

# Experimental temperature measurements of Fe-bearing silicate minerals and glasses to 1.6 TPa

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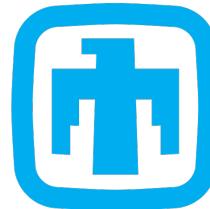
# Collaborators



**UCDAVIS**

 Lawrence Livermore National Laboratory

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ROCHESTER



Sandia  
National  
Laboratories



# Funding



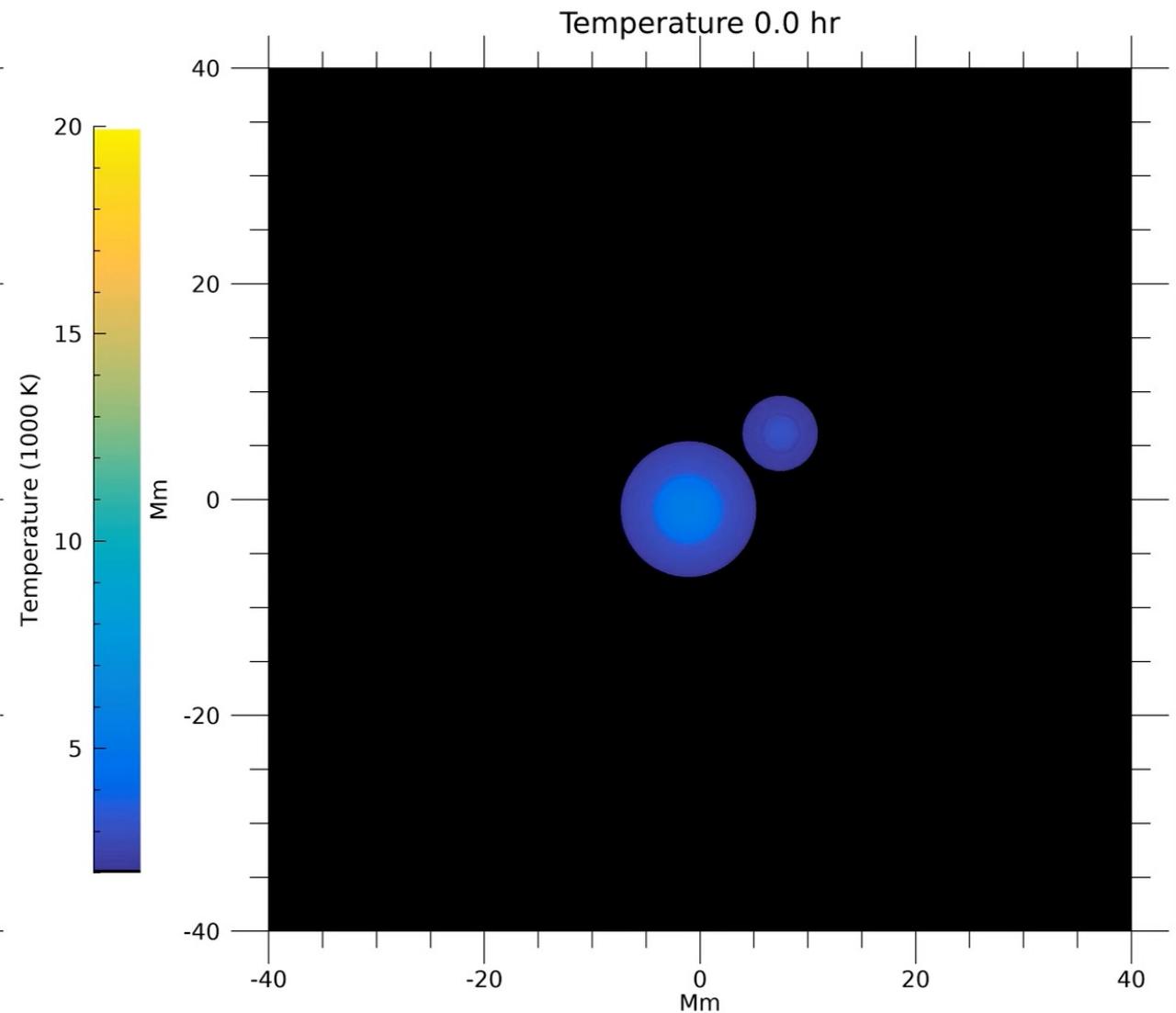
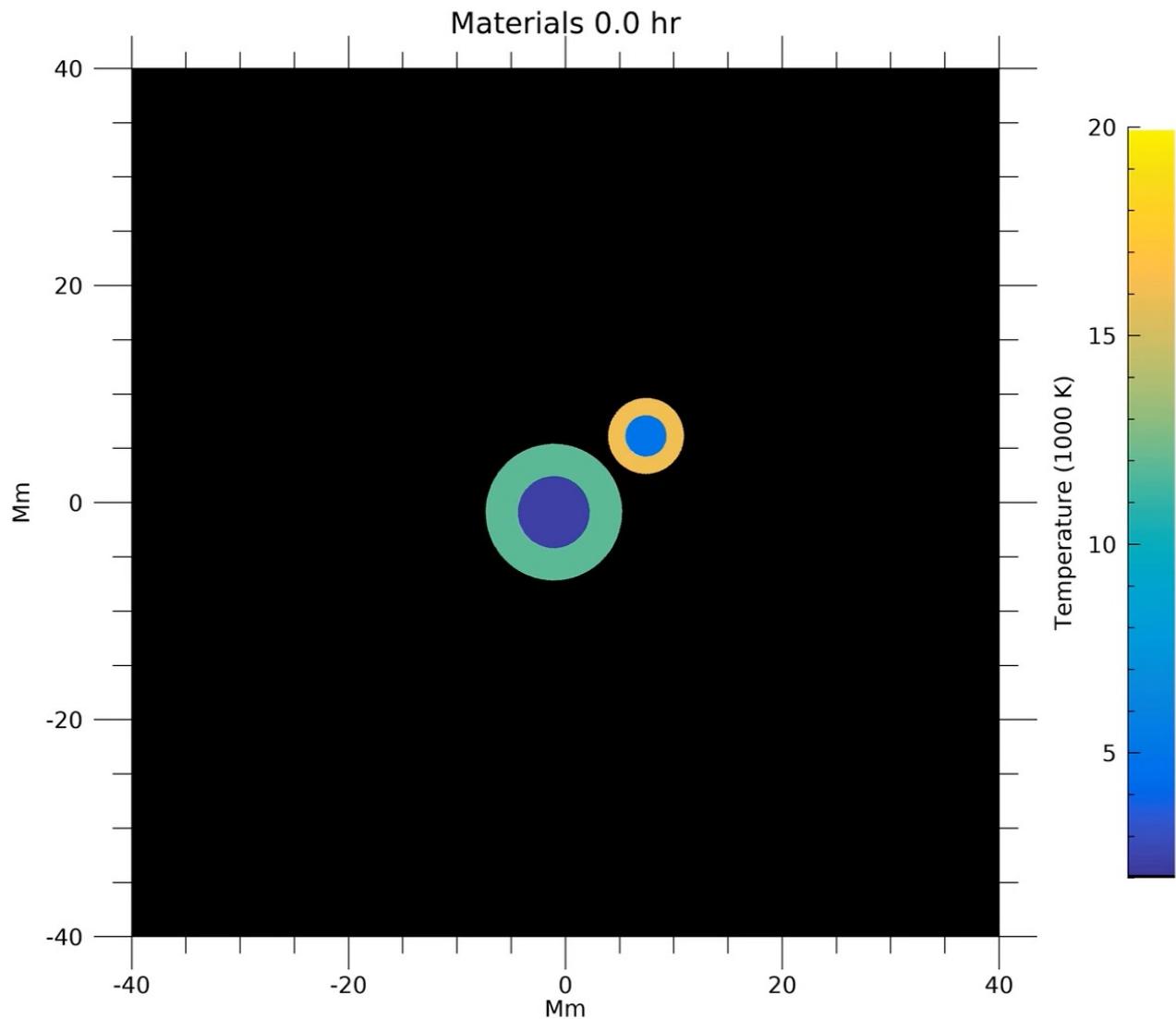
Z FUNDAMENTAL SCIENCE



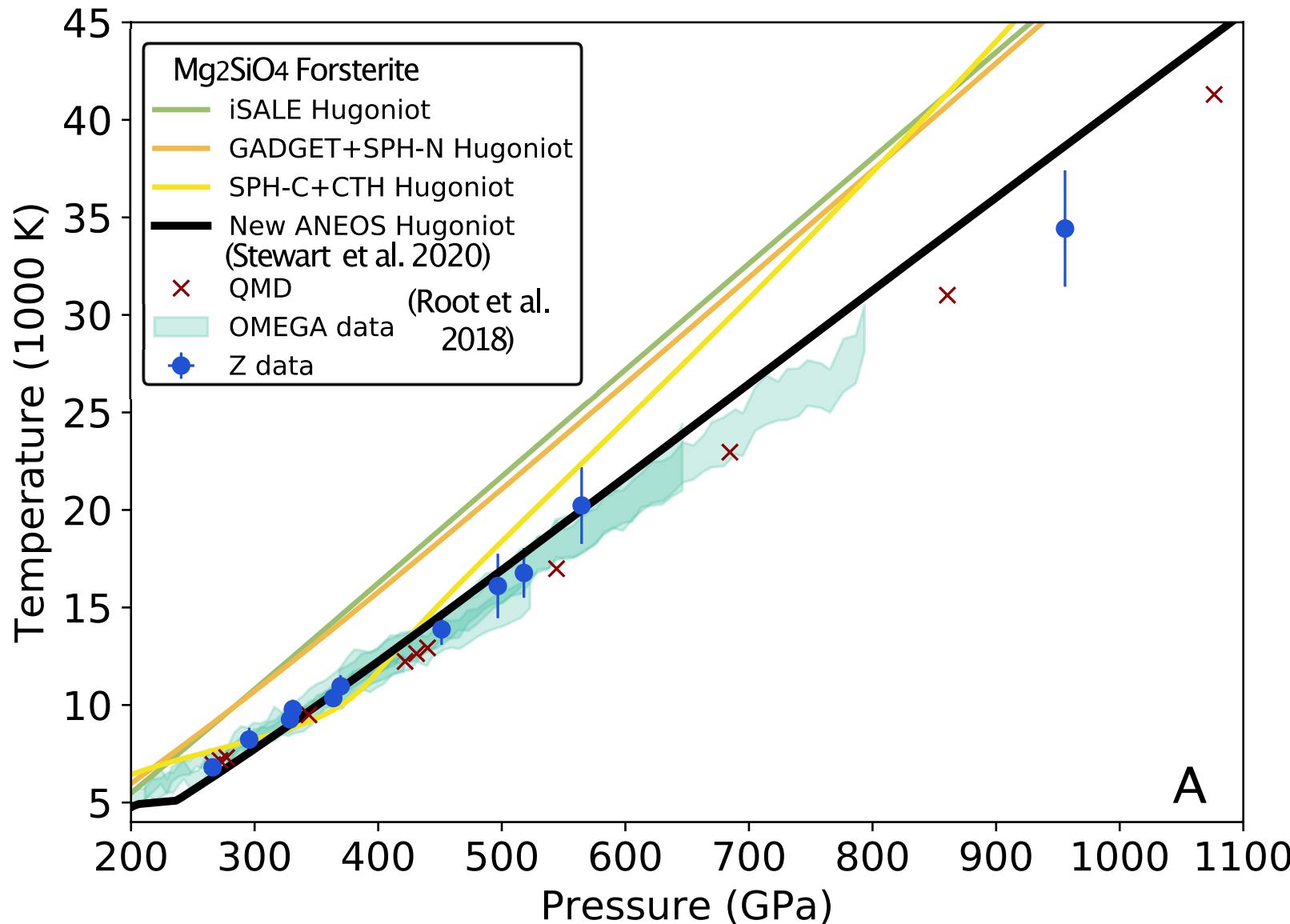
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*This work describes objective technical results and analysis. Any subjective views or opinions that might be expressed in the work do not necessarily represent the views of the U.S. Department of Energy or the United States Government.*

# Moon-forming giant impact



# Shock temperatures are a particular problem

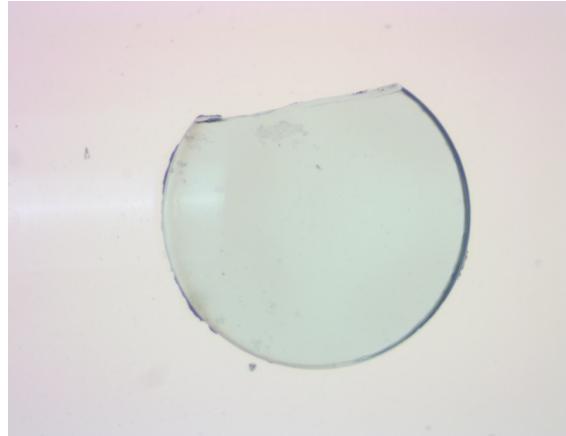




$(\text{Mg},\text{Fe})\text{SiO}_3$   
(Enstatite/Bronzite)

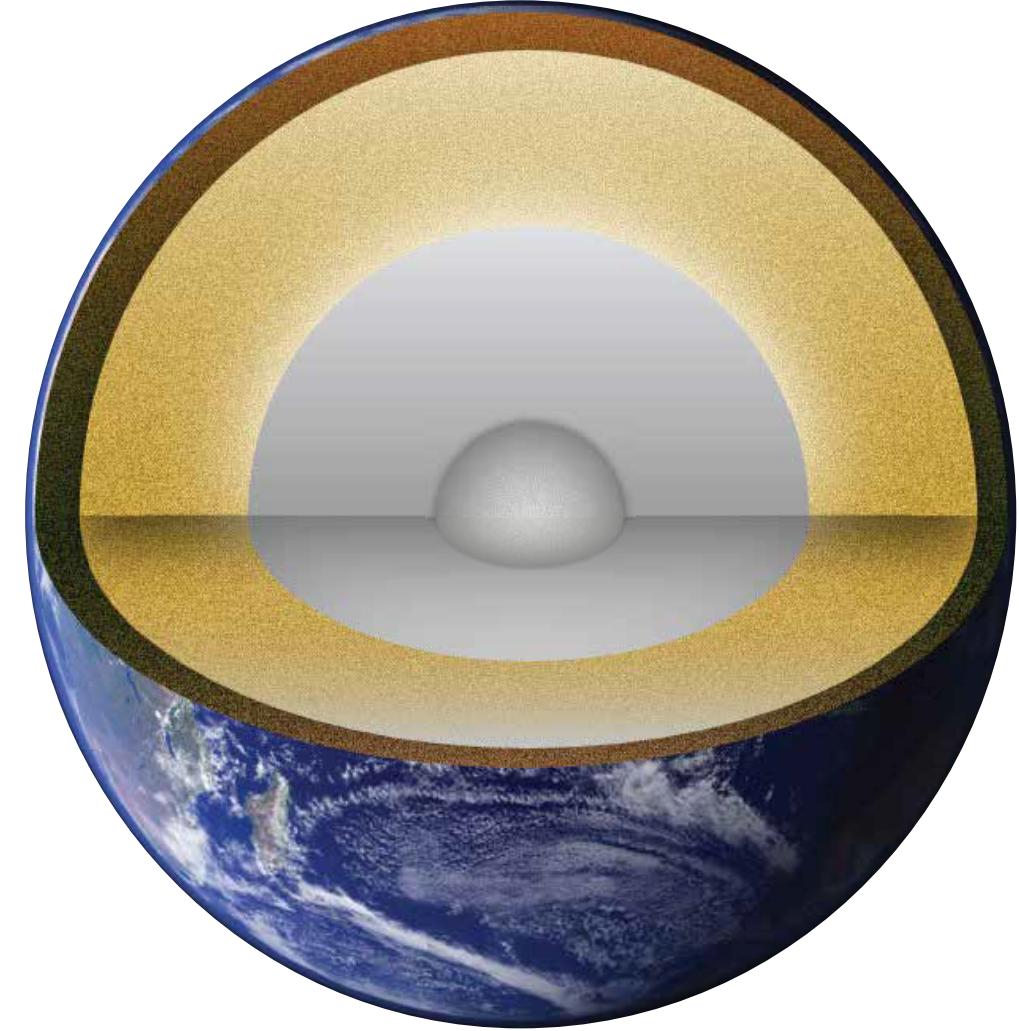


$(\text{Mg},\text{Fe})_2\text{SiO}_4$   
(Forsterite/Olivine)

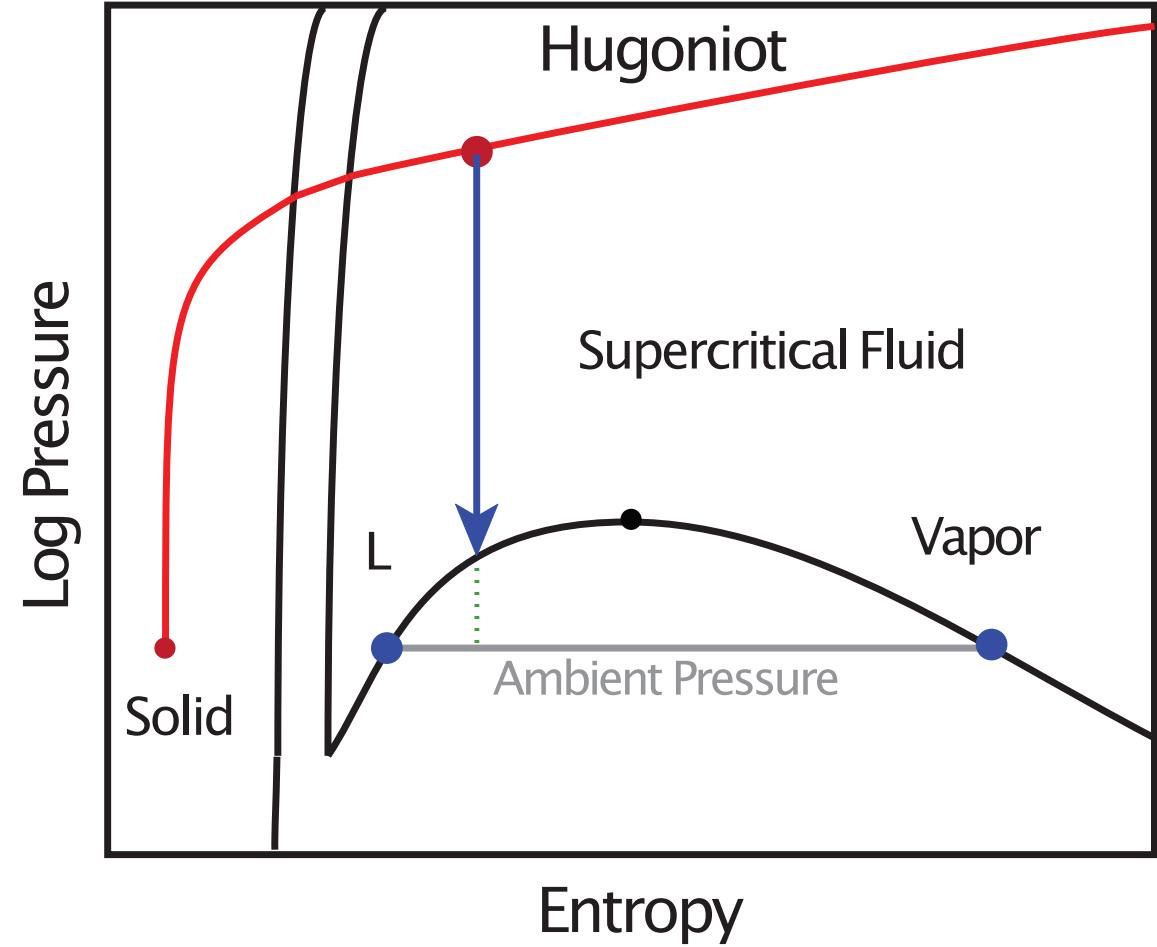
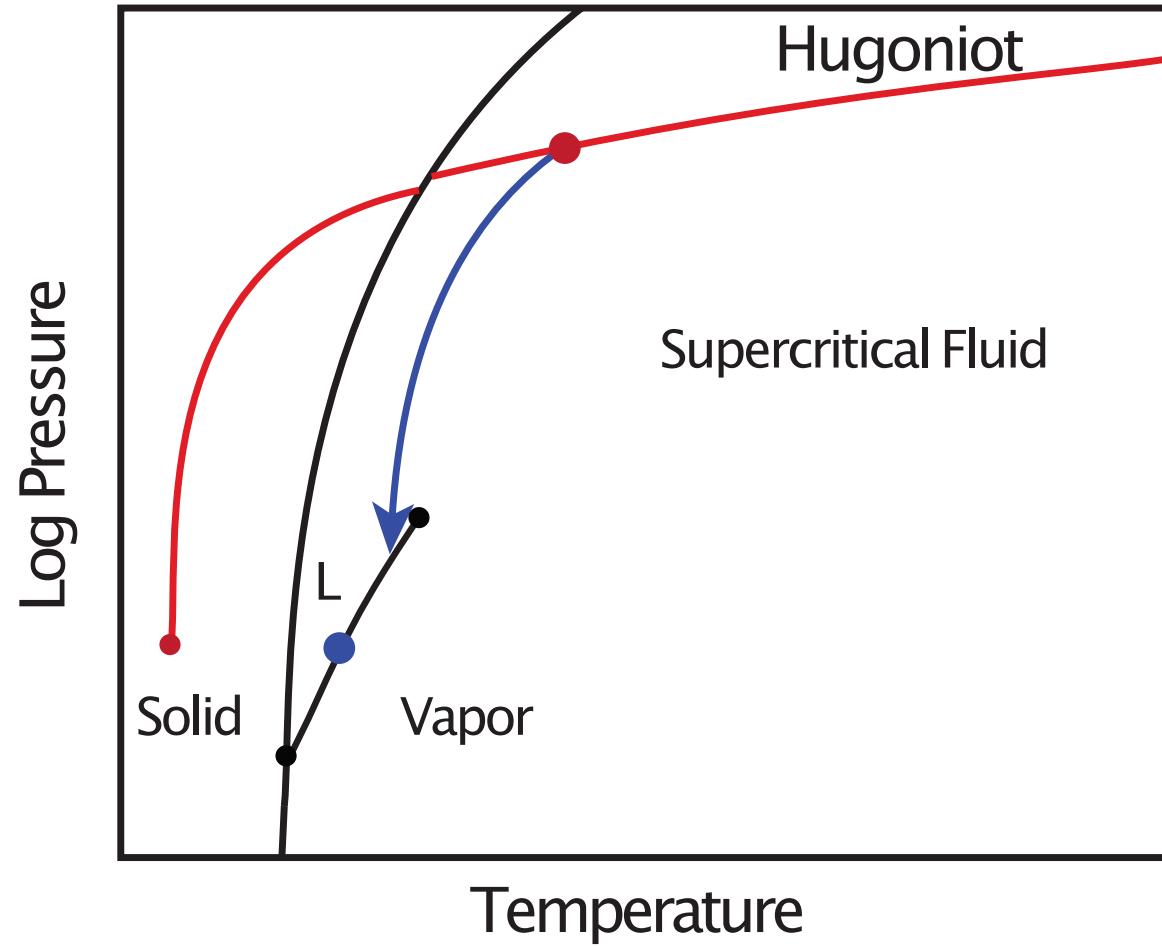


Pyrolite glass

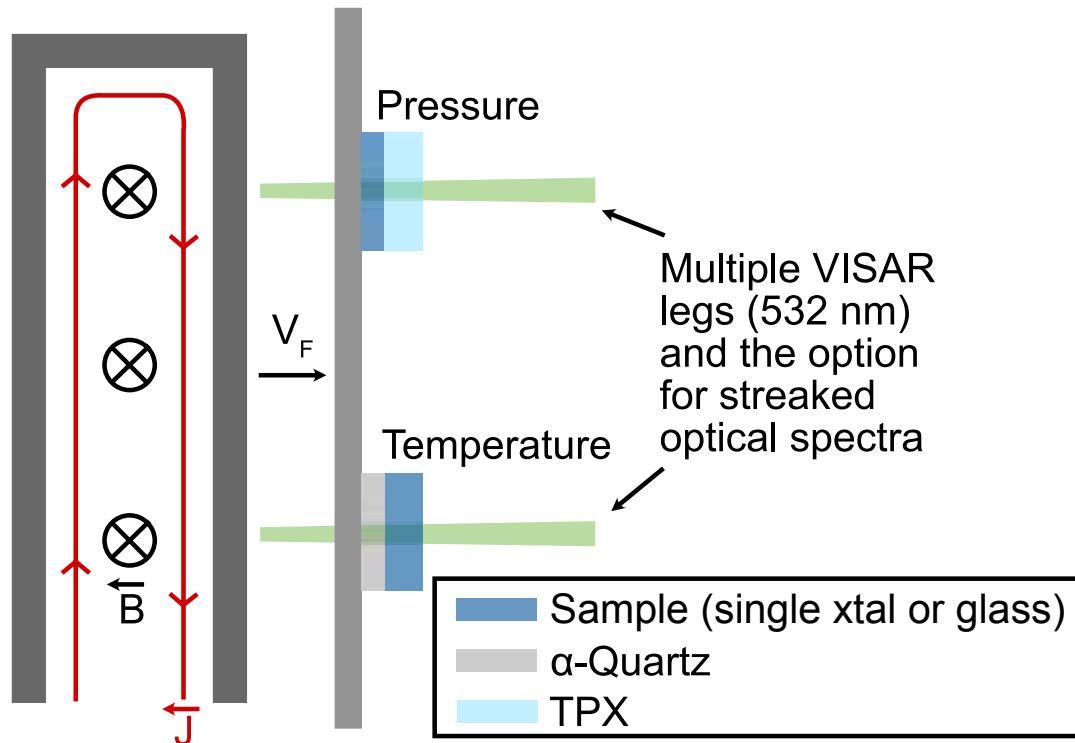
Fratanduono et al., *PRB*, 2018  
Root et al., *GRL*, 2018  
Chidester et al. *GRL*, 2021



# Shock Physics Experiments: Giant Impact P-T Paths



# Shock compression experiments at Z Machine

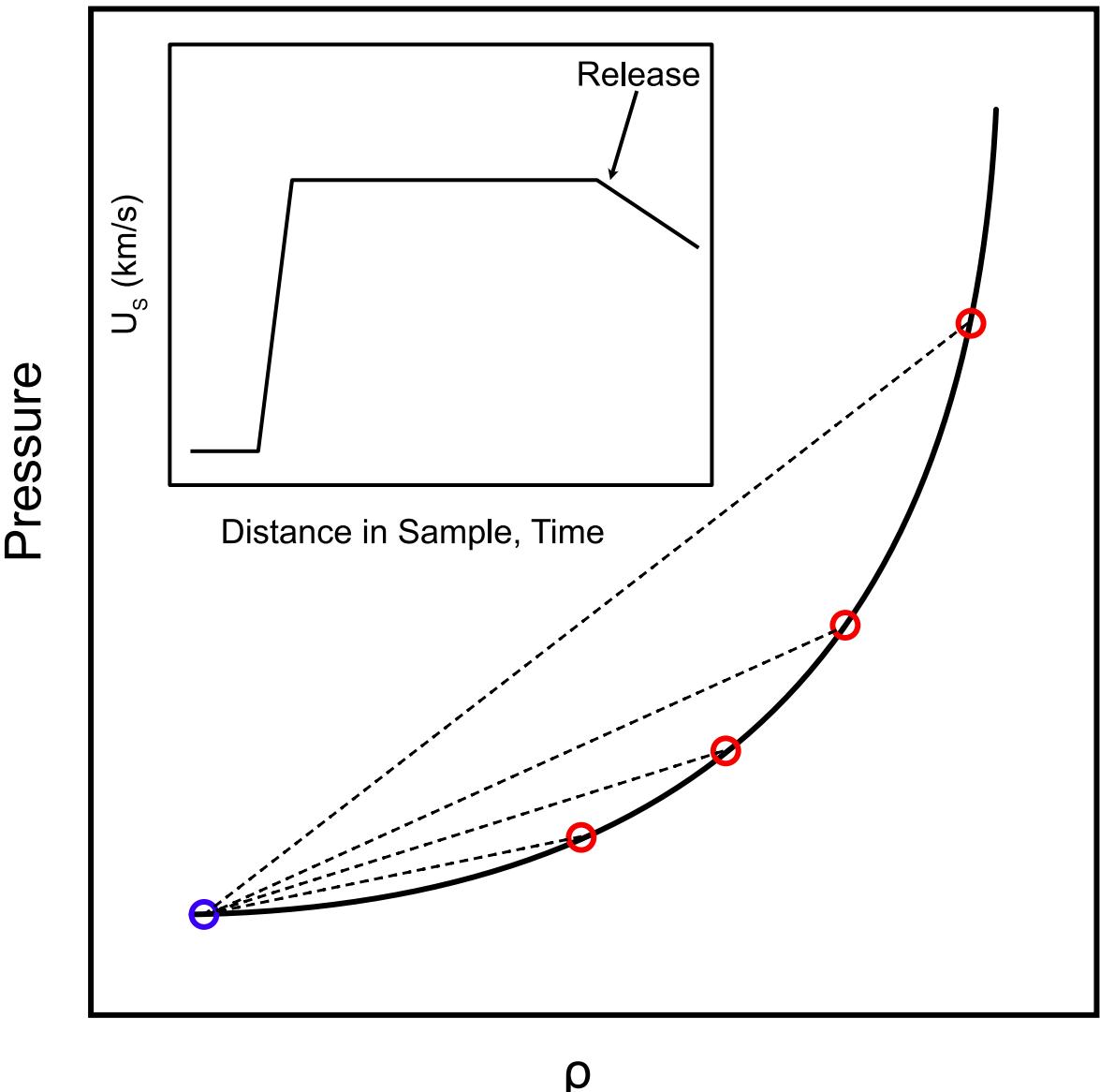


Rankine-Hugoniot jump condition

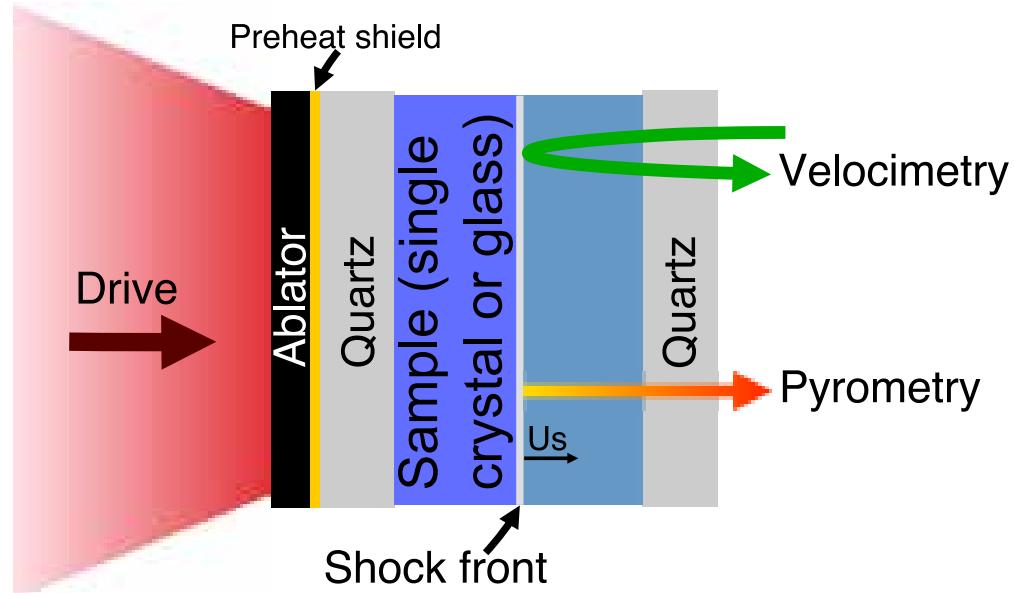
$$E - E_0 = \frac{1}{2} (P + P_0)(V_0 - V)$$

$$P - P_0 = \rho_0 (U_s - u_{p,0}) (u_p - u_{p,0})$$

$$\rho_0 / \rho = 1 - (u_p - u_{p,0}) / (U_s - u_{p,0})$$



# Shock compression experiments at OMEGA EP

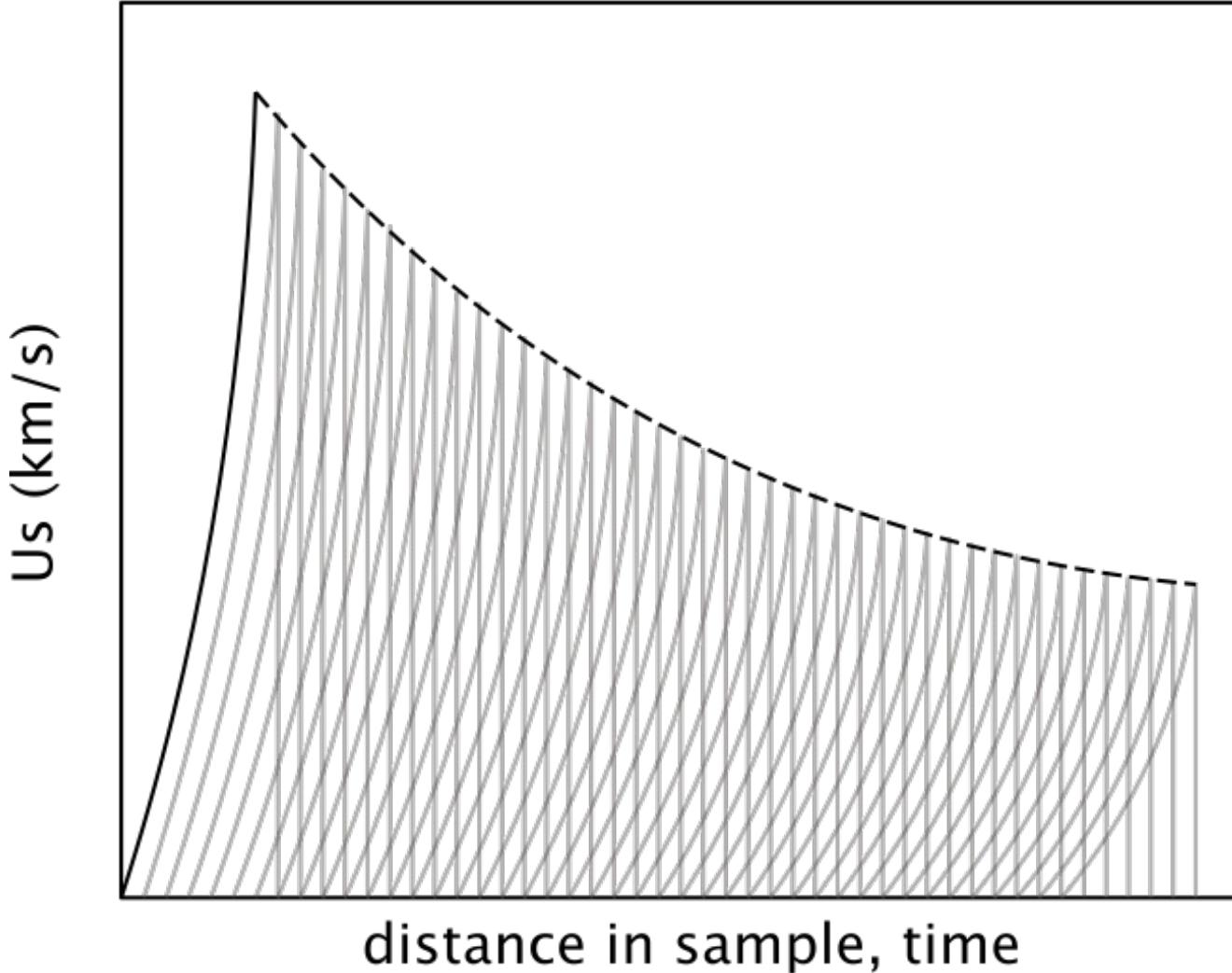


## Rankine-Hugoniot jump condition

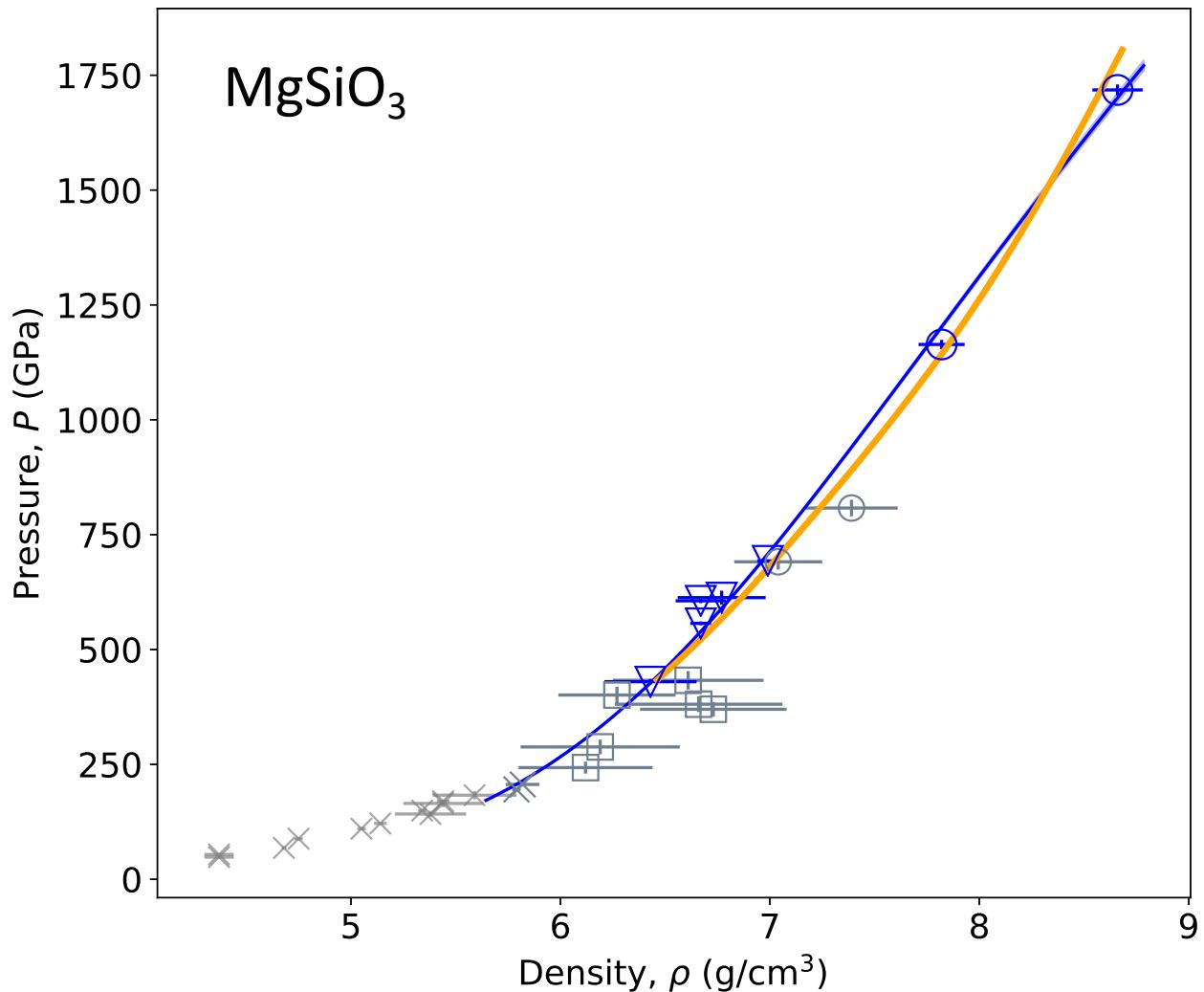
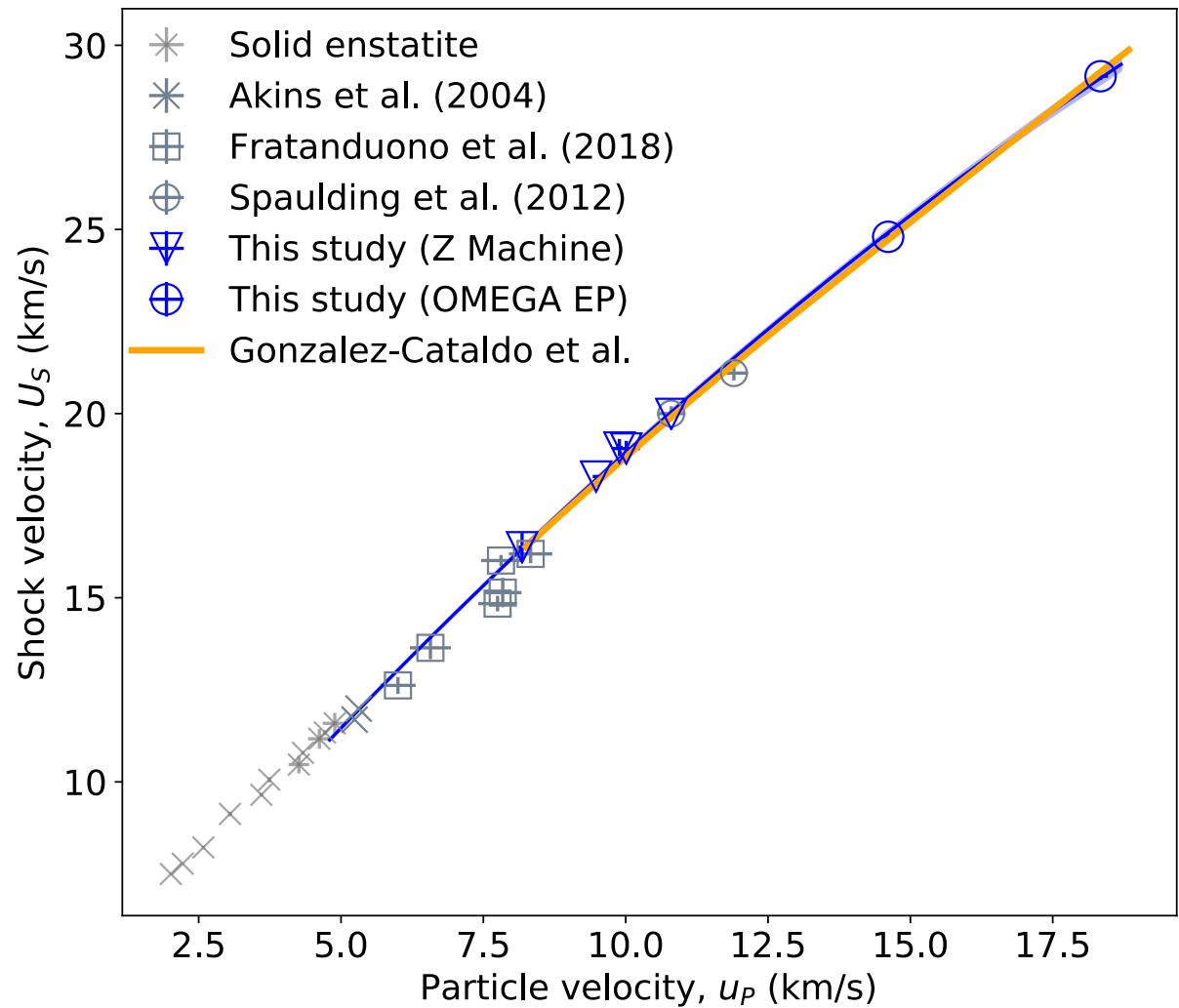
$$E - E_0 = \frac{1}{2} (P + P_0)(V_0 - V)$$

$$P - P_0 = \rho_0 (U_s - u_{p,0}) (u_p - u_{p,0})$$

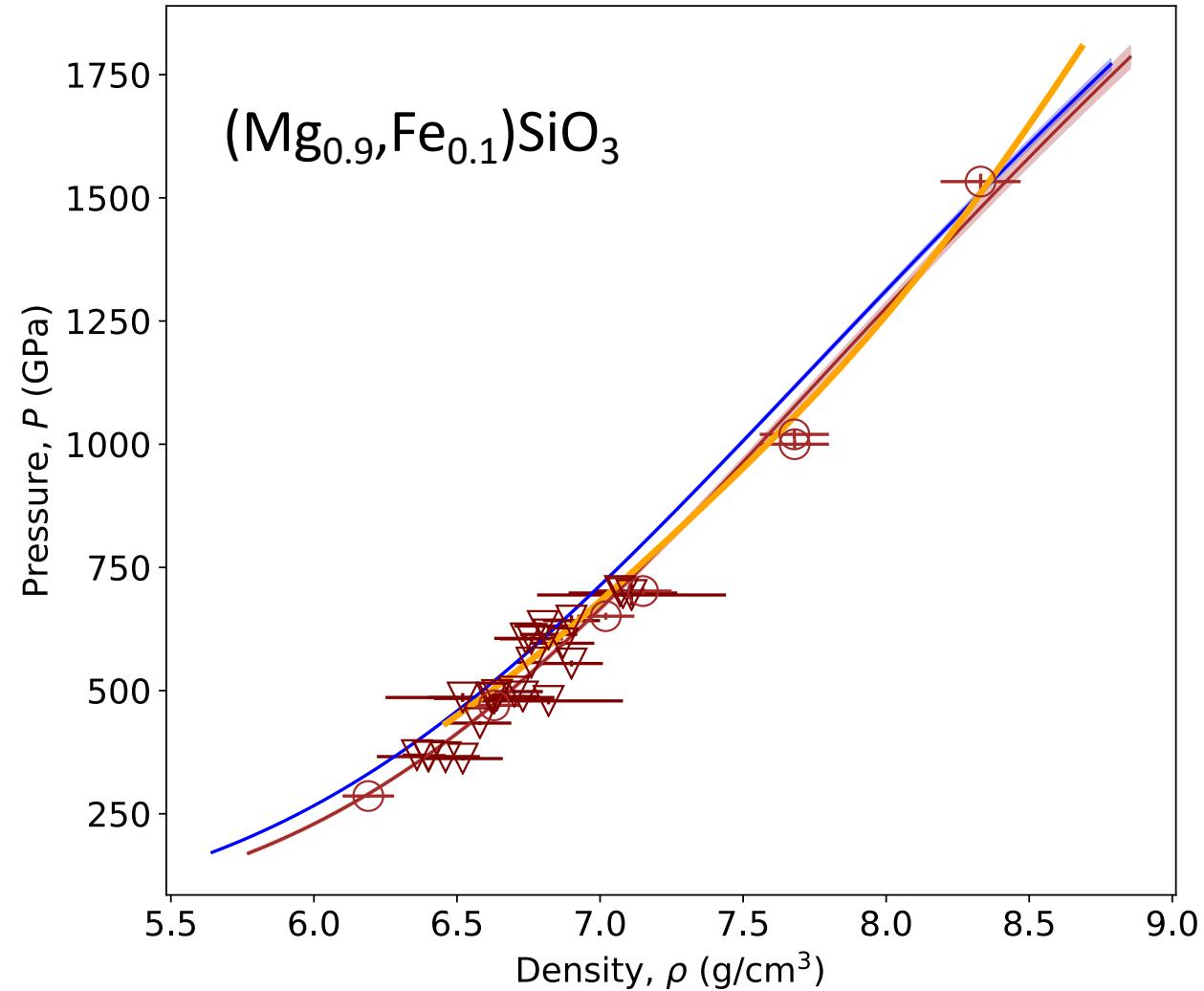
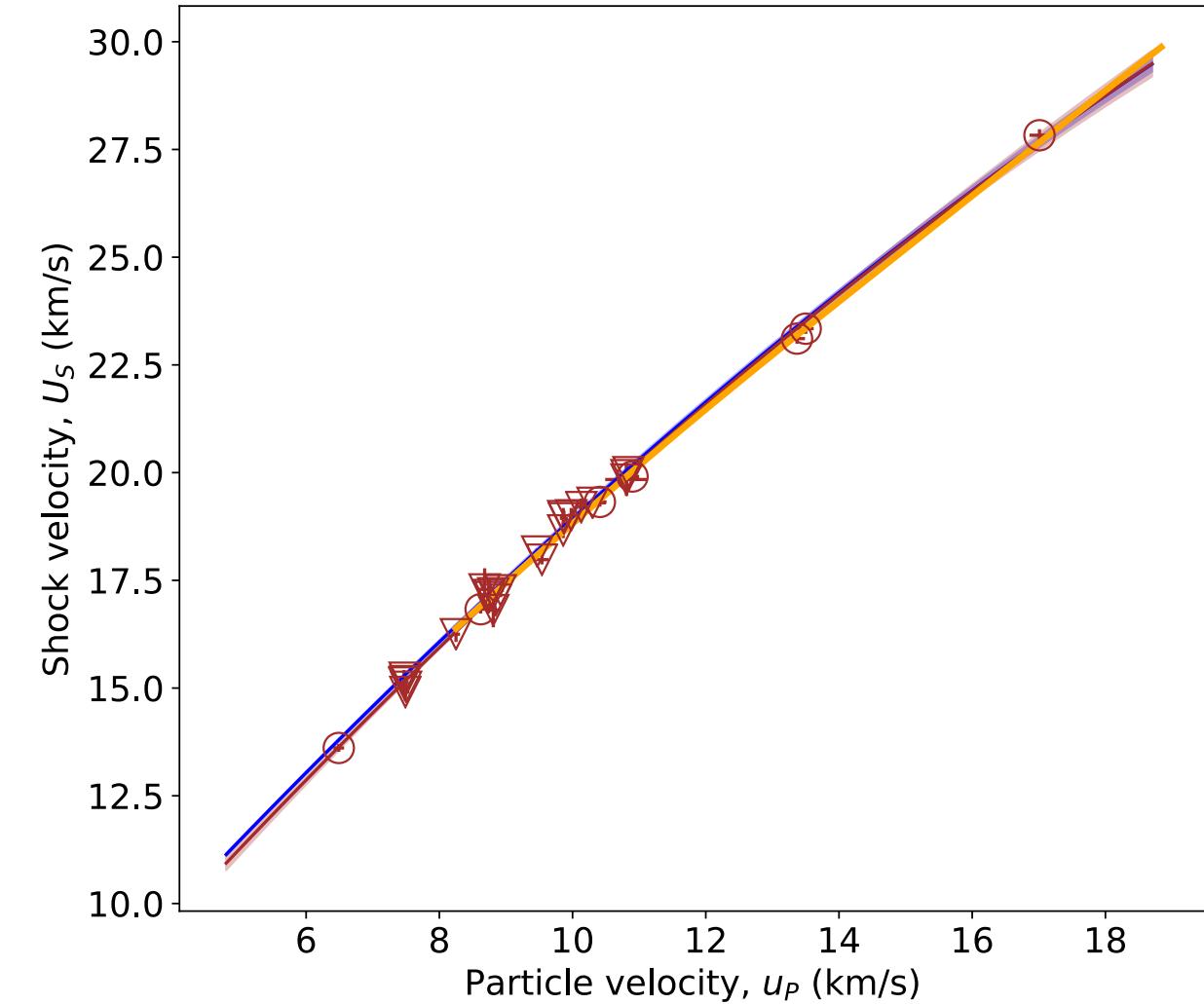
$$\rho_0 / \rho = 1 - (u_p - u_{p,0}) / (U_s - u_{p,0})$$



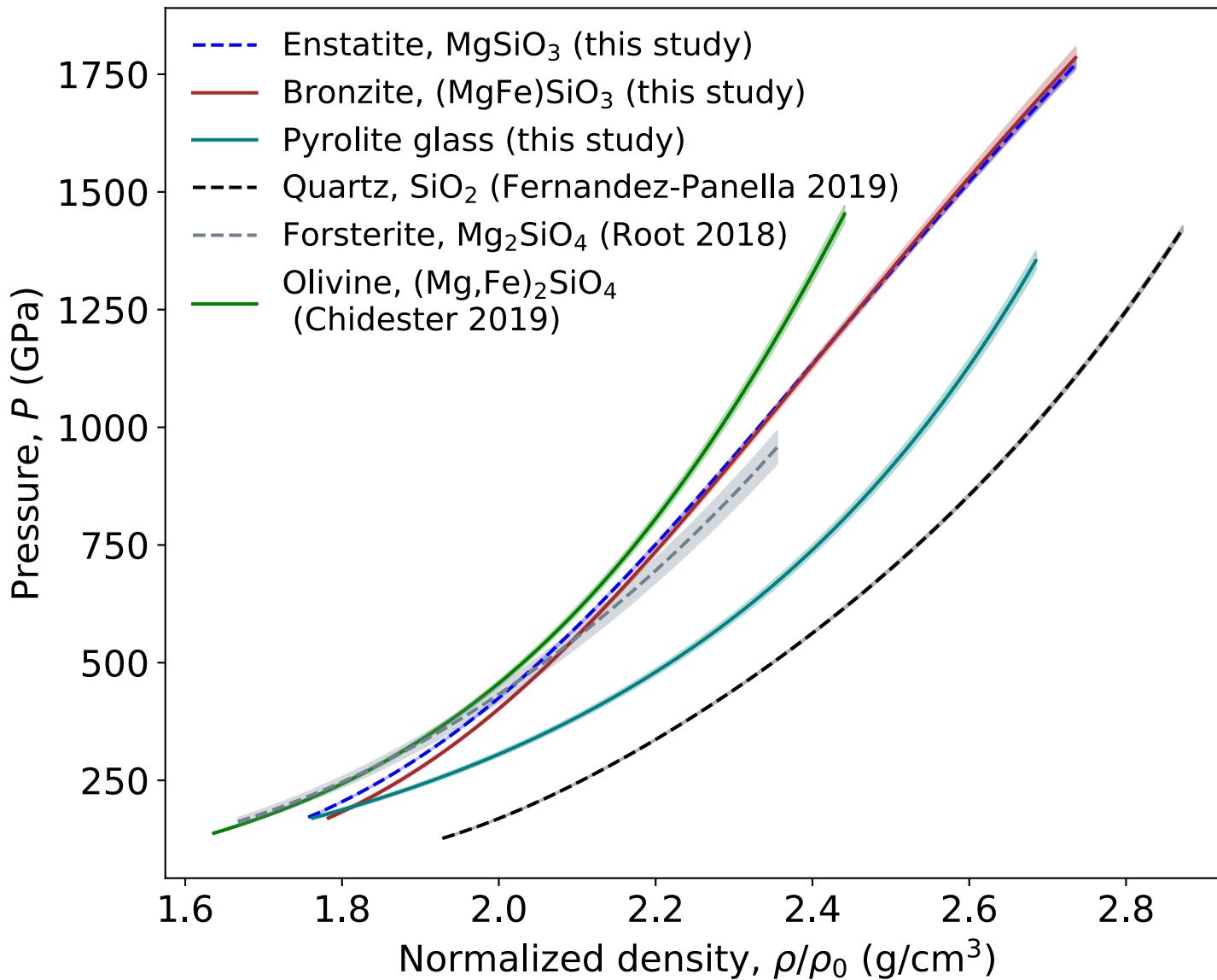
# Enstatite compressibility results



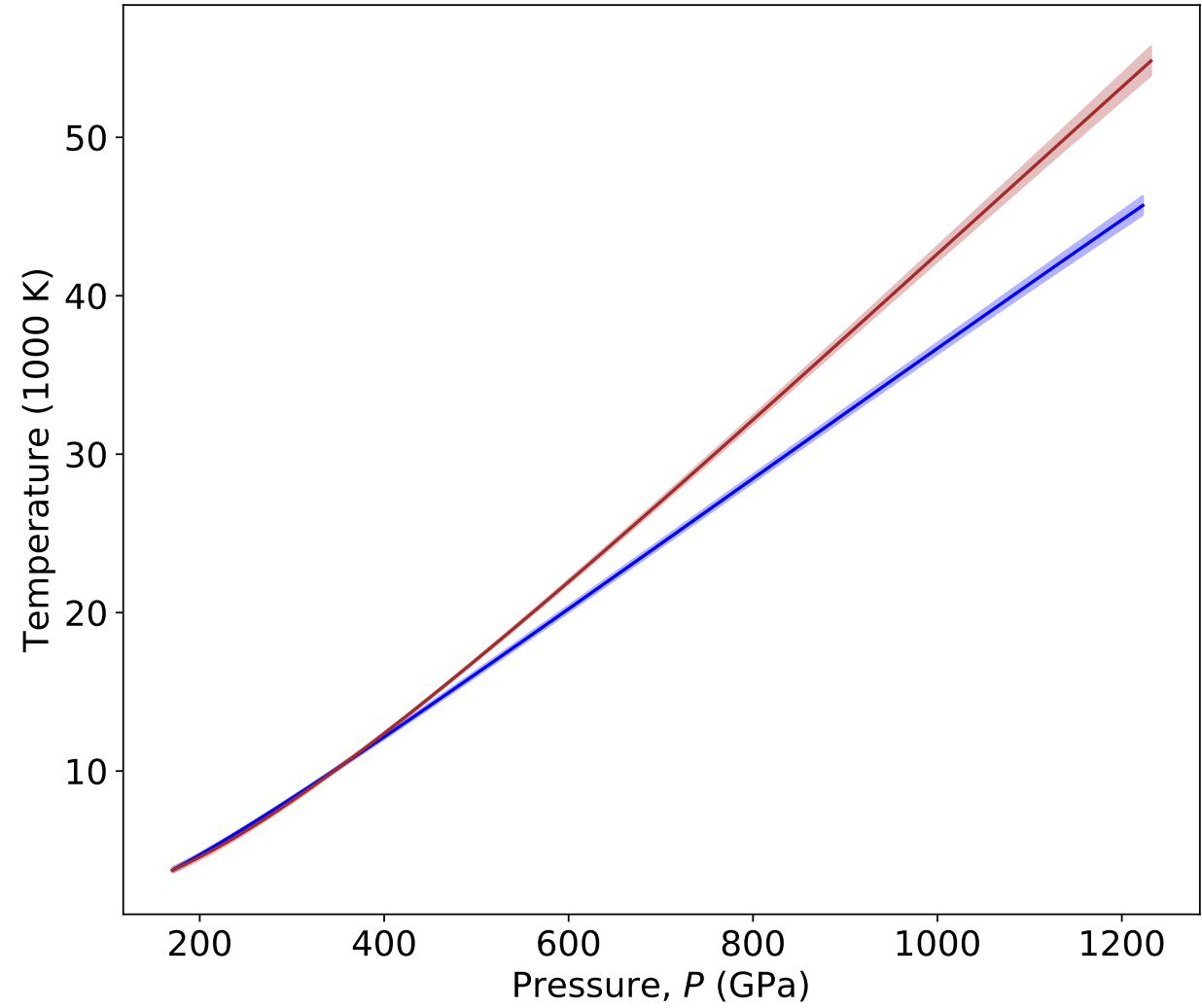
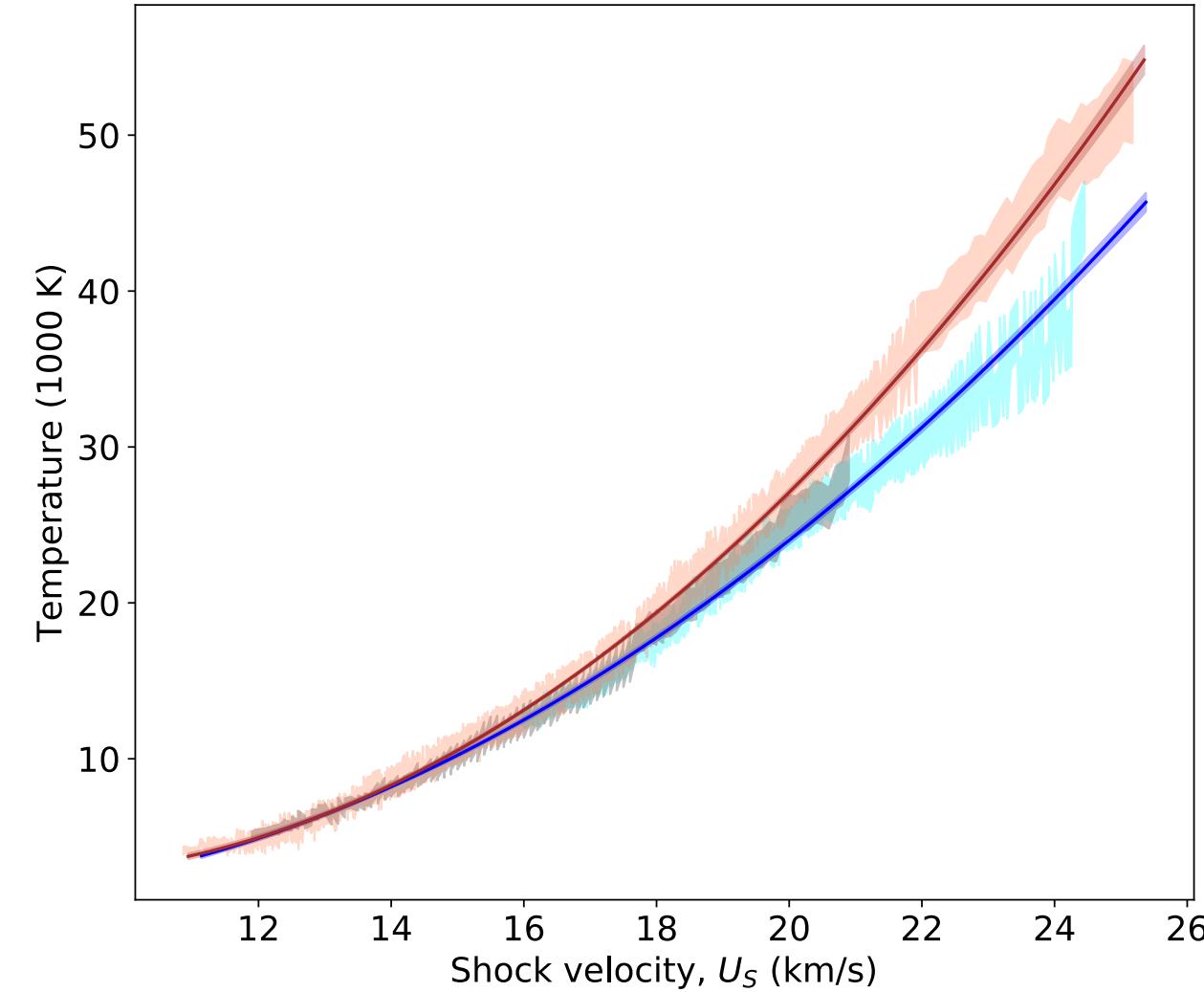
# Bronzite compressibility results



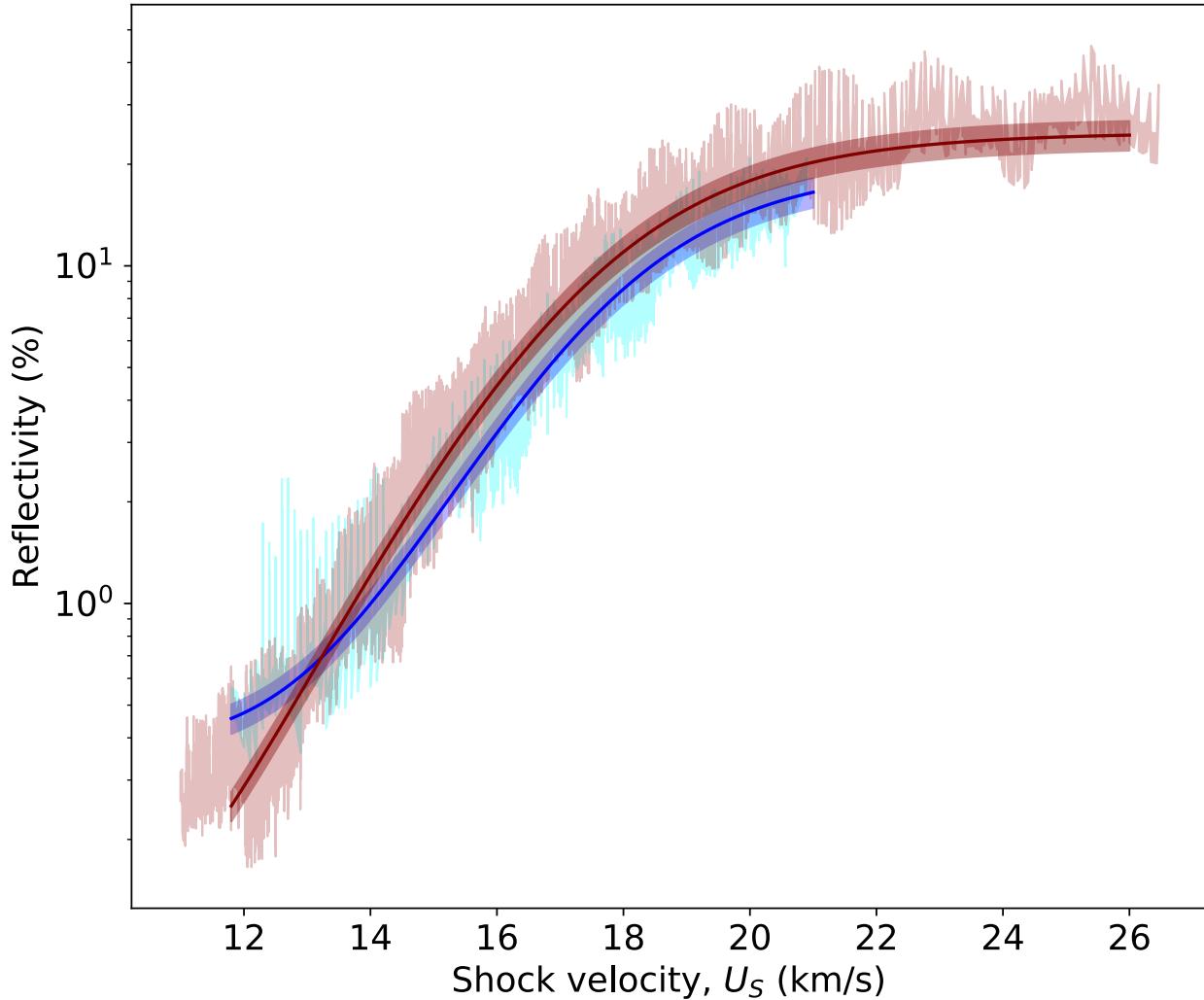
# Comparison to other silicates



# Temperature



# Reflectivity



Silicates become conductive at high temperatures!

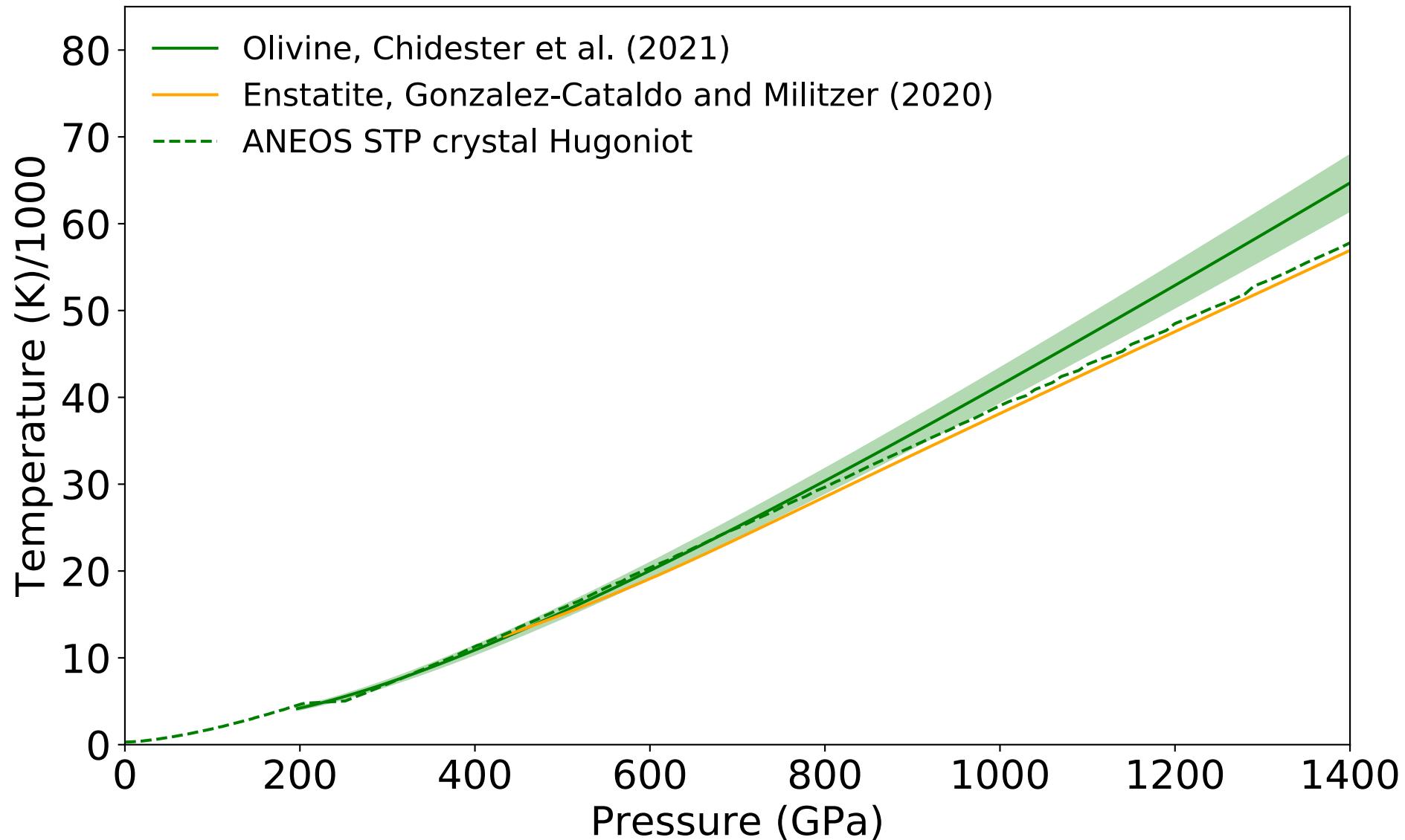
Fe atoms likely increase the effect.

## Conclusions

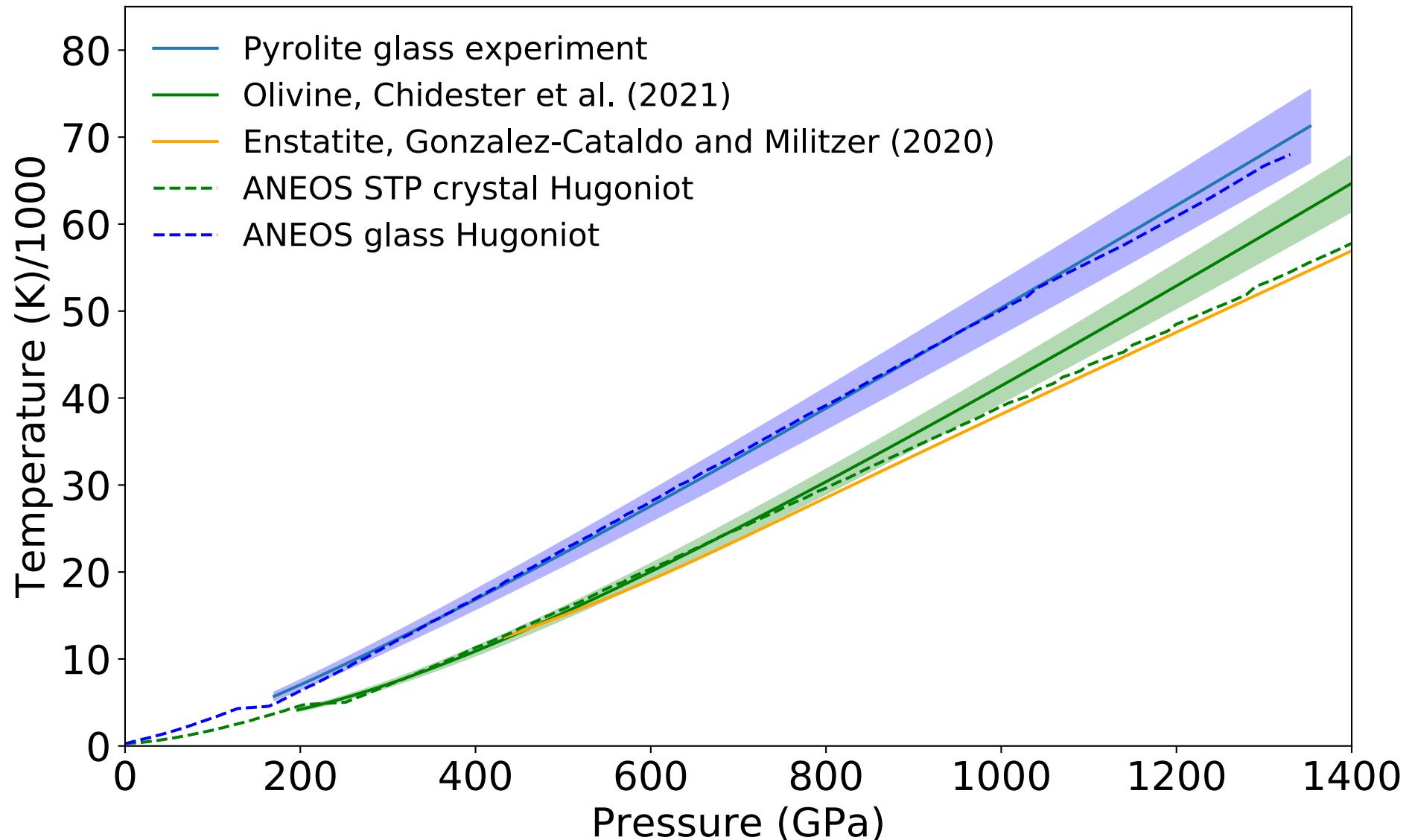
- Fe-bearing silicates seem to shock to higher temperatures than their Fe-free endmembers
- This is related to the change in reflectivity, likely a different electronic contribution to the heat capacity as the material becomes more conductive

Might stop here, but if I have time I'll  
continue on to other slides

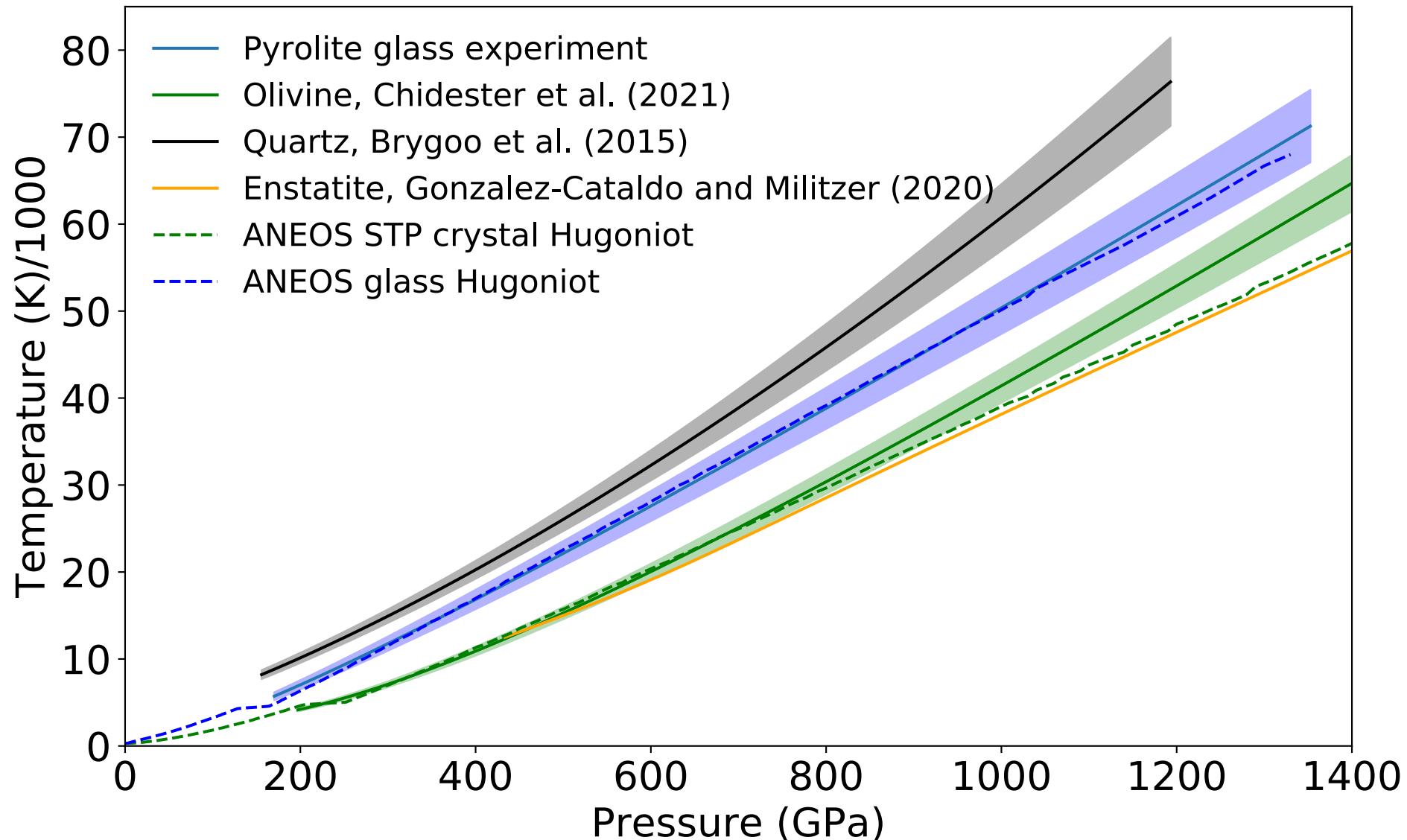
# Analytical EOS constrained by experimental results



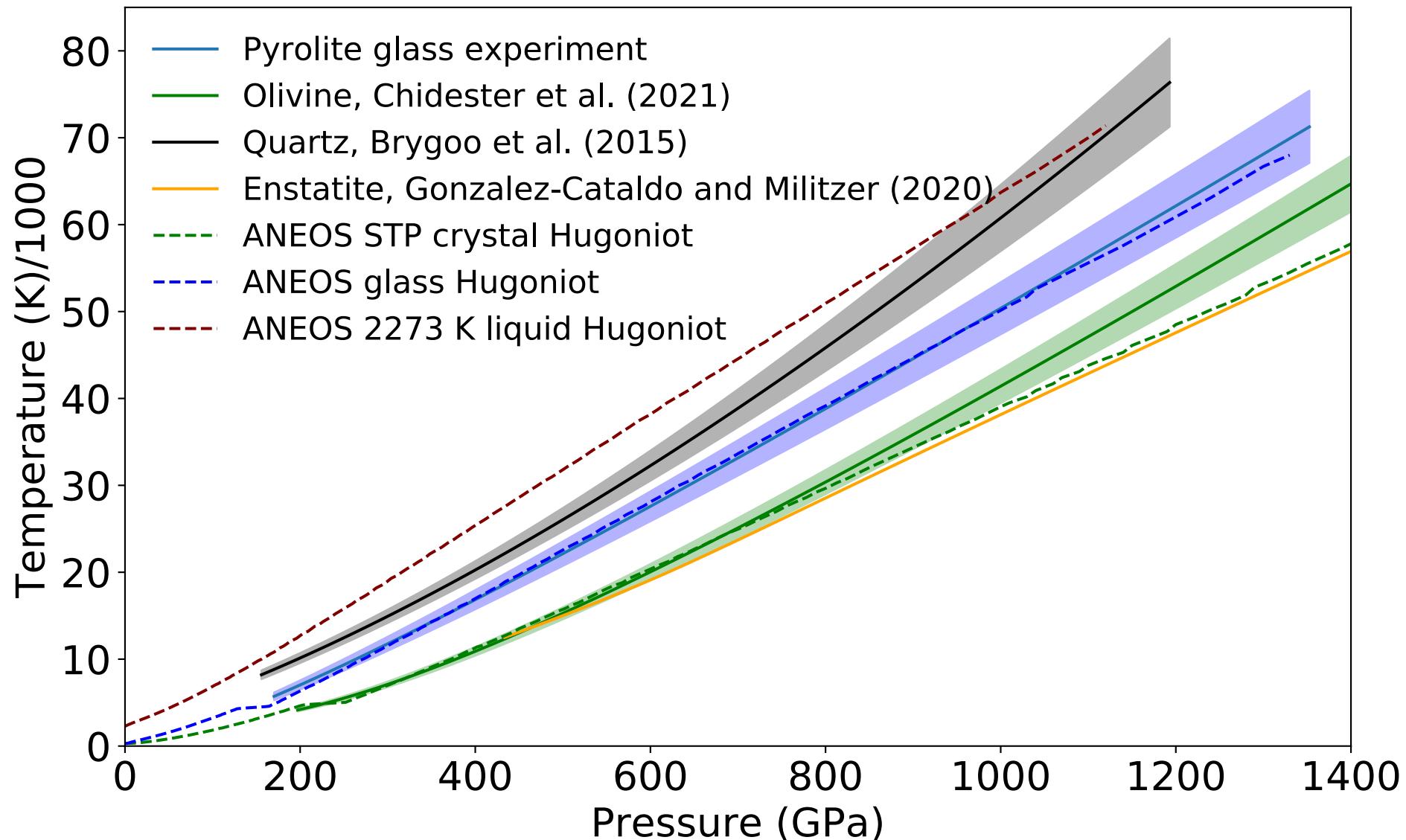
# Analytical EOS constrained by experimental results



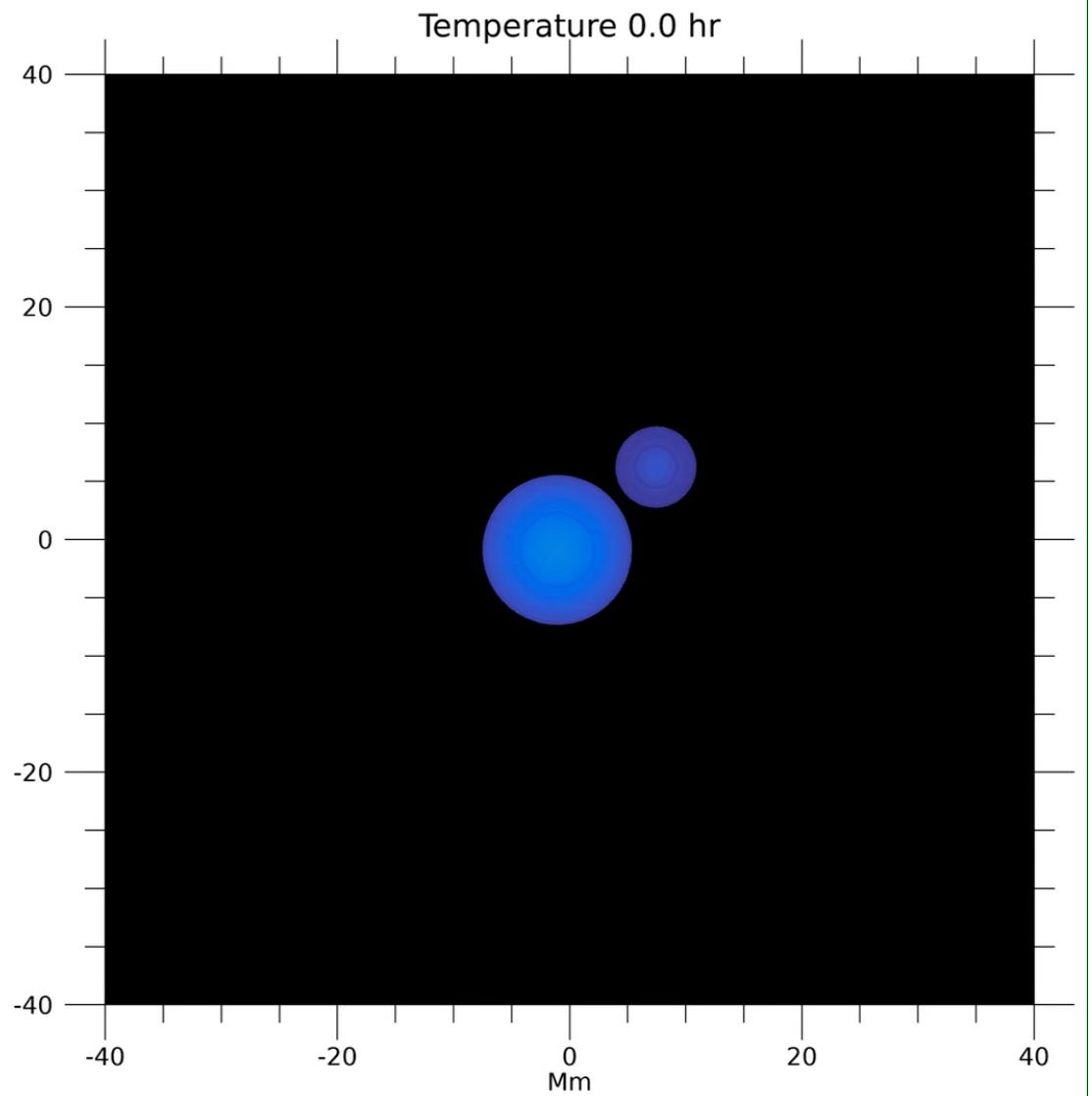
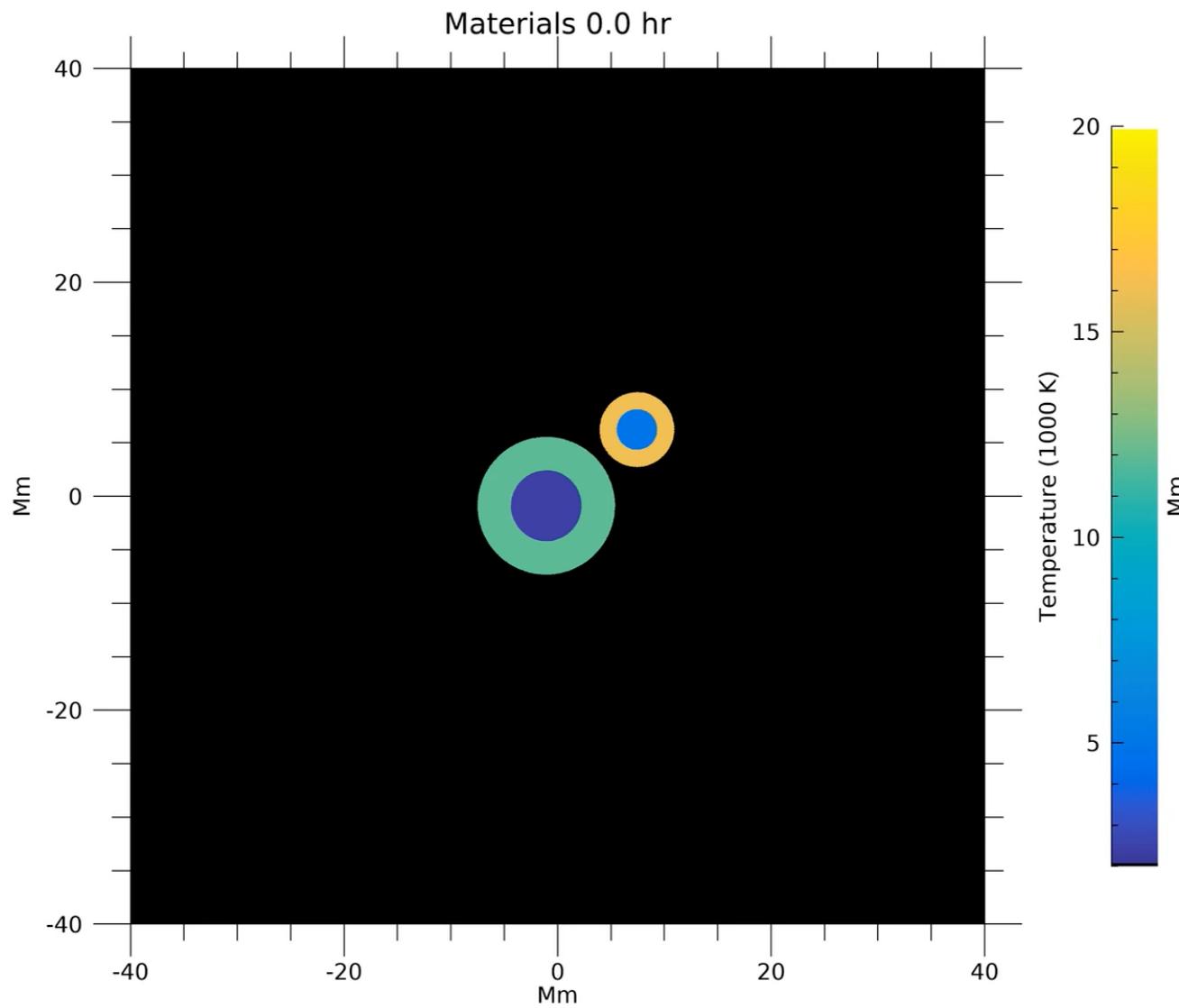
# Analytical EOS constrained by experimental results



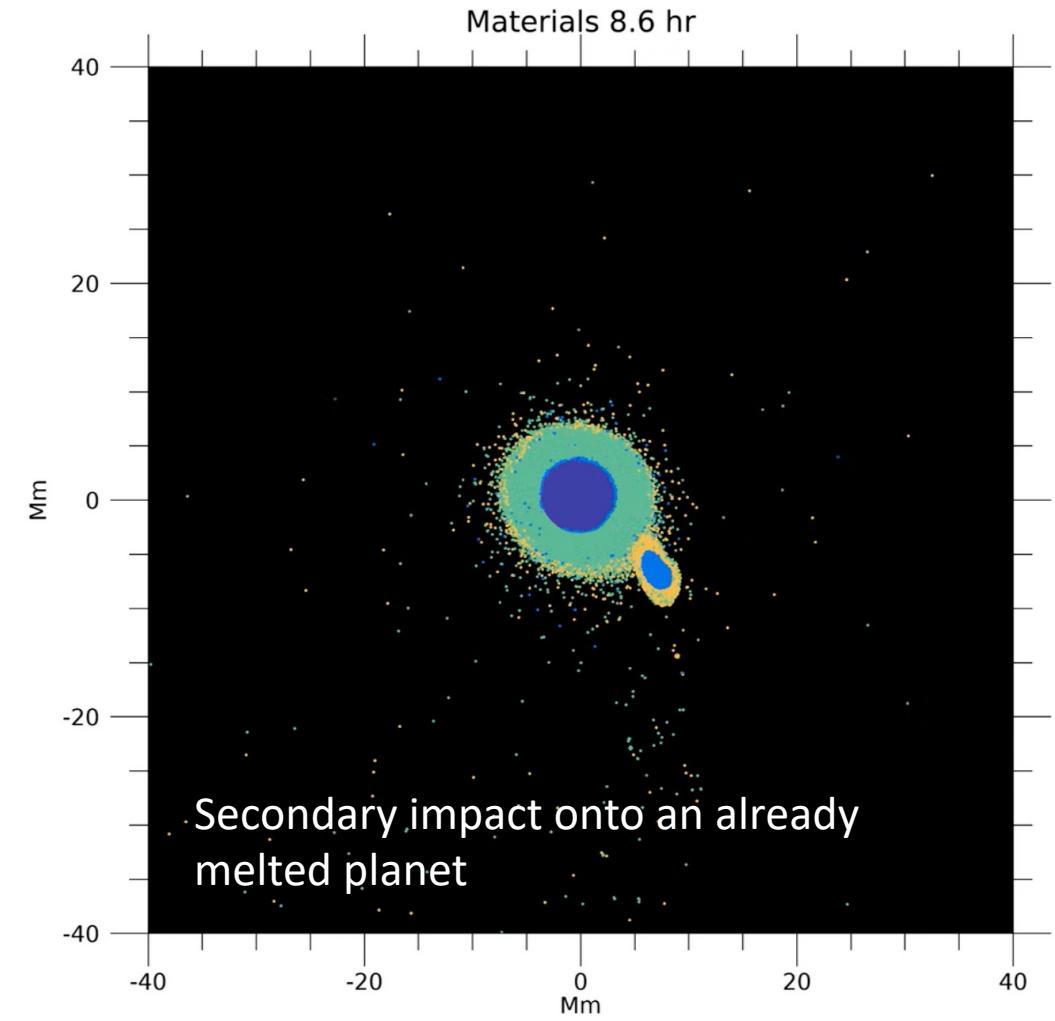
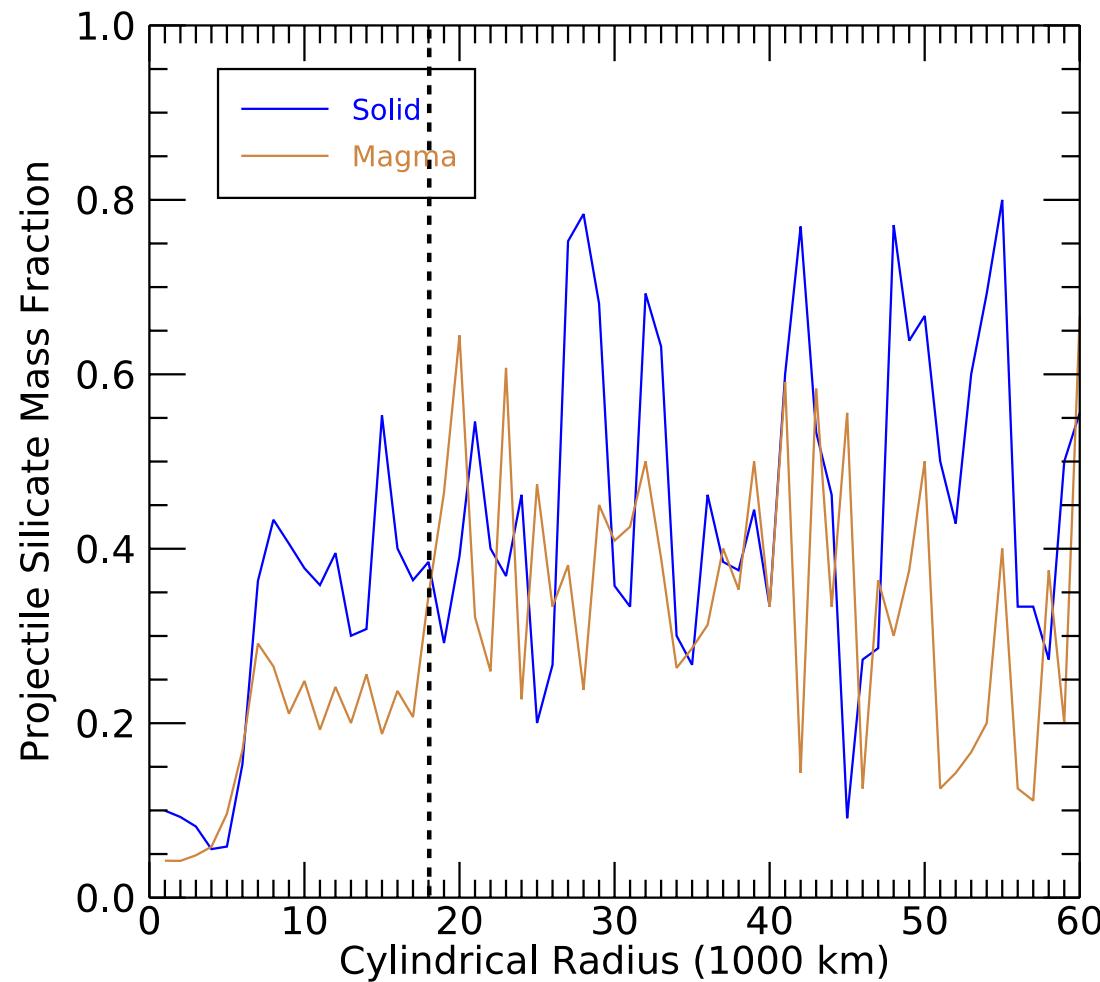
# Analytical EOS constrained by experimental results



# Magma ocean proto-Earth

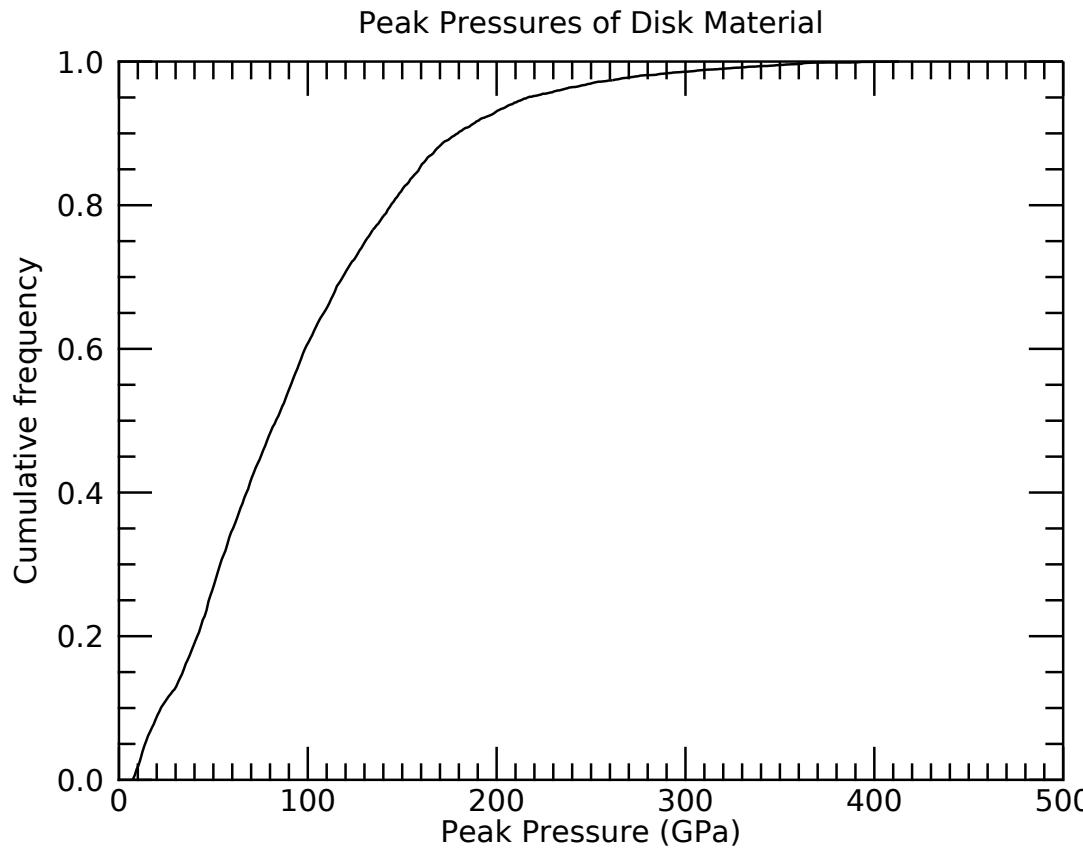


# Result 1: Giant impacts onto a magma ocean are not different than giant impacts onto a solid planet

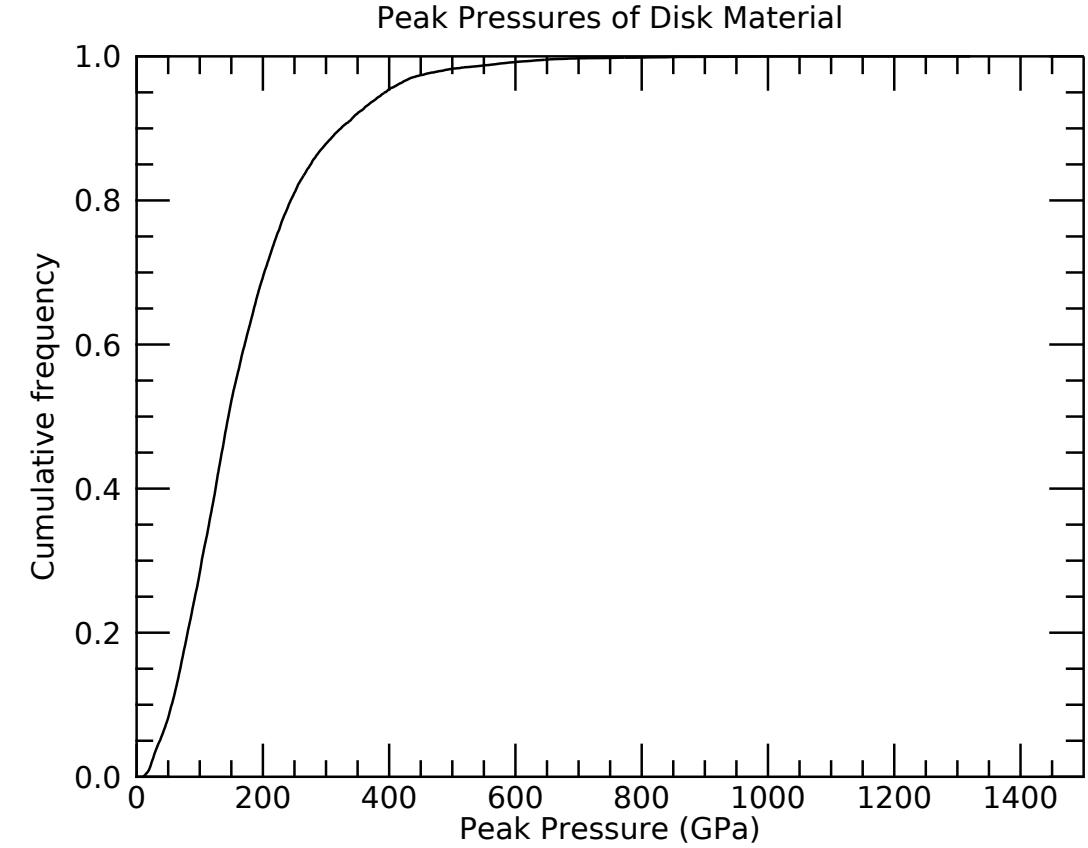


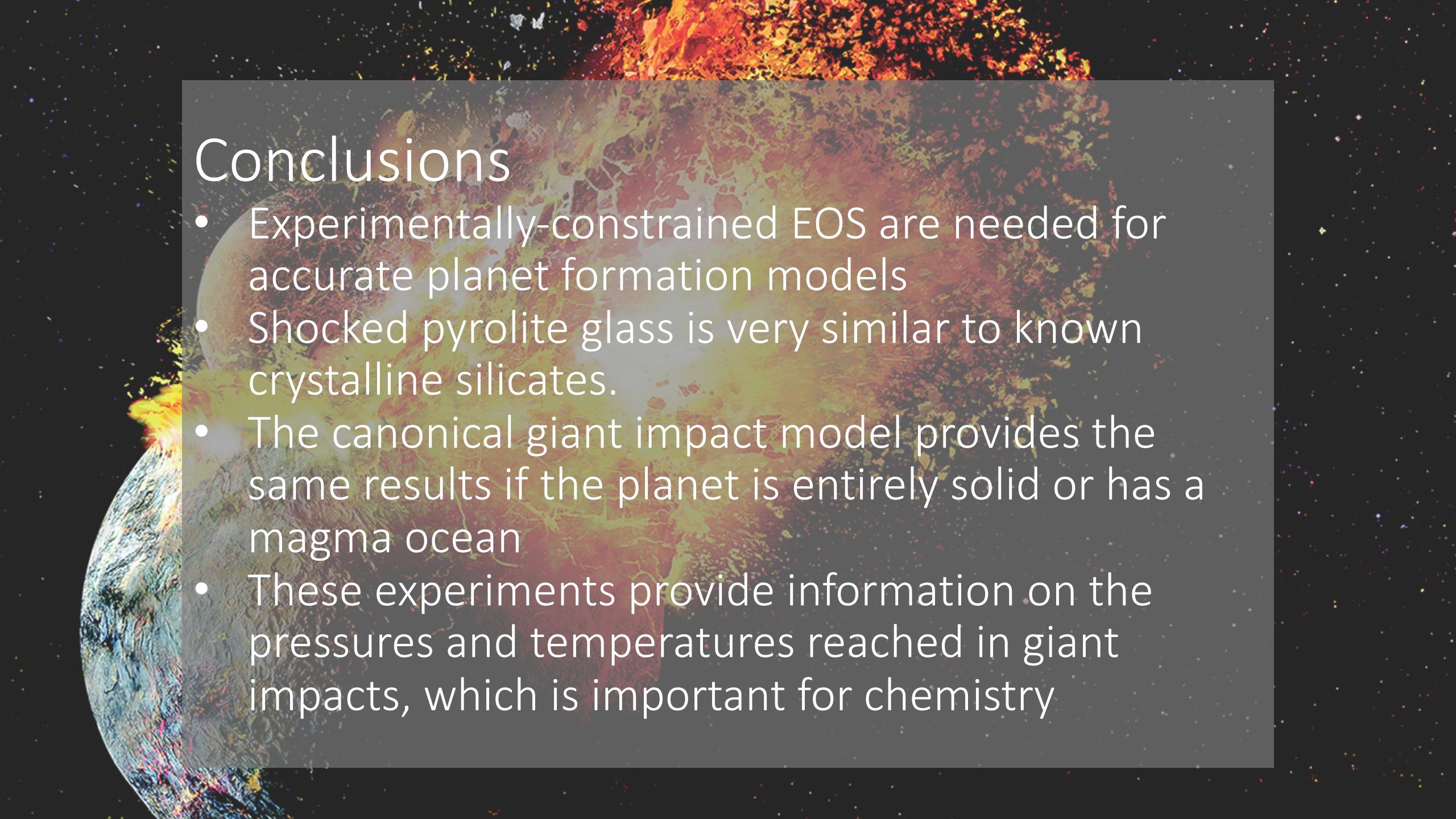
# Result 2: The disk is heavily shocked

## Mars-Mass Giant Impact



## Two Half-Earths Colliding





## Conclusions

- Experimentally-constrained EOS are needed for accurate planet formation models
- Shocked pyrolite glass is very similar to known crystalline silicates.
- The canonical giant impact model provides the same results if the planet is entirely solid or has a magma ocean
- These experiments provide information on the pressures and temperatures reached in giant impacts, which is important for chemistry