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*Results of Investigations at the
Zunil Geothermal Field, Guatemala:
Well Logging and Brine Geochemistry*

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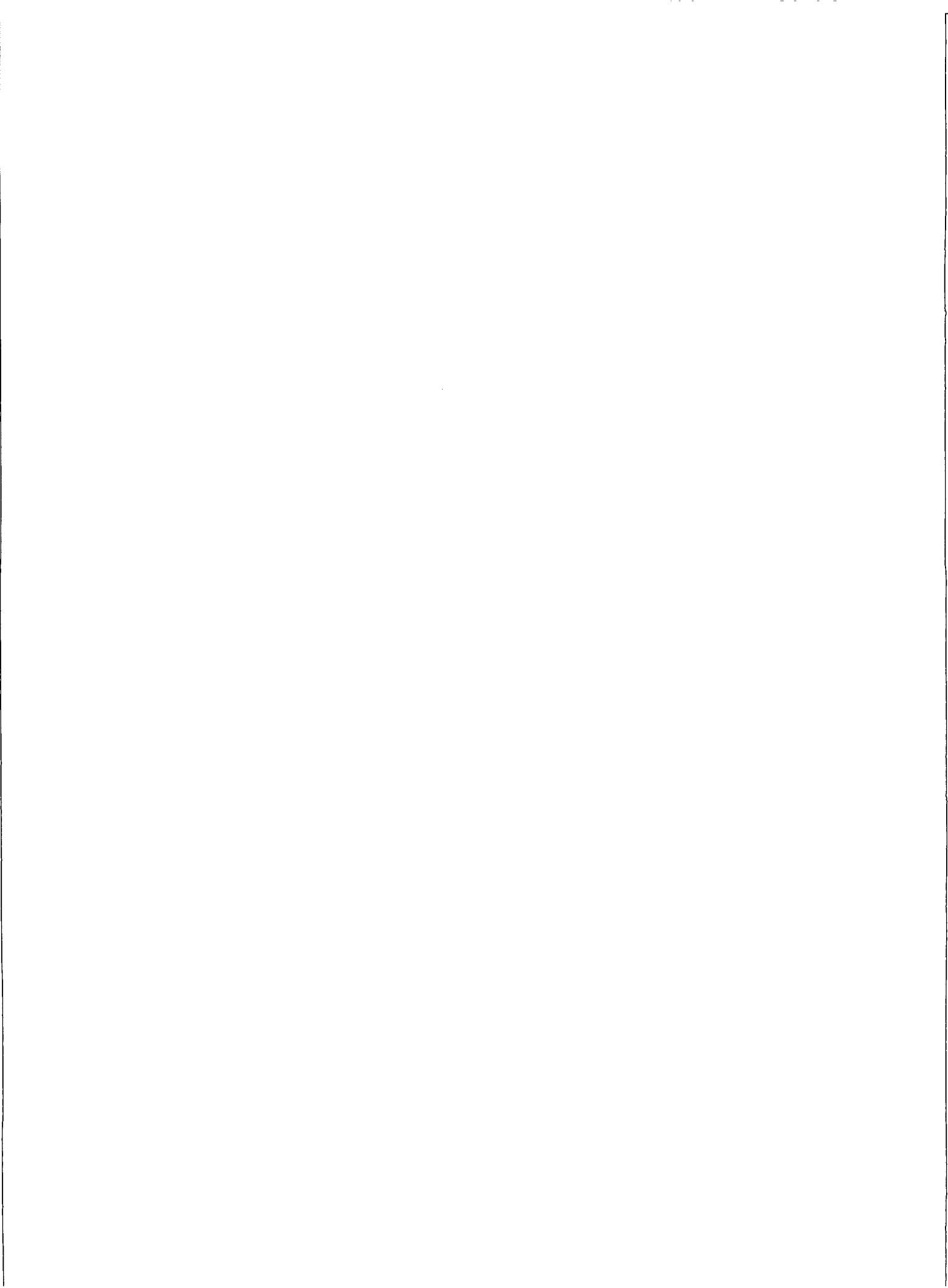
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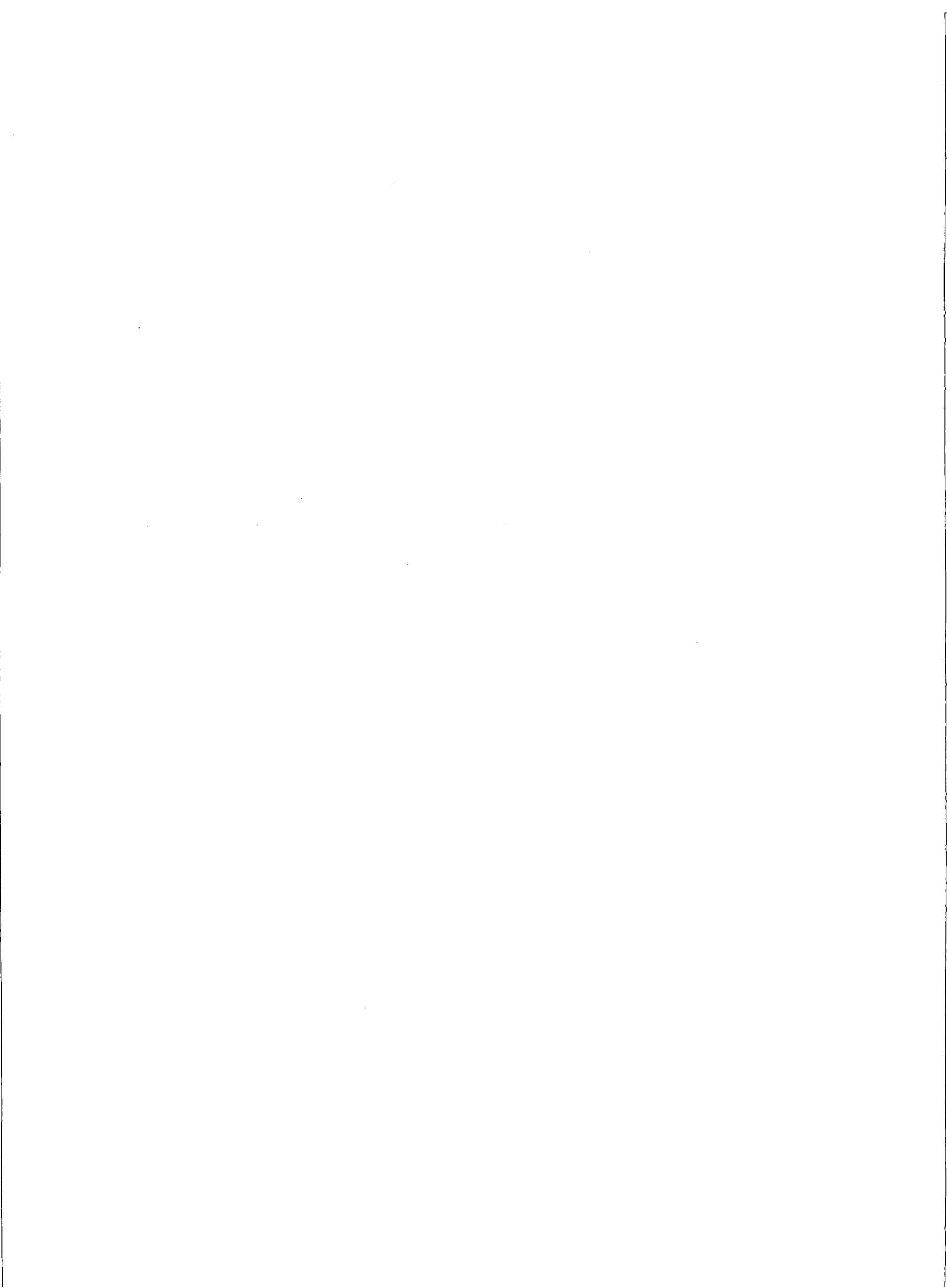
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EXECUTIVE SUMMARY

The Zunil geothermal prospect has been the subject of investigation by the electric utility of Guatemala (Instituto Nacional de Electrificación, or INDE) for the last 15 years. The initial stages of the investigation were supported by the Japan International Cooperation Agency (JICA) and Electroconsult (ELC) of Italy. During the initial stages a number of small-diameter holes were drilled (in the mid-1970s) to quantify the resource. These holes led to the drilling of six deep large-diameter exploration wells (numbered ZCQ-1 through ZCQ-6) to confirm the reservoir existence (Fig. *i*). This drilling was performed in 1980 and 1981. Testing and evaluation of these wells were performed from 1981 until 1986.

In 1987, INDE, using funding provided by Banco Interamericano de Desarrollo (BID), funded Cordón y Mérida Ing., MK-Ferguson, and MK-Engineers, Inc. to develop and engineer a 15-MW plant based on existing or possible new wells. Two reports have been published to date that summarize the status of this work (MK-Ferguson Co. 1988, 1989).

In 1988, Los Alamos National Laboratory (LANL), as part of its Central American Energy and Resources Project, was commissioned to determine the condition of wells ZCQ-3 through ZCQ-6 by using the high-temperature tools developed for other geothermal programs (see Appendix A for a description of the tools). Additionally, LANL was to collect water/steam and gas samples downhole and on the surface to extend the geochemical information data base. Wells ZCQ-1 and ZCQ-2 were not investigated because they were unproductive, possibly from well damage during drilling.

Operational difficulties slowed the well logging activities initiated in the fall of 1988, making it impossible to complete the operation in one trip and necessitating a second trip in 1989. Between the two trips, numerous logs were executed that defined the condition of each well. The thermodynamic condition of the flowing wells was difficult to assess because the fluids in the wells did not reach a state of thermal equilibrium during the entire well logging program. The wells had been shut in and therefore static for several years. During the running of a number of fluid spinner/temperature/pressure (STP) logs in two

flowing wells, it became evident that the liquid level, and thus the liquid/vapor interface, continued to move down the flowing well. A stable thermal equilibrium condition could not be achieved.

Fluid samples were collected in wells ZCQ-3, ZCQ-4, and ZCQ-5 from depths of 670, 775, and 900 m, respectively. Weirbox samples were collected from wells ZCQ-3, ZCQ-4, ZCQ-6, and ZC-11 (a slimhole direct-use well). Gas samples were collected from numerous wells and fumaroles. Additional fluid samples were collected in the Zunil-Quetzaltenango region to provide further information on recharge, mixing, dilution, and age of the reservoir water. Results from the downhole fluid sample chemistry analysis did indicate that the Zunil reservoir is geochemically inhomogeneous. Stable isotope data indicate that recharge to the field is from the northeast volcanic area, and tritium data indicate that the reservoir water may be only 500 to 700 years old. Previous reports on the Zunil field appear in Fournier *et al.* (1982), Giggenbach (1986, 1988), Adams et al (1990), MK-Ferguson Co. (1988, 1989), and others.

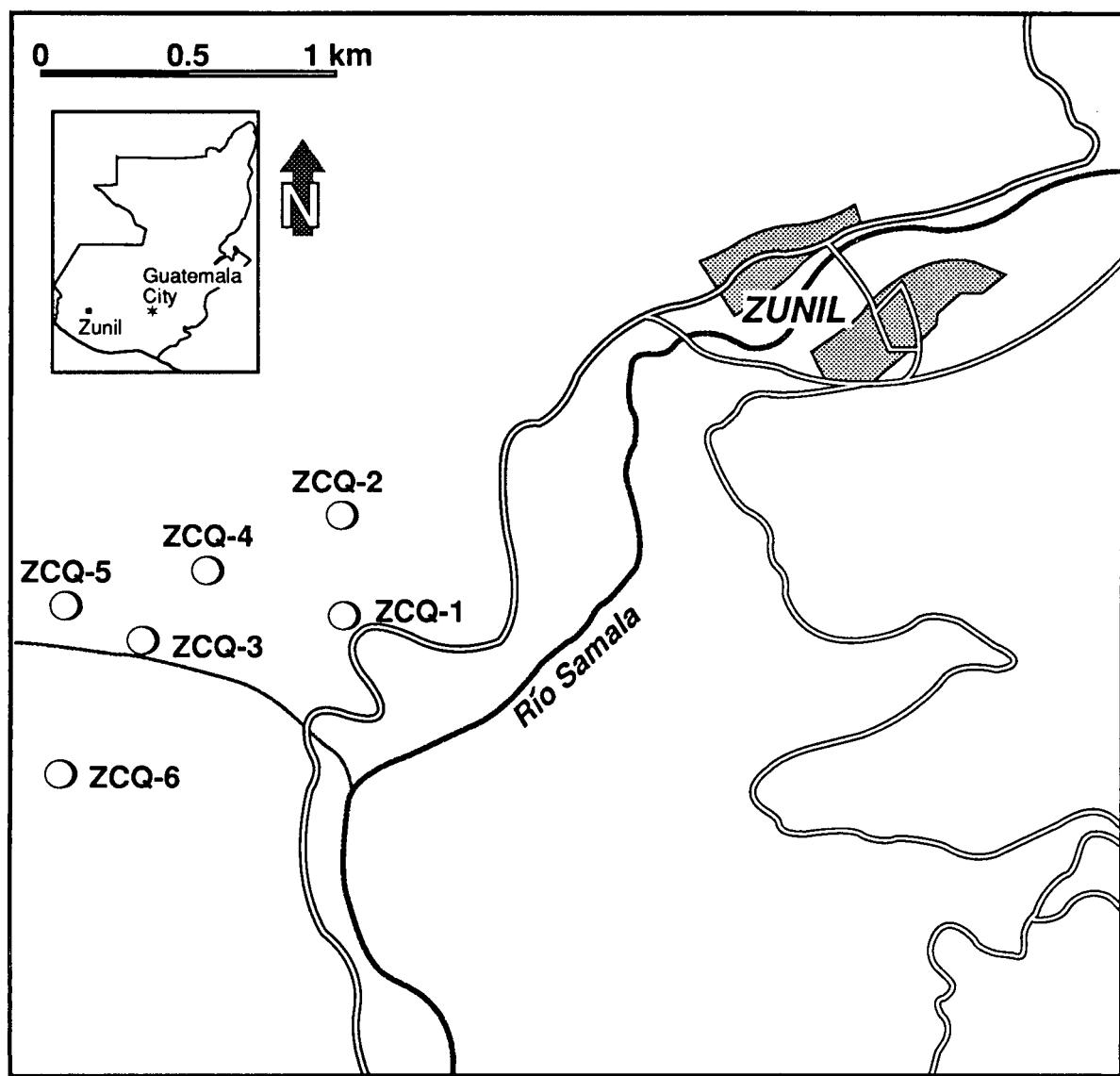


Fig. i. Location map of geothermal wells in the Zunil geothermal prospect.



RESUMEN EJECUTIVO

El potencial geotérmico de Zunil ha sido objeto de investigación por la compañía eléctrica de Guatemala (Instituto Nacional de Electrificación, o INDE) durante los últimos 15 años. Las etapas iniciales de investigación recibieron el respaldo de la compañía japonesa "Japan International Cooperation Agency" (JICA) y la compañía italiana "Electroconsult" (ELC). Al comienzo de la investigación, un número de pozos de pequeños diámetros fueron perforados (a mediados de la década de los 70) para determinar la cantidad de los recursos geotérmicos existentes. Estos pozos fueron los precursores de la perforación de seis pozos exploratorios profundos de gran diámetro (identificados ZCQ-1 al ZCQ-6) para confirmar la existencia del depósito (Figura *ii*). Esta perforación se llevó a cabo durante los años 1980 y 1981. Los análisis y evaluaciones de estos pozos se efectuaron durante los años 1981 y 1986.

Durante el año siguiente, el INDE, usando fondos suministrados por el Banco Interamericano de Desarrollo (BID) financió a Cordón y Mérida Ing., MK-Ferguson, y MK-Engineers, Inc., para desarrollar y construir una planta de 15 MW, basándose en pozos existentes y los potenciales. A la fecha se han publicado dos reportes, los cuales resumen el estado de este trabajo (MK-Ferguson Co. 1988, 1989).

Durante 1988, Los Alamos National Laboratory (LANL), como parte de su Proyecto Centroamericano de Energía y Recursos Naturales, fue comisionado a determinar las condiciones de los pozos ZCQ-3 al ZCQ-6, utilizando los instrumentos de medición que había desarrollado para otros programas geotérmicos (para descripción de los instrumentos, véase Anexo A). Además, LANL debía colectar muestras del agua, vapor y gas del subsuelo y la superficie de los pozos para extender la información geoquímica a la base de datos. Los pozos ZCQ-1 y ZCQ-2 no se investigaron por no ser productivos, posiblemente debido a daños causados a los pozos durante su perforación.

Dificultades operacionales atrasaron las actividades iniciadas en el registro cronológico de la perforación de pozos durante el otoño del año 1988, haciendo imposible terminar la operación en una sola visita, requiriéndose un segundo viaje en el año 1989. Entre los dos viajes, numerosos registros se llevaron

a cabo, produciendo éstos resultados categóricos de las condiciones de cada pozo. Las condiciones del flujo de los pozos fue difícil de catalogar porque los fluidos en los pozos no alcanzaron un estado de equilibrio termal durante todo el período del programa analítico de pozo. Los pozos se cerraron y por lo tanto estuvieron inactivos por varios años. Durante varios de los registros de flujo/temperatura/presión (STP) de los pozos fluyentes, se hizo obvio que el nivel del líquido, a su vez el intercambio del líquido y vapor, causara movimiento continuo hacia abajo en el pozo fluyente. Condición de equilibrio termal no se pudo alcanzar.

Muestras del flujo se colectaron de los pozos ZCQ-2, ZCQ-4 y ZCQ-5, desde profundidades de 670, 775 y 900 metros, respectivamente. Muestras Weirbox se obtuvieron de los pozos ZCQ-3, ZCQ-4, ZCQ-6 y ZC-11 (un pozo angosto de uso directo). Muestras de gas fueron colectadas de varios pozos y fumarolas. Muestras adicionales se colectaron en las regiones de Zunil y Quetzaltenango, proveyendo mayor información sobre recarga, mezcla, dilución y la era del depósito. Los resultados de las muestras fluidas de los análisis químicos del pozo abajo indicaron que la reserva de Zunil no es geoquímicamente homogénea. Los datos isotópicos estables indican que el recargo en el campo proviene desde el área volcánica noreste, y los datos de tritio indican que el depósito de agua puede existir desde hace ya 500 a 700 años. Reportes anteriores sobre el campo de Zunil aparecen en escritos de Fournier y otros (1982), Giggenbach (1986, 1988), Adams y otros (1990), MK-Ferguson Co., y otros.

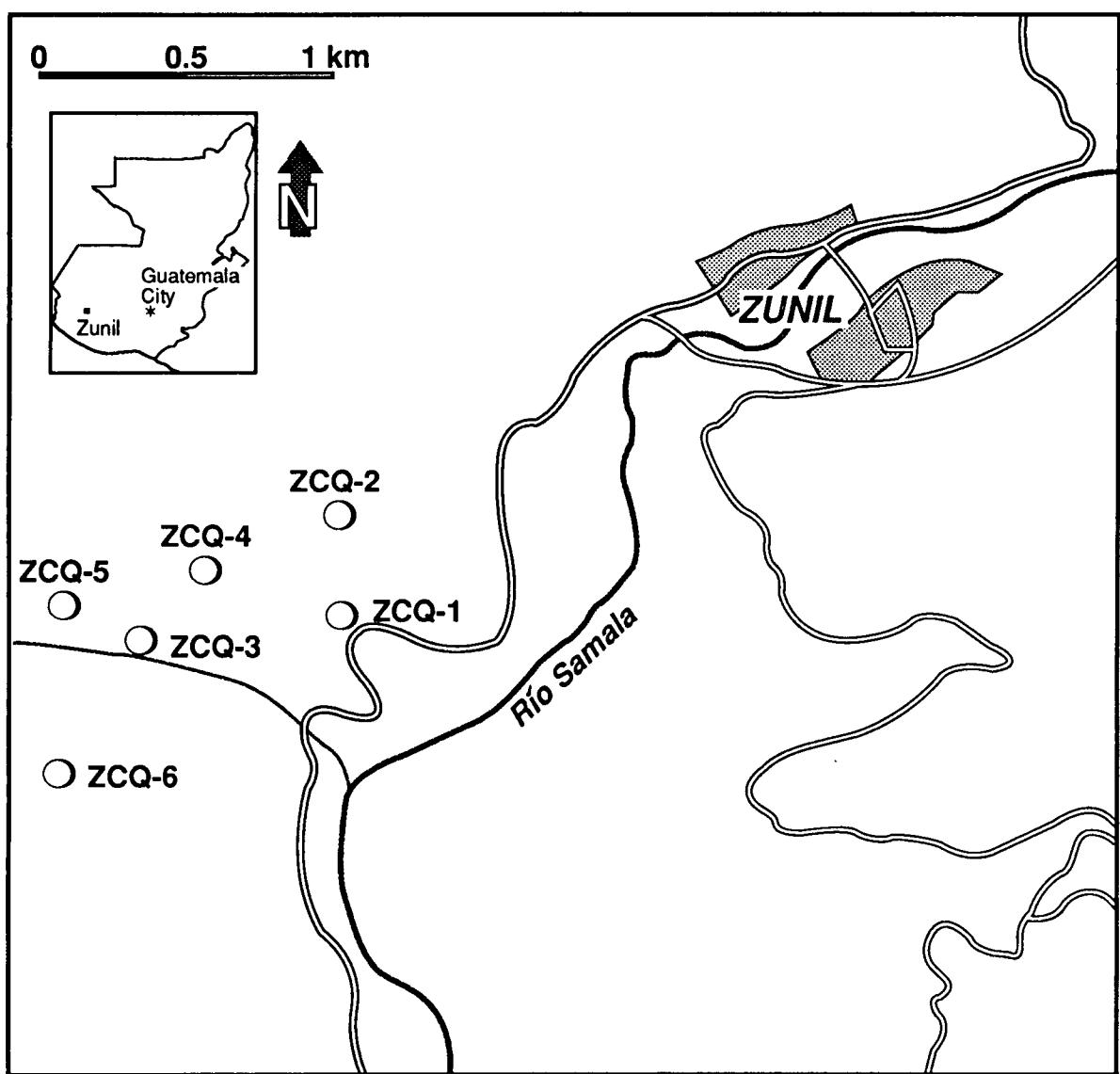


Fig. *ii*. Mapa de localización de pozos potenciales geotérmicos de Zunil.

RESULTS OF INVESTIGATIONS AT THE ZUNIL GEOTHERMAL FIELD, GUATEMALA: WELL LOGGING AND BRINE GEOCHEMISTRY

by

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ABSTRACT

The well logging team from Los Alamos and its counterpart from Central America were tasked to investigate the condition of four producing geothermal wells in the Zunil Geothermal Field. The information obtained would be used to help evaluate the Zunil geothermal reservoir in terms of possible additional drilling and future power plant design. The field activities focused on downhole measurements in four production wells (ZCQ-3, ZCQ-4, ZCQ-5, and ZCQ-6). The teams took measurements of the wells in both static (shut-in) and flowing conditions, using the high-temperature well logging tools developed at Los Alamos National Laboratory.

Two well logging missions were conducted in the Zunil field. In October 1988 measurements were made in well ZCQ-3, ZCQ-5, and ZCQ-6. In December 1989 the second field operation logged ZCQ-4 and repeated logs in ZCQ-3. Both field operations included not only well logging but the collecting of numerous fluid samples from both thermal and nonthermal waters.

I. INTRODUCTION

The Zunil geothermal field, as noted in Fig. *i*, is 120 miles northwest of Guatemala City. Preliminary work there resulted in the drilling of 6 wells in 1980 and 1981. Of the six, two were damaged and unproductive (wells ZCQ-1 and ZCQ-2), two produced superheated steam (wells ZCQ-4 and ZCQ-5), and two are currently commercially productive at acceptable wellhead pressures (wells ZCQ-3 and

ZCQ-6). The location of these wells relative to a schematic of the structure is shown in Fig. 1a. Figures 1b and 1c show the temperature schematics from early measurements.

Sections II and III describe the logging performed and the geochemical results obtained during our efforts.

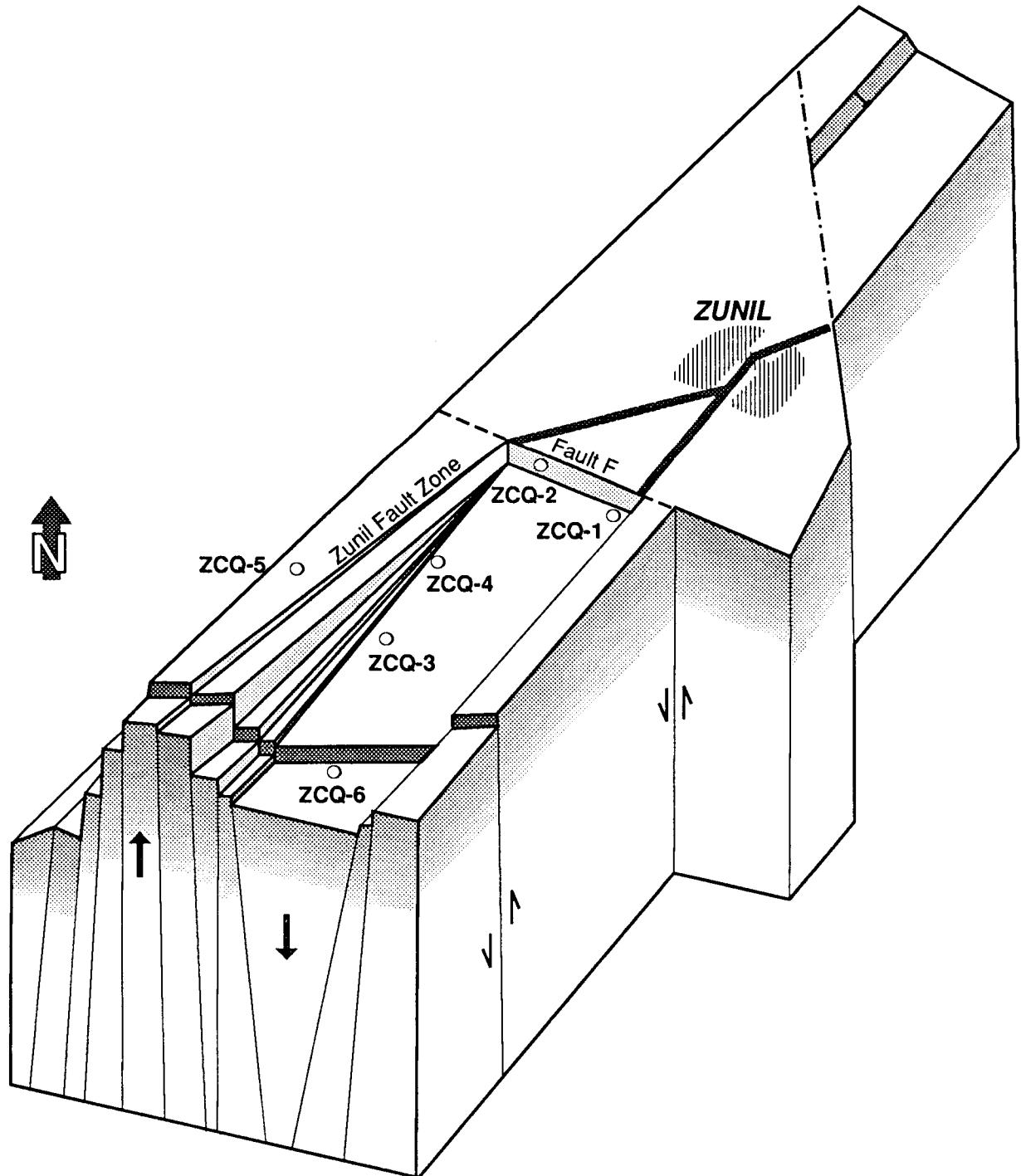


Fig. 1a. Three-dimensional block diagram of Zunil area.

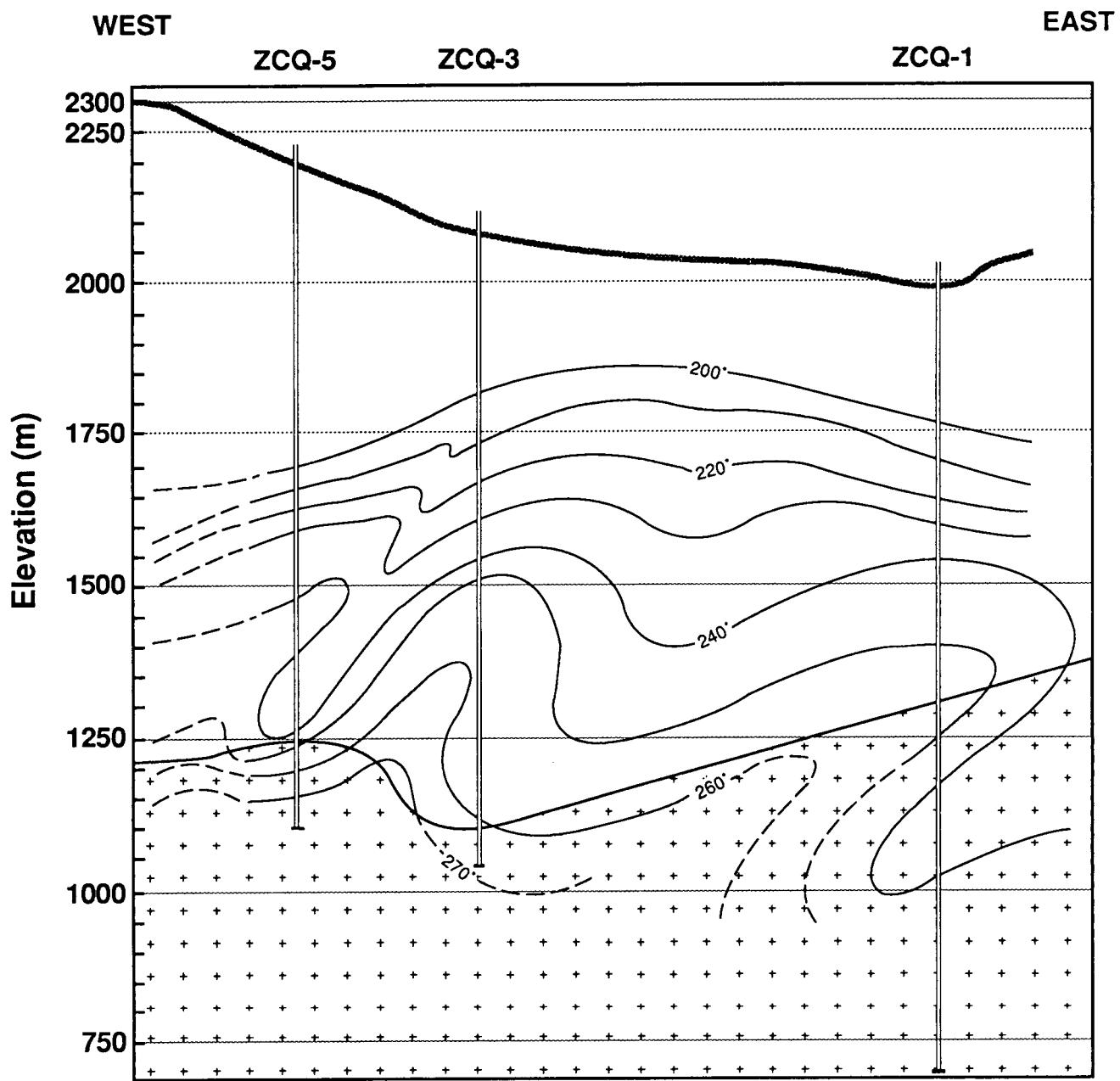


Fig. 1b. Temperature cross section, wells ZCQ-5, ZCQ-3, and ZCQ-1 (MK-Ferguson 1988).

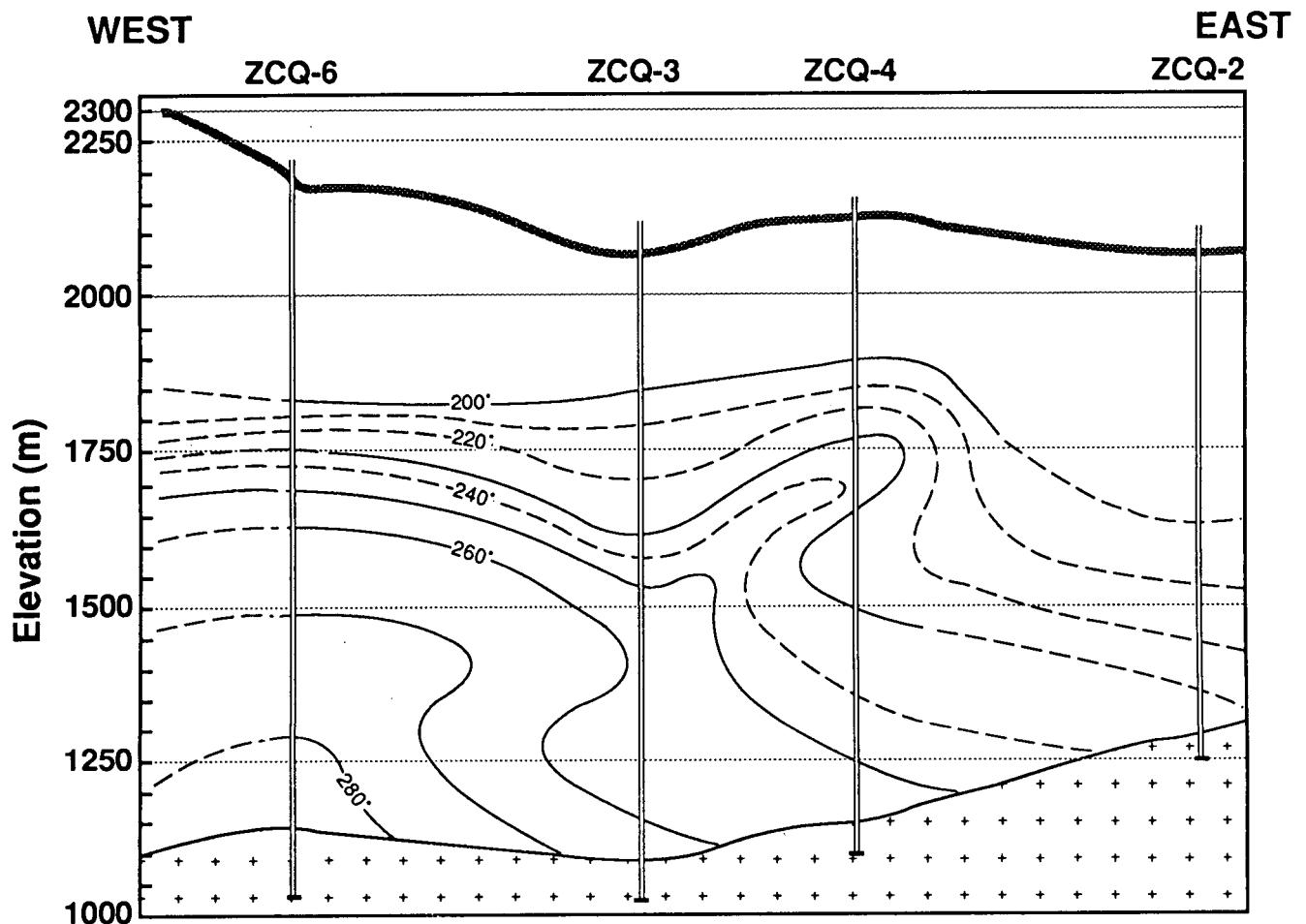


Fig. 1c. Temperature cross section, wells ZCQ-6, ZCQ-3, ZCQ-4, and ZCQ-2.

II. WELL LOGGING OPERATIONS

A. ZCQ-3

Well ZCQ-3 was drilled to a depth of 1040 m and completed with 9 5/8-in. casing to 570 m. A 7 5/8-in. slotted liner was hung from 560 m to a depth of 1030 m. Earlier surveys had reported the maximum temperature at 260°C near the bottom of the well and liquid level at 600 m. Wellhead pressure with the well shut in was 668 psi, which dropped to 150 psi about 1 hour after the system was opened to flow.

The first attempts to run the well logging tools in ZCQ-3 in October 1988 encountered several problems. The slots in the liner were wider than dimensions typically used in Central America, and the tips

on the three-arm caliper tool were not wide enough to bridge the slots. The teams ran the caliper log in the casing only to avoid damage to the caliper arms. Also, pressurized air was necessary to lift this well. However, the lack of working equipment (compressor) during the first field operation limited logging operations to static well conditions.

A shut-in temperature/rabbit survey was made to a depth of 1030 m. The maximum temperature measured 260°C at about 670 m. This temperature run was followed by a fluid sample taken at 671 m. Then the three-arm caliper was run from 557 m to the surface and showed no chemical deposits and very little, if any, scaling.

For the second phase of the Zunil field operations (December 1989) we modified the tips on the caliper tool to allow completion of the caliper log through the slotted liner from 1010 m to 500 m. Figure 2 shows data from the 1988 log and the 1989 log. When a temperature/rabbit survey of the well was repeated, the collar locator in the temperature tool found the slotted liner at 568 m, establishing agreement with the caliper log. Figure 3 summarizes data from all temperature surveys made in ZCQ-3.

Data of more interest resulted from a sequence of STP measurements of ZCQ-3. The STP tool was rigged up and set in the pressure lock. While the well was shut in, pressurized air lifted the well. Then flow from the well was allowed through the vent lines to the mufflers until a significant flow rate was established (about 30 min). The well was then shut in and the first STP survey was run from the surface to 1005 m. The liquid/vapor interface appeared to be between 460 and 500 m although the well was in a transient condition (Fig. 4). After this shut-in log, we made two additional STP surveys with the well flowing. The first flowing log ran from 1005 m up to 580 m with the bypass valve open to full flow. It is quite obvious from the pressure and spinner data that the liquid/vapor interface had moved below 1005 m during the third log (Fig. 5). The entire wellbore contained two-phase flow and was still in a highly transient state. It is probable that in a relatively short time, only two-phase fluid will intersect this well. The drop in fluid temperature and pressure, measured during the third log, confirms the absence of liquid in this region of the well.

B. ZCQ-4

Well ZCQ-4 was a production well drilled to a depth of 1025 m. The 7 5/8-in. slotted liner was hung in the 9 5/8-in. casing at 420 m and extended to a depth of 1020 m. The reported water level with the well shut-in was about 650 m. Wellhead pressure was 573 psi with the wellhead valves closed and dropped to 110 psi under flowing conditions. Maximum fluid temperature was 240°C near the bottom of the well.

We ran the three-arm caliper tool in the well from 1000 m to the surface. The slotted liner section of the wellbore from 1000 m up to 420 m showed evidence of some chemical build-up inside the liner

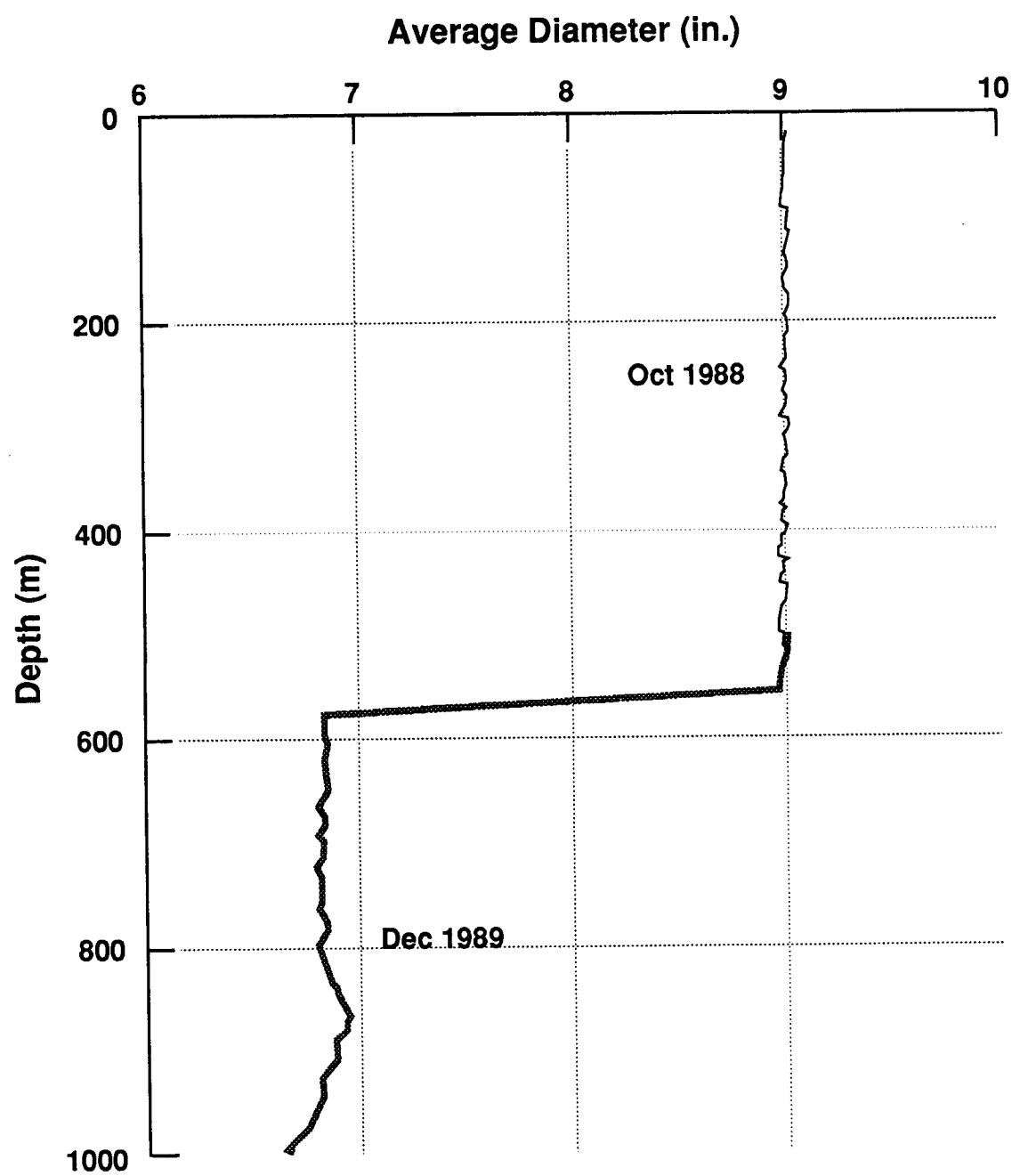


Fig. 2. Three-arm caliper survey of ZCQ-3.

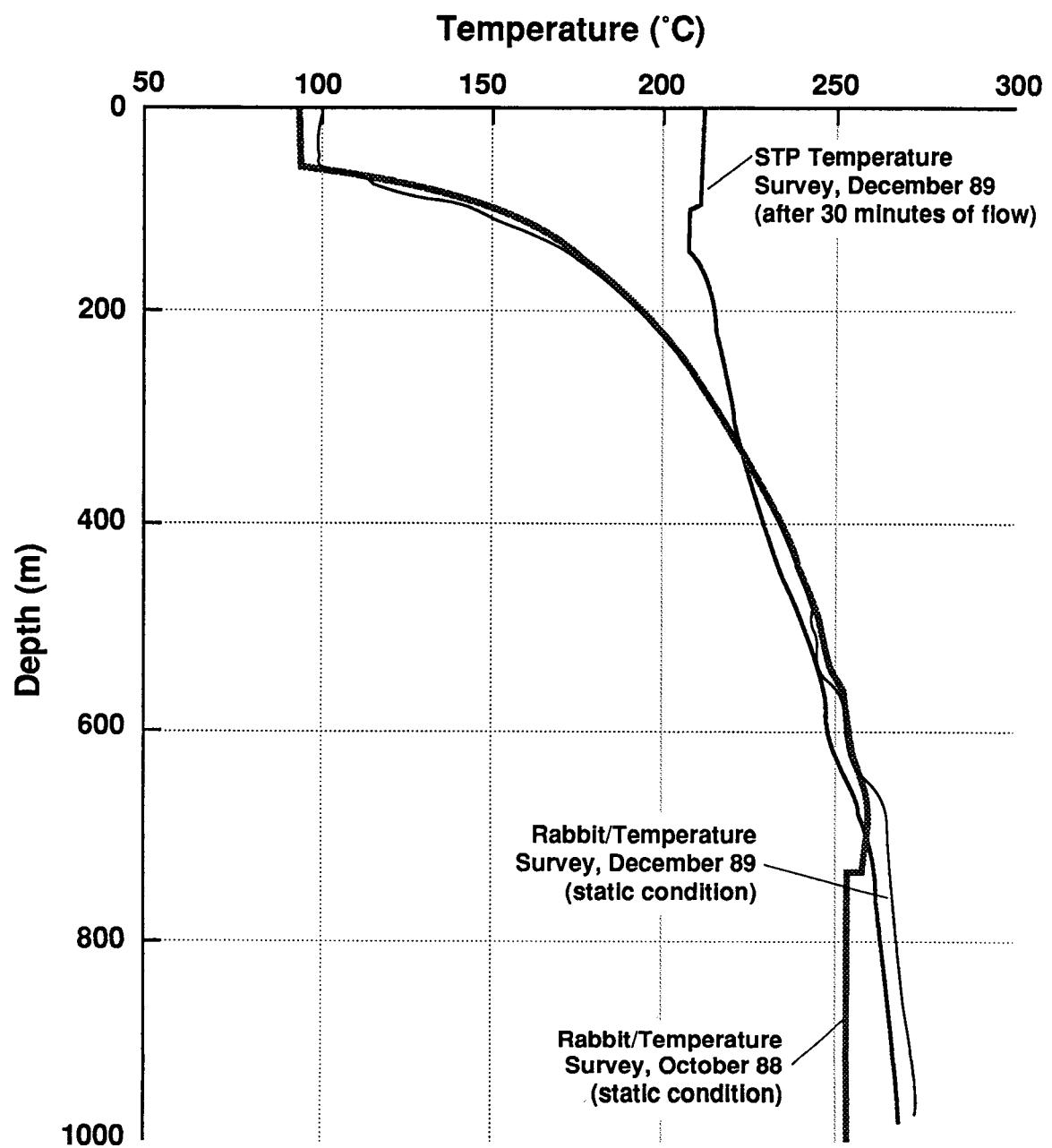


Fig. 3. Temperature surveys of ZCQ-3.

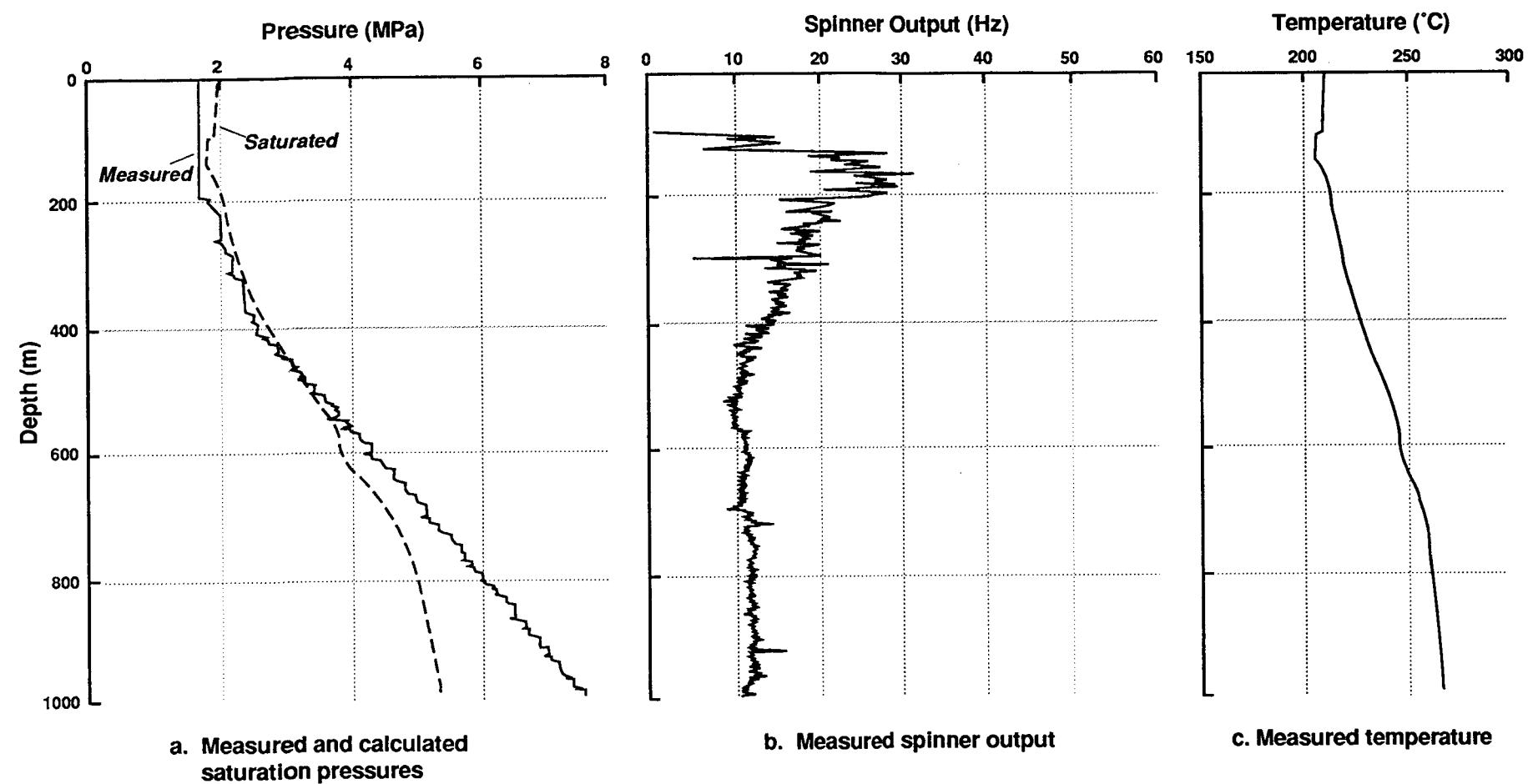


Fig. 4. STP results for well ZCQ-3, well shut-in after flowing 30 min, December 1989.

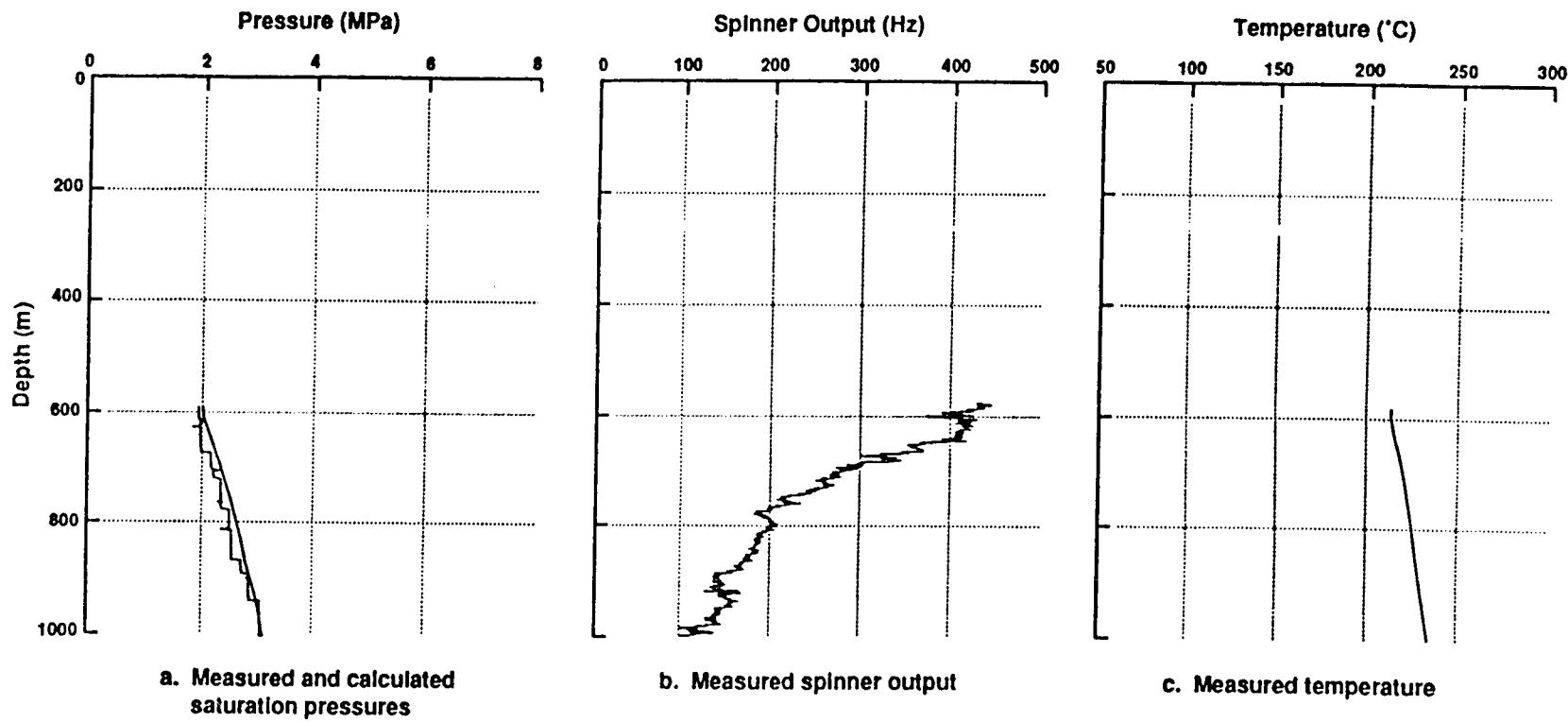


Fig. 5. STP results for well ZCQ-3, well flowing with bypass valve open to full flow.

(Fig. 6). There was also evidence of some slight chemical build-up in the 9 5/8-in. casing from the liner hanger up to about 200 m. The caliper log was followed by a fluid sample taken at 900 m.

The STP tool was rigged up to run a series of logs in this well. The first survey ran from the surface to 1000 m with the well shut-in after the well was allowed to flow for an undetermined period of time. The STP data from this log (Fig. 7) show that the liquid level was detected at 470 m. The tool velocity from the surface to 650 m was 35 m/min. Increasing the velocity to 50 m/min from 650 to 1000 m resulted in the step increase in the spinner output at 650 m. The spinner data taken from 870 to 1000 m indicate a circulation zone in which fluid entered the well below 1000 m, flowed up the well, and exited at 870 m.

The second STP survey was run from a depth of 1000 m up to 460 m with the wellhead bypass valve cracked open (1/8 of full flow). The resulting low flow allowed in the well provided some interesting results, notably that the circulation zone was noticeably diminished and the liquid/vapor interface moved down the well to about 515 m (Fig. 8). The data were essentially unchanged on the third STP survey (Fig. 9), which ran from about 1000 m to 460 m under conditions the same as those for the second log.

Then the STP tool was returned to the 1000 m depth and the wellhead bypass valve was opened to 33% of full flow. We recorded data from 1000 m to 200 m that showed the flowing conditions in a high transient state (Fig. 10). The measured pressure, higher than the saturation pressure from about 760 m to 1000 m (liquid region), showed that the liquid/vapor interface was rapidly moving down the well. The spinner output data showed fluid entering the well at about 920 m. The pressure gradient, between 900 m and 1000 m during this log, was 0.015 MPa/m, which was 1.8 times a normal steady-state hydrostatic pressure gradient. Because well conditions were in a such a transient state, no thermodynamic analysis could be calculated. Figure 11 shows a comparison of the temperature data for the four STP logs. For the shut-in and very low flow rate conditions, the temperatures throughout the wellbore were essentially identical. For the larger flow rate (STP log 4), the vapor or steam temperatures dropped as expected, although the temperatures in the liquid from 760 m to the bottom of the well remained the same. The drop in vapor temperatures may give some indication of the vapor quality of the two-phase fluid flow.

C. ZCQ-5

Well ZCQ-5 was a nonproductive well with ambient pressure at the wellhead. ZCQ-5 was drilled to a depth of 1080 m. The 9 5/8-in. casing extended to a depth of 751 m and a 7 5/8-in. slotted liner was hung from 737 m to 951 m. The maximum reported temperature was about 260°C at the bottom of the well.

Because this well had no differential pressure at the wellhead and was nonproductive, the pressure lock and cable pack-off assembly were unnecessary. The fluid sampler was deployed to 775 m, and the measured temperature of the borehole fluid at the sample depth was 260°C.

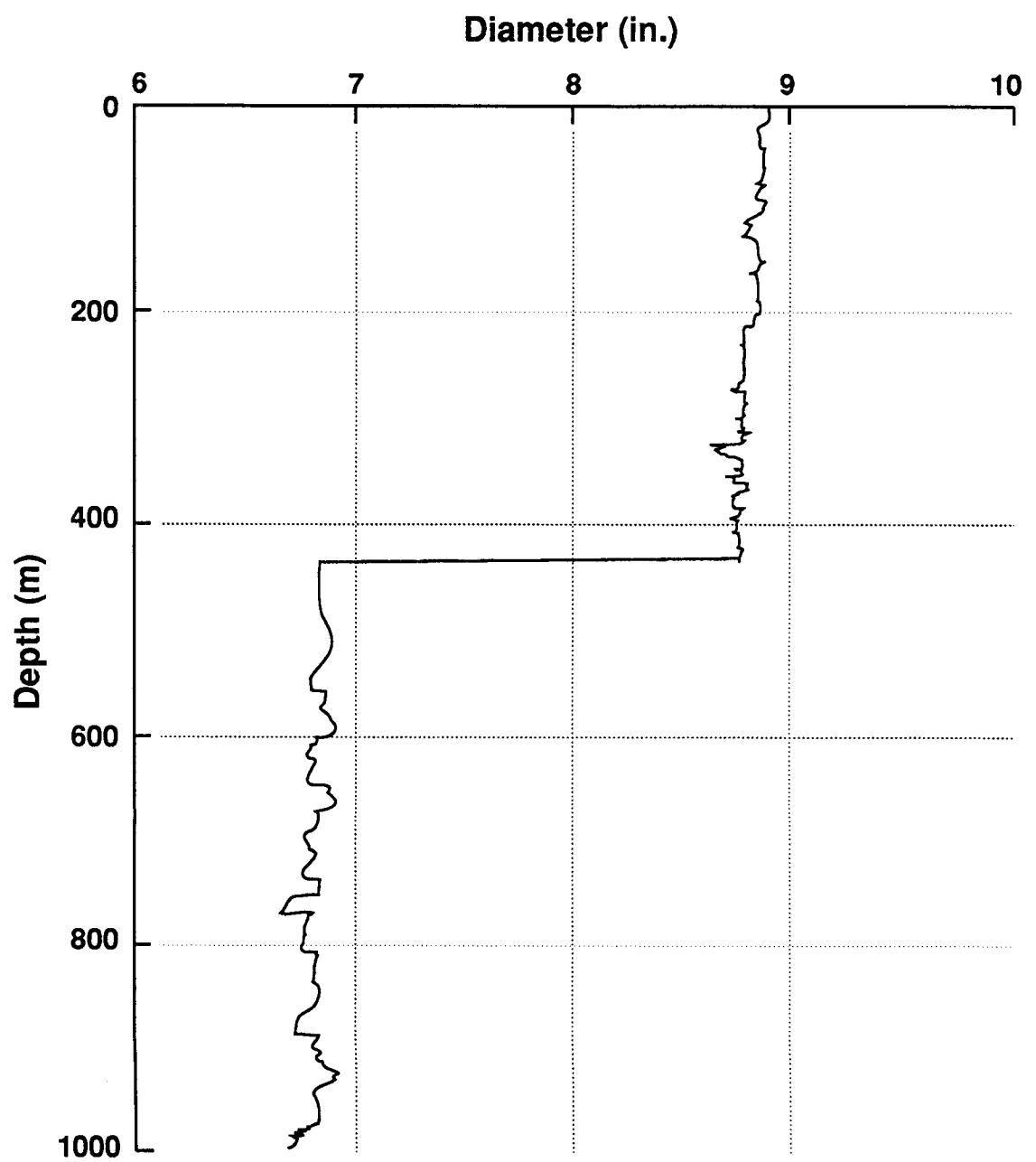


Fig. 6. Caliper survey of well ZCQ-4, December 1989.

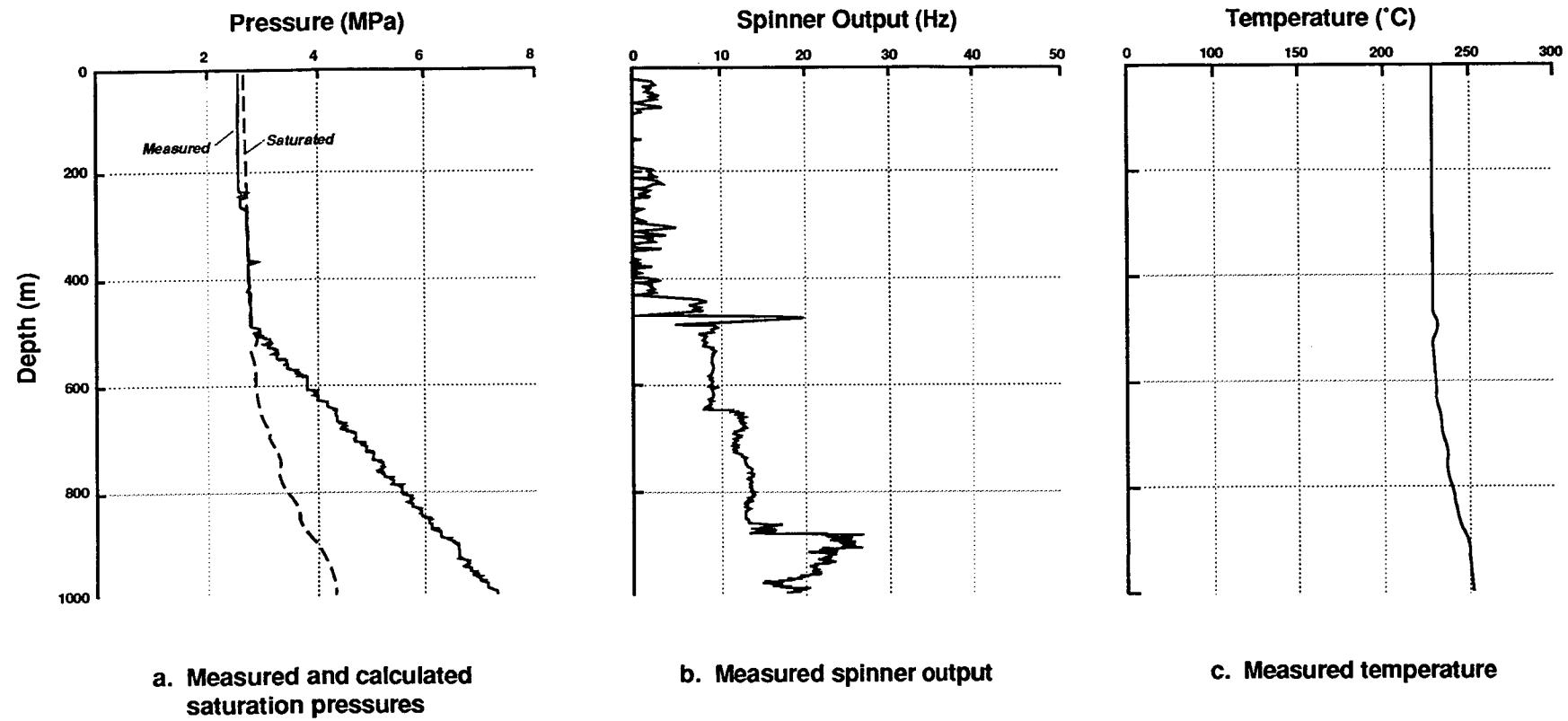


Fig. 7. STP log #1 in well ZCQ-4, December 11, 1989 (log-in). Well shut in.

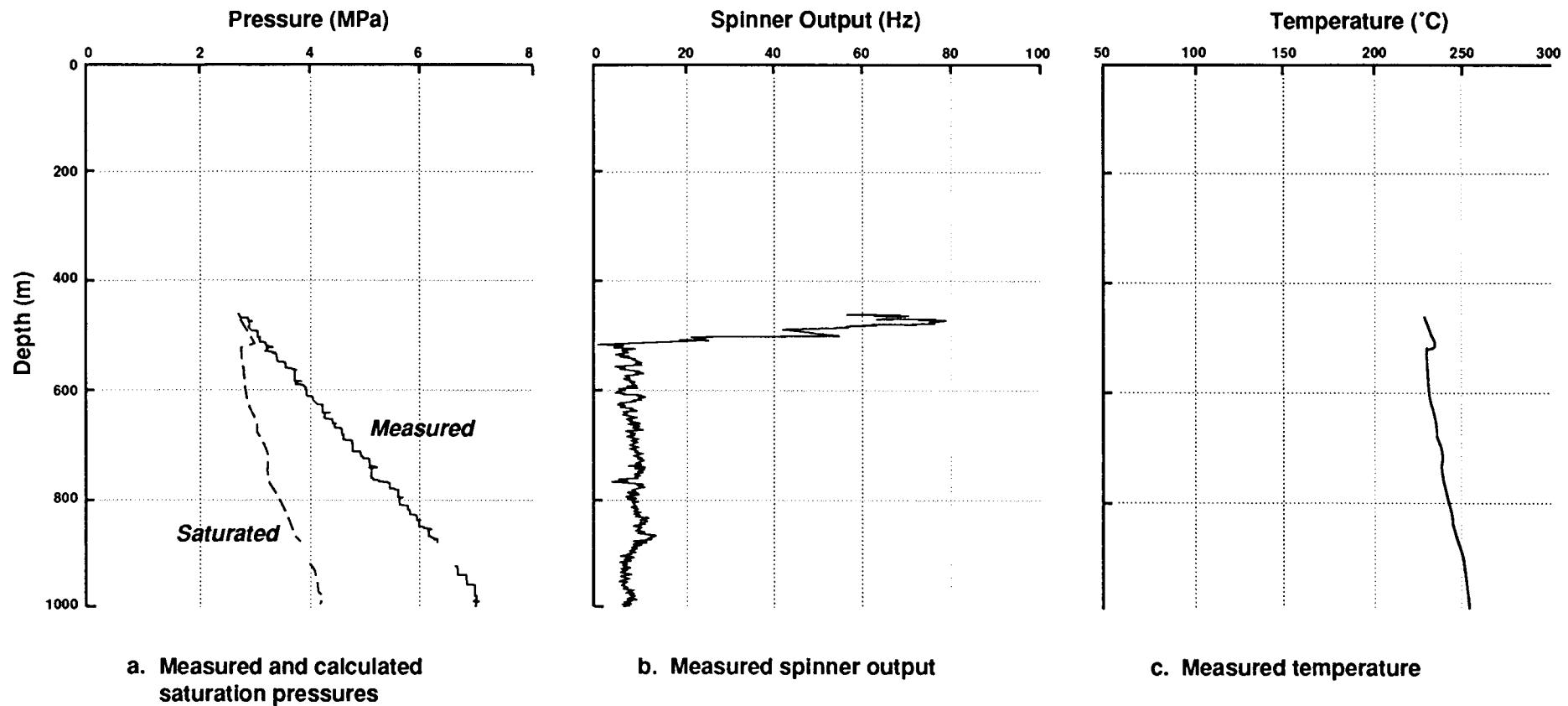


Fig. 8. STP log #2 in well ZCQ-4, December 11, 1989 (log-out).

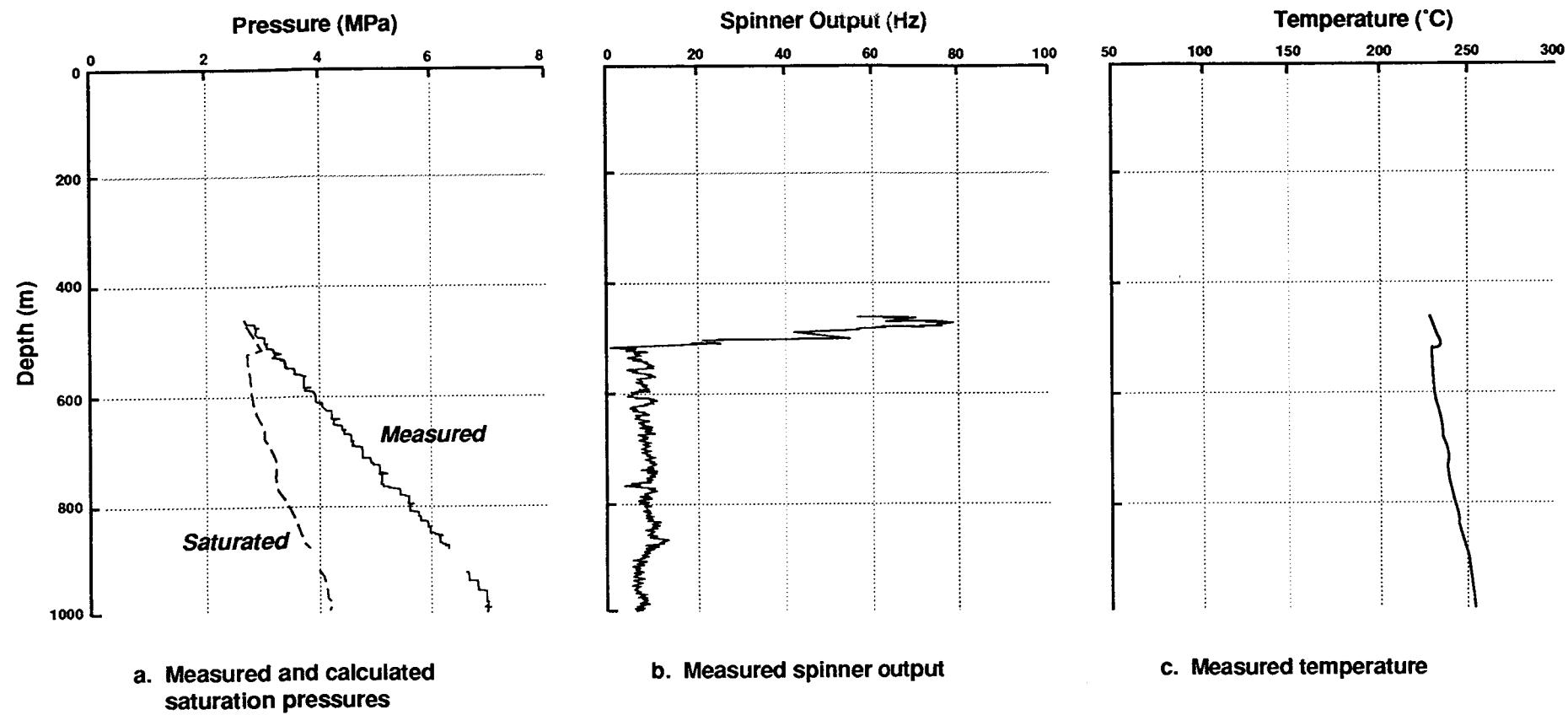


Fig. 9. STP log #3 in well ZCQ-4, December 11, 1989 (log-in).

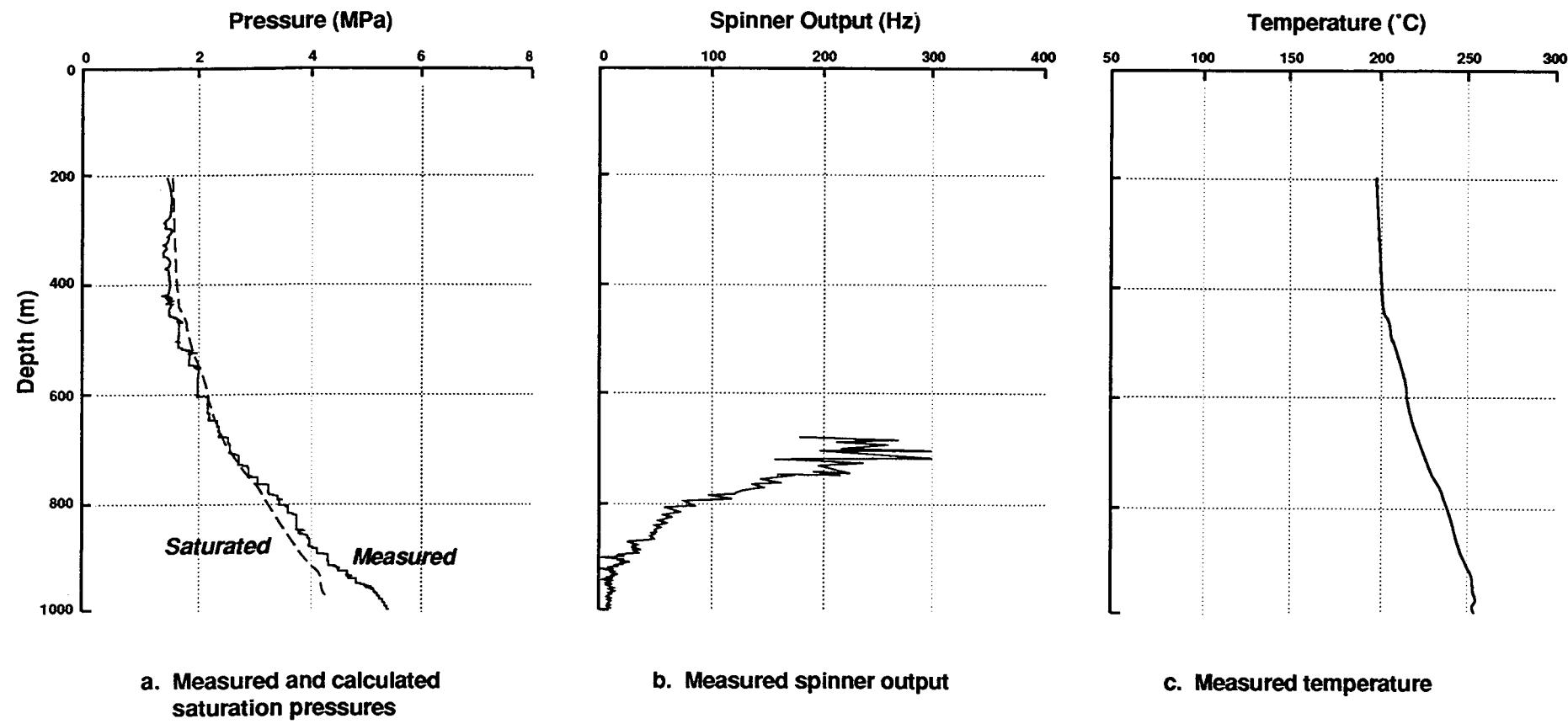


Fig. 10. STP log #4 in well ZCQ-4, December 11, 1989 (log-out). Flowing well.

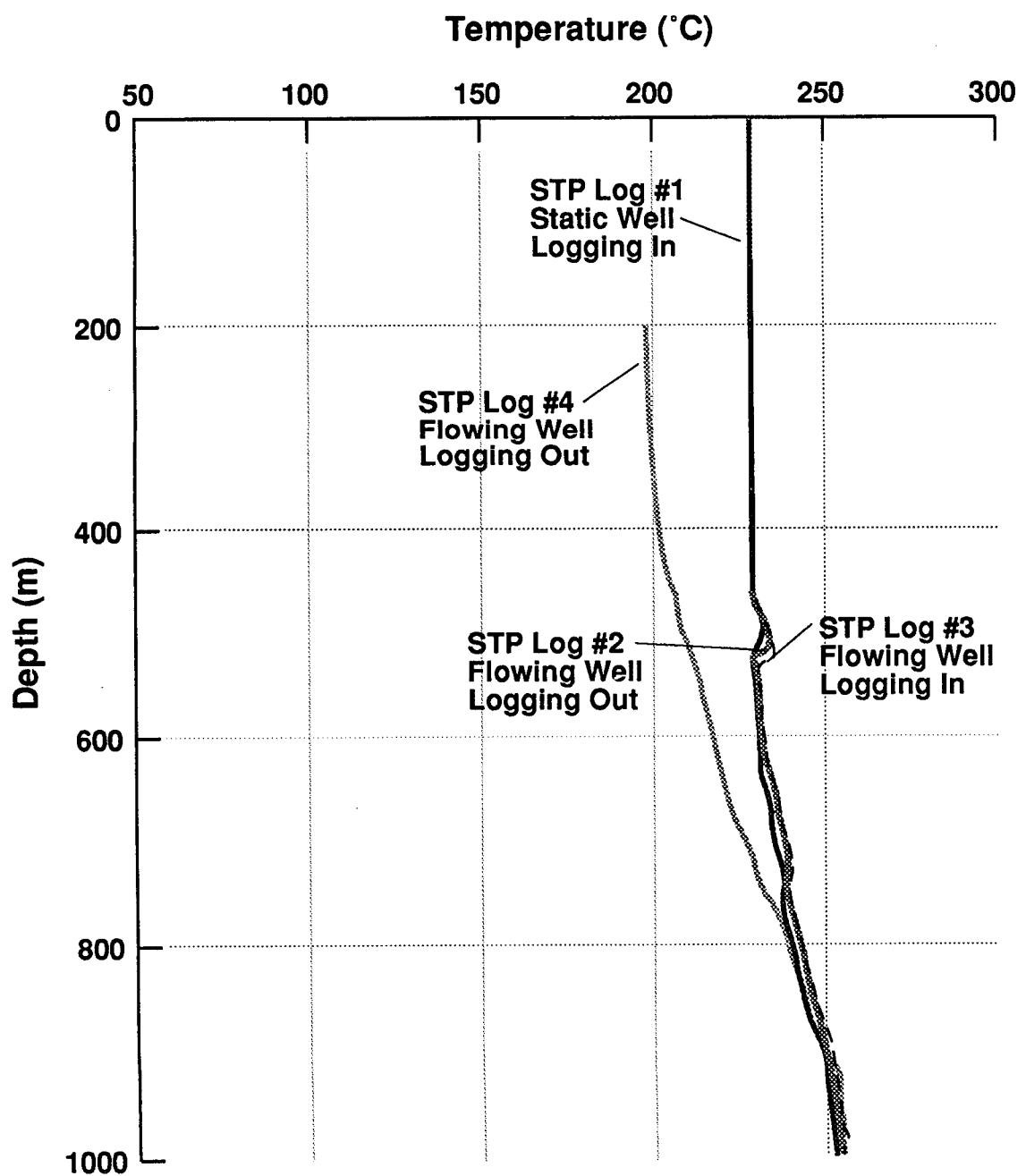


Fig. 11. Temperature data collected in well ZCQ-4.

A static STP log was run in this well to the bottom of the liner primarily for the temperature profile (Fig. 12). The water level in this well was near 200 m, with some boiling taking place (Fig. 13.)

D. ZCQ-6

Another production well, ZCQ-6, was drilled to a depth of 1140 m. This well was cased with 9 5/8-in. casing to 548 m, and the slotted liner was hung at 540 to 1135 m. The static water level was 800 m; the maximum fluid temperature was 270°C. The shut-in wellhead pressure, 180 psi, dropped to 63 psi under flowing conditions.

The pressure lock and Bowen cable pack-off were rigged up and the temperature/rabbit tool was inserted into the well to run a static temperature log. It was standard practice to run the temperature/rabbit tool first to ensure a safe operating region in the borehole before deploying the more expensive well logging tools. We planned to run a temperature log to the bottom of the liner (1140 m), but the tool became set at 794 m. INDE personnel at the site reported a "dog leg" at about 800 m. They had encountered similar problems when running Kuster tools in the well several years before. To get by the obstruction, it had been necessary to flow the well.

ZCQ-6 had been shut-in for at least the past three years, and when the flow control valve to the mufflers was opened for our operation, the wellhead pressure suddenly increased to over 500 psi according to the wellhead pressure gauge. The tool and cable weight indicator also recorded a drop from 700 to 500 lb. The pressure finally decreased to about 150 psi, and the tool was started downhole from 740 m but again became set at 794 m. After several attempts to continue past the obstruction, we aborted the log, shut in the well, and started removing the tool from the hole. However, the tool caught between 615 and 794 m. Repeated runs up the well at various cable velocities failed to free the tool, and the cable finally pulled loose from the cable head, leaving the tool in the well. We assumed that the slotted liner was damaged from the pressure surge when the well was opened to flow. Temperature data recorded in ZCQ-6 during this sequence of events appear in Fig. 14. No other tools were used in this well.

III. HYDROGEOCHEMICAL INVESTIGATIONS

A. FLUID COMPOSITIONS

Field data and analytical data for the thermal and nonthermal waters in the Zunil region appear in Tables I, II, and III. Figure 15 shows the sample locations. Zunil geothermal fluids (<1500 mg/kg chlorine, flash-corrected) occur on the dilute side of the typical range of values reported for geothermal brines (500 to 10,000 mg/kg chlorine; Fournier 1981). Table II shows that the fluids contain relatively high

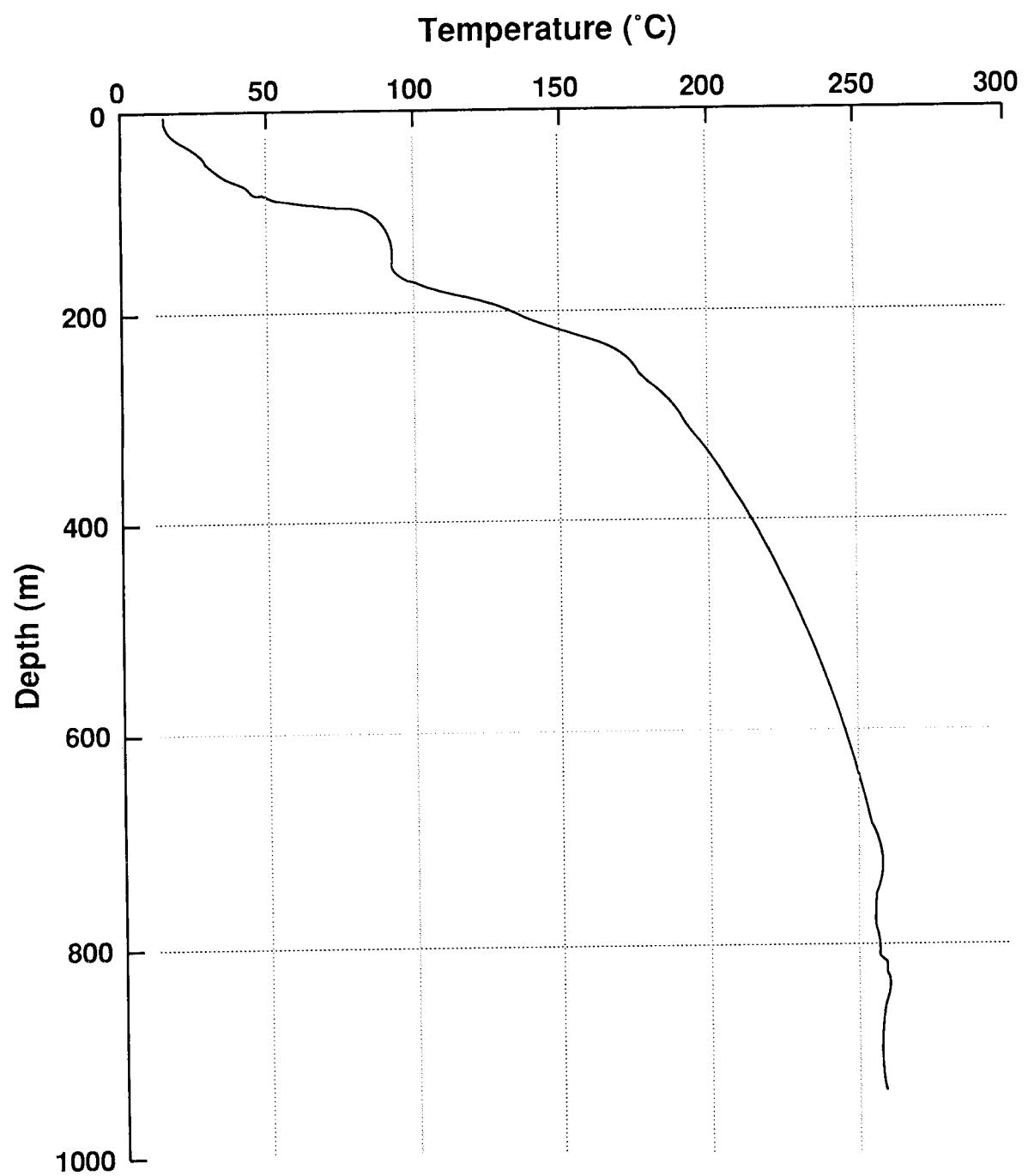


Fig. 11. Temperature data collected in well ZCQ-4.

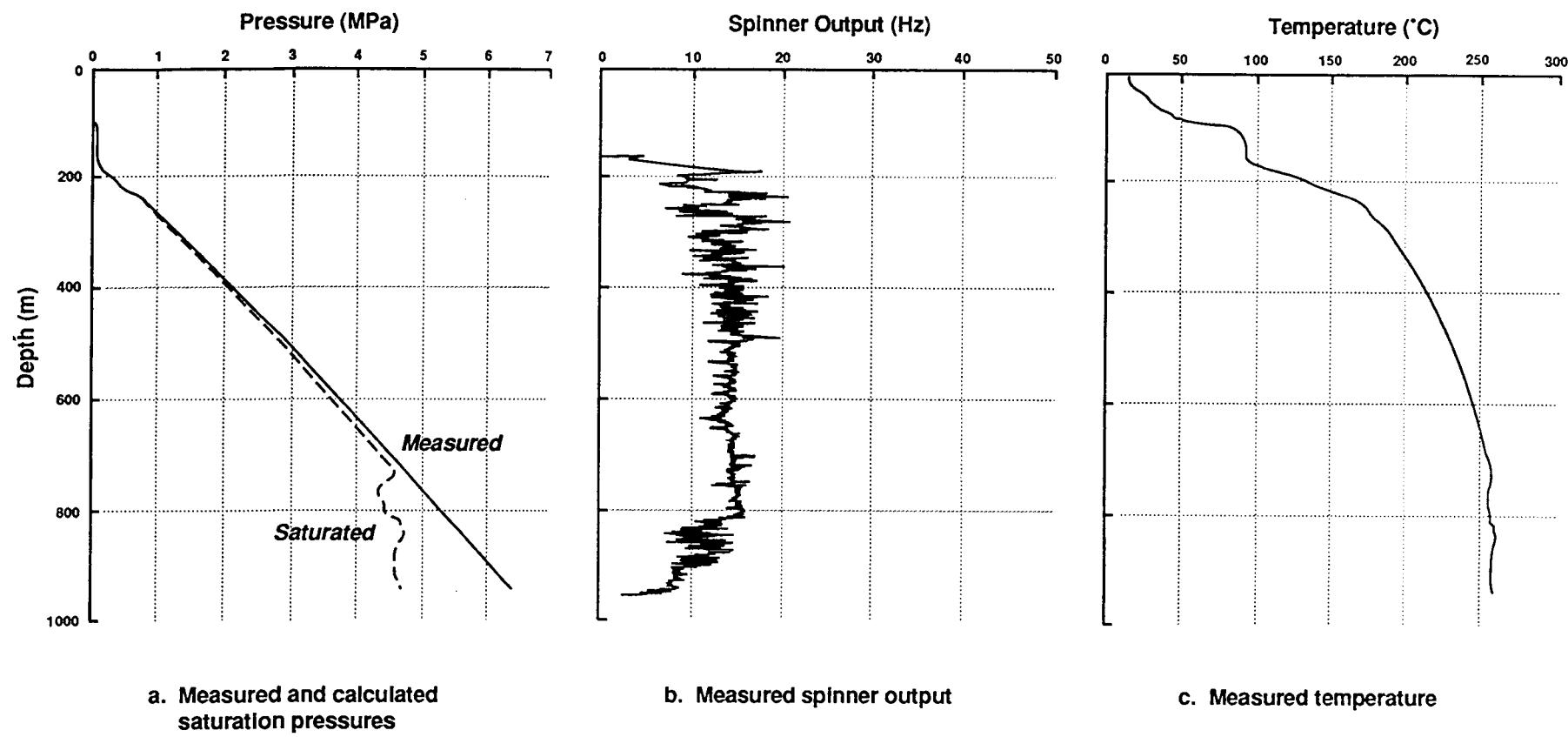


Fig. 13. STP log in well ZCQ-5, November 2, 1988. Static log.

concentrations of lithium, arsenic, boron, and bromine. Because of these high concentrations it is relatively easy to trace mixing patterns of the geothermal fluids as they combine with other types of waters found in the area. Trace element chemistry (Table III) shows the same relationships. For example, sample ZTG-28, a hot spring on the Río Samala north of the town of Zunil, is an obvious mix of thermal water and shallow groundwater. The chemistry of the cold spring (ZTG-12) near Aguas Amargas (ZTG-10, ZTG-11) and the cold spring (ZTG-26) near Baños Sulfurosos (ZTG-25, ZTG-27) shows a thermal influence from the nearby hot springs (Tables II and III) and indicates mixed waters. However, ZTG-19, located 1 km SW of ZTG-28, is composed primarily of thermal water that shows little mixing with shallow groundwater (Figs. 15-19). Compared with geothermal brines, these springs display high HCO_3/Cl , suggesting that subsurface boiling and subsequent loss of CO_2 occurs locally in the reservoir. The CO_2 is added to near-surface groundwater making the springs rich in bicarbonate (Fournier *et al.* 1982).

The compositions of the *in situ* samples are generally similar to those of the other deep reservoir samples (flash-corrected). The waters plot on a relatively crude mixing line. However, a striking range in compositions (770 to 1270 mg/kg chlorine) indicates that producing wells tap horizons of different chemistry. The Zunil field does not consist of one homogeneous aquifer extending throughout the subsurface. Fournier *et al.* (1982) suggests that a deeper, hotter (290°C to 305°C) source reservoir may underlie the current production zones. If so, the reservoir must exist in the granitic rocks that underlie the volcanic section. Exploration of producing horizons in the granitic rocks is one of the current objectives of MK-Ferguson (1988).

B. GAS COMPOSITIONS

Gas compositions of the fumaroles and wells that were sampled are listed in Table IV. Although not many samples were obtained, the samples were relatively rich in CO_2 and H_2S (water-free basis) as is typical of geothermal gases. The data suggest an enrichment in H_2S in fluids east of the main production zone. When our data are combined with the data of Giggenbach (1986, 1988), we find that this enrichment of H_2S is more apparent. The $\text{CO}_2/\text{H}_2\text{S}$ ratio is sometimes used in geothermal exploration to look for upflow zones (Hedenquist *et al.*). The lowest ratio in a suite of samples is often closest to the upflow zone. A general west to east decrease in $\text{CO}_2/\text{H}_2\text{S}$ is evident when all gas data are plotted (Fig. 20), indicating

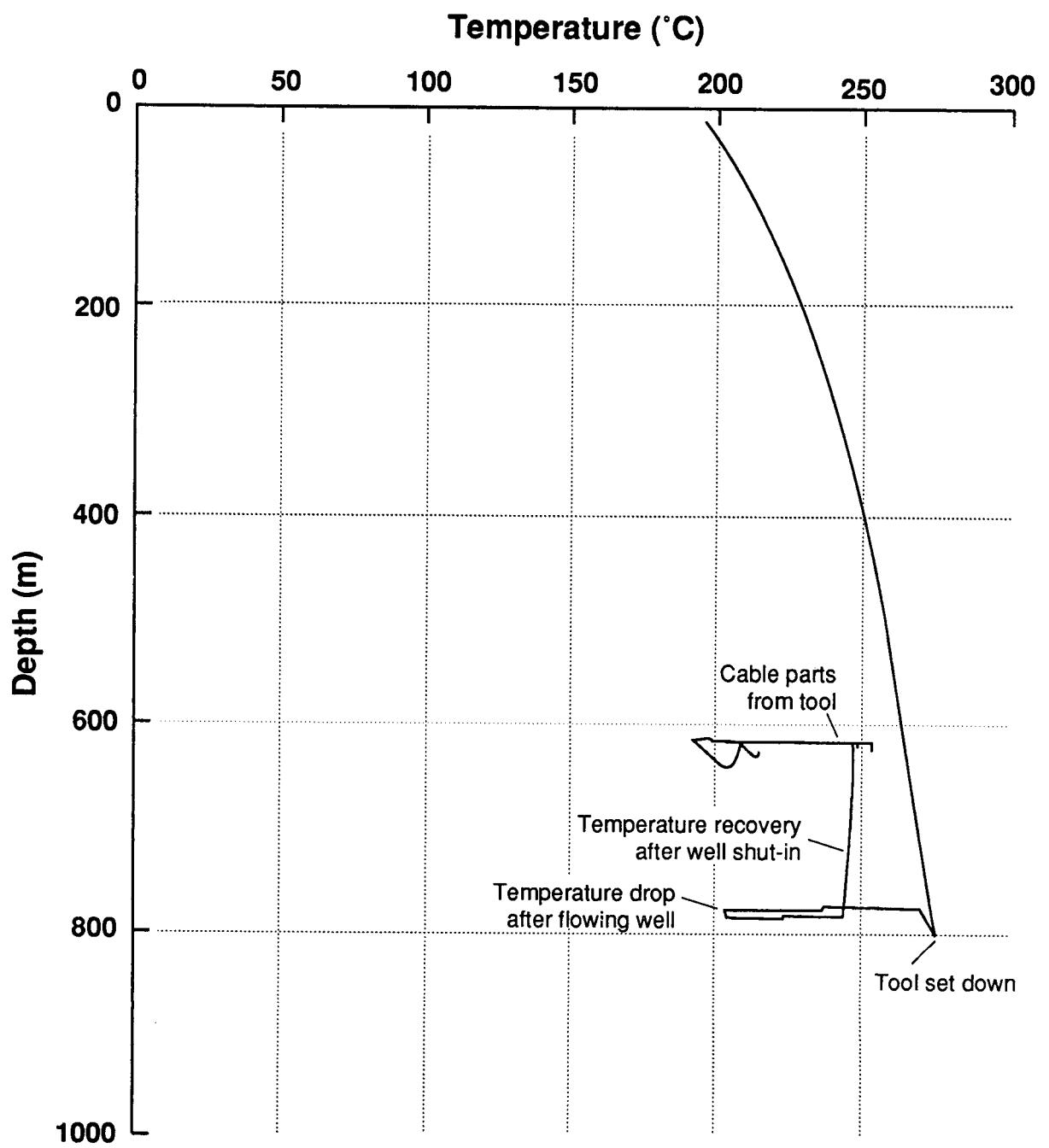


Fig. 14. Measured temperatures in well ZCQ-6, November 1988.

TABLE I. Field Data for All Samples Collected in the Zunil-Quetzaltenango Region, Guatemala

Sample	Description	Date	Temp (°C)	pH (F)	Flow Rate (l/min)	Vol. of Sample Bottle (l)	Vol. of Sample (l)	Depth (m)
ZTG-1	Warm spring (Z-58), 100 m W of ZCQ-5	10/26/88	28.8 ^a	6.5	10	---	---	---
ZTG-2	Cold spring (Z-47B), 100 m S10E of ZTG-1	10/26/88	24.5 ^a	6.5	64	---	---	---
ZTG-3	Hot spring (Z-47), from iron pipe	10/27/88	90.2 ^a	3.5	40	---	---	---
ZTG-4	Hot spring, 5 m SW of ZTG-3	10/27/88	60.8 ^a	3.5	5	---	---	---
ZTG-5	Fuentes Georgines (Z-34)	10/28/88	51.7 ^a	2.0	15	---	---	---
ZTG-6	Fuentes Georgines (Z-31), lower pool	10/28/88	40.2 ^a	2.0	10	---	---	---
ZTG-7	Steam condensate, Las Fresas	10/28/88	94.2 ^a	5.5	---	---	---	---
ZTG-8	Steam condensate (ZF-38)	10/28/88	94.2 ^a	6.5	---	---	---	---
ZTG-9	Fumarole Negro (ZF-36)	10/30/88	94.7 ^a	4.75	---	---	---	---
ZTG-10	Aguas Amargas (Z-19)	10/30/88	53.3 ^a	2.0	---	---	---	---
ZTG-11	Aguas Amargas (Z-19), 20 m NNW of ZTG-10	10/30/88	68.4 ^a	2.0	---	---	---	---
ZTG-12	Aguas Amargas (Z-19A), cold spring	10/30/88	18.3 ^a	3.5	---	---	---	---
ZTG-13	Downhole at ZCQ-3 (671 m)	10/31/88	258 ^b	5.5	---	3.911	2.387	671
ZTG-14	Cold spring (Z-42)	10/31/88	17.3 ^a	6.5	10	---	---	---
ZTG-15	Cold spring, Chica	11/01/88	14.4 ^a	6.5	1	---	---	---
ZTG-16	Weirbox at ZCQ-6	11/01/88	278 ^b	8.5	---	---	---	---
ZTG-17	Rain water at ZCQ-1	11/02/88	---	---	---	---	---	---
ZTG-18	Downhole at ZCQ-5 (775 m)	11/02/88	255 ^b	5.5	---	3.911	2.579	775
ZTG-19	Hot spring (Z-17)	11/03/88	74.4 ^a	7.75	100	---	---	---
ZTG-20	Cerro Quemado, rock sample	11/03/88	---	---	---	---	---	---
ZTG-21	Fumarole Azufalito, gas sample	02/27/89	92.8 ^a	3.00	---	---	---	---
ZTG-22	Fumarole near Georgina, gas sample	02/28/89	94.0 ^a	3.50	---	---	---	---
ZTG-23a	Steam condensate, Fumarole Negro	02/28/89	94.0 ^a	3.25	---	---	---	---
ZTG-23b	Gas sample, Fumarole Negro	02/28/89	94.0 ^a	3.25	---	---	---	---
ZTG-24	Agua Tibia, 1.5 km NE of Ostuncalco	12/06/89	21.2 ^a	7.00	200	---	---	---
ZTG-25	Baños Sulfurosos El Recreo	12/07/89	41.3 ^a	6.00	15	---	---	---
ZTG-26	3 m E of ZTG-22 cold spring	12/07/89	20.1 ^a	6.50	3	---	---	---
ZTG-27	Baños Sulfurosos Well	12/07/89	45.4 ^a	6.50	pumped	---	---	---
ZTG-28	Hot spring on NW bank of Río Samala	12/07/89	61.4 ^a	7.00	10	---	---	---
ZTG-29	ZC-11 direct use well from weirbox	12/08/89	255 ^b	7.50	33	---	---	---
ZTG-30	Make up water for ZC-11, warm spring	12/08/89	30.5 ^a	6.80	5	---	---	---
ZTG-31	Steam side of separator at ZC-11	12/09/89	255 ^b	5.30	---	---	---	---
ZTG-32	Total flow from separator at ZC-11	12/09/89	255 ^b	5.00	---	---	---	---
ZTG-33	Downhole at ZCQ-4 (900 m)	12/11/89	258 ^b	5.00	---	3.911	3.043	900
ZTG-34	Hot well 0.5 km S of Totonicapán	12/11/89	46.0 ^a	6.50	---	---	---	---
ZTG-35	Cold spring 0.7 km W of Totonicapán	12/11/89	19.1 ^a	6.00	7	---	---	---
ZTG-36	Weirbox at ZCQ-4	12/11/89	258 ^b	7.50	1080	---	---	---
ZTG-37	Steam condensate, Baños Los Vahos	12/12/89	26.0 ^a	6.00	---	---	---	---
ZTG-38	Cold spring (Z-73)	12/12/89	14.9 ^a	6.00	seep	---	---	---
ZTG-39	ZP-6 exploration well	12/12/89	94.6 ^a	8.00	15	---	---	---
ZTG-40	Cold spring off cliff above Río Zarmaca	12/12/89	15.3 ^a	6.00	30	---	---	---
ZTG-41	Weirbox at ZCQ-3	12/12/89	258 ^b	8.30	2017	---	---	---

^aCollection temperature.^bFormation temperature of wells from temperature logs is listed if known.

TABLE II. Major Element Chemistry for Wells, Hot/Cold Springs and Fumaroles in the Zunil-Quetzaltenango Region, Guatemala^a (all values in mg/kg, except where noted)

Sample	Description	Temp °C	pH (Lab)	As	B	Br	Ca	Cl	CO ₃	F	HCO ₃	K	Li	Mg	Na	SiO ₂	SO ₄	Sr	TDS	Cation Σ	Anion Σ
Fumaroles, Warm and Hot Springs																					
ZTG-1	Warm spring (Z-58), 100 m W of ZCQ-5	28.8	6.92	<0.05	<0.1	<0.1	20.9	4.1	0	0.46	168	4.7	0.02	16.9	15.4	98	15.0	0.12	351.8	3.233	3.348
ZTG-3	Hot spring (Z-47), from iron pipe	90.2	2.78	<0.05	<0.1	<0.1	41.9	5.4	0	0.38	0	7.9	0.03	23.2	19.2	159	377	0.11	656.2	8.076	8.083
ZTG-4	Hot spring, 5 m SW of ZTG-3	60.8	2.85	<0.05	<0.1	<0.1	47.5	5.7	0	0.41	0	8.0	0.04	25.6	19.2	168	455	0.11	764.7	10.360	9.732
ZTG-5	Fuentes Georgines (Z-34)	51.7	2.44	<0.05	1.7	<0.1	38.4	5.4	0	0.39	0	27	0.05	18.5	79	182	818	0.10	1200.4	17.355	17.409
ZTG-6	Fuentes Georgines (Z-31), lower pool	40.2	2.26	<0.05	1.7	<0.1	64.9	4.6	0	0.42	0	34	0.08	25.5	121	276	1459	0.10	2057.3	31.166	30.805
ZTG-7	Steam condensate, Las Fresas	---	4.59	<0.05	<0.1	<0.1	<0.1	<0.5	0	0.20	22.0	1.1	<0.01	0.27	0.9	<0.2	2.5	<0.01	27.7	0.131	0.424
ZTG-8	Steam condensate (ZF-38)	---	6.45	<0.05	<0.1	<0.1	1.1	0.7	0	0.49	19.5	0.7	<0.01	0.45	1.2	3	5.2	<0.01	35.8	0.282	0.474
ZTG-9	Fumarole Negro, mud pot (ZF-36)	94.7	4.23	<0.05	<0.1	<0.1	49.7	7.3	0	0.24	0	11.6	<0.01	22.5	50	160	344	0.31	658.5	7.530	7.441
ZTG-10	Aguas Amargas (Z-19)	53.3	2.35	<0.05	1.4	<0.1	32.9	4.7	0	0.44	0	23	0.04	10.6	71	191	999	0.18	1391.3	21.618	21.173
ZTG-11	Aguas Amargas (Z-19), 20 m NNW of ZTG-1	68.4	2.28	<0.05	2.3	<0.1	32.7	4.9	0	0.46	0	33	0.05	10.7	84	183	1180	0.18	1593.4	25.215	25.022
ZTG-19	Hot spring (Z-17)	74.4	7.41	0.34	5.6	0.5	40.2	186	0	0.74	665	32	0.54	40.0	313	195	202	0.41	1682.8	19.798	20.639
ZTG-23	Steam condensate, Fumarole Negro	94.0	---	---	<0.05	<0.05	---	0.6	---	0.13	---	0.1	---	---	0.1	---	5.3	---	---	---	
ZTG-25	Baños Sulfurosos, El Recreo	41.3	6.91	<0.05	0.64	0.08	21.0	23.9	0	0.25	193	7.8	0.11	10.46	76	161	64.0	0.08	559.4	5.475	5.232
ZTG-28	Hot spring on NW Bank of Río Samala	61.4	7.60	0.27	2.42	0.23	13.8	70.8	0	0.40	210	14.4	0.28	8.78	116	22	73.2	0.08	670.4	6.869	7.168
ZTG-30	Make-up water for ZC-11, warm spring	30.5	7.70	<0.05	0.17	<0.05	26.0	5.06	53.5	0.19	136	5.7	0.05	22.6	29	535	20.4	0.15	414.6	4.580	4.735
ZTHG-37	Steam condensate, Baños Los Vahos	26.0	6.06	<0.05	0.12	<0.05	<0.1	1.11	0	<0.02	18.3	0.4	0.01	<0.01	0.8	1	0.16	0.04	22.3	0.050	0.341
Cold Springs																					
ZTG-2	Cold spring (Z-47B), 100 m S10E of ZTG-1	24.5	7.19	<0.05	<0.1	<0.1	17.6	4.4	0	0.32	121	3.7	0.02	12.9	10.2	90	15.4	0.09	285.8	2.486	2.608
ZTG-12	Aguas Amargas (Z-19A), cold spring	18.3	2.95	<0.05	0.2	<0.1	25.2	5.7	0	0.33	0	11.2	0.02	8.01	36	118	358	0.14	588.1	7.758	7.722
ZTG-14	Cold spring (Z-42)	17.3	7.11	<0.05	<0.1	<0.1	12.0	2.4	0	0.27	89.1	3.5	<0.01	6.82	9.3	66	1.7	0.09	202.3	1.664	1.759
ZTG-15	Cold spring, Chicua	14.4	7.06	<0.05	<0.1	<0.1	19.6	4.4	0	0.23	98.8	4.5	<0.01	10.8	8.5	68	12.7	0.12	250.4	2.360	2.385
ZTG-24	Agua Tibia, 1.5 km NE of Ostuncalco	21.2	7.50	<0.05	0.11	<0.05	11.5	1.70	0	0.06	78.1	3.7	0.01	5.93	9.0	82	5.07	0.07	205.8	1.557	1.586
ZTG-26	3 m E of ZTG-25 cold spring	20.1	7.23	<0.05	0.37	<0.05	77.7	41.0	0	0.13	153	7.8	0.12	41.4	64	98	84.4	0.26	833.5	10.289	9.737
ZTG-35	Cold spring 0.7 km W of Totonicapán	19.1	7.10	<0.05	<0.05	<0.05	11.0	3.71	0	0.04	80.5	1.2	0.02	5.93	11.9	73	7.93	0.08	204.6	1.594	1.744
ZTG-38	Cold spring (Z-73)	14.9	6.86	<0.05	0.11	<0.05	92.9	5.35	0	0.42	72.0	2.0	0.02	11.6	13.5	73	230	0.05	510.8	6.236	6.347
ZTG-40	Cold spring off cliff on Río Zarmaca	15.3	7.26	<0.05	<0.05	<0.05	11.1	3.55	0	0.03	62.2	3.2	<0.01	4.22	6.9	57	9.60	0.05	163.6	1.289	1.412
Geothermal and Other Wells																					
ZTG-13	Downhole at ZCQ-3 (671 m)	258	6.18	1.21	23.9	4.0	8.1	1065	0	2.49	62.2	122	4.48	0.35	653	460	28.1	0.15	2407.0	29.907	33.178
ZTG-16 ^b	Weirbox at ZCQ-6	278	8.27	3.61	16.27	2.59	6.07	768	36.9	2.08	0	82.9	3.77	0.29	433	291	40	0.15	4221.5	55.485	60.322
ZTG-18	Downhole at ZCQ-5 (775 m)	255	5.70	<0.05	27.8	5.1	18.2	1367	0	3.93	50.0	128	8.12	<0.01	805	470	18.9	0.16	2928.2	40.407	40.995
ZTG-27	Baños Sulfurosos well	45.4	7.11	<0.05	1.02	0.09	22.5	27.8	0	0.15	231	9.8	0.15	11.1	93	172	57.9	0.09	638.9	6.390	6.024
ZTG-29 ^b	ZC-11 direct use well from weirbox	255	8.08	7.33	35.32	4.62	23.34	1269	20.5	1.42	0	71.2	4.57	0.1	822	321	83.5	0.36	3906.3	57.469	58.658
ZTG-31	Steam side of separator at ZC-11	255	6.19	0.14	2.76	0.30	1.5	79.6	0	0.14	34.2	5.2	0.31	0.17	51	22	5.65	0.03	211.9	2.783	3.101
ZTG-32	Total flow from separator at ZC-11	255	6.55	<0.05	<0.05	<0.05	0.2	0.85	0	<0.02	34.2	0.3	0.01	0.10	1.2	3	0.43	<0.01	47.1	0.372	0.620
ZTG-33	Downhole at ZCQ-4 (900 m)	258	5.87	0.16	24.3	3.45	11.2	903	0	2.85	53.7	113	5.26	0.14	601	565	32.0	0.13	2298.5	30.525	28.219
ZTG-34	Hot well 0.5 km S of Totonicapán	46.0	7.63	<0.05	1.20	0.11	11.0	23.5	0	2.03	160	9.4	0.52	2.04	75	111	55.7	0.17	458.8	4.668	4.615
ZTG-36 ^b	Weirbox at ZCQ-4	258	8.25	6.6	26.35	3.56	17.94	931	0	1.36	38.8	94.03	5.59	0.08	603	374	33.2	0.21	3234.4	45.866	43.681
ZTG-39 ^b	ZP-6 exploration well	94.6	8.27	12.34	33.4	4.25	41.141084	19.36	3.37	0	29.6	5.35	0.14	774	320	106	0.71	2370.6	41.509	40.094	
ZTG-41 ^b	Weirbox at ZCQ-3	258	8.41	7.26	27.9	4.24	11.371081	11.8	3.13	0	139	5.75	0.08	718	540	19.5	0.20	3954.7	55.288	51.037	

^a Analyses by P. E. Trujillo (Los Alamos National Laboratory).^b Chemistry corrected for flashing according to formation temperatures (see Table I).

TABLE III. Trace Element Chemistry of Wells, Hot/Cold Springs, and Fumaroles in the Zunil-Quetzaltenango Region, Guatemala (all values in mg/kg)

Sample No.	Description	Ag	Al	Ba	Cd	Co	Cr	Cs	Cu	Fe	Hg	I	Mn	Mo	NH ₄	Ni	NO ₃	Pb	PO ₄	Rb	Sb	Se	S ₂ O ₃	U	Zn
Fumaroles and Warm and Hot Springs																									
ZTG-1	Warm spring (Z-58), 100 m W of ZCQ-5	<0.001	<0.1	0.04	<0.002	<0.002	<0.002	<0.005	0.002	<0.1	<0.1	<0.05	<0.01	<0.002	0.06	0.006	8.3	<0.002	0.2	0.014	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-3	Hot spring (Z-47), from iron pipe	<0.001	4.9	0.02	<0.002	<0.002	<0.002	<0.005	<0.002	11.9	<0.1	<0.05	0.80	<0.002	2.19	0.005	<0.1	<0.002	<0.1	0.042	<0.1	<0.1	<0.05	<0.1	0.07
ZTG-4	Hot spring, 5 m SW of ZTG-3	<0.001	12.2	0.02	<0.002	0.008	0.002	<0.005	<0.002	17.0	<0.1	<0.05	0.90	<0.002	2.43	0.018	0.2	<0.002	<0.1	0.043	<0.1	<0.1	<0.05	<0.1	0.16
ZTG-5	Fuentes Georgines (Z-34)	<0.001	16.1	0.03	<0.002	0.008	0.002	0.011	0.036	4.75	<0.1	<0.05	0.83	<0.002	0.34	0.009	0.1	0.002	<1	0.18	<0.1	<0.1	<0.05	<0.1	0.10
ZTG-6	Fuentes Georgines (Z-31), lower pool	<0.001	33.6	0.06	<0.002	<0.002	0.003	0.013	0.004	18.8	<0.1	<0.05	1.41	0.003	0.68	0.005	<0.1	0.003	<1	0.23	<0.1	<0.1	<0.05	<0.1	0.22
ZTG-7	Steam condensate, Las Fresas	<0.001	<0.1	<0.01	<0.002	<0.002	<0.002	0.008	<0.002	0.07	<0.1	<0.05	<0.01	<0.002	0.67	<0.002	<0.1	<0.002	<0.1	0.008	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-8	Steam condensate (ZF-38)	<0.001	<0.1	<0.01	<0.002	<0.002	0.030	<0.005	0.005	0.27	<0.1	<0.05	<0.01	<0.002	2.90	0.046	<0.1	<0.002	<0.1	<0.005	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-9	Fumarole Negro, mud pot (ZF-36)	<0.001	0.4	0.07	<0.002	<0.002	0.004	0.006	0.004	0.33	<0.1	<0.05	0.32	<0.002	11.7	0.015	0.3	<0.002	<0.1	0.045	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-10	Aguas Amargas (Z-19)	<0.001	36.3	0.03	<0.002	<0.002	0.002	<0.005	0.019	8.43	<0.1	<0.05	0.74	<0.002	0.52	0.009	0.3	0.002	<1	0.11	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-11	Aguas Amargas (Z-19), 20 m NNW of ZTG-10	<0.001	35.1	0.03	<0.002	<0.002	0.002	<0.005	0.007	10.4	<0.1	<0.05	0.98	<0.002	0.54	0.009	0.9	0.003	<1	0.12	<0.1	<0.1	<0.05	<0.1	0.12
ZTG-19	Hot spring (Z-17)	<0.001	<0.1	<0.01	<0.002	<0.002	<0.002	0.033	<0.002	<0.01	0.24	0.19	0.018	0.07	<0.002	<0.1	<0.002	<0.1	0.073	<0.1	<0.1	<0.05	<0.1	<0.01	
ZTG-23	Steam condensate, Fumarole Negro	---	---	---	---	---	---	---	---	---	---	---	---	---	0.49	---	<0.05	---	<0.05	---	---	---	---	---	
ZTG-25	Baños Sulfurosos, El Recreo	<0.001	<0.1	0.13	<0.001	<0.002	<0.002	<0.002	<0.002	0.08	<0.1	0.02	0.63	<0.002	0.28	0.005	<0.05	<0.002	0.41	0.023	<0.05	<0.1	<0.01	<0.1	0.03
ZTG-28	Hot spring on NW bank of Río Samala	<0.001	<0.1	<0.01	<0.001	<0.002	<0.002	0.019	<0.002	<0.01	<0.1	0.08	0.04	0.010	0.05	<0.002	3.88	<0.002	<0.1	0.048	<0.05	<0.1	<0.01	<0.1	0.01
ZTG-30	Make up water for ZC-11, warm spring	<0.001	<0.1	0.05	<0.001	<0.002	<0.002	0.011	<0.002	0.04	<0.1	<0.01	0.01	0.003	0.05	<0.002	8.0	<0.002	0.82	0.018	<0.05	<0.1	<0.01	<0.1	<0.01
ZTG-37	Steam condensate, Baños Los Vahos	<0.001	<0.1	<0.01	<0.001	<0.002	<0.002	<0.002	<0.002	0.02	<0.1	<0.01	<0.01	<0.002	<0.05	<0.002	0.06	<0.002	<0.1	<0.002	<0.05	<0.1	<0.01	<0.1	0.04
Cold Springs																									
ZTG-2	Cold Spring (Z-47B), 100 m S10E of ZTG-1	<0.001	<0.1	0.02	<0.002	<0.002	<0.002	<0.005	<0.002	<0.01	<0.1	<0.05	<0.01	<0.002	0.06	0.004	9.6	<0.002	0.2	0.012	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-12	Aguas Amargas (Z-19A), cold spring	<0.001	18.9	0.03	<0.002	<0.002	<0.002	<0.005	0.004	3.16	<0.1	<0.05	0.29	<0.002	0.07	0.015	1.4	<0.002	<1	0.045	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-14	Cold spring (Z-42)	<0.001	<0.1	0.05	<0.002	<0.002	<0.002	<0.005	<0.002	0.01	<0.1	<0.05	<0.01	<0.003	0.12	<0.002	11.2	<0.002	<0.1	0.006	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-15	Cold spring, Chicua	<0.001	<0.1	0.11	<0.002	<0.002	<0.002	<0.007	<0.002	<0.01	<0.1	<0.05	<0.01	<0.002	0.08	0.002	22.5	<0.002	<0.1	0.009	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-24	Aguas Tibia, 1.5 km NE of Ostuncalco	<0.001	<0.1	0.02	<0.001	<0.002	<0.002	<0.002	0.006	0.05	<0.1	<0.01	0.02	<0.002	0.05	<0.002	8.7	<0.002	0.11	0.014	<0.05	<0.1	<0.01	<0.1	<0.01
ZTG-26	3 m east of ZTG-25 cold spring	<0.001	<0.1	0.08	<0.001	<0.002	<0.002	<0.002	<0.002	<0.01	<0.1	<0.01	0.02	<0.002	0.06	<0.002	265	<0.002	0.59	0.020	<0.05	<0.1	<0.01	<0.1	0.01
ZTG-35	Cold spring 0.7 km W of Totonicapán	<0.001	<0.1	0.02	<0.001	<0.002	<0.002	<0.002	0.002	0.02	<0.1	<0.01	<0.002	0.06	<0.002	9.4	<0.002	<0.1	0.003	<0.05	<0.1	<0.01	<0.1	<0.01	
ZTG-38	Cold spring (Z-73)	<0.001	<0.1	<0.01	<0.001	<0.002	<0.002	<0.002	0.003	<0.1	<0.01	0.02	<0.002	0.06	0.004	9.9	<0.002	0.08	0.007	<0.05	<0.1	<0.01	<0.1	0.03	
ZTG-40	Cold spring off cliff on Río Zarmaca	<0.001	<0.1	<0.01	<0.001	<0.002	<0.002	0.005	<0.002	0.02	<0.1	<0.01	0.004	0.07	0.002	5.41	<0.002	0.09	0.010	<0.05	<0.1	<0.01	<0.1	<0.01	
Geothermal and other Wells																									
ZTG-13	Downhole at ZCQ-3 (671 m)	<0.001	0.5	0.02	<0.002	<0.002	0.003	0.58	<0.002	0.47	<0.1	1.79	0.05	0.015	1.96	0.005	<0.1	<0.002	<0.1	0.77	<0.1	<0.1	<0.05	<0.1	<0.01
ZTG-16a	Weirbox at ZCQ-6	<0.001	0.69	<0.01	<0.002	<0.002	<0.002	0.59	<0.002	0.45	<0.1	2.06	<0.01	0.028	0.86	0.016	0.75	<0.002	<0.1	0.885	0.38	<0.1	<0.05	<0.1	<0.01
ZTG-18	Downhole at ZCQ-5 (775 m)	<0.001	<0.1	0.13	<0.002	<0.002	0.007	0.92	<0.002	0.62	<0.1	2.73	0.08	0.004	1.65	0.009	<0.1	<0.002	<0.1	1.11	<0.1	<0.1	2.90	<0.1	<0.01
ZTG-27	Baños Sulfurosos well	<0.001	<0.1	0.13	<0.001	<0.002	<0.002	0.007	<0.002	<0.01	<0.1	0.02	0.54	<0.002	0.28	<0.002	11.2	<0.002	0.27	0.027	<0.05	<0.1	<0.01	<0.1	0.01
ZTG-29a	ZC-11 direct use well from weirbox	<0.001	0.27	0.03	<0.001	<0.002	<0.002	0.007	<0.002	<0.01	<0.1	2.27	0.01	0.023	0.49	0.003	<0.1	<0.002	<0.2	0.38	0.18	<0.1	0.77	<0.1	0.01
ZTG-31	Stream side of separator at ZC-11	<0.001	<0.1	0.08	<0.001	<0.002	0.002	0.073	<0.002	0.35	<0.1	0.16	0.02	<0.002	4.95	0.004	0.43	0.002	<0.1	0.028	<0.05	<0.1	2.32	<0.1	0.05
ZTG-32	Total flow from separator at ZC-11	<0.001	<0.1	0.03	<0.001	<0.002	<0.002	<0.002	0.004	<0.1	<0.01	0.01	<0.002	5.22	0.006	0.14	0.003	<0.1	<0.002	<0.05	<0.1	1.35	<0.1	0.02	
ZTG-33	Downhole at ZCQ-4 (900 m)	<0.001	0.5	0.03	<0.001	<0.002	0.003	0.70	<0.002	1.08	<0.1	1.60	0.20	<0.002	1.59	0.009	0.08	<0.002	<0.1	1.03	<0.1	<0.1	6.49	<0.1	0.02
ZTG-34	Hot well 0.5 km S of Totonicapán	<0.001	<0.1	0.06	<0.001	<0.002	<0.002	0.13	0.003	6.50	<0.1	0.02	0.17	<0.002	0.24	<0.002	0.10	<0.002	<0.1	0.081	<0.05	<0.1	<0.01	<0.1	0.01
ZTG-36a	Weirbox at ZCQ-4	<0.001	0.46	0.03	<0.001	<0.002	<0.002	0.64	<0.002	0.3	<0.1	1.75	0.02	0.019	0.44	<0.002	<0.1	<0.002	<0.2	0.9	0.53	<0.1	0.58	<0.1	0.02
ZTG-39a	ZP-6 exploration well	<0.001	0.18	<0.01	<0.001	<0.002	<0.002	0.84	<0.002	<0.02	<0.1	2.09	0.02	0.019	0.37	<0.002	<0.1	<0.002	<0.2	0.032	0.63	<0.1	0.56	<0.1	0.07
ZTG-41a	Weirbox at ZCQ-3	<0.001	0.59	0.01	<0.001	<0.002	<0.002	0.93	<0.002	0.05	<0.1	2.01	<0.01	0.034	0.26	<0.002	<0.1	<0.002	<0.2	1.37	0.52	<0.1	0.23	<0.1	0.02

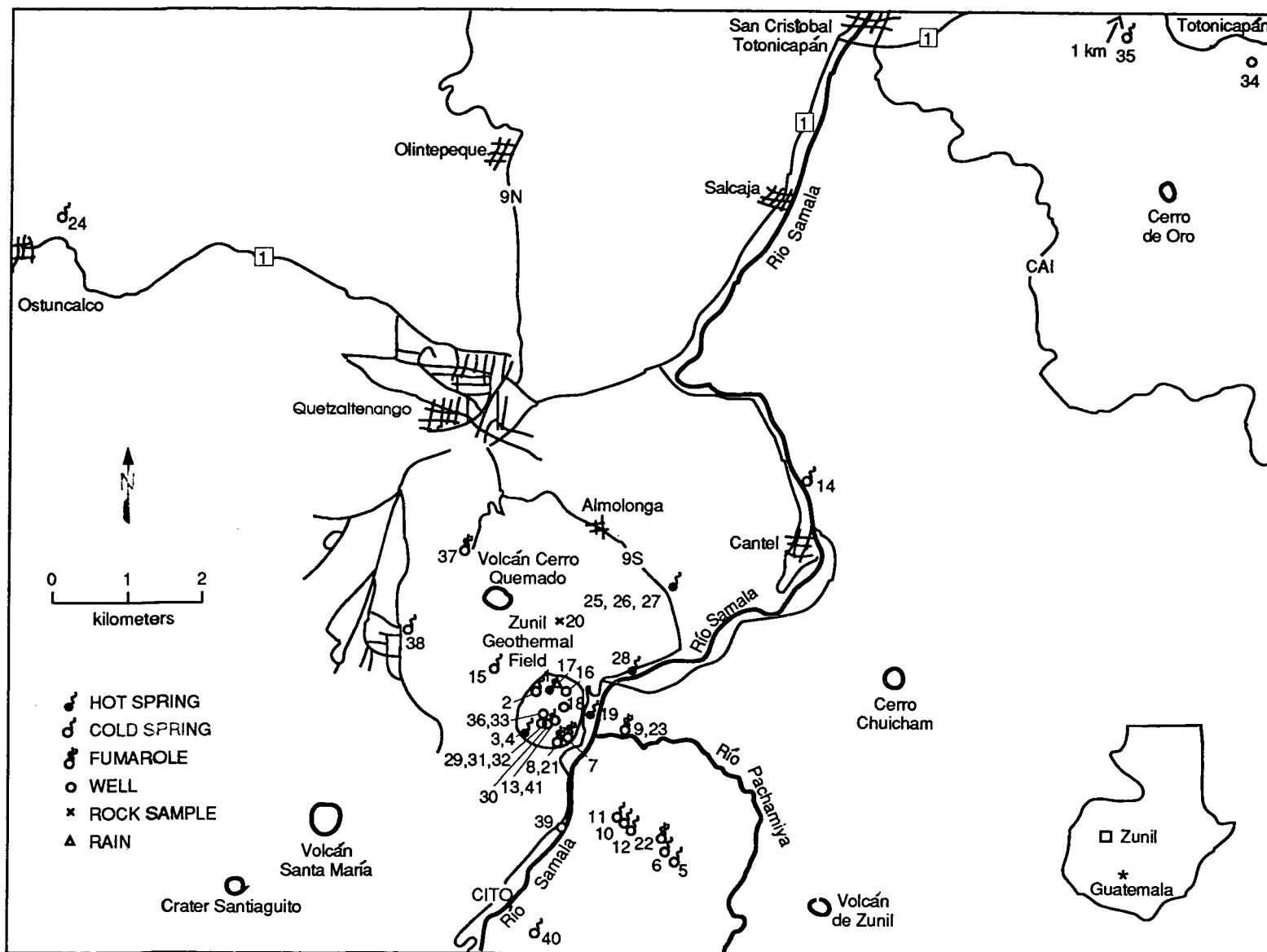


Fig. 15. Location map of Zunil-Quetzaltenango region, Guatemala, showing sample locations.

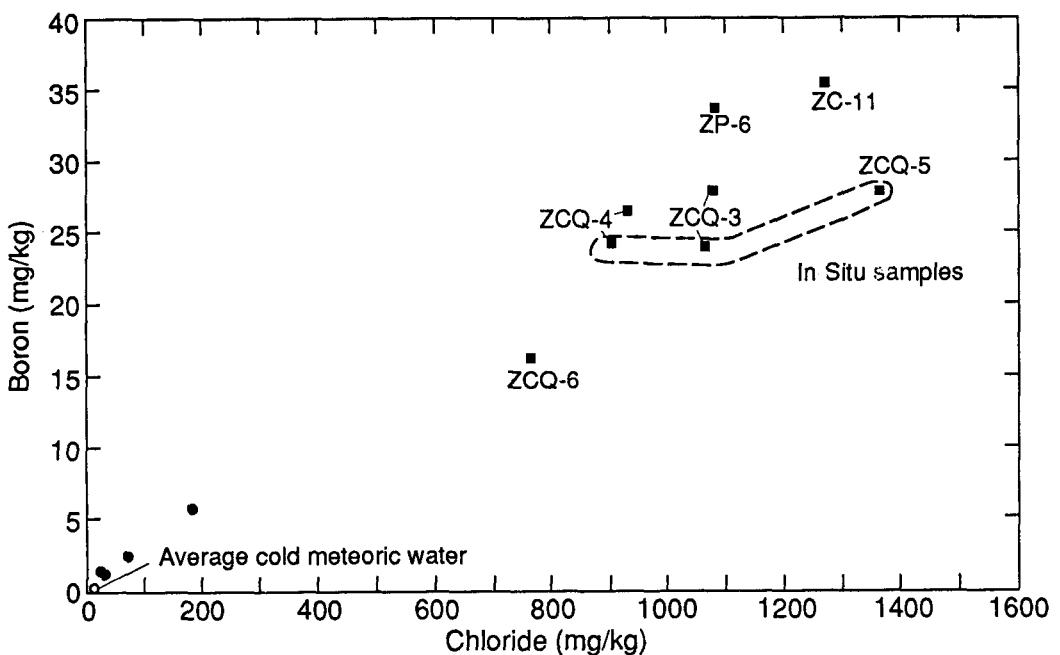


Fig. 16. Boron versus chloride for reservoir fluids, hot springs, and average values of the cold waters sampled, Zunil geothermal field, Guatemala. All data are flash-corrected where appropriate. Well and spring numbers can be found in Table I. ○ = average cold springs; ● = hot springs; ■ = well samples.

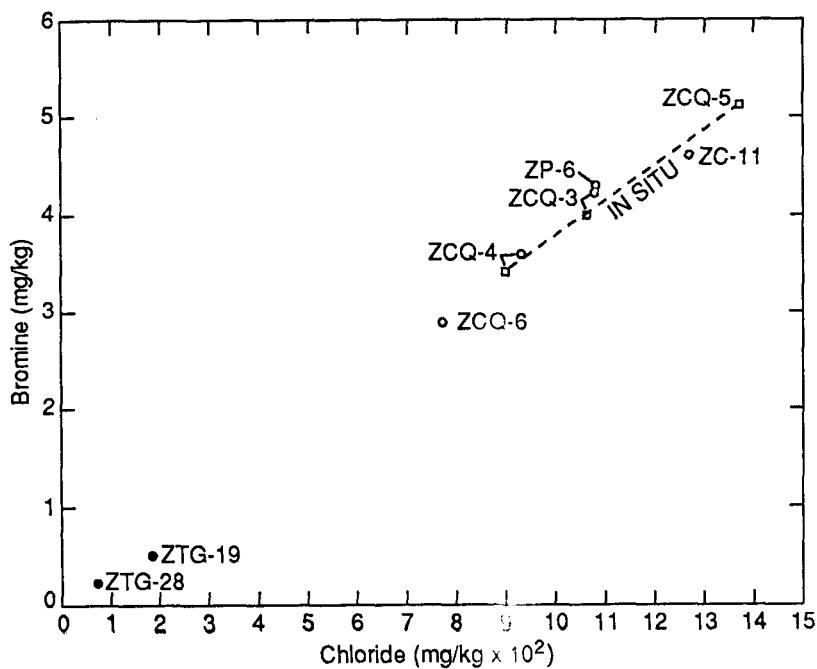


Fig. 17. Bromine versus chloride for reservoir fluids and hot springs waters sampled, Zunil-Quetzaltenango region, Guatemala. All data are flash-corrected where appropriate. Well and spring numbers can be found in Table I. ○ = well samples, weirboxes; ■ = well samples *in situ*; ● = hot springs.

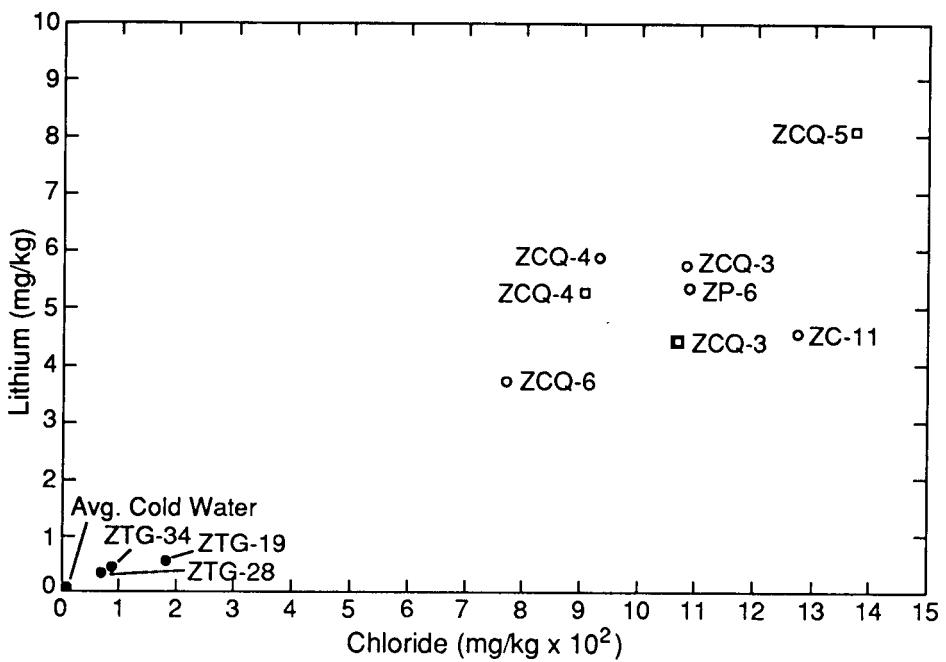


Fig. 18. Lithium versus chloride for reservoir fluids, hot springs, and average values of the cold waters sampled, Zunil-Quetzaltenango region, Guatemala. All data are flash-corrected where appropriate. Well and spring numbers can be found in Table I. \circ = well waters, weirbox; \bullet = hot springs; \square = well waters *in situ*.

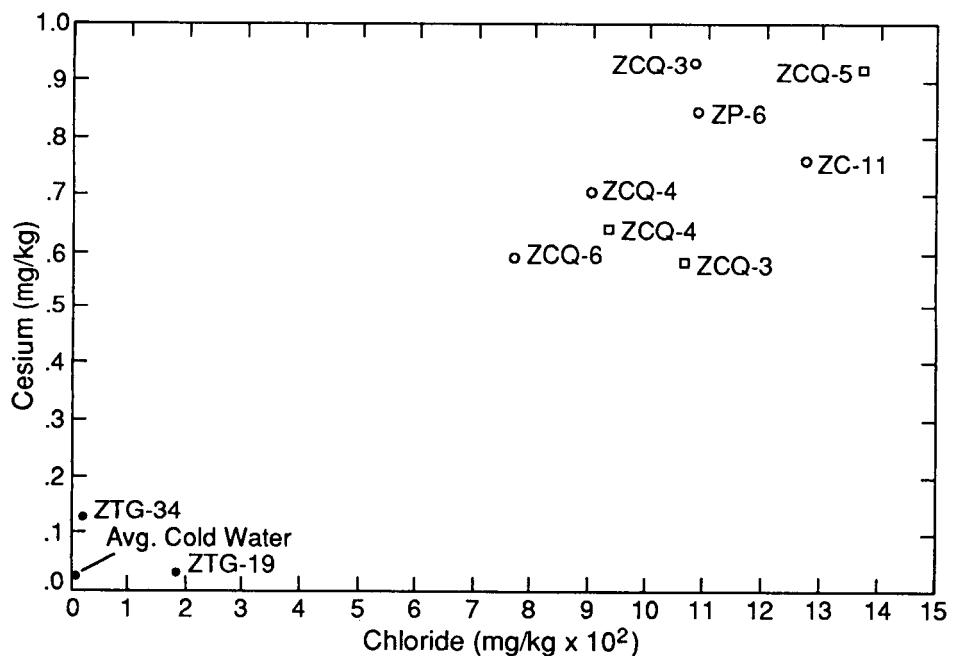


Fig. 19. Cesium versus chloride for reservoir fluids, hot springs, and average values of the cold waters sampled, Zunil-Quetzaltenango region, Guatemala. All data are flash-corrected where appropriate. Well and spring numbers can be found in Table I. \bullet = springs; \circ = wells, weirbox; \square = well samples *in situ*.

that a thermal plume originates from beneath the zone of acid-sulfate springs to the east of the production wells. This figure suggests that the Zunil II area might be a better geothermal prospect than Zunil I. Giggenbach (1988) proposed a similar type of model.

C. STABLE ISOTOPES

Figure 21 and Table V present stable isotope variations of the Zunil area fluids. With the exception of the acid-sulfate waters, all cold and hot springs plot on or near the isotopic composition of meteoric waters for the area and have a δD range of -73 to -87. The *in situ* fluid samples have a very restricted δD range of -76.6 to -78.2 compared to flash-corrected reservoir samples and probably represent the isotopic composition of the unflushed reservoir fluid. The rainwater sample is not typical of the local groundwaters having a δD of -17.5 and $\delta^{18}O$ of -3.04. This probably represents a storm of unusual composition.

The trio of *in situ* well samples is shifted to the right of the world meteoric water line as are the weirbox samples indicating that significant oxygen-18 enrichment has occurred between water and rock in a high-temperature environment. Oxygen-18 shifts of 2‰ to 12‰ are characteristic of geothermal reservoirs $>200^{\circ}\text{C}$ and the 3‰ shift of the *in situ* samples is quite typical. According to the isotopic composition of a limited number of cold springs and wells, the mountains to the north and east of the field are logical recharge areas for the geothermal reservoir.

D. CHEMICAL GEOTHERMOMETRY

Table VI lists calculated subsurface equilibration temperatures for well waters and bicarbonate-rich hot spring waters. Measured temperatures in the feed zones of the wells logged by LANL range from 255°C to 278°C . There is reasonably good agreement between measured temperatures and quartz-conductive, Na/K, and Na-K-Ca geothermometers. Temperatures calculated by the Na/Li and K/Mg geothermometers are consistently lower than the measured temperatures. From the K-Mg relation and other arguments, Giggenbach (1988) suggests that the fluids have reequilibrated to lower temperatures from their original temperatures. Differences in chemistry between *in situ* samples and weirbox samples for a given well and the resulting differences in geothermometer estimates probably indicate that the wells produce fluids from several horizons of chemistry slightly different from that reflected by the single feed zones sampled *in situ*. The unusually dilute chemistry of well ZCQ-6 and the relatively low geothermometer estimates based on its chemistry may indicate that this well has low permeability and is an excess enthalpy well.

TABLE IV. Gas Analyses of Well and Fumarole Samples, Zunil, Guatemala^a (all values in vol% dry gas)

Sample	Date	Description	CO ₂	H ₂ S	NH ₃	Ar	N ₂	O ₂	CH ₄	H ₂	He	Total
ZTG-21 ^b	02/27/89	Fumarole Azufralito	96.99	1.23	0	0.0075	1.65	0.0018	0.049	0.016	0.0018	99.94
ZTG-22	02/28/89	Fumarole near Georgina	87.74	9.48	0	0.0337	2.58	0.0081	0.043	0.068	0.0203	99.97
ZTG-23	02/28/89	Fumarole Negro	93.93	2.78	0	0.0510	2.67	0.0398	0.085	0.391	0.0178	99.87
ZTG-31	12/11/89	Well ZC-11, mini-separator	97.43	0.59	0.19	0.0167	1.43	0.0	0.025	0.372	0.0073	100.06
ZTG-33	12/11/89	Well ZCQ-4 downhole @ 900 m	82.09	<0.05	--	0.13	16.84	0.0	<0.05	0.58	<0.01	98.93

^a Analyses by P. Trujillo, LANL.

^b Analysis for ZTG-21 includes 0.0051 vol% C₂H₆.

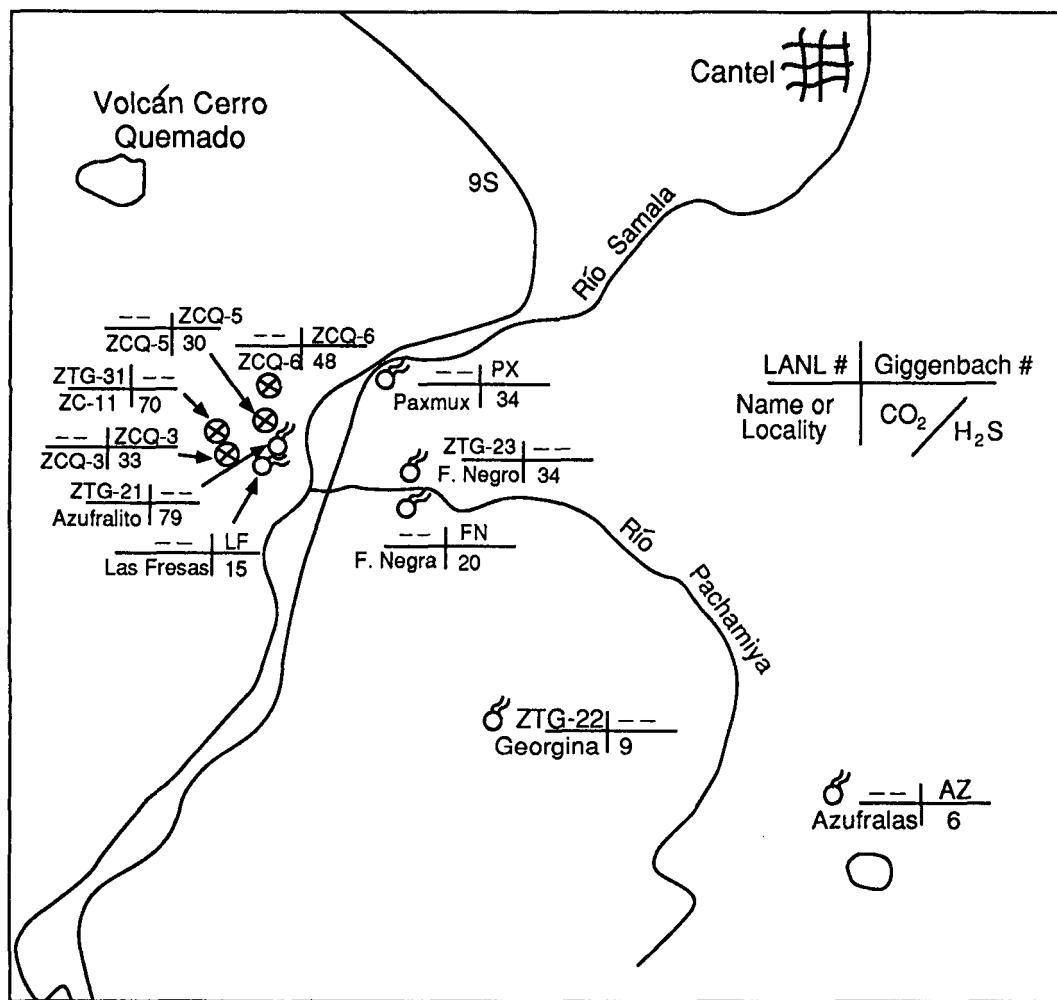


Fig. 20. Map showing CO₂/H₂S for wells and fumaroles in the Zunil-Quetzaltenango region, Guatemala.

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TABLE V. Isotope Data for Geothermal Wells, Hot/Cold Springs and Fumaroles in the Zunil-Quetzaltenango Region, Guatemala

Sample	Description	Date	δD (‰)	$\delta^{18}O$ (‰)	Tritium (T.U.)
ZTG-1	Warm spring (Z-58), 100 m W of ZCQ-5	10/26/88	-87.2	-12.19	2.9
ZTG-2	Cold spring (Z-47B), 100 m S10E of ZTG-1	10/26/88	-87.6	-12.06	1.71
ZTG-3	Hot spring (Z-47), from iron pipe	10/27/88	-84.0	-11.47	3.47
ZTG-4	Hot spring, 5 m SW of ZTG-3	10/27/88	-81.5	-11.42	---
ZTG-5	Fuentes Georgines (Z-34)	10/28/88	-79.4	-10.58	6.23
ZTG-6	Fuentes Georgines (Z-31), lower pool	10/28/88	-78.4	-9.40	---
ZTG-9	Fumarole Negro (ZF-36)	10/30/88	-64.8	-7.18	---
ZTG-10	Aguas Amargas (Z-19)	10/30/88	78.6	-9.57	---
ZTG-11	Aguas Amargas (Z-19), 20 m NNW of ZTG-10	10/30/88	-74.8	-9.04	5.24
ZTG-12	Aguas Amargas (Z-19A), cold spring	10/30/88	-74.4	-11.03	---
ZTG-13	Downhole at ZCQ-3 (671 m)	10/31/88	76.7	-7.92	5.25
ZTG-14	Cold spring (Z-42)	10/31/88	-80.1	-11.12	---
ZTG-15	Cold spring, Chicua	11/01/88	-86.7	-11.86	6.35
ZTG-16	Weirbox at ZCQ-6	11/01/88	-72.3 ^a	-7.40 ^a	0.32
ZTG-17	Rainwater at ZCQ-1	11/02/88	-17.5	-3.04	3.89
ZTG-18	Downhole at ZCQ-5 (775 m)	11/02/88	-78.2	-8.14	14.8
ZTG-19	Hot spring (Z-17)	11/03/88	-83.0	-10.94	0.16
ZTG-23	Steam condensate, Fumarole Negro	02/28/89	-86.5	-12.75	---
ZTG-24	Agua Tibia, 1.5 km NE of Ostuncalco	12/06/89	-81.7	-11.43	---
ZTG-25	Baños Sulfurosos El Recreo	12/07/89	-81.8	-11.23	---
ZTG-26	3 m E. of ZTG-22 cold spring	12/07/89	-78.5	-10.35	---
ZTG-27	Baños Sulfurosos well	12/07/89	-82.0	-11.24	---
ZTG-28	Hot spring on NW bank of Río Samala	12/07/89	-80.3	-11.02	---
ZTG-29	ZC-11 direct use well from weirbox	12/08/89	-61.7 ^a	-6.44 ^a	---
ZTG-30	Make up water for ZC-11, warm spring	12/08/89	-88.3	-12.57	---
ZTG-31	Steam side of separator at ZC-11	12/09/89	-74.1	-9.42	---
ZTG-32	Total flow from separator at ZC-11	12/09/89	-77.8	-9.45	---
ZTG-33	Downhole at ZCQ-4 (900 m)	12/11/89	-77.1	-8.22	0.02
ZTG-34	Hot well 0.5 km S of Totonicapán	12/11/89	-77.3	-10.38	---
ZTG-35	Cold spring 0.7 km W of Totonicapán	12/11/89	-77.3	-9.94	---
ZTG-36	Weirbox at ZCQ-4	12/11/89	-78.9 ^a	-7.94 ^a	0.06
ZTG-37	Baños Los Vahos, steam condensate	12/12/89	-82.0	-11.24	---
ZTG-38	Cold spring (Z-73)	12/12/89	-84.4	-11.56	---
ZTG-39	ZP-6 exploration well	12/12/89	-74.8	-7.40	---
ZTG-40	Cold spring off cliff above Río Zarmaca	12/12/89	-75.8	-10.82	---
ZTG-41	Weirbox at ZCQ-3	12/12/89	-76.8 ^a	-7.93 ^a	0.11

^aIsotopes corrected for flashing according to formation temperatures (see Table I).

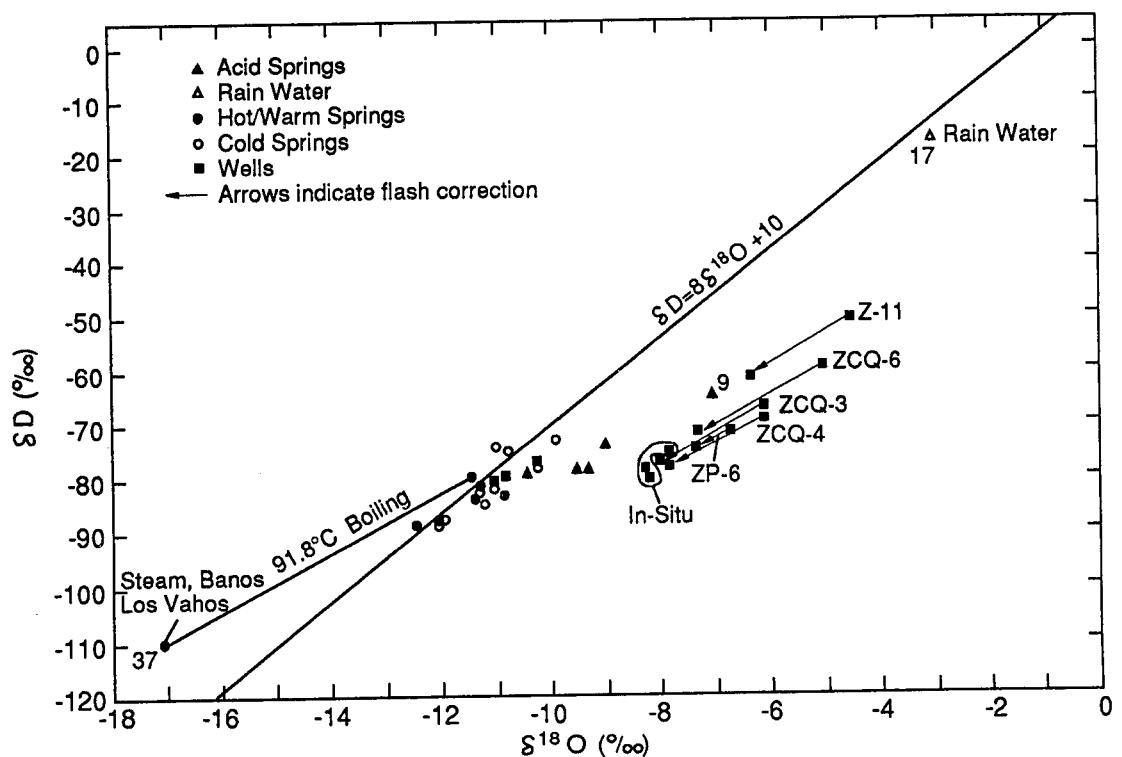


Fig. 21. Plot of δD versus $\delta^{18}O$ for thermal and nonthermal waters of Zunil-Quetzaltenango region, Guatemala.

TABLE VI. Calculated "Reservoir" Temperatures Based on Several Chemical Geothermometers

Sample	Description	Temp. ^a (°C)	Cl (mg/kg)	Total HCO_3 (mg/kg)	Silica ^b		Na-K-Ca ^b			Na/Li ^c	K/Mg ^d	
					Chal.	Qtz-C	Na/K ^b	$\beta=1/3$	$\beta=4/3$	Mg-Corr.		
<u>Geothermal Wells</u>												
ZTG-13	ZCQ-3; <i>in situ</i>	258	1065	62	235	245	277	258	--	--	221	200
ZTG-41	ZCQ-3; weirbox	258	1081 ^e	37	253	259	281	264	--	--	238	251
ZTG-33	ZCQ-4; <i>in situ</i>	258	903	54	259	264	278	253	--	--	248	217
ZTG-36	ZCQ-4; weirbox	258	931 ^e	59	214	227	258	239	--	--	255	232
ZTG-18	ZCQ-5; <i>in situ</i>	255	1367	50	238	247	260	242	--	--	266	—
ZTG-16	ZCQ-6; weirbox	278	768 ^e	120	190	207	280	259	--	--	248	197
ZTG-29	ZC-11; weirbox	255	1269 ^e	61	199	215	205	205	--	--	200	212
ZTG-39	ZP-6 exploration well	90.4	1084 ^e	44	209	223	146	152	--	--	222	163
ZTG-27	Baños Sulfurosos Well	45.4	27.8	231	147	170	222	—	95	43	104	66
ZTG-34	Well near Totonicapán	46.0	23.5	160	117	143	237	183	--	99	222	85
<u>Warm/Hot Springs</u>												
ZTG-1	Spring near ZCQ-5	28.8	4	168	109	136	336	—	54	--	91	46
ZTG-19	Spring Z-17	74.4	186	665	156	178	219	187	--	"cool"	108	78
ZTG-25	Spring, Baños Sulfurosos	41.3	24	193	142	166	219	—	87	44	97	61
ZTG-28	Spring, Rio Samala	61.4	71	210	35	67	236	189	—	31	131	77
ZTG-30	Spring near ZC-11	30.5	5	136	252	259	283	—	62	"cool"	108	47

^a Formation temperatures are listed where known.

^b Equations of Fournier (1981).

^c Fouillac and Michard (1981).

^d Giggenbach *et al.* (1986b).

^e Flash-corrected values assuming temperature of flashing is 93.5°C.

Samala is composed primarily of much older water (0.16 T.U.). Four tritium samples of Zunil production well waters have tritium values of 0.02 to 0.32 indicating they are also primarily composed of "old" water. Assuming that the Zunil reservoir is well mixed, the maximum mean age of the deep well water ranges from 500 to 7500 years (Goff *et al.* 1987b; Person and Truesdell 1978).

TABLE VII. Calculated "Reservoir" Temperatures Based on Gas Geothermometers

Field	Site	Date	Temp Collection	CO ₂	H ₂ S	Geothermometers	
						t gas ^a	t CO ₂ -CH ₄ ^b
ZTG-21	Fumarole Azufalito	02/27/89	92.8	96.99	1.23	214	275
ZTG-22	Fumarole near Georgina	02/28/89	94	87.74	9.48	213	276
ZTG-23	Fumarole Negro	02/28/89	94	93.93	2.78	234	254
ZTG-31	Well ZC-11, mini-separator	12/11/89	164.5	97.43	0.59	222	299

^a D'Amore and Panichi 1980.

^b Norman and Bernhart 1981.

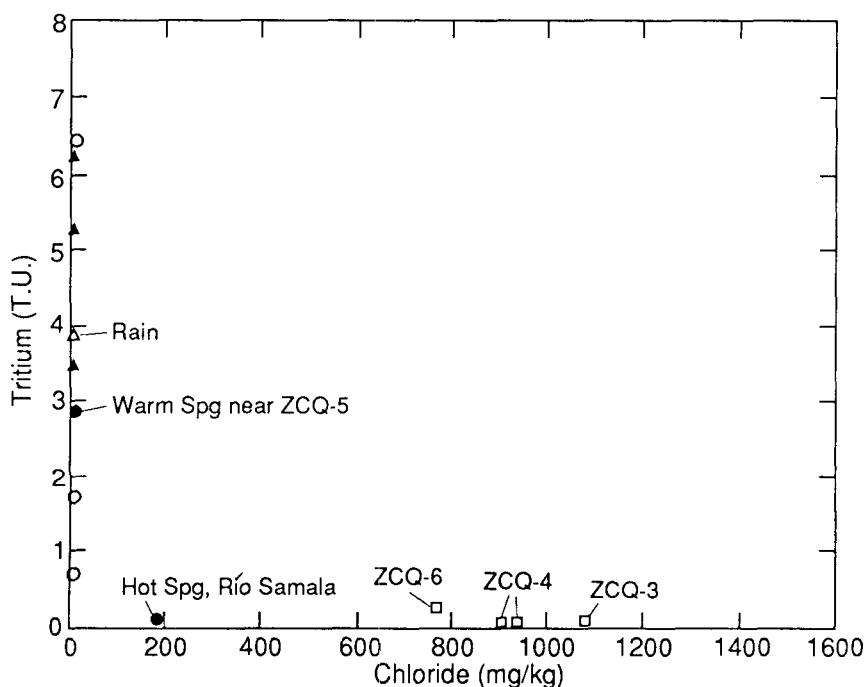


Fig. 22. Plot of tritium versus chloride for thermal and nonthermal waters of the Zunil-Quetzaltenango region, Guatemala. Δ = rainwater; \bullet = hot/warm springs; \square = well samples; \blacktriangle = acid springs; \circ = cold springs.

G. WELL ZC-11

We gave special attention to slimhole ZC-11 because it is being used to supply steam to the direct-use, vegetable/fruit drying facility adjacent to production well ZCQ-3. Well ZC-11 was drilled during the early 1976–1980 exploration program at Zunil to a depth of 700 m. The well took six months to drill and, in subsequent years, had mechanical problems with the casing and wellhead assembly. These problems were partially corrected and in 1987 the well was deemed suitable for use in the joint INDE/LANL direct-use project (Tobias 1987).

A steam separator constructed by INDE is connected to the side of a T-fitting at the wellhead (Tobias 1987). The Los Alamos heat exchanger and heat loop is connected to the steam side of the separator.* The other side of the T-fitting is connected to a "New Zealand" style muffler and a weirbox. Partial chemical analyses and some pressure/flow rate data are given in Tobias (1987). Additional chemical data can be found in Giggenbach (1988) and in Tables II and III. Our objective is to estimate the total flow (gas plus liquid) from the well and to obtain a chemical analysis of the downhole fluid composition.

In January 1990, the flow rate of ZC-11 was measured at the weirbox while flow was partially directed toward the heat extraction loop and while flow was directed only to the muffler. When the heat extraction loop was operating the flow rate at the weirbox was $28.6 \pm 0.9 \text{ Q/min}$ and when not operating, the flow rate was $34.2 \pm 1.5 \text{ Q/min}$. According to the casing schedule and temperature log, ZC-11 produced water from an open-hole section at 560 m and 255°C. Chemical geothermometry suggests a lower temperature fluid may have been entering the wellbore but we have no physical means to verify this deduction (see above). Assuming the formation temperature was 255°C, the steam fraction, y , of the total mass flow is 0.314 (consult Henley *et al.* 1984, Chap. 1, and steam tables). Therefore, the total mass flow of ZC-11 was about 49.5 Q/min when the heat extraction loop received no fluid, but when the loop was operating, only 41.1 Q/min passed through the weirbox. This means that roughly 8 Q/min (or 17%) of the fluid was directed toward the heat extraction loop when the drying plant was operating. For various reasons, it was not possible to collect samples off the steam separator during any of our sampling trips. Therefore, it was impossible for us to assess the quality of steam separation or to present chemical data on the steam that eventually entered the heat exchanger during drying operations. Chemical data (flash-corrected) in the tables and figures were based on samples collected either at the weirbox or from a portable steam separator mounted at the wellhead.

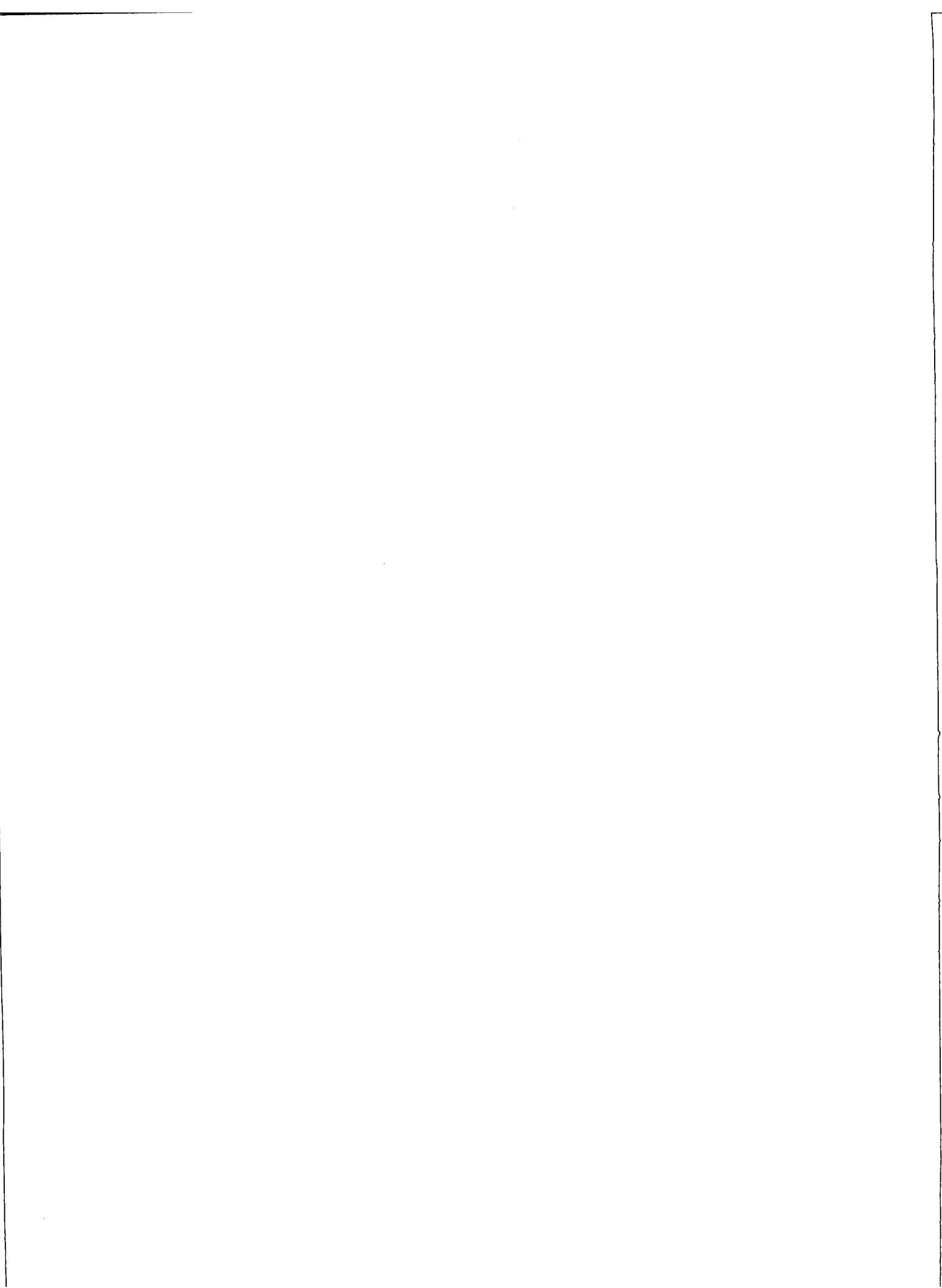
* Diagrams are available from G. Thayer, A-4, MS B299, Los Alamos National Laboratory, Los Alamos, NM 87545.

H. DISCUSSION

Although our isotopic data are far from comprehensive, they suggest that recharge to the Zunil reservoir may come from the north and east. The different chemistries displayed by each well and differences in chemistry observed in weirbox and *in situ* samples suggest that the Zunil reservoir is not homogeneous and is poorly connected geochemically. The maximum mean age of water in different horizons may be as much as 500 to 7500 years.

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APPENDIX A

DESCRIPTION OF LOGGING TOOLS

Los Alamos National Laboratory furnished a new logging truck for the Central American Energy and Resources Program, complete with auxiliary generators, a data acquisition system, hydraulic-powered draw works, and associated controls. The draw works were equipped with 3000 m (10,000 ft) of armored cable that contains seven electrical conductors with TFE Teflon® insulation. The cable armor package is improved galvanized plow steel, and the cable is rated for continuous service at temperatures up to 320°C.

Los Alamos National Laboratory also equipped the logging truck with wellhead apparatus required to "rig up" for deployment of the downhole instrument. (The downhole instrument is commonly referred to as the logging tool.) This apparatus includes the upper and lower sheaves, pressure lock with cooling jacket, Bowan control head (pack-off), and cable cooling system. All downhole instrument systems were designed, fabricated, and tested at Los Alamos National Laboratory. The set of logging tools consists of cableheads, a temperature probe with a casing collar locator (CCL) and ring gauge, a downhole fluid sampler, a three-arm caliper, and a fluid velocity (spinner)/temperature/pressure (STP) tool.

THE CABLEHEAD

The cablehead (Fig A.1) is designed specifically for long-term operation in a geothermal environment where fluid temperatures exceed 300°C. The cablehead permits termination of the seven-conductor armored cable; it completes the transition of the logging tool while ensuring watertight integrity in the high-temperature and high-pressure geothermal fluids. The cablehead can be disconnected from the armored cable should the tool stick in a wellbore.

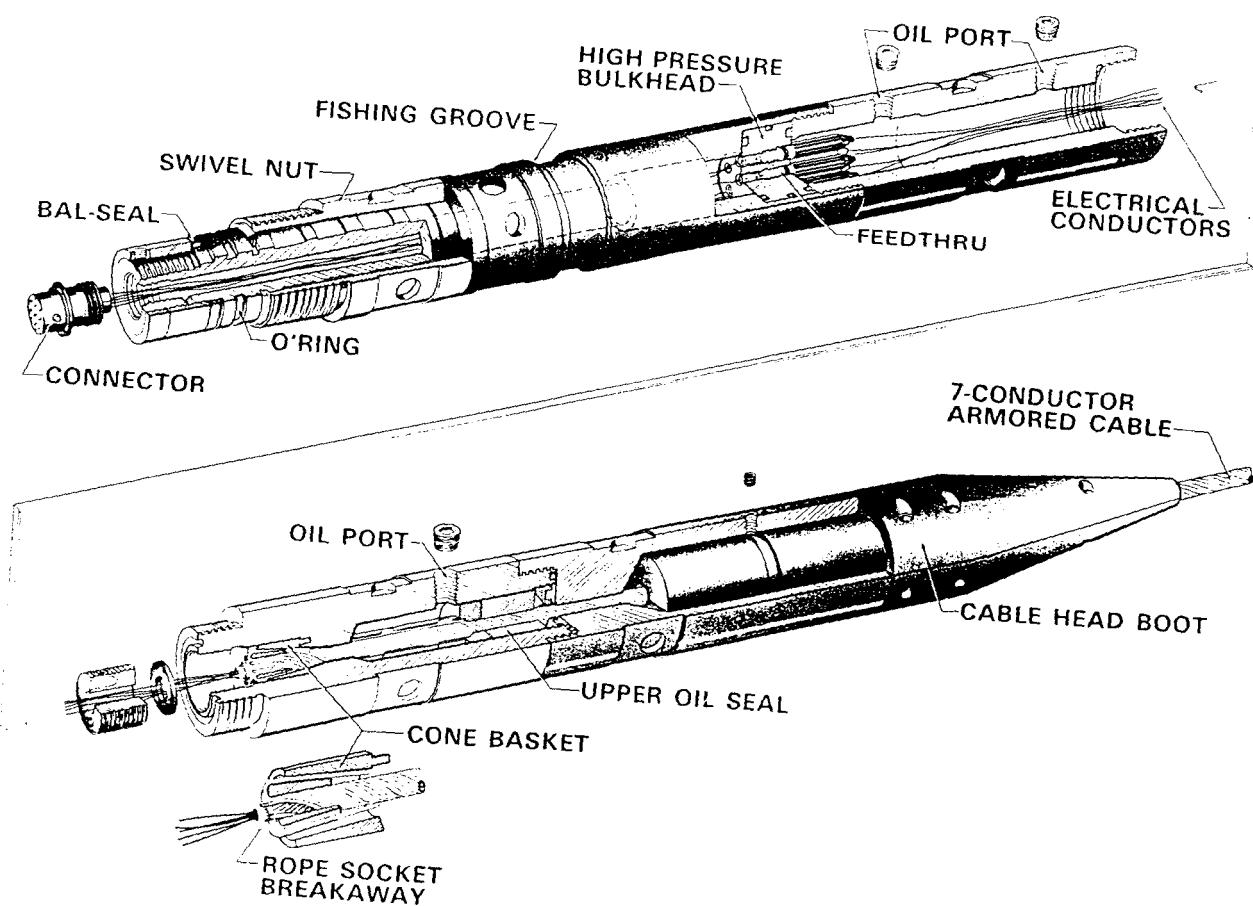


Fig. A.1. The cablehead assembly.

THE TEMPERATURE PROBE

Borehole temperature surveys determine thermal gradients along the borehole under both static and flowing conditions. Temperature anomalies in regions where fluid flows into or out of the borehole are easily detected, so that flow zones and well lining condition can be assessed. The temperature tool includes the CCL to detect casing signatures to correct tool depth, which can be exaggerated by cable stretch. The ringgauge or "rabbit" is attached to the pressure housing to gauge the borehole diameter (Fig. A.2). The temperature probe with the CCL and rabbit is run first in every wellbore before the borehole is logged with the more expensive, complex instrument packages.

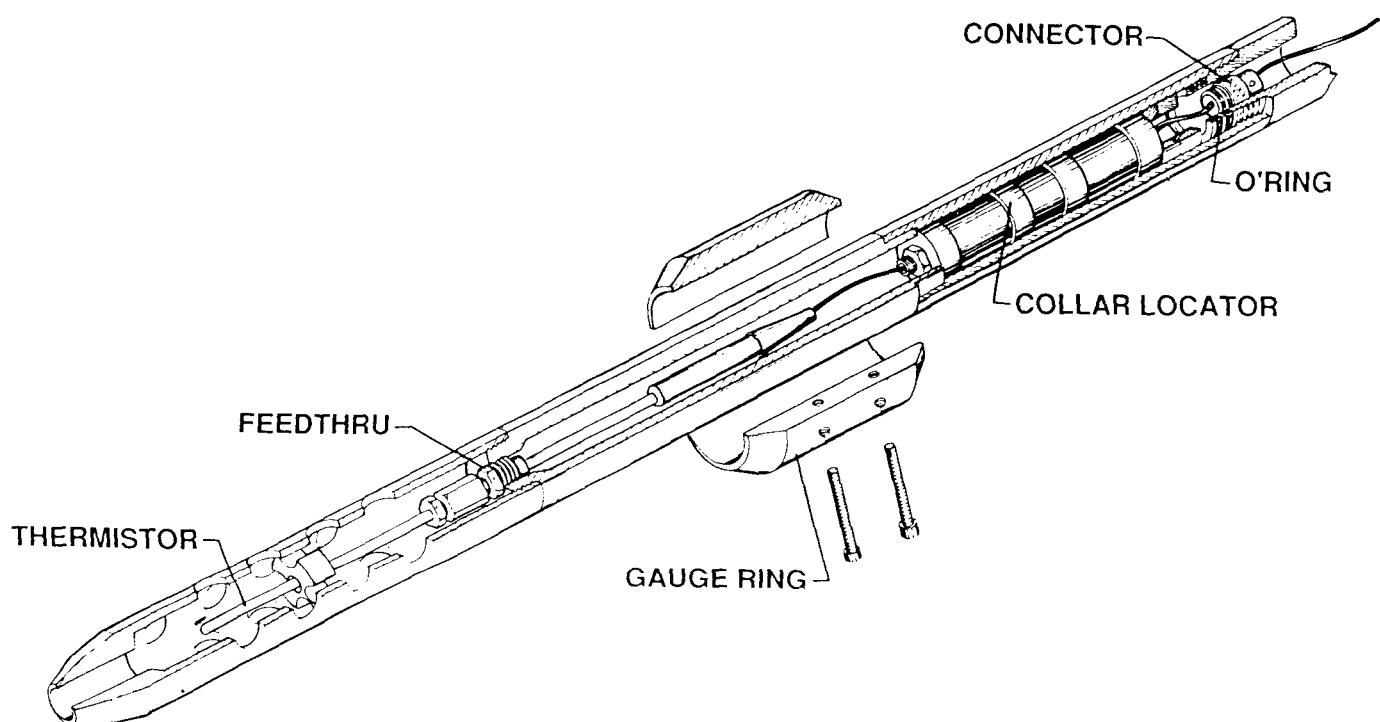


Fig. A.2. The temperature/rabbit tool.

LOS ALAMOS *IN SITU* SAMPLER

The Los Alamos *in situ* sampler is described in Archuleta *et al.* (1978) and, its field application is reported in Goff *et al.* (1987a) and Grigsby *et al.* (1989) (Fig. A.3). The sample bottle is evacuated at the surface and the tool is then lowered to the desired depth by a seven-conductor armored cable. The sample chamber valve is opened by a temperature-hardened motor that is controlled electrically from the surface. The sample chamber is then closed by reversing polarity on the circuit. When the sampler returns to the surface, it is cooled to $<70^{\circ}\text{C}$ to prevent any loss of steam before the pressurized sample is removed. A temperature well in the valve assembly cap extends into the sample bottle to provide a means of monitoring the internal temperature of the sample. A gas extraction system (GES) (Fig. A.4) is used to quantitatively remove the gaseous portion of the sample from the sample bottle. Goff *et al.* (1987a) gives a more complete description of GES procedures.

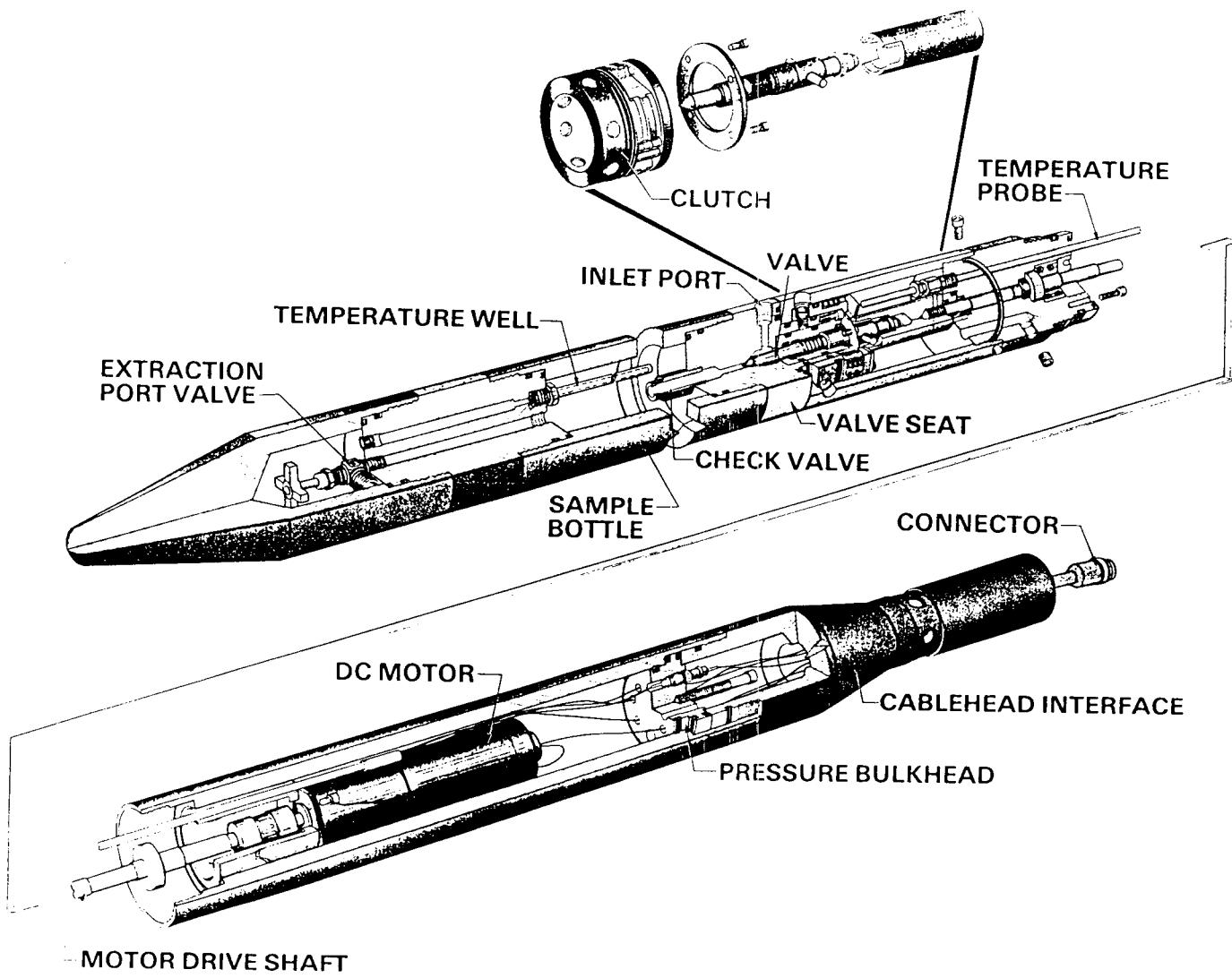


Fig. A.3. Schematic diagram of the Los Alamos fluid sampler used during this investigation.

After extraction of the gaseous portion of the sample, the tool is opened and the geothermal brine is poured into a tared beaker and the mass weighed on a top-loading balance. In this case, the brine contained less than 10,000 mg/kg chloride; thus, the brine density was not calculated in the field. The density is assumed to be 1.0 at 25°C. An aliquot of sample is used to determine pH and any other parameters that a geoscientist may wish to measure in the field. The remainder of the sampler is split and preserved for various chemical and isotopic analyses. Procedures for collecting the other Zunil samples discussed in this report vary depending upon the situation (spring, fumarole, weirbox, etc.). Trujillo *et al.* (1987) discusses methods of collection and preservation of samples.

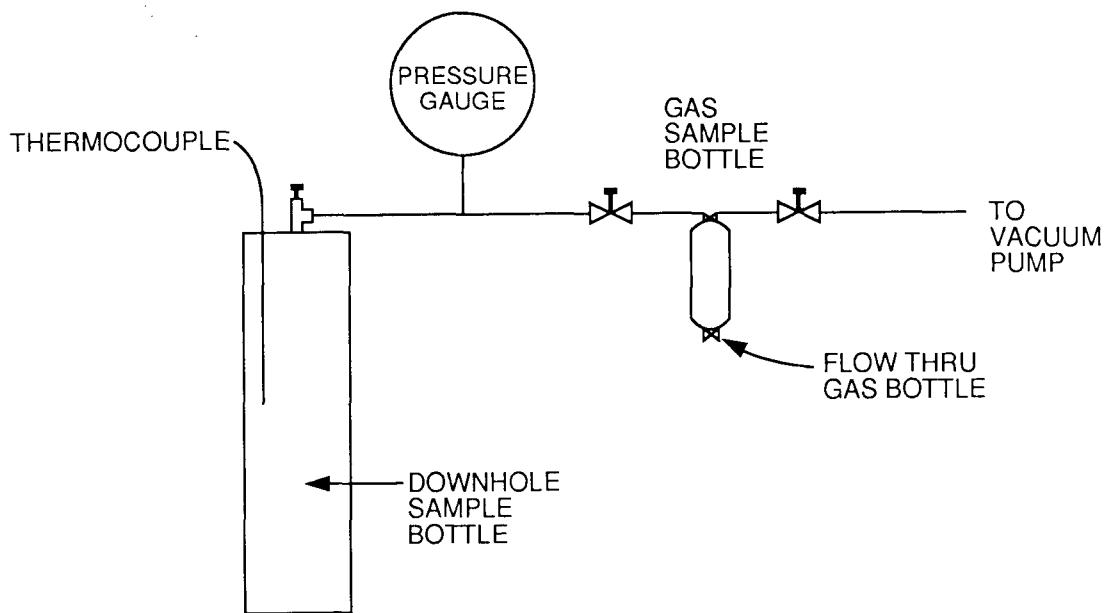


Fig. A.4. Gas extraction system used in operations at the Zunil geothermal field, Guatemala.

THE THREE-ARM CALIPER

Borehole conditions, such as open-hole or casing diameters, contour, wear, scale accumulations, breakouts, etc., can be measured with the three-arm caliper tool. The three arms, which are spaced 120° apart on the circumference, operate independently (Fig. A.5). Normally, the caliper tool is deployed in the borehole, with the arms retracted, to the lowest depth of interest. The arms are then extended by applying current to the downhole dc motor and associated drive. When fully extended, the arms provide a moderate spring-activated force against the borehole or casing wall. The tool is then pulled up the borehole, and the motion of each arm, as it follows the contour, is transformed to a rotational motion sensed by a cosine-type potentiometer. The output signal of each of the three potentiometers is a function of the radius from the center line of the centralized tool to the tip of the arm. Caliper calibration is verified before each log.

1. D.C. MOTOR, HIGH TEMP.
2. BAL-SEALS NO. 505
3. BALL BEARINGS NDH
4. LEAD SCREW (17-4 SST)
5. CENTRALIZERS (12)
6. ARM (3)
7. BEAD CHAIN
8. MAGNET, ALNICO 5-7 (12)
9. POTENTIOMETER (3)
10. O-RING
11. COVER
12. BULKHEAD FEED THROUGH

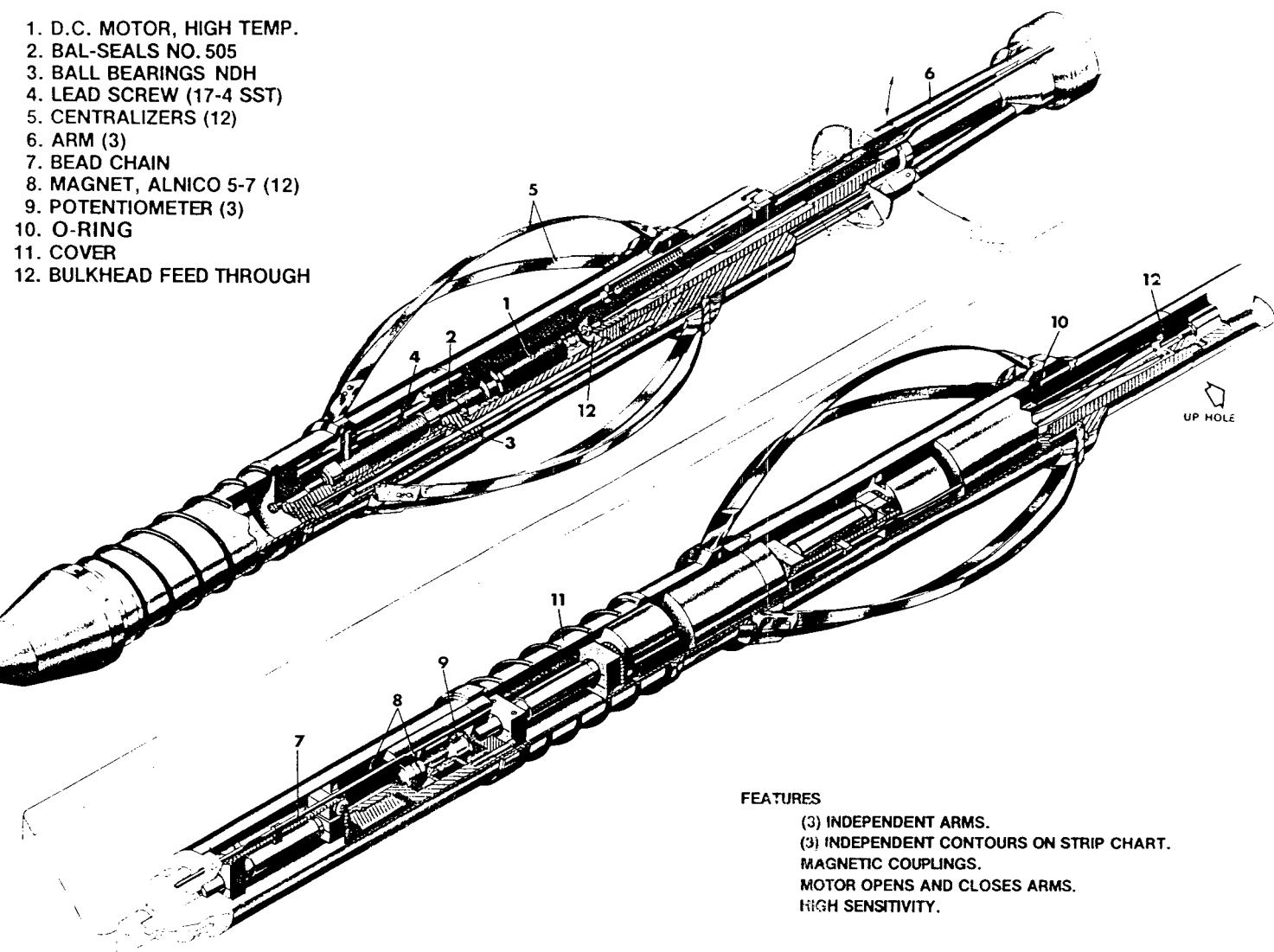


Fig. A.5. The three-arm caliper tool.

SPINNER/TEMPERATURE/PRESSURE TOOL

To determine the thermodynamic state of a flowing borehole, simultaneous measurements of temperature, pressure, and fluid velocity are necessary. A high-temperature (300°C) well logging tool was developed specifically to measure these parameters in the production wells in Central America. This tool is referred to as the spinner/temperature/pressure or STP tool. The fluid velocity transducer (spinner) incorporates a rotating impeller with hardened steel pivot bearings (the spinner end of the tool is shown in Fig A.6). The rotating shaft operates a reed switch that transforms the rotational speed of the impeller to pulses recorded as frequency in hertz. The rotational speed in hertz is proportional to the velocity of the fluid relative to the logging tool. A proportionality constant is determined by logging the liquid-filled region of the borehole. The temperature sensor is a thermistor that has been calibrated to an accuracy of 0.10°C up to 300°C. This temperature sensor provides very fast response and exceptional resolution. The pressure transducer provides accurate pressure measurement in the geothermal fluids when meticulous calibration producers are used. Pressure measurements are made in the STP tool with a 0- to 34.5-MPa (5000 psi) potentiometer gauge pressure transducer. The constant current excitation is sensed at the pressure gauge so that the power supply at the surface can compensate for line losses over the 3000-m armored cable.

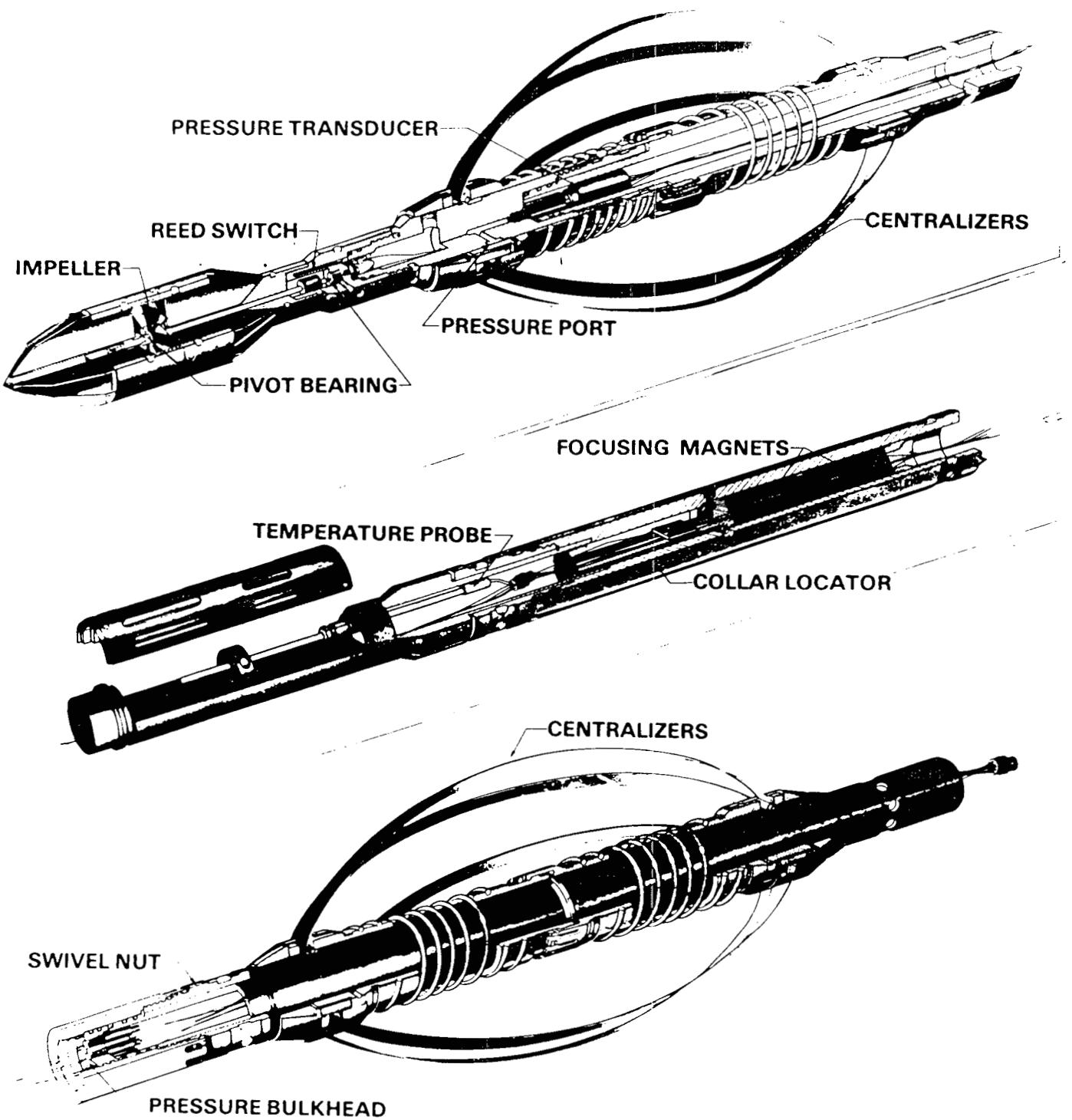


Fig. A.6. The spinner/temperature/pressure tool.

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