4I . \quad (8a)$$

$B_{ij}^{\circ}$  and  $B_{ij}^1$  are adjustable constants.

The above expressions for  $g_i$  and  $g_j$  (eqs 5-8) when introduced into eq 4 give the following expression for  $\phi$

$$\phi = 1.0 + \frac{-2SG(I) + \sum_i \sum_j B_{ij}^{\phi} [i][j]}{\sum_i [i] + \sum_j [j]} , \quad (9)$$

$$G(I) = (1 + I^{1/2}) - \frac{1}{(1 + I^{1/2})} - 2 \ln(1 + I^{1/2}) , \quad (9a)$$



$$B_{ij}^{\phi} = B_{ij}^{\circ} + B_{ij}^1 \exp(-2I^{1/2}) \quad (9b)$$

Since four kinds of ions are distinguished, two cations ( $H^+$  and  $NH_4^+$ ) and two anions ( $X^-$  and  $X^{2-}$ ), there are four pairwise combinations of cations with anions and 8 adjustable parameters. Six of the following values were assigned by fitting the appropriate expressions above to the known osmotic and activity coefficients of pure  $H_2SO_4$  and  $(NH_4)_2SO_4$  solutions given by Robinson and Stokes<sup>15</sup> (Fig. 4), and the observed dissociation of  $HSO_4^-$  ion in pure sulfuric acid solutions (Young et al<sup>16</sup>).

i	j	$B_{ij}^{\circ}$	$B_{ij}^1$
$H^+$	$x^-$	0.272	3.080
$H^+$	$x^{2-}$	0.050	-4.005
$NH_4^+$	$x^-$	0.0*	0.319*
$NH_4^+$	$x^{2-}$	-0.049	-1.186

The remaining two (with asterisks) were adjusted to give approximately the observed solubility of  $(NH_4)_2SO_4$  in  $H_2SO_4$  solutions (Seidell,<sup>17</sup> Vol. II p. 756), Fig. 5.

## 2. Initial Estimate of Quantities Needed for the Calculation of the Plume Composition

The initial estimates of the amount of condensed water  $g_w$  and of the ion concentrations  $[H^+]$ ,  $[NH_4^+]$ ,  $[x^2]$  and  $[x^{2-}]$  are made by assuming that the only ions present in solution are  $H^+$ ,  $NH_4^+$ ,  $HSO_4^-$ , and  $SO_4^{2-}$  and that all the  $NH_3$  is present as  $NH_4^+$ . The following expressions then apply;

$$[NH_4^+] = C(NH_3)/g_w, \quad (10)$$

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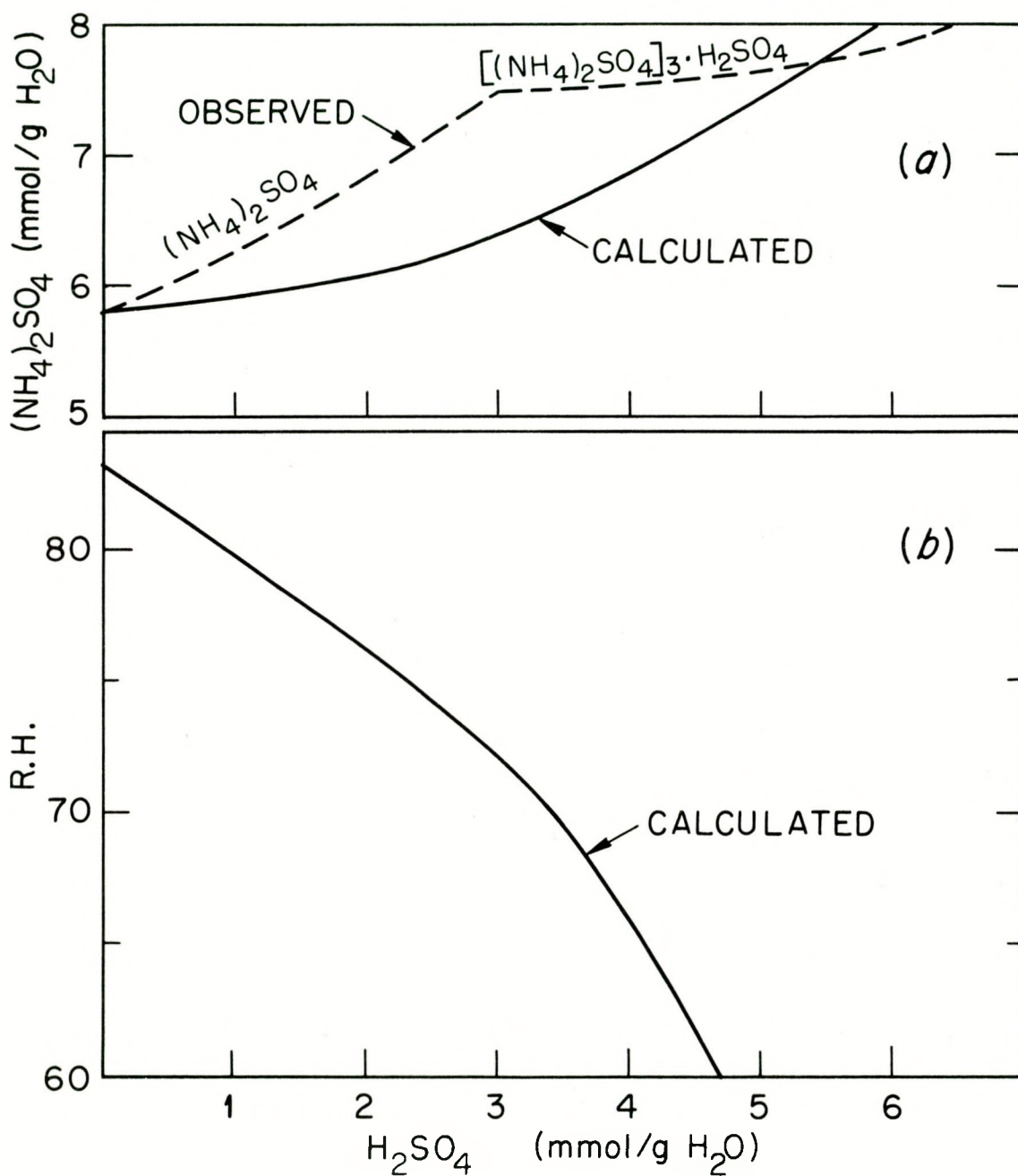


Fig. 5 The sulfuric acid - ammonium sulfate - water system at 25°C.

$$[x^-] = [HSO_4^-] = \frac{[H^+]}{[H^+] + Q_4} \cdot \frac{C(S(VI))}{g_w}, \quad (11)$$

$$[x^{2-}] = [SO_4^{2-}] = \frac{Q_4}{[H^+] + Q_4} \cdot \frac{C(S(VI))}{g_w}, \quad (12)$$

where  $Q_4$  is the concentration quotient from  $K_4$  (Table I),

$$Q_4 = \frac{[H^+][SO_4^{2-}]}{[HSO_4^-]} = K_4 \cdot \frac{g_{x^-}}{g_{H^+}g_{x^{2-}}}. \quad (13)$$

Introducing Eqs. 10 to 13 into the charge balance equation

$$[H^+] + [NH_4^+] = [HSO_4^-] + 2[SO_4^{2-}], \quad (14)$$

a quadratic in  $[H^+]$  is obtained with the solution

$$[H^+] = -B + (B^2 - C)^{1/2}, \quad (15)$$

wherein

$$B = [Q_4 + [C(NH_3) - C(S(VI))]/g_w]/2, \quad (15a)$$

$$C = Q_4[C(NH_3) - 2C(S(VI))]/g_w. \quad (15b)$$

The only unknowns are  $g_w$  and the ion activity coefficients.

The calculation begins by solving the following equation for  $g_w$ ,

$$[C(H_2O) - g_w/0.018] \frac{RT}{P_w^\circ} = \frac{g_w}{g_w + 3 \cdot C(S(VI)) \cdot 0.018}. \quad (16)$$

Each side is an expression of the activity of water, the right side being an approximation of the mole fraction of water in solution assuming three moles of ions are produced by a mole of S(VI). The solution for  $g_w$  is

$$g_w = [-B' + (B'^2 - C')^{1/2}] \cdot 0.018, \quad (17)$$

$$B' = [P_w^\circ/RT - C(H_2O) + 3 C(S(VI))]/2, \quad (17a)$$

$$C' = -3 C(S(VI)) \cdot C(H_2O). \quad (17b)$$

The activity coefficients are set initially at unity. The ion concentrations thus obtained from Eqs. 10-15 are used to estimate activity coefficients in S. GAMAW, and the calculation is repeated by cycling between S. EVAL AND S. GAMAW until the activity coefficients become constant (only a few interactions are needed). Then the water activity given by S. GAMAW is compared with that given by the amount of gaseous water,

$$a_w = [C(H_2O) - g_w/0.018]RT/P_w^\circ \quad (18)$$

and  $g_w$  is adjusted by repeated cycling between S. NONLIN and S. EVAL until a good estimate of  $g_w$  (and ion concentrations and activity coefficients) is obtained.

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