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# Investigating carbon stabilization in soils via mineral adsorption

What are we missing when we define  
mineral associated organic matter.

September 2023

Qian Zhao  
Odeta Qafoku,  
Kenton Rod

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# **Investigating carbon stabilization in soils via mineral adsorption**

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Qian Zhao  
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Kenton Rod

Prepared for  
the U.S. Department of Energy  
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## Abstract

Understanding carbon (C) fluxes released from soil to atmosphere is critical to regulating global climate change. Soil minerals play a crucial role in stabilizing C, based on recent studies which have found that mineral associated C can be stored in soils for decades to centuries longer than non-mineral associated C. We aim to maximize C sequestration via mineral adsorption in soils. To achieve this objective, it is important to understand the potential impacts of C stabilization via adsorption to mineral surfaces. Particularly, investigating the quantitative impacts of C sorption to minerals is novel to understanding and implementing multi-scale biogeochemical processes of C stabilization in soils. We will conduct a literature review on potential impacts of C sorption to minerals, thereby contributing to the net C storage in soils. Our findings will allow us to quantitatively understand soil C changes and durability at molecular and ecosystem scales, fulfilling existing scientific gaps in the community as well as interests of sponsors, such as DOE-BER.

## Summary

This project provides insights into the role of minerals on soil carbon (C) storage, sequestration, and contribution to the global C cycle. The outcomes of this study demonstrated how mineral adsorption processes might enhance C storage in terrestrial ecosystems. We have conducted an outline of a manuscript that focuses on the diverge quantification methods of mineral associated organic matter (MAOM) pool in soil could lead to an uncertain estimation of global mineral associated C stock in models. Meanwhile, we also build up a body of literatures as well as a dataset of MAOM-C stock and the fraction of MAOM-C to total soil organic carbon (SOC) across different ecosystems from literatures. Along with MAOM pool size data, this dataset also includes site metadata, including mean annual temperature (MAT), mean annual precipitation (MAP), soil moisture, soil pH, etc. We are continuously building this dataset by synthesize data from literatures. Information from this LDRD project will be used to develop a literature review manuscript and potential DOE BER proposals.

## Acknowledgments

This research was supported by the **EBSD Mission Seed**, under the Laboratory Directed Research and Development (LDRD) Program at Pacific Northwest National Laboratory (PNNL). PNNL is a multi-program national laboratory operated for the U.S. Department of Energy (DOE) by Battelle Memorial Institute under Contract No. DE-AC05-76RL01830.

## Acronyms and Abbreviations

soil organic matter (SOM)

mineral associated organic matter (MAOM)

carbon (C)

soil organic carbon (SOC)

mean annual temperature (MAT)

mean annual precipitation (MAP)

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## 1.0 Introduction

Understanding carbon (C) fluxes released from soil to atmosphere is critical to regulating global climate change. Soil minerals play a crucial role in stabilizing C, based on recent studies which have found that mineral associated C can be stored in soils for decades to centuries longer than non-mineral associated C (Lützow et al 2006, Mikutta et al. 2019, Schmidt et al. 2011). For example, microbial-derived C can be stabilized in soil over years through adsorption to amorphous Fe hydroxides (e.g., ferrihydrite), Al oxides, and Ca-bearing minerals (e.g., calcite), due to their high surface areas, the prevalence of surface hydroxyl groups, and bridging capacities. This project probes the impact of organo-mineral and microbe-mineral interactions on C assimilation and protection in soil to guide soil C sequestration. Thus, our findings allow us to quantitatively understand the persistent soil C that associates with soil minerals at molecular and ecosystem scales, fulfilling existing scientific gaps in the community as well as interests of sponsors, such as the U.S. Department of Energy (DOE)-Biological and Environmental Research (BER). This effort aligns well with the directorate objectives of PNNL Earth and Biological Sciences Directorate (EBSD) on Durable Soil Carbon Storage for Atmospheric Carbon Dioxide Removal.

## 2.0 Current understanding of C budget in the MAOM pools

Second State of the Carbon Cycle Report (SOCCR2) has identified the soil C stock in top 1m soil from North American is 400 PgC (Lajtha et al. 2018). Estimates of global soil organic carbon (SOC) stocks vary from 684 to 724 Pg in the surface 0.3 m (Batjes 1996), the Intergovernmental Panel of Climate Change (IPCC) standard sampling depth (Aalde et al. 2006), and (Sokol et al., 2019) 1462 to 1548 Pg in the top meter and  $2,060 \pm 220$  Pg C to 2 m (Batjes 1996). Diverse C inputs result in heterogeneity of the SOM pool. Studies showed both microbial-derived C and plant-derived C are major sources of mineral-associated organic matter (MAOM). The formation and accumulation of SOC are primarily derived from plant- and microbial C (Angst et al., 2021; Ma et al., 2018; Sokol & Bradford, 2019). While the significance of plant residues to SOM persistence is well recognized due to the slow cycling of plant debris (e.g., lignin and phenols) (Angst et al., 2021; Schmidt et al., 2011), the importance of microbial products and residues (hereafter ‘microbial necromass’) in the slowly cycling SOC pool has only recently been appreciated (Fan et al., 2021; Kallenbach et al., 2016; Liang et al., 2019; Wang et al., 2021; Wu et al., 2023). Microbial necromass contributes to 33-62% of total SOC in various types of ecosystems (measured as amino sugars) (Liang et al., 2019; Wang et al., 2021). However, the long-term persistence of microbial necromass in the SOM pool depends on the associations with soil minerals (Kästner et al. 2021), even though the specific necromass-mineral interactive mechanisms are still unclear.

The estimations of MAOM pool in soils have been largely studied in the last decade. Studies showed the proportion of MAOM to total SOM varies significantly, ranging from 25%-90% (Table 1). The proportion of the MAOM pool varies across different ecosystem types with 60% in forest, 72% in grassland, and 79% in cropland. Lugato et al (2021) did size fractionation on 400 soils and quantified MAOM with a machine learning (ML) approach. The model was then used to predict the C and N in the MAOM fraction for the 9,229. The average mineral associated C stocks in European grasslands and forest soils varied between 15 and 38 g C kg<sup>-1</sup> soil (Figure 1).

**Table 1. Percentage of Fe-OC to total OC in natural soil and sediment environments.**

Site or ecosystem type	Proportion of MAOM to total SOM	Methods	Reference
14 Forest soils in the USA	0.6-57.8% (average 37.8%)	Dithionite extraction	<a href="#">Zhao et al. 2016</a>
4 wetland soils in China	7-91%	Size and density	<a href="#">Liu et al. 2023</a>
Mangrove soils in the Philippines	~15 %	Dithionite extraction	<a href="#">Dicen et al., 2019</a>
Arable soils in China	6.2–31.2 % Agriculture soils – grassland and arid	Dithionite extraction	<a href="#">Wan et al., 2019</a>
10 Peatlands in China	1.64–5.94 %	Dithionite extraction	<a href="#">Huang et al., 2021a</a>
186 grassland and forest soils in European	25-85% (60% in forest, 72% in grassland, 79% in cropland)	Size	<a href="#">Cotrufo et al., 2019</a>
1451 all ecosystem types	65%	Size	<a href="#">Georgiou et al. 2022</a>

**Grasslands**

Tibetan alpine grasslands in China	15.8 ± 12.0 %	Dithionite extraction	<a href="#">Fang et al., 2019</a>
Meadow soils in the Qinghai-Tibetan Plateau	4.1–25.6 %	Dithionite extraction	<a href="#">Mu et al., 2020</a>

**Sediments**

Marine sediments in Mexican and Indian margins, the Southern Ocean, the St. Lawrence estuary and gulf, and the Black Sea	21.5 ± 8.6 %	Dithionite extraction	<a href="#">Lalonde et al., 2012</a>
East China Sea sediments	2.77–31.5 %	Dithionite extraction	<a href="#">Ma et al., 2018</a>
Changjiang estuary sediments in China	7.4 ± 3.5 %	Dithionite extraction	<a href="#">Zhao et al., 2018</a>
Wax Lake Delta sediments in the USA	~15.0 %	Dithionite extraction	<a href="#">Shields et al., 2016</a>
Sediments in Eurasian Arctic Shelf	0.5-22%	Dithionite extraction	<a href="#">Salvadó et al., 2015</a>
Saanich, Arabian Sea, Mexican margin, and St. Lawrence estuary	25.7–62.6 %	Dithionite extraction	<a href="#">Barber et al., 2017</a>

**Permafrost**

Permafrost soils in northern Alaska	13.68±2.31 %	Dithionite extraction	<a href="#">Joss et al., 2022</a>
Permafrost soils in the Qinghai-Tibetan Plateau	19.5 ± 12.3 %	Dithionite extraction	<a href="#">Mu et al., 2016</a>
Discontinuous permafrost region in Sweden	9.9–14.8%	Dithionite extraction	<a href="#">Patzner et al., 2020</a>

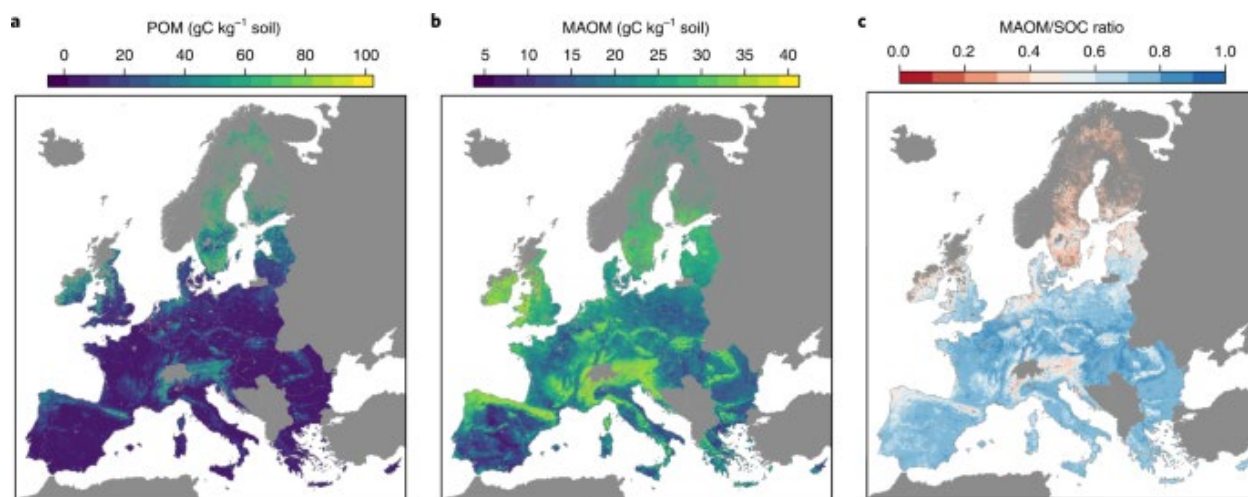


Figure 1. **a–c**, Carbon content (gC kg<sup>-1</sup> soil) in POM (**a**) and MAOM (**b**), and ratio between MAOM carbon and total SOC (**c**) from Lugato et al (2021).

### 3.0 Different parameters impact on MAOM pool size

#### 3.1 Quantification methods

There are diverse methodologies for quantifying the MAOM pool in soil. Abramoff et al (2021) compiled 402 laboratory sorption experiments and quantified, for the first time, the sorption capacity of mineral soils to DOC for six soil orders. They find that mid- and low-latitude soils and subsoils have a greater capacity to store DOC by sorption compared to high-latitude soils and topsoils. The global additional DOC sorption is estimated to be  $107 \pm 13$  Pg C to 1 m depth, projecting a 7% increase in the existing total carbon stock.

Another two major MAOM quantification methods are chemical extraction and size fractionation. Chemical extraction includes sodium pyrophosphate, sodium dithionite, oxalate acid, and HCl. These solvents target to dissolve specific minerals from soils, thereby releasing OC that associates with these minerals. Among these solvents, sodium dithionite has been widely used to quantify mineral bound OC in addition to quantify extracted Fe, Al, and other metals (Lalonde et al 2012, Zhao et al. 2016). Meanwhile, size fractionation separates the MAOM pool by less than  $53\mu\text{m}$  particle size. This approach is easy and high throughput so that can be applied to a large number of sample sites. Dithionite extraction could underestimate the MAOM pool as it only extracts a portion of minerals from soils but not all, whereas size fractionation most likely overestimates the pool as there could be non-mineral bound C within particles less than  $53\mu\text{m}$ . These different quantification approaches result in large variations in estimating the MAOM pool size in soils. For instance, dithionite extraction estimates about 25% of total SOC as MAOM C, whereas size fractionation estimates an average of 65% of total SOC as MAOM C pool. Such different estimations on MAOM pool result in diverge projections in continental scale MAOM pool. Given the total C stock, we estimated 60-260PgC in the MAOM pool in top 1m soil in North America. This diverge estimations on the MAOM pool size will impact on the accuracy of C cycling predictions by biogeochemical process-based models.

#### 3.2 Other factors

MAOC to reactive Fe ratios (OC:Fe) can indicate the mechanism of Fe-OC interactions. A mass ratio over 0.22 is indicative for Fe-OC associations predominantly formed by co-precipitation or chelation (Wagai and Mayer, 2007). Below an OC:Fe mass ratio of 0.22, OC is mainly assumed to be sorbed onto Fe minerals. Around 86% of the permafrost samples exceed an OC:Fe mass ratios of 0.22 (Joss et al. 2022). Other soil parameters, such as the bulk density, pH, and moisture content, were correlated with MAOM-C (Mu et al. 2022). Soil with higher pH results in higher fraction of MAOM-C, whereas higher soil moisture contents result in lower MAOM-C due to high microbial activities and consumption of available C.

## 4.0 Implications and discussion on C cycling modeling

Ungeneralizable MAOM fractionation approach results in a large uncertainty to the estimation of MAOM pool, thereby impacting the estimation of C persistence in soil. Consensus on the experimental method of defining MAOM pool is necessary to the soil community. The lab-scale quantification of C storage via minerals is applicable to ecosystem scale and global scale. Soil mineralogy is a major factor to predict ecosystem-C behavior in terms of microbe-mineral interactions and nutrient-mineral interactions. The quantitative data from this project can be incorporated into process-based models, such as Earth and Environmental Systems (ESS), and/or artificial intelligence and machine learning (AI/ML) based models to improve model prediction on C storage by considering organo-mineral protection.

## 5.0 Conclusion

This project provides insights into the role of minerals on soil carbon (C) storage, sequestration, and contribution to the global C cycle. The outcomes of this study demonstrated how mineral adsorption processes might enhance C storage in terrestrial ecosystems. We have conducted an outline of a manuscript that focuses on the diverge quantification methods of MAOM pool in soil could lead to an uncertain estimation of global mineral associated C stock in models. Meanwhile, we also build up a body of literatures (Figure 2) as well as a dataset of MAOM-C stock and the fraction of MAOM-C to total SOC across different ecosystems from literatures (Appendix A). Along with MAOM pool size data, this dataset also includes site metadata, including MAT, MAP, soil moisture, soil pH, etc. We are continuously building this dataset by synthesize data from literatures. Information from this LDRD project will be used to develop a literature review manuscript and potential DOE BER proposals.

The screenshot displays the EndNote 20 interface. The main window shows a list of references under the title 'Mineral-C Lit Rev'. The references are sorted by author and year. The first reference is 'Zhao, Qian; Poulton, T. A. et al. 2016. Iron-bound organic carbon in forest soils: quantification and characterization'. The preview window on the right shows the first page of this paper, which includes a map of the United States and a bar chart of Fe-OC concentrations in forest soils.

Author	Year	Title
Zhao, Qian; Poulton, T. A. et al.	2016	Iron-bound organic carbon in forest soils: quantification and characterization
Zafar, Rabia; Waheed, A. et al.	2017	Organic compound-mineral interactions: Using flash pyrolysis to monitor
Yan, Zhifeng; Bo, J. et al.	2018	A moisture function of soil heterotrophic respiration that incorporates
Whalen, Emily D.; et al.	2022	Clarifying the evidence for microbial- and plant-derived soil organic matter
Wagai, Rota; Kaji, Y. et al.	2020	Iron and aluminum association with microbially processed organic matter
Van Der Kellen, I.; et al.	2022	Recent weathering promotes C storage inside large phyllosilicate particles
Tang, Ni; Siebers, R. et al.	2022	Implications of Free and Occluded Fine Colloids for Organic Matter Partitioning
Strawn, Daniel G.	2021	Sorption mechanisms of chemicals in soils
Sokol, Noah W.; et al.	2022	Global distribution, formation and fate of mineral-associated soil organic carbon
Sokol, Noah W.; et al.	2022	Global distribution, formation and fate of mineral-associated soil organic carbon
Sokol, Noah W.; et al.	2019	Pathways of mineral-associated soil organic matter formation: Integrating

Figure 2. The list of literatures of quantifying mineral-associated OC in soil.

## 6.0 References

- Aalde, H., P. Gonzalez, M. Gytarsky, T. Krug, W. Kurz, R. D. Lasco, D. Martino et al. "Generic Methodologies Applicable to Multiple Land-use Categories. IPCC Guidelines for National Greenhouse Gas Inventories." *Institute for Global Environmental Strategies (IGES), Hayama, Japan* (2006).
- Abramoff, Rose Z., Katerina Georgiou, Bertrand Guenet, Margaret S. Torn, Yuanyuan Huang, Haicheng Zhang, Wenting Feng et al. "How much carbon can be added to soil by sorption?." *Biogeochemistry* 152, no. 2-3 (2021): 127-142.
- Angst, G., Mueller, K. E., Nierop, K. G., & Simpson, M. J. (2021). Plant-or microbial-derived? A review on the molecular composition of stabilized soil organic matter. *Soil Biology and Biochemistry*, 108189.
- Batjes, Niels H. "Total carbon and nitrogen in the soils of the world." *European journal of soil science* 47, no. 2 (1996): 151-163.
- Fan, X., Gao, D., Zhao, C., Wang, C., Qu, Y., Zhang, J., & Bai, E. (2021). Improved model simulation of soil carbon cycling by representing the microbially derived organic carbon pool. *The ISME Journal*, 15(8), 2248-2263.
- Joss, Hanna, Monique S. Patzner, Markus Maisch, Carsten W. Mueller, Andreas Kappler, and Casey Bryce. "Cryoturbation impacts iron-organic carbon associations along a permafrost soil chronosequence in northern Alaska." *Geoderma* 413 (2022): 115738.
- Kallenbach, C. M., Frey, S. D., & Grandy, A. S. (2016). Direct evidence for microbial-derived soil organic matter formation and its ecophysiological controls. *Nature Communications*, 7(1), 13630.
- Kästner, M., A. Miltner, S. Thiele-Bruhn, and C. Liang. "Microbial necromass in soils—linking microbes to soil processes and carbon turnover." *Frontiers in Environmental Science* 9 (2021): 597.
- Lajtha, Kate, Vanessa L. Bailey, and Karis McFarlane. The Second State of the Carbon Cycle Report-Chapter 12. Soils. No. LLNL-TR-757064. Lawrence Livermore National Lab.(LLNL), Livermore, CA (United States), 2018.
- Lalonde, Karine, Alfonso Mucci, Alexandre Ouellet, and Yves Gélina. "Preservation of organic matter in sediments promoted by iron." *Nature* 483, no. 7388 (2012): 198-200.
- Liang, C., Amelung, W., Lehmann, J., & Kästner, M. (2019). Quantitative assessment of microbial necromass contribution to soil organic matter. *Global Change Biology*, 25(11), 3578-3590.
- Lugato, Emanuele, Jocelyn M. Lavalley, Michelle L. Haddix, Panos Panagos, and M. Francesca Cotrufo. "Different climate sensitivity of particulate and mineral-associated soil organic matter." *Nature Geoscience* 14, no. 5 (2021): 295-300.
- Lützow, M. V., Kögel-Knabner, I., Ekschmitt, K., Matzner, E., Guggenberger, G., Marschner, B., & Flessa, H. (2006). Stabilization of organic matter in temperate soils: mechanisms and their



relevance under different soil conditions—a review. *European journal of soil science*, 57(4), 426-445.

Ma, T., Zhu, S., Wang, Z., Chen, D., Dai, G., Feng, B., Su, X., Hu, H., Li, K., & Han, W. (2018). Divergent accumulation of microbial necromass and plant lignin components in grassland soils. *Nature Communications*, 9(1), 3480.

Mikutta, R., Turner, S., Schippers, A., Gentsch, N., Meyer-Stüve, S., Condon, L. M., ... & Guggenberger, G. (2019). Microbial and abiotic controls on mineral-associated organic matter in soil profiles along an ecosystem gradient. *Scientific reports*, 9(1), 1-9.

Mu, Cuicui, Feng Zhang, Mei Mu, Xu Chen, Zhilong Li, and Tingjun Zhang. "Organic carbon stabilized by iron during slump deformation on the Qinghai-Tibetan Plateau." *Catena* 187 (2020): 104282.

Schmidt, M. W. I., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., Janssens, I. A., Kleber, M., Kögel-Knabner, I., Lehmann, J., Manning, D. A. C., Nannipieri, P., Rasse, D. P., Weiner, S., & Trumbore, S. E. (2011). Persistence of soil organic matter as an ecosystem property. *Nature*, 478, 49-56. <https://doi.org/doi:10.1038/nature10386>

Schmidt, M. W., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., Janssens, I. A., ... & Trumbore, S. E. (2011). Persistence of soil organic matter as an ecosystem property. *Nature*, 478(7367), 49-56.

Sokol, N. W., & Bradford, M. A. (2019). Microbial formation of stable soil carbon is more efficient from belowground than aboveground input. *Nature Geoscience*, 12(1), 46-53.

Sokol, N. W., Sanderman, J., & Bradford, M. A. (2019). Pathways of mineral-associated soil organic matter formation: Integrating the role of plant carbon source, chemistry, and point of entry. *Global Change Biology*, 25(1), 12-24.

Wagai, Rota, and Lawrence M. Mayer. "Sorptive stabilization of organic matter in soils by hydrous iron oxides." *Geochimica et Cosmochimica Acta* 71, no. 1 (2007): 25-35.

Wang, B., An, S., Liang, C., Liu, Y., & Kuzyakov, Y. (2021). Microbial necromass as the source of soil organic carbon in global ecosystems. *Soil Biology and Biochemistry*, 162, 108422.

Wu, H., Wan, S., Ruan, C., Wan, W., Han, M., Chen, G., Liu, Y., Zhu, K., Liang, C., & Wang, G. (2023). Soil microbial necromass: The state-of-the-art, knowledge gaps, and future perspectives. *European Journal of Soil Biology*, 115, 103472.

Zhao, Qian, Simon R. Poulson, Daniel Obrist, Samira Sumaila, James J. Dynes, Joyce M. McBeth, and Yu Yang. "Iron-bound organic carbon in forest soils: quantification and characterization." *Biogeosciences* 13, no. 16 (2016): 4777-4788.



# Appendix A – A working dataset of MAOM-C

Author	Year	Lat	Lon	MAT	MAP	Soil.orde	Mineral	Top.dep	Bottom.	Vegetat	frc_mel	Clay	Silt	Sand	SiltClay	P Bulk.C	SiltClayC	POM.C	source	MAOM.	f MAOM
Amelung	1998	52.92	-105.8	0.9	456	Mollisols	HM	0	10	Grassland	Particle_Size	29	22	49	51	58.2	39.88	13.1168	Synthesis	45.0832	0.7746254
Amelung	1998	52.19	-106.17	1.6	343	Mollisols	HM	0	10	Grassland	Particle_Size	16	19	65	35	42.3	27.63	11.0297	Synthesis	31.2703	0.7392506
Amelung	1998	50.17	-107.5	3.2	380	Mollisols	HM	0	10	Grassland	Particle_Size	22	27	51	49	42.3	27.01	7.5469	Synthesis	34.7531	0.8215863
Amelung	1998	45.35	-95.55	6.1	565	Mollisols	HM	0	10	Grassland	Particle_Size	34	31	35	65	64	49.23	10.9116	Synthesis	53.0884	0.8295063
Amelung	1998	46.5	-100.54	5	419	Mollisols	HM	0	10	Grassland	Particle_Size	21	20	59	41	31.6	22.84	7.8396	Synthesis	23.7604	0.7531114
Amelung	1998	48.33	-109.41	6.1	300	Mollisols	HM	0	10	Grassland	Particle_Size	17	17	66	34	16.4	9.92	4.7943	Synthesis	11.0557	0.7076646
Amelung	1998	44.5	-105.51	7.2	400	Andisols	HM	0	10	Grassland	Particle_Size	24	34	42	58	18.6	12.57	3.3882	Synthesis	15.2118	0.8178387
Amelung	1998	41.22	-104.63	8.9	400	Mollisols	HM	0	10	Grassland	Particle_Size	14	12	74	26	8.1	6.99	1.7353	Synthesis	6.3647	0.7857654
Amelung	1998	41.43	-105.7	9	400	Mollisols	HM	0	10	Grassland	Particle_Size	12	8	80	20	12.3	7.41	3.7291	Synthesis	8.5709	0.6968211
Amelung	1998	40.1	-103.13	10.8	375	Mollisols	HM	0	10	Grassland	Particle_Size	18	15	67	33	16.9	9.92	5.2655	Synthesis	11.1435	0.688432
Amelung	1998	40.26	-99.22	11.6	660	Mollisols	HM	0	10	Grassland	Particle_Size	24	23	53	47	29.9	23.12	9.3886	Synthesis	20.5114	0.686
Amelung	1998	38.53	-99.2	12.2	573	Mollisols	HM	0	10	Grassland	Particle_Size	28	17	55	45	27.5	18.69	6.965	Synthesis	20.535	0.7467273
Amelung	1998	40.48	-96.42	10.9	792	Mollisols	HM	0	10	Grassland	Particle_Size	32	28	40	60	40.5	29.39	7.785	Synthesis	32.715	0.8077778
Amelung	1998	39.11	-96.35	12.4	791	Mollisols	HM	0	10	Grassland	Particle_Size	33	27	40	60	30.7	24.24	8.6536	Synthesis	22.0464	0.7181238
Amelung	1998	37.2	-95.16	14.2	1000	Aflisols	HM	0	10	Grassland	Particle_Size	18	36	46	54	26.8	19.84	6.1404	Synthesis	20.6596	0.7708806
Amelung	1998	27.57	-98.54	23.4	440	Andisols	HM	0	10	Grassland	Particle_Size	28	22	50	50	22.5	22.22	3.1572	Synthesis	8.1428	0.7206018
Amelung	1998	33.06	-97.21	19.4	865	Vertisols	HM	0	10	Grassland	Particle_Size	45	32	23	77	59.7	41.76	7.1408	Synthesis	52.5592	0.8803886
Amelung	1998	29.42	-96.33	20	1030	Mollisols	HM	0	10	Grassland	Particle_Size	25	14	61	39	23.7	15.12	4.9984	Synthesis	18.7016	0.789097
Amelung	1998	30.05	-94.06	20.3	1308	Mollisols	HM	0	10	Grassland	Particle_Size	26	20	54	46	24.7	16.06	6.313	Synthesis	18.387	0.744413
Amelung	1998	27.45	-98.04	22.2	700	Mollisols	HM	0	10	Grassland	Particle_Size	26	7	67	33	16	12.56	6.6429	Synthesis	9.5751	0.5848188
Anderson	1998	50	-103.5	3	450	Mollisols	HM	0	10	Grassland	Particle_Size	37	34	29	71	17.1	16.04	2.6826	Synthesis	19.8174	0.8807733
Anderson	1981	50	-103.5	3	450	Mollisols	HM	0	10	Grassland	Particle_Size	37	34	29	71	17.1	16.04	0.64	Synthesis	16.46	0.9625731
Angers	1993	NA	NA	NA	NA	NA	NA	0	NA	Cropland	Particle_Size	23	35	38	62	33.2	30.7	1.26	Synthesis	31.94	0.9620483
Angers	1993	NA	NA	NA	NA	NA	NA	0	NA	Cropland	Particle_Size	23	50	27	73	27	17.07	9.8275862	Synthesis	17.172414	0.6360155
Angers	1993	NA	NA	NA	NA	NA	NA	0	NA	Cropland	Particle_Size	23	50	27	73	27.5	17.12	9.9099099	Synthesis	17.90909	0.6396396
Angers	1991	NA	NA	NA	NA	Spodosols	HM	0	15	Cropland	Particle_Size	23	50	27	73	25.8	13.98	10.423729	Synthesis	15.376271	0.5959795
Angers	1991	NA	NA	NA	NA	Spodosols	HM	0	15	Cropland	Particle_Size	13	55	31	69	35.2	29.87	6.545	Synthesis	53.227	0.7584871
Angers	1991	NA	NA	NA	NA	Spodosols	HM	0	15	Cropland	Particle_Size	12	55	33	67	39.5	29.68	8.308	Synthesis	26.892	0.7639773
Balabane	2004	49	2	10	640	Aflisols	HM	0	15	Fallow	Particle_Size	16	23	61	39	4.7	3.91	7.689	Synthesis	31.811	0.8053418
Balabane	2004	49	2	10	640	Aflisols	HM	0	15	Fallow	Particle_Size	17	22	61	39	11	8.73	NA	Synthesis	NA	NA
Balabane	2004	49	2	10	640	Aflisols	HM	0	15	Fallow	Particle_Size	16	58	26	74	4.7	4.22	NA	Synthesis	NA	NA
Balabane	2004	49	2	10	640	Aflisols	HM	0	15	Fallow	Particle_Size	17	50	33	67	11	9.26	NA	Synthesis	NA	NA
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	30	Temperate Fc	Particle_Size	21	40	39	61	52.6	47.4	1.90248	Synthesis	50.69752	0.9638312
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	26	Cropland	Particle_Size	19	43	38	62	30.9	25.8	1.6841	Synthesis	29.2159	0.9454984
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	26	Cropland	Particle_Size	18	40	42	58	17.8	14.08	1.22536	Synthesis	16.57464	0.9311596
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	30	Temperate Fc	Particle_Size	21	40	39	61	52.6	47.4	1.1648	Synthesis	51.7794	0.9877741
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	26	Cropland	Particle_Size	19	43	38	62	30.9	25.8	1.22122	Synthesis	29.87878	0.9604783
Balesdent	1998	43	0.5	13	1200	Inceptisols	HM	0	26	Cropland	Particle_Size	18	70	12	88	17.8	14.38	0.92136	Synthesis	16.87864	0.9482382
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Grassland	Particle_Size	19	54	21	79	58.8	30.7	0.6	Synthesis	58.2	0.9897595
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Grassland	Particle_Size	17	64	29	71	31.9	79.6	21.4	Synthesis	10.5	0.3291536
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Cropland	Particle_Size	24	39	37	63	16.9	8.9	1.2	Synthesis	15.7	0.9289941
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Grassland	Particle_Size	19	73	8	92	58.8	30.9	0.4	Synthesis	58.4	0.9931973
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Grassland	Particle_Size	17	78	5	95	31.9	98.6	2.4	Synthesis	29.5	0.9247649
Balesdent	1998	NA	NA	NA	NA	NA	NA	0	NA	Cropland	Particle_Size	24	64	12	88	16.9	9.3	0.8	Synthesis	16.1	0.9526627
Barthes	2008	-4	13.3	25	1100	Oxisols	LM	0	10	Savanna	Particle_Size	NA	NA	14	86	35	23.7	6.2	Synthesis	28.8	0.8228571
Barthes	2008	-4	13.3	25	1100	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	22	78	21.8	16.7	3.6	Synthesis	18.2	0.8348624
Barthes	2008	-4	13.3	25	1100	Oxisols	LM	0	10	Savanna	Particle_Size	NA	NA	24	76	36.4	27.8	7.7	Synthesis	28.7	0.7884615
Barthes	2008	-4.1	13.3	25	1400	Oxisols	LM	0	10	Savanna	Particle_Size	NA	NA	37	63	42.5	22.4	17	Synthesis	25.5	0.6
Barthes	2008	-4.1	13.3	25	1400	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	23	77	18.8	14.4	4	Synthesis	14.8	0.787234
Barthes	2008	6.24	2.2	27	1200	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	74	26	5.1	2.9	1.6	Synthesis	3.5	0.6862745
Barthes	2008	6.24	2.2	27	1200	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	86	14	6.5	3.3	2.3	Synthesis	4.2	0.6461538
Barthes	2008	6.24	2.2	27	1200	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	83	17	11.4	5.3	3.6	Synthesis	7.8	0.6842105
Barthes	2008	6.24	2.2	27	1200	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	81	19	8.5	4.1	2.5	Synthesis	6	0.7058824
Barthes	2008	-16	-49.3	23	1500	Oxisols	LM	0	10	Savanna	Particle_Size	NA	NA	53	47	22.6	15.9	5.3	Synthesis	17.3	0.7654867
Barthes	2008	-16	-49.3	23	1500	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	43	57	21.4	17.1	2.7	Synthesis	18.7	0.8738318
Barthes	2008	-16	-49.3	23	1500	Oxisols	LM	0	10	Pasture	Particle_Size	NA	NA	52	48	22	16.6	3.2	Synthesis	18.4	0.8545455
Barthes	2008	-21.27	-48.03	23	1150	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	40	60	20.7	12.6	6.4	Synthesis	11.3838	0.9289213
Barthes	2008	-21.36	-48.22	23	1500	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	81	19	7.4	5.2	2.2	Synthesis	5.2	0.7027027
Barthes	2008	-21.12	-47.35	23	1500	Ultisols	LM	0	10	Cropland	Particle_Size	NA	NA	92	8	7	4.1	2.9	Synthesis	4.1	0.5857143
Barthes	2008	-23.23	-51.11	21	1600	Oxisols	LM	0	10	Tropical Fore	Particle_Size	NA	NA	23	77	30.8	17.3	7.8	Synthesis	23	0.7467532
Barthes	2008	-23.23	-51.11	21	1600	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	13	87	17.8	12.9	2.7	Synthesis	15.1	0.8483146
Barthes	2008	-23.23	-51.11	21	1600	Oxisols	LM	0	10	Cropland	Particle_Size	NA	NA	17	83	23	15.7	3.4	Synthesis	19.6	0.8512739
Bates	1960	7.4	4	26.7	1230	Ultisols	LM	0	10	Tropical Fore	Particle_Size	1	9	90	10	39.5	7.06	0.244	Synthesis	15.1	0.3822785
Bernard	2001	49	4	10	620	Inceptisols	HM	0	10	Cropland	Particle_Size	38	3	59	41	15.2	8.58	3.79803	Synthesis	11.40197	0.7501296
Bernard	2001	49	4	10	620	Inceptisols	HM	0	10	Cropland	Particle_Size	30	20	50	50	25.9	12.16	12.6228	Synthesis	13.2772	0.5126332
Bernard	2001	49	4	10	620	Inceptisols	HM	0	10	Cropland	Particle_Size	29	13	58	42	52.2	18.57	27.8236	Synthesis	24.3764	0.4669808
Bernard	2001	49	4	10	620	Inceptisols	HM	0	10	Cropland	Particle_Size	34	20	46	54	28.5	12.3	14.7706	Synthesis	16.2562	0.8313868
Bernard	2001	49	4	10	620	Inceptisols	HM	0	10	Cropland	Particle_Size	38	13								

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