

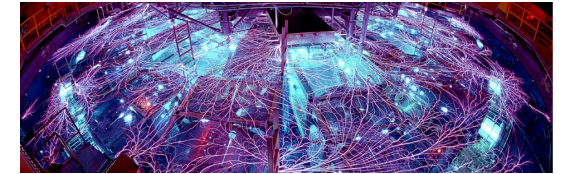
Experimental temperature measurements of Fe-bearing silicate minerals and glasses to 1.6 TPa

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Collaborators



Funding



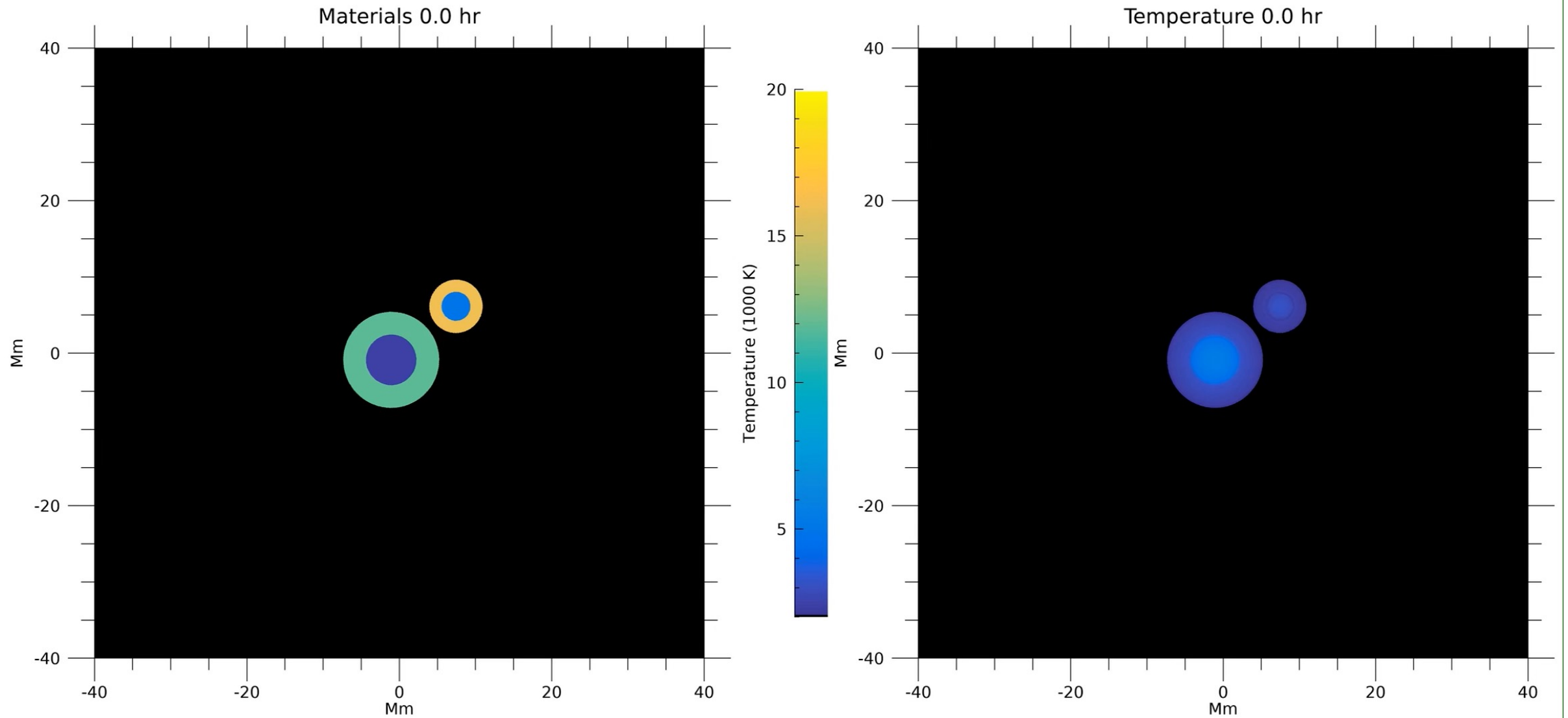
Z FUNDAMENTAL SCIENCE



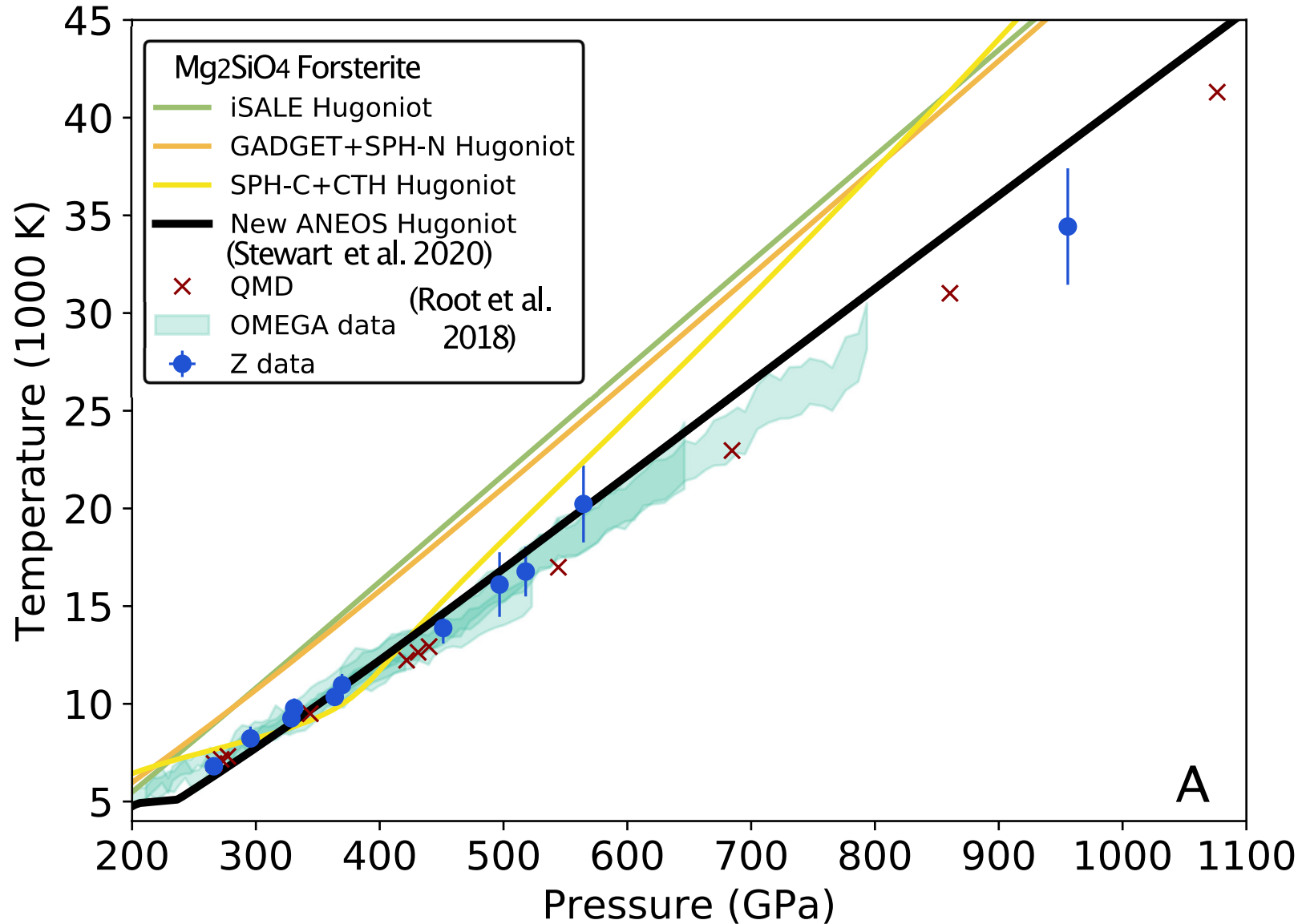
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This work describes objective technical results and analysis. Any subjective views or opinions that might be expressed in the work do not necessarily represent the views of the U.S. Department of Energy or the United States Government.

Moon-forming giant impact



Shock temperatures are a particular problem



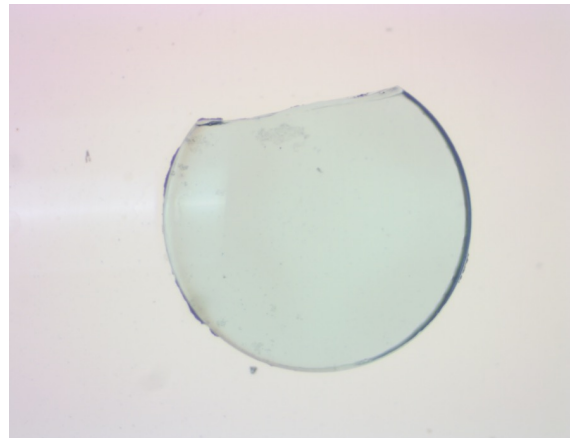
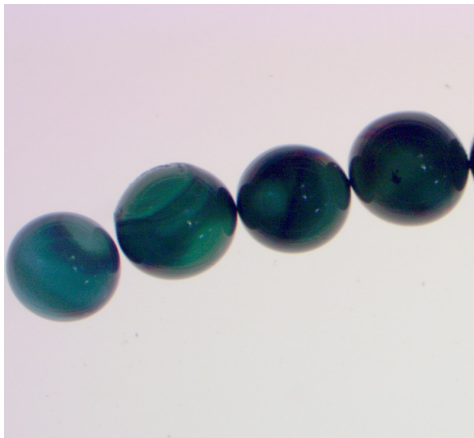


$(\text{Mg,Fe})\text{SiO}_3$
(Enstatite/Bronzite)



$(\text{Mg,Fe})_2\text{SiO}_4$
(Forsterite/Olivine)

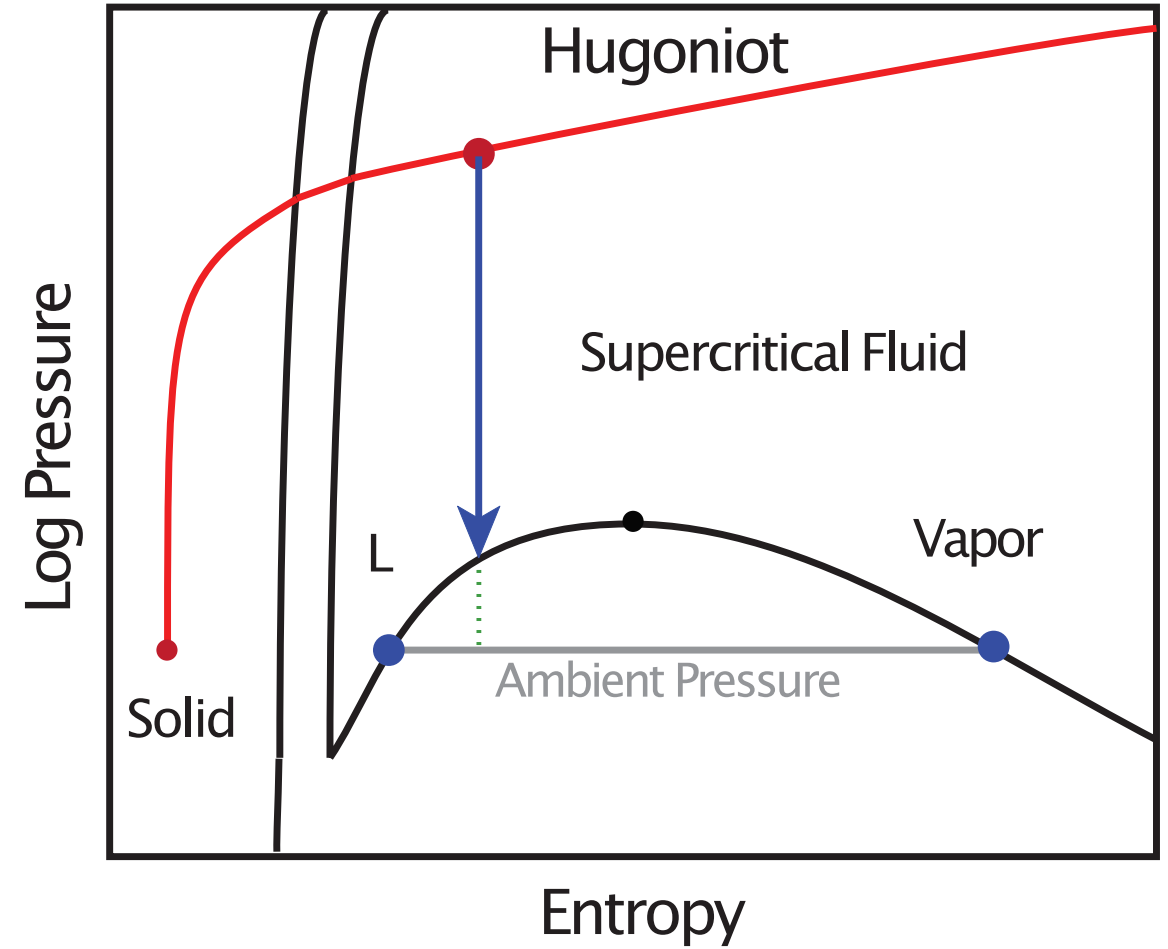
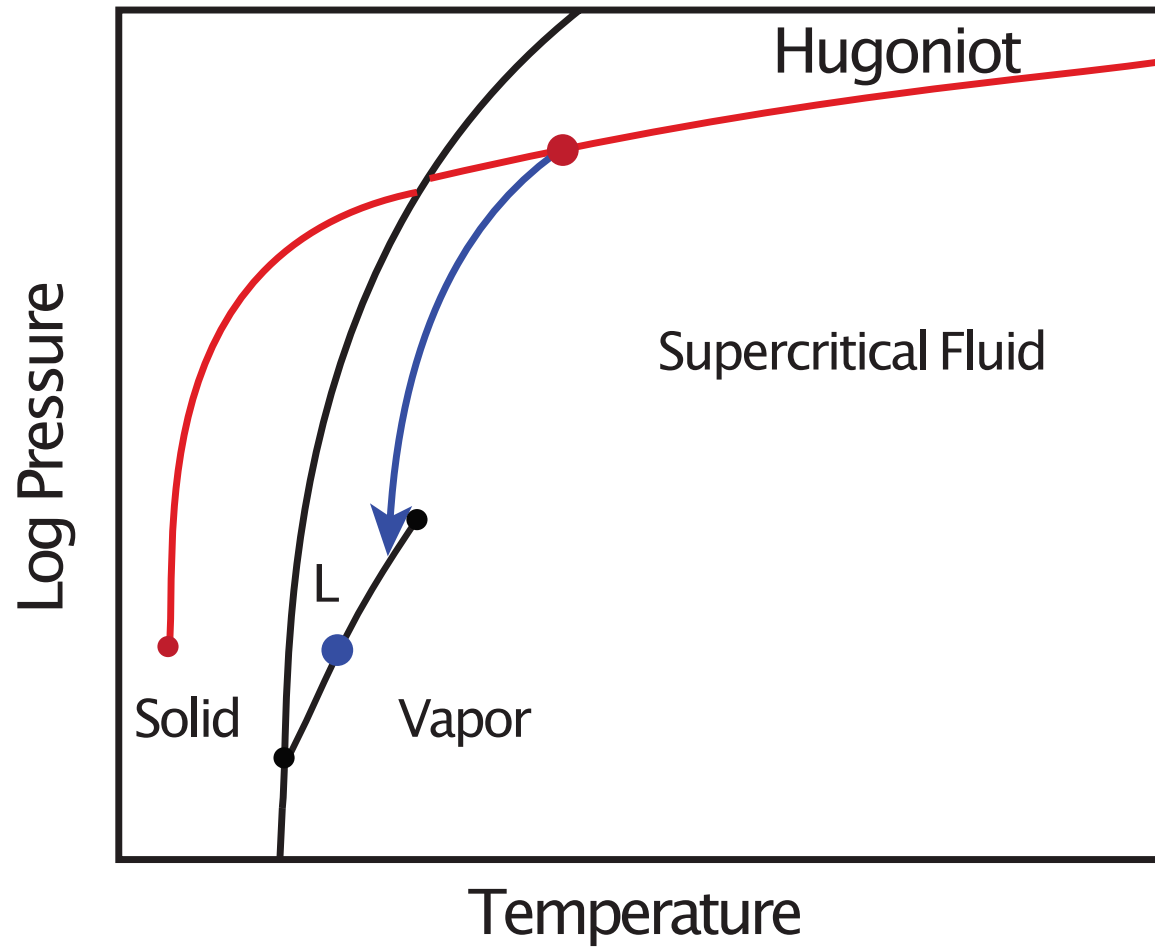
Fratanduono et al., *PRB*, 2018
Root et al., *GRL*, 2018
Chidester et al. *GRL*, 2021



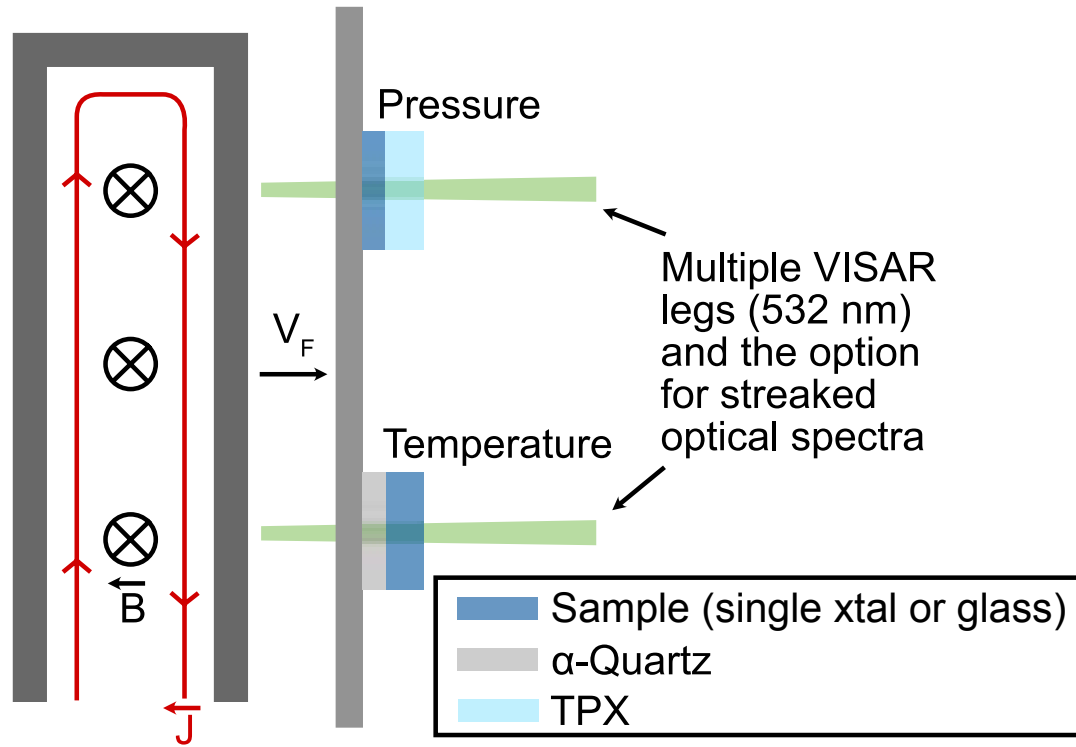
Pyrolite glass



Shock Physics Experiments: Giant Impact P-T Paths



Shock compression experiments at Z Machine

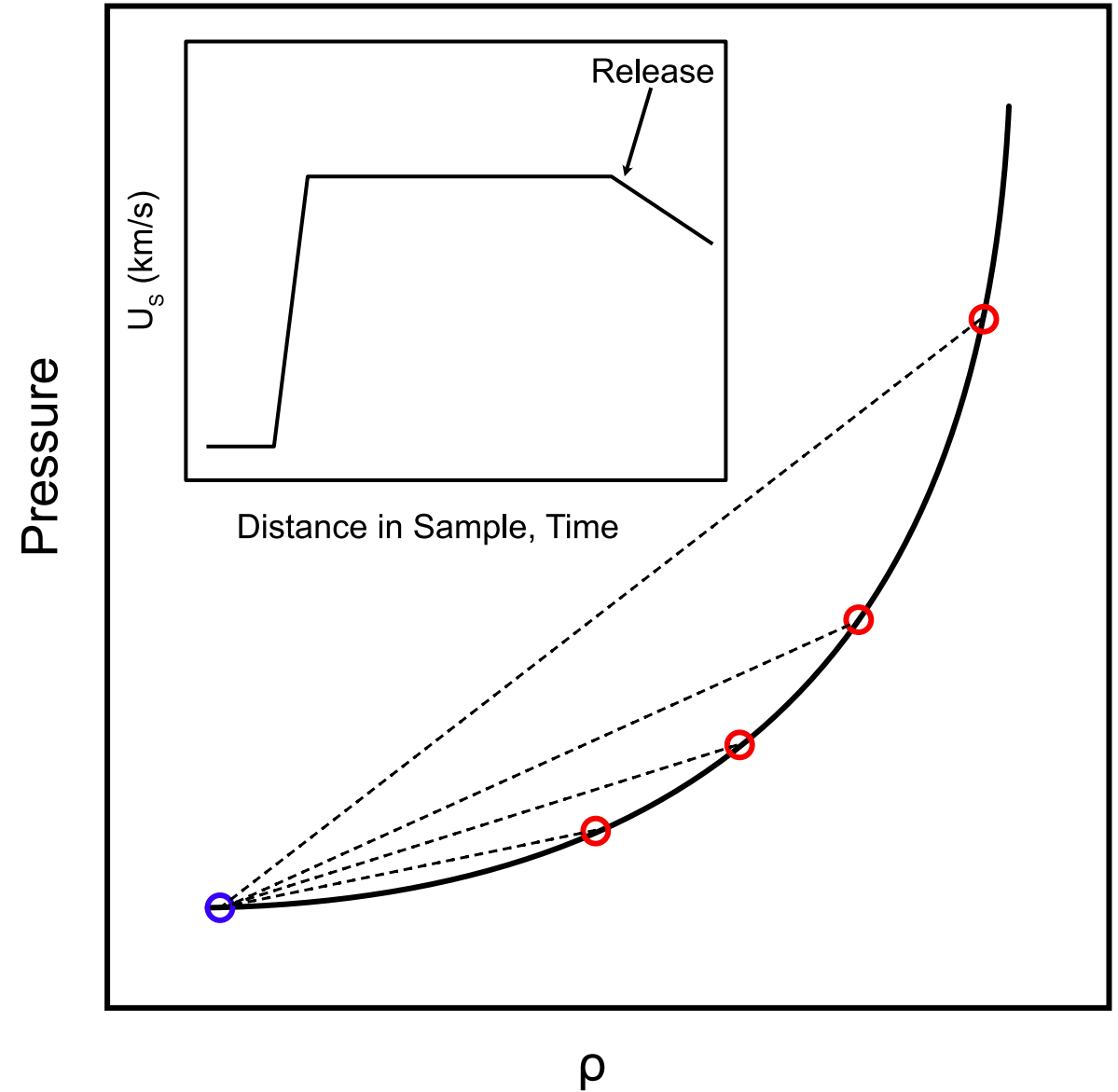


Rankine-Hugoniot jump condition

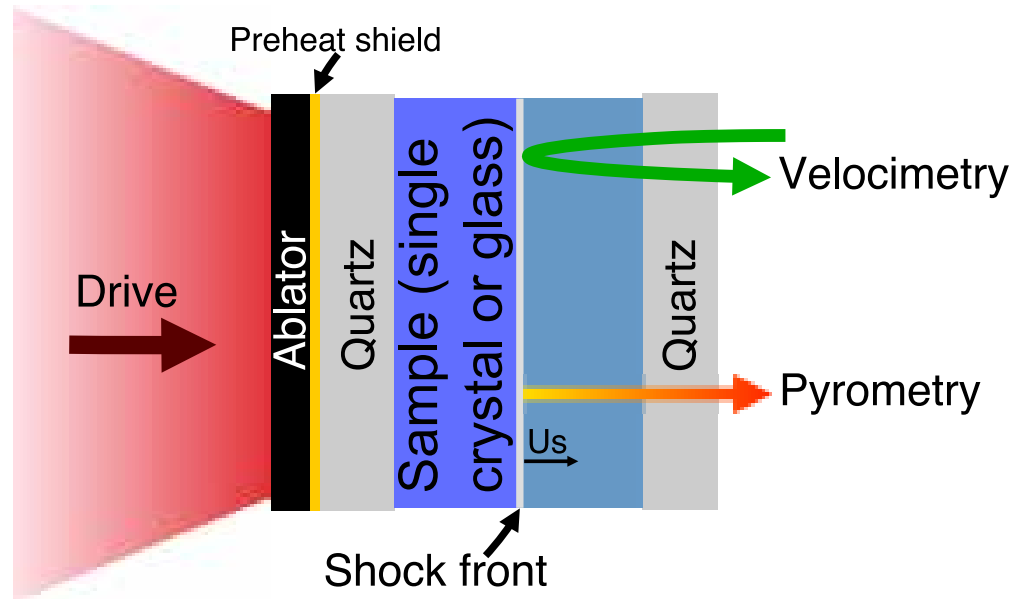
$$E - E_0 = \frac{1}{2}(P + P_0)(V_0 - V)$$

$$P - P_0 = \rho_0(U_s - u_{p,0})(u_p - u_{p,0})$$

$$\rho_0/\rho = 1 - (u_p - u_{p,0})/(U_s - u_{p,0})$$



Shock compression experiments at OMEGA EP

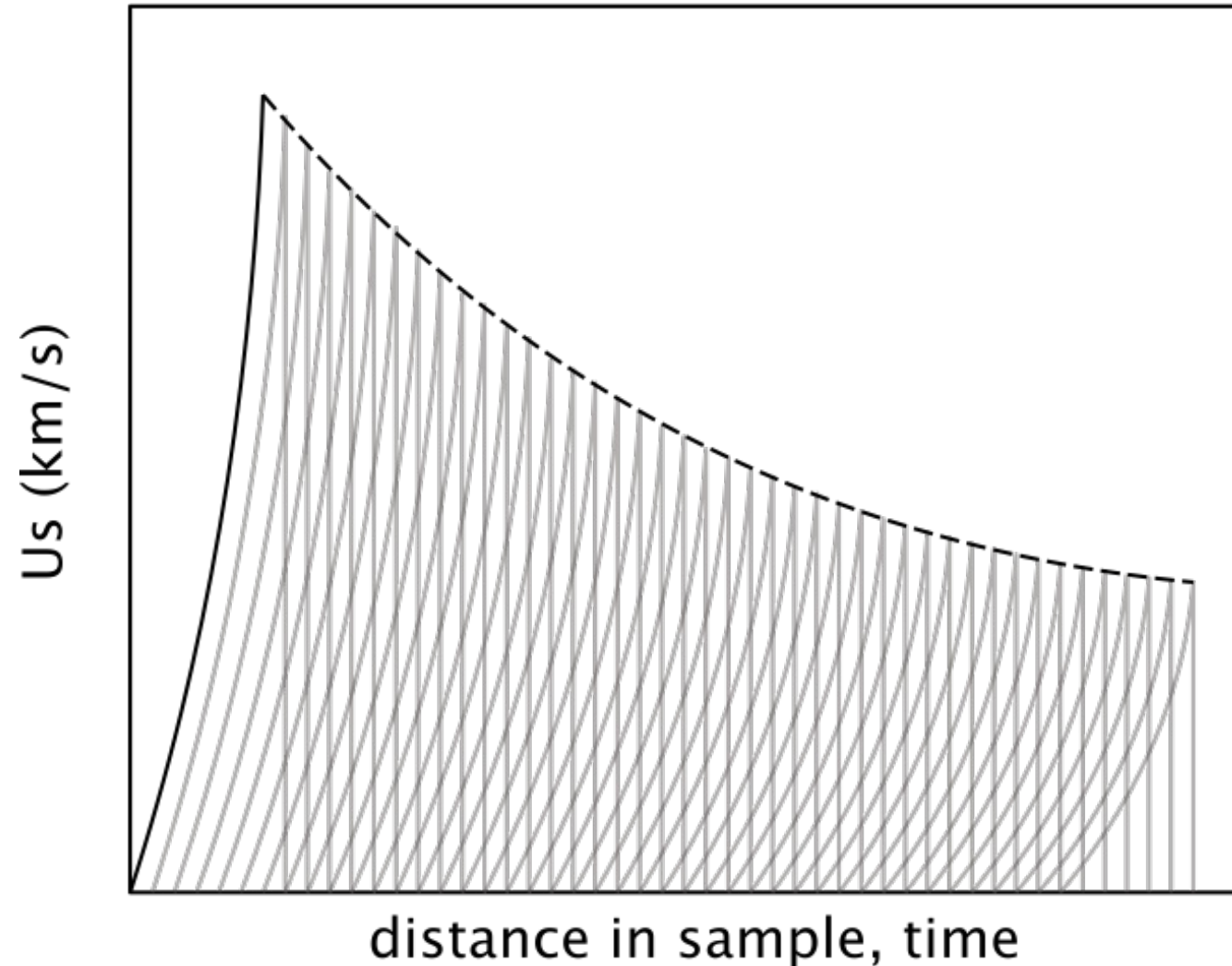


Rankine-Hugoniot jump condition

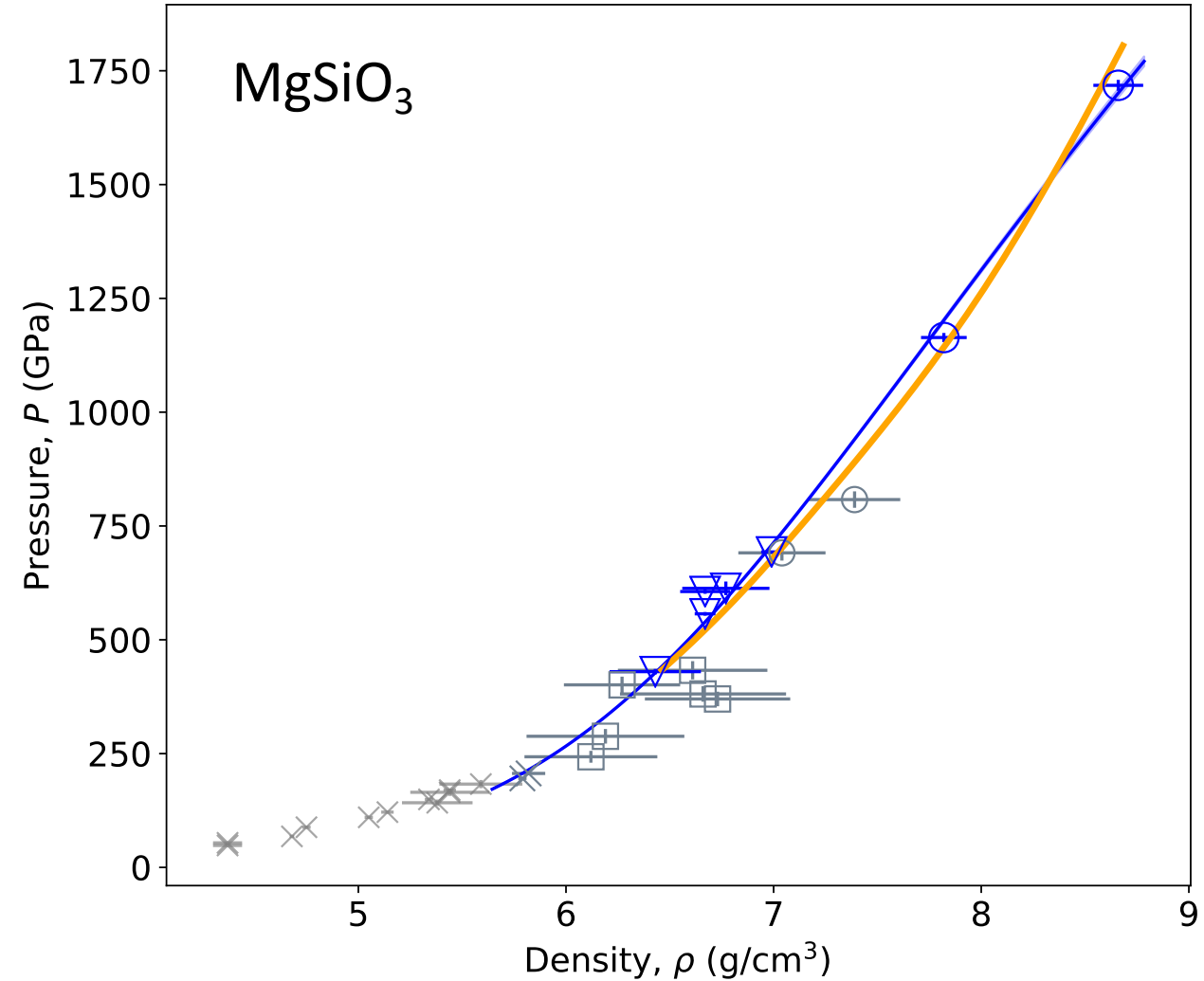
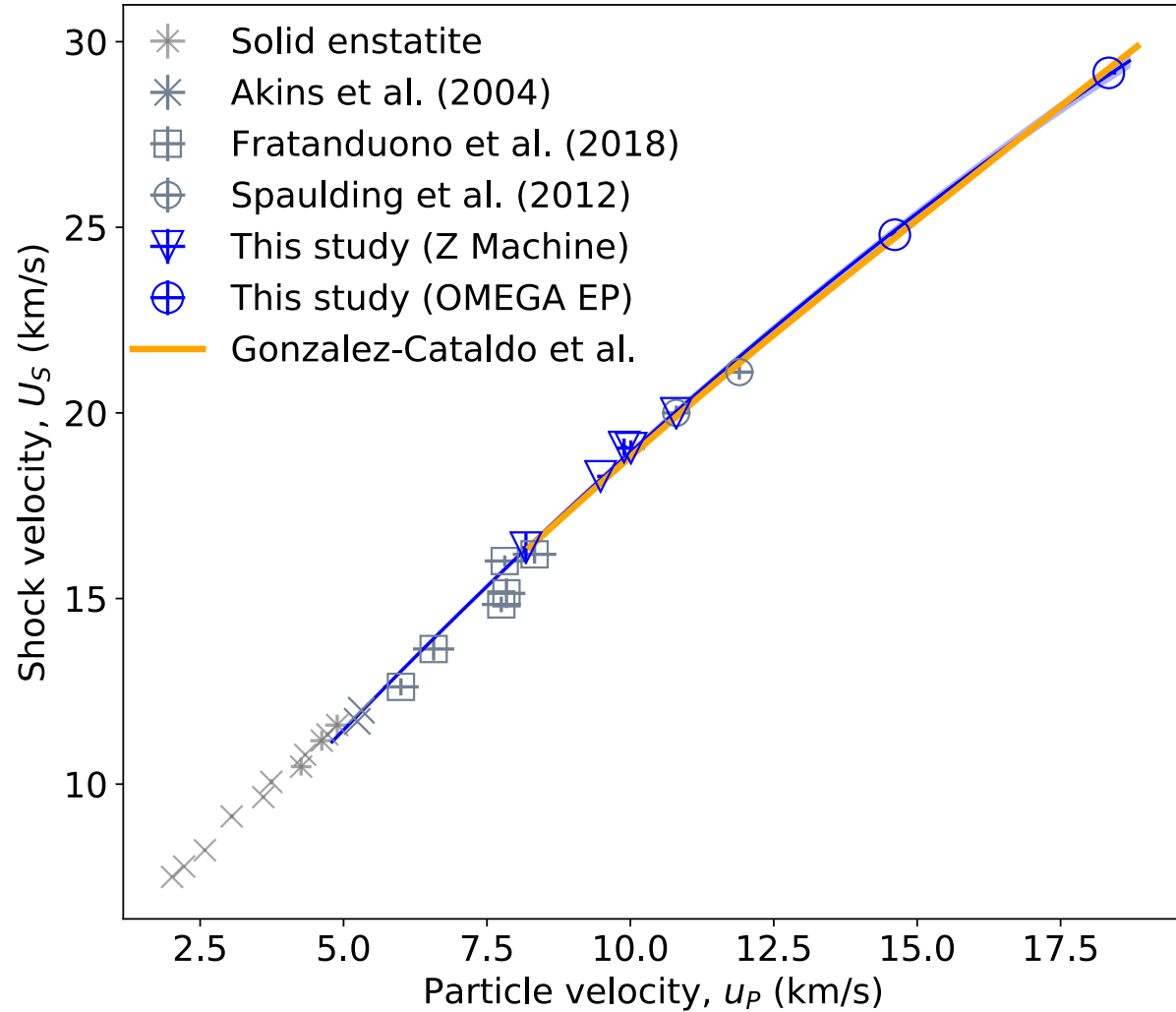
$$E - E_0 = \frac{1}{2} (P + P_0)(V_0 - V)$$

$$P - P_0 = \rho_0 (U_s - u_{p,0})(u_p - u_{p,0})$$

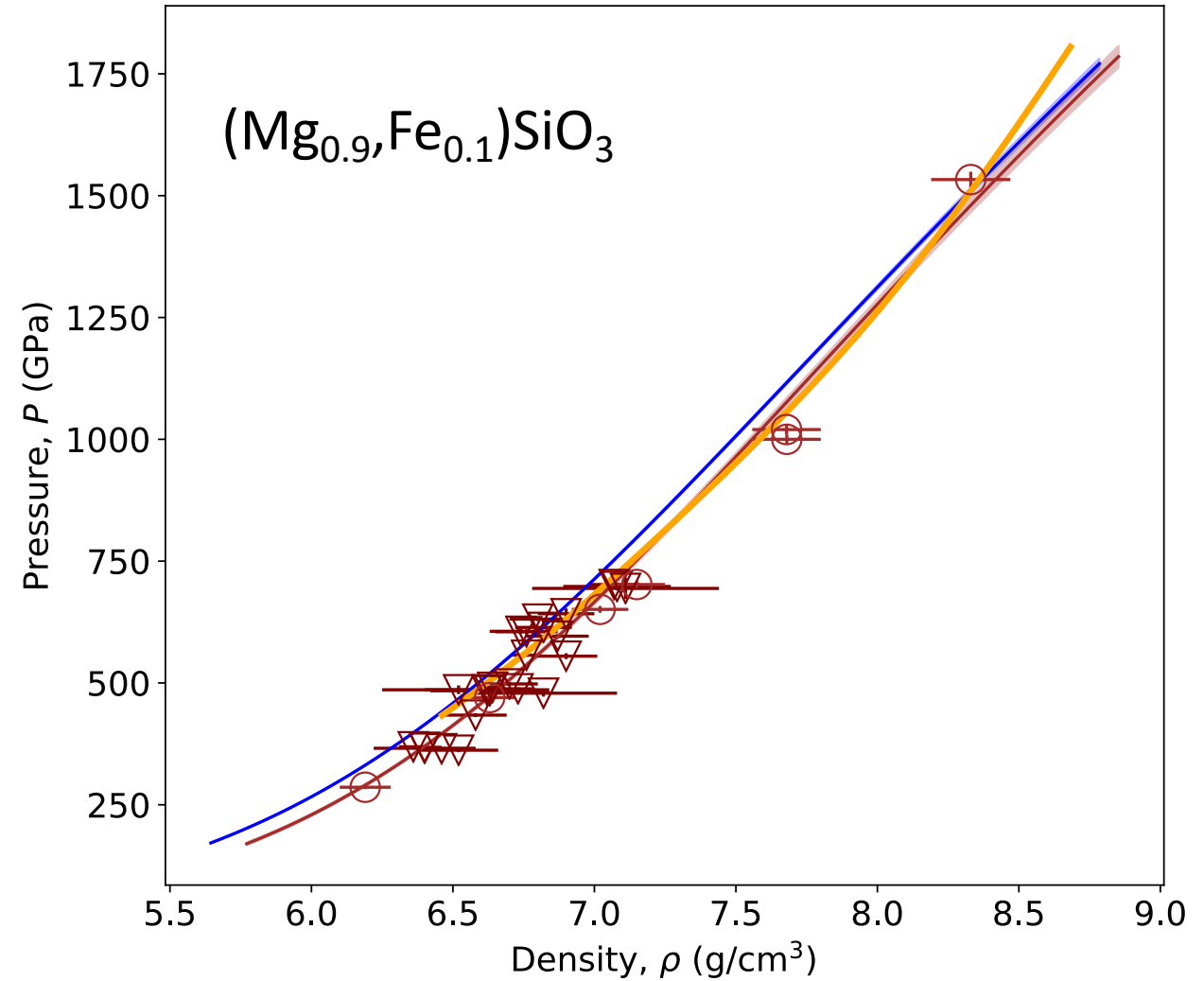
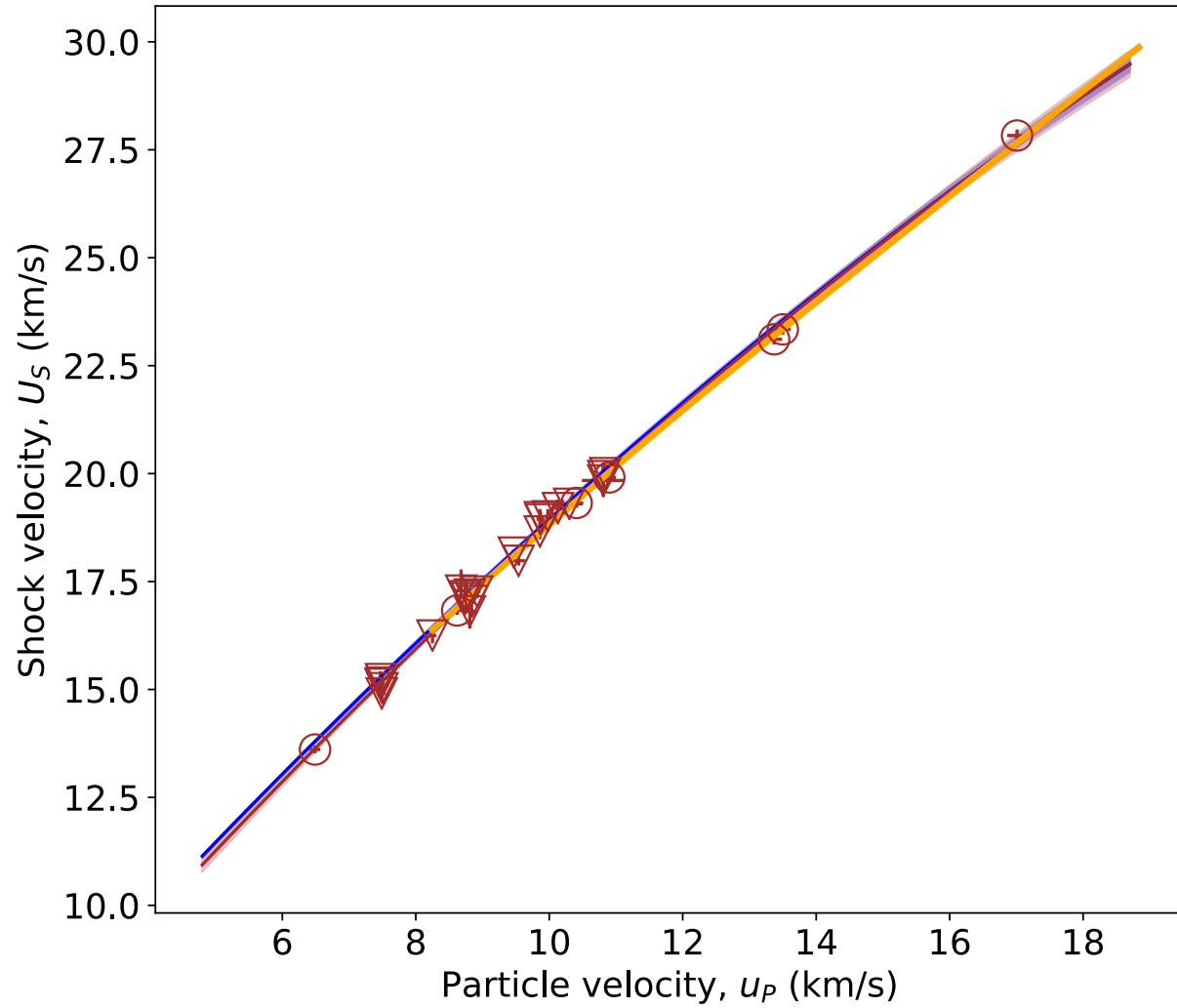
$$\rho_0 / \rho = 1 - (u_p - u_{p,0}) / (U_s - u_{p,0})$$



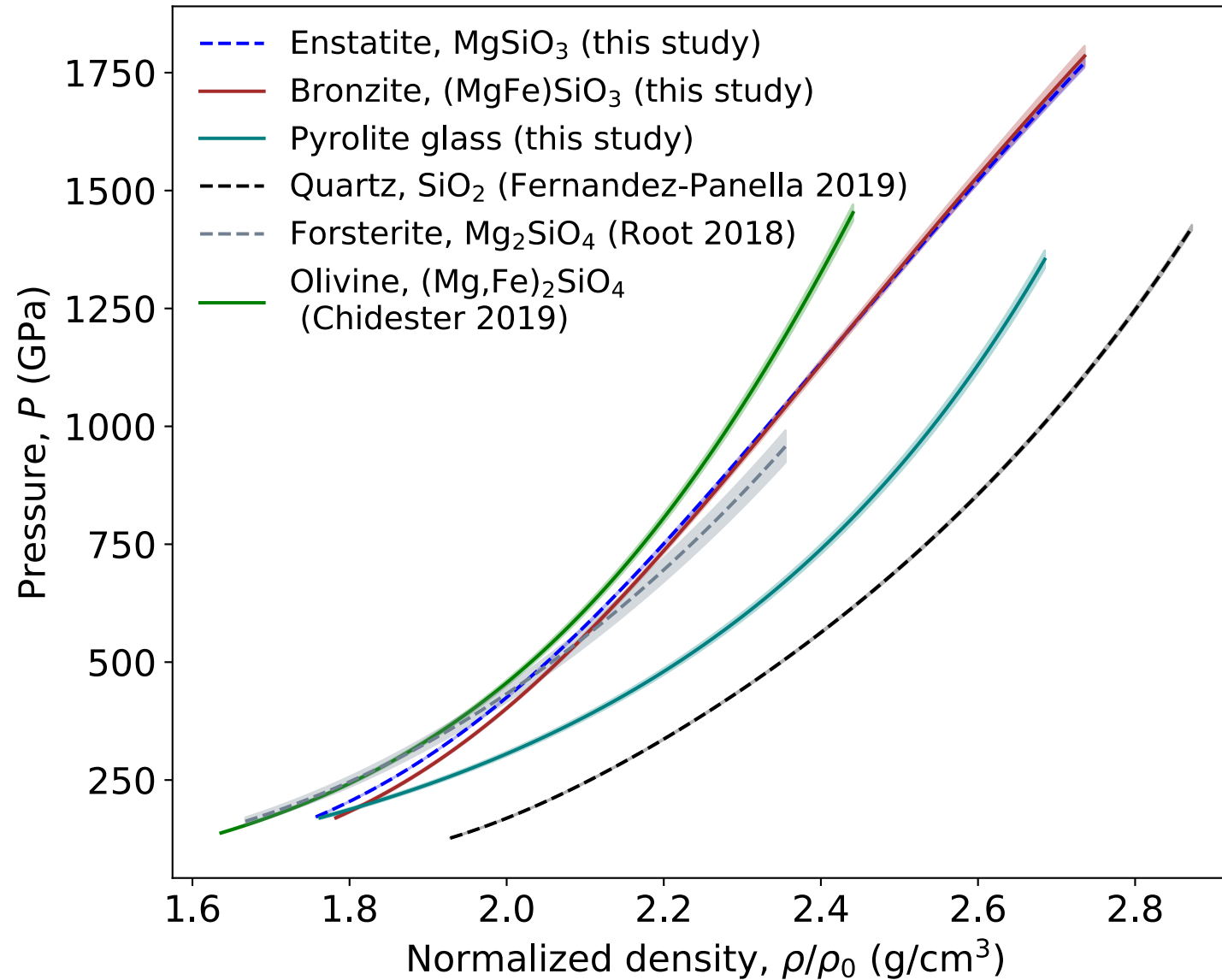
Enstatite compressibility results



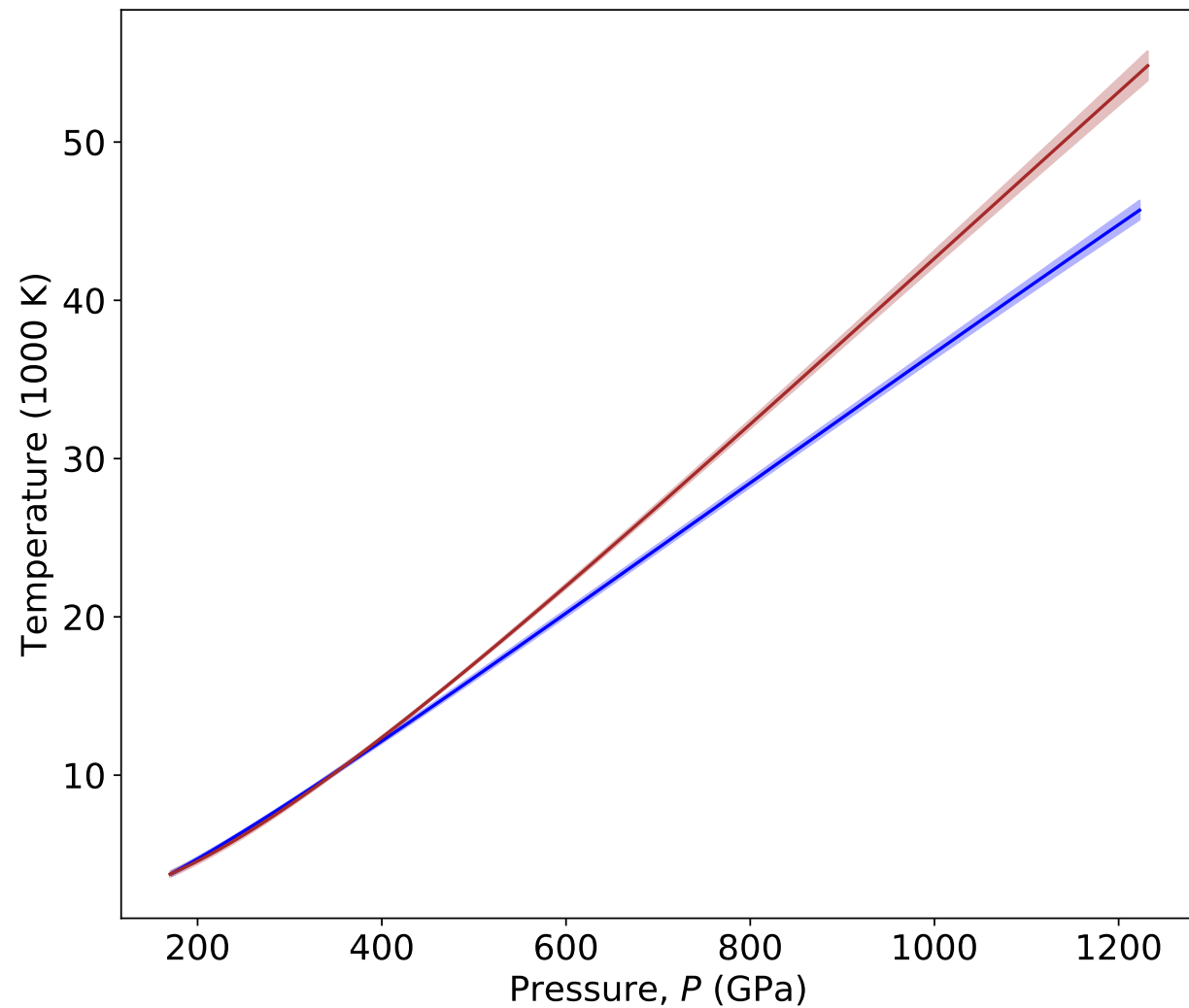
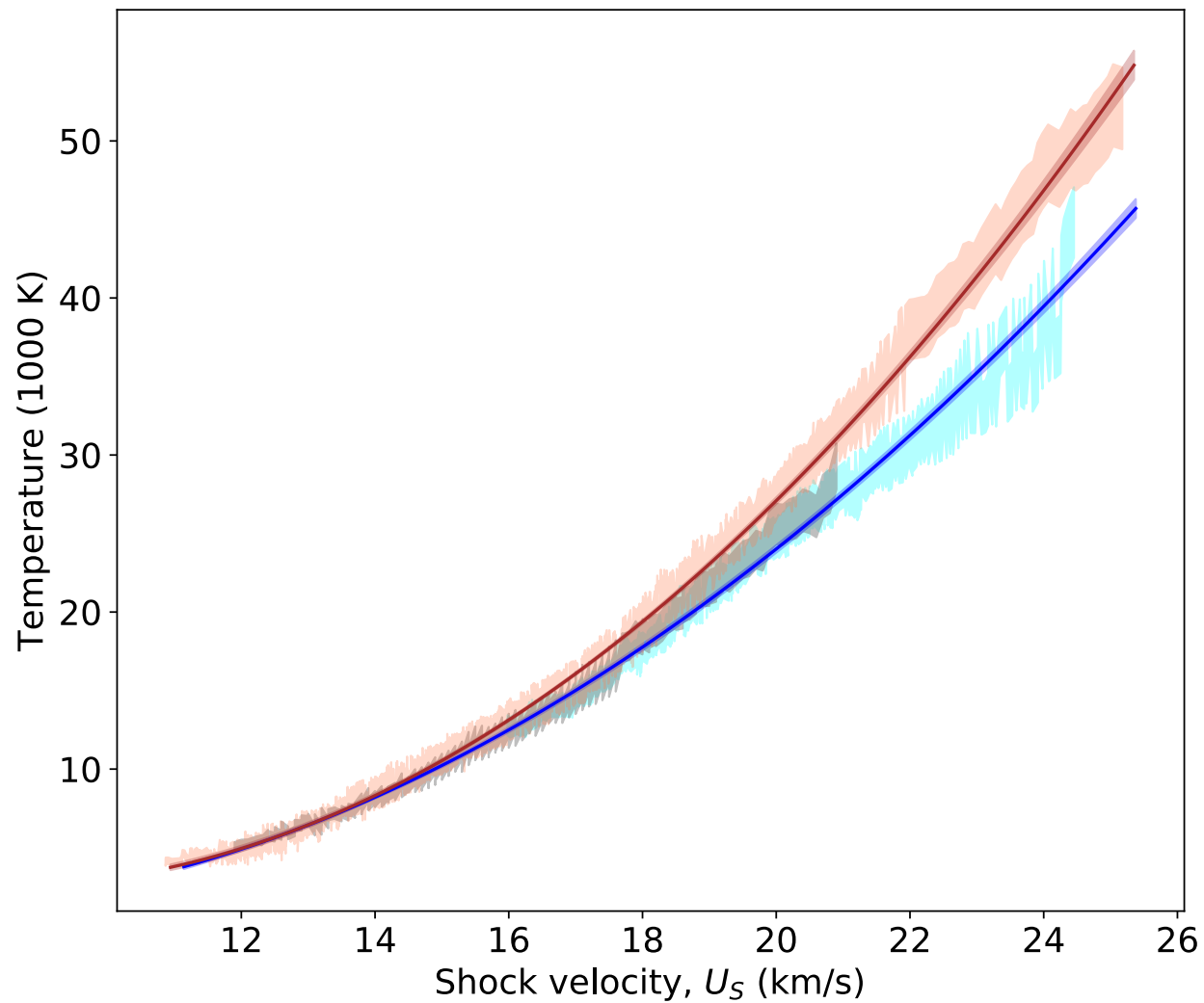
Bronzite compressibility results



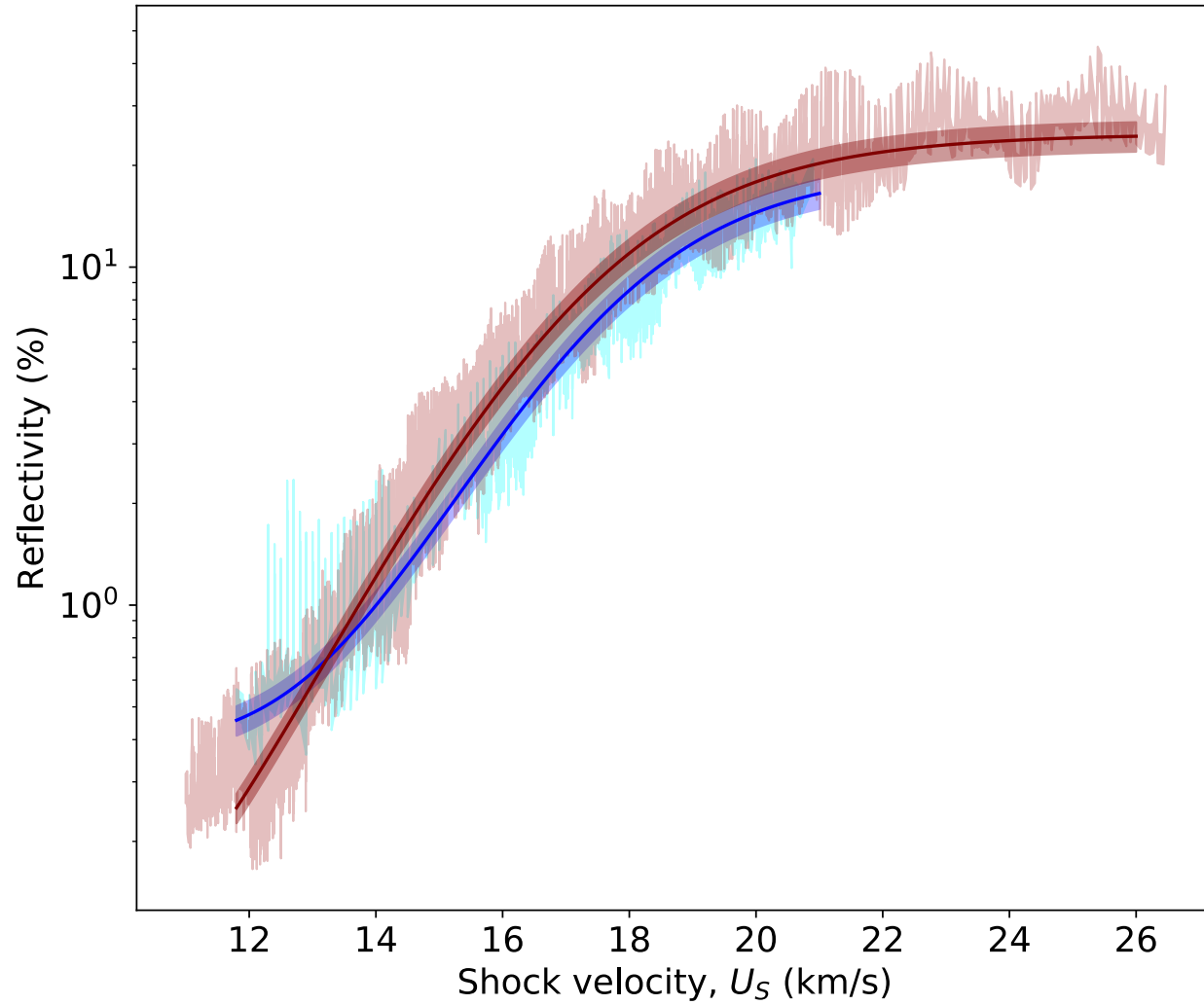
Comparison to other silicates



Temperature



Reflectivity



Silicates become conductive at high temperatures!

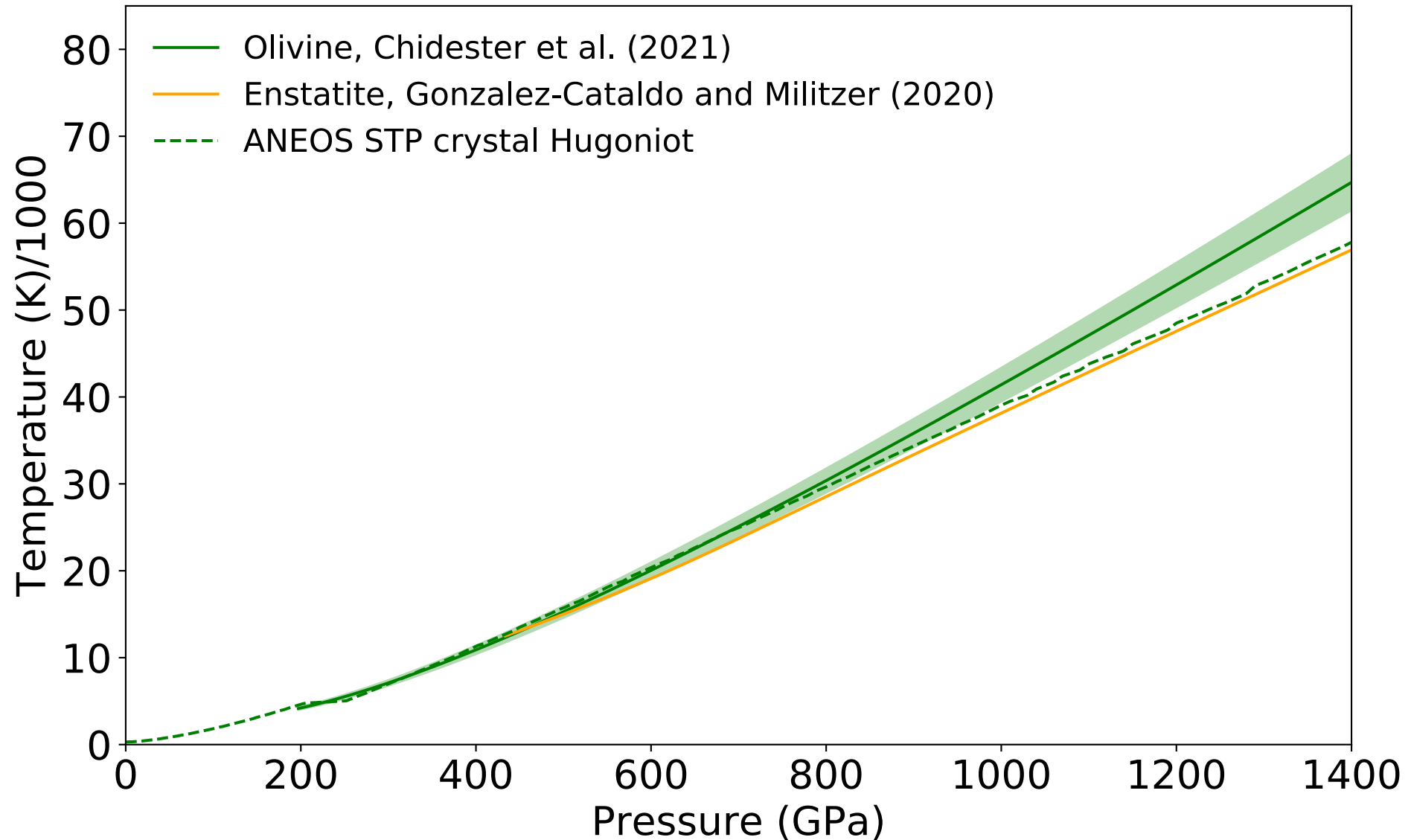
Fe atoms likely increase the effect.

Conclusions

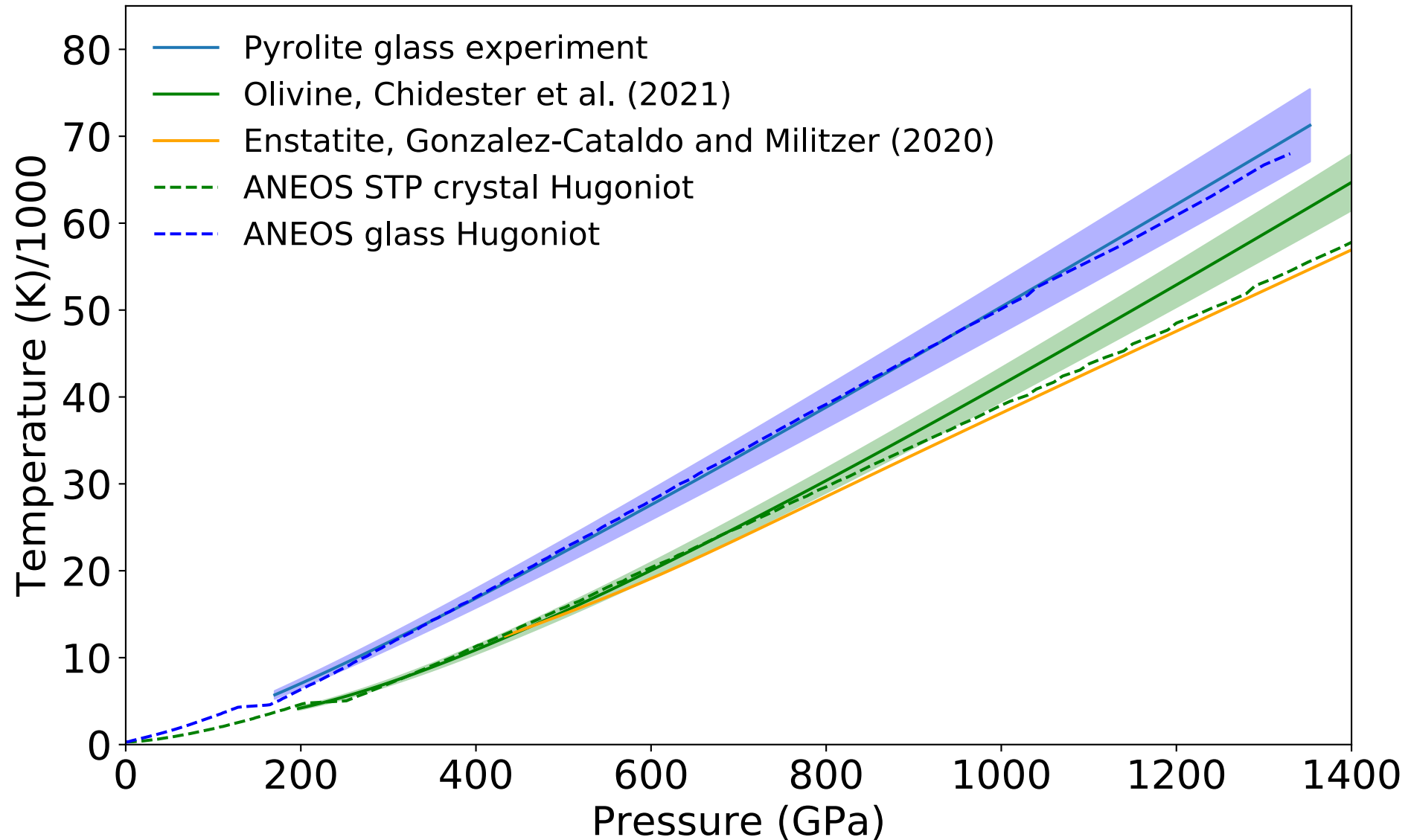
- Fe-bearing silicates seem to shock to higher temperatures than their Fe-free endmembers
- This is related to the change in reflectivity, likely a different electronic contribution to the heat capacity as the material becomes more conductive

Might stop here, but if I have time I'll
continue on to other slides

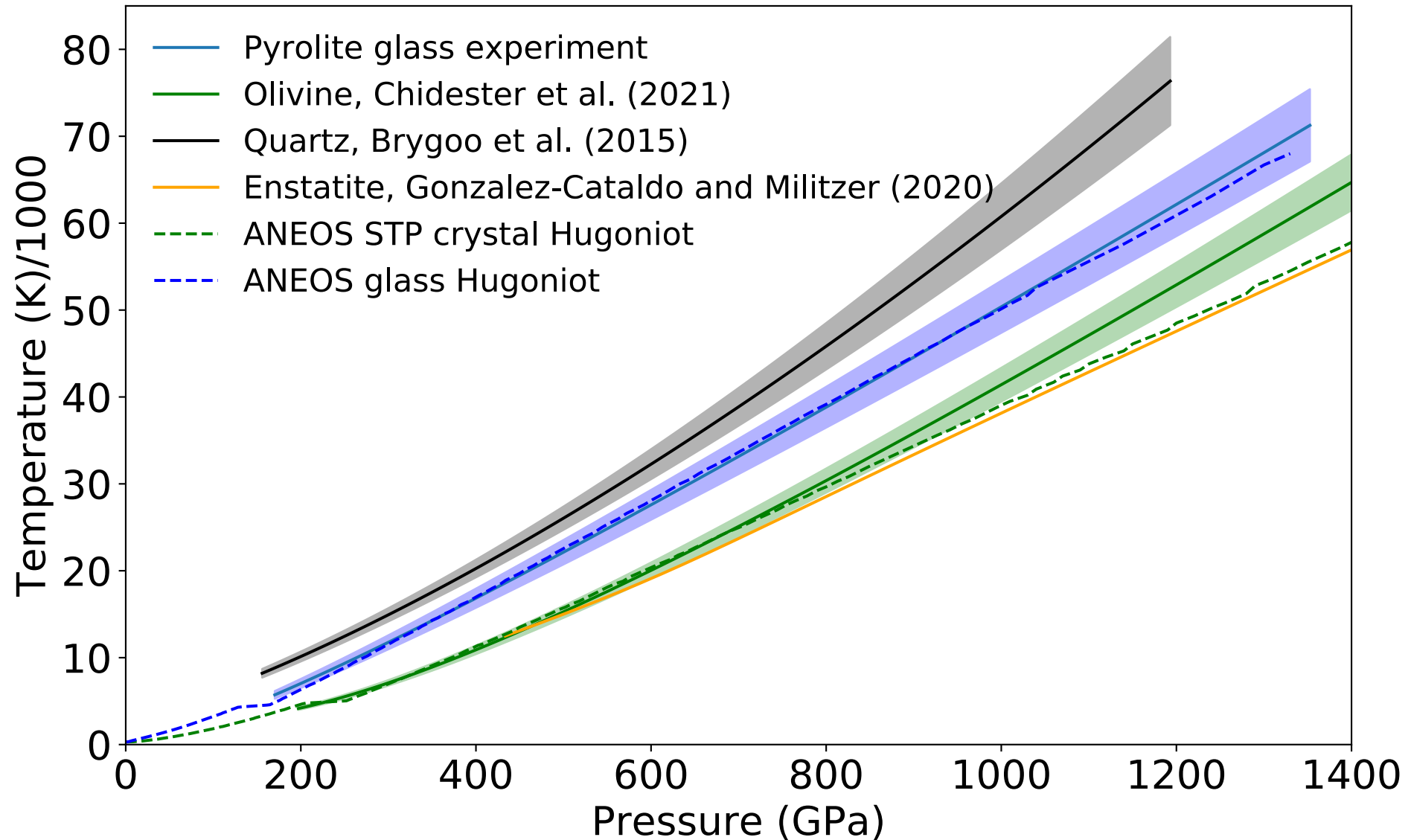
Analytical EOS constrained by experimental results



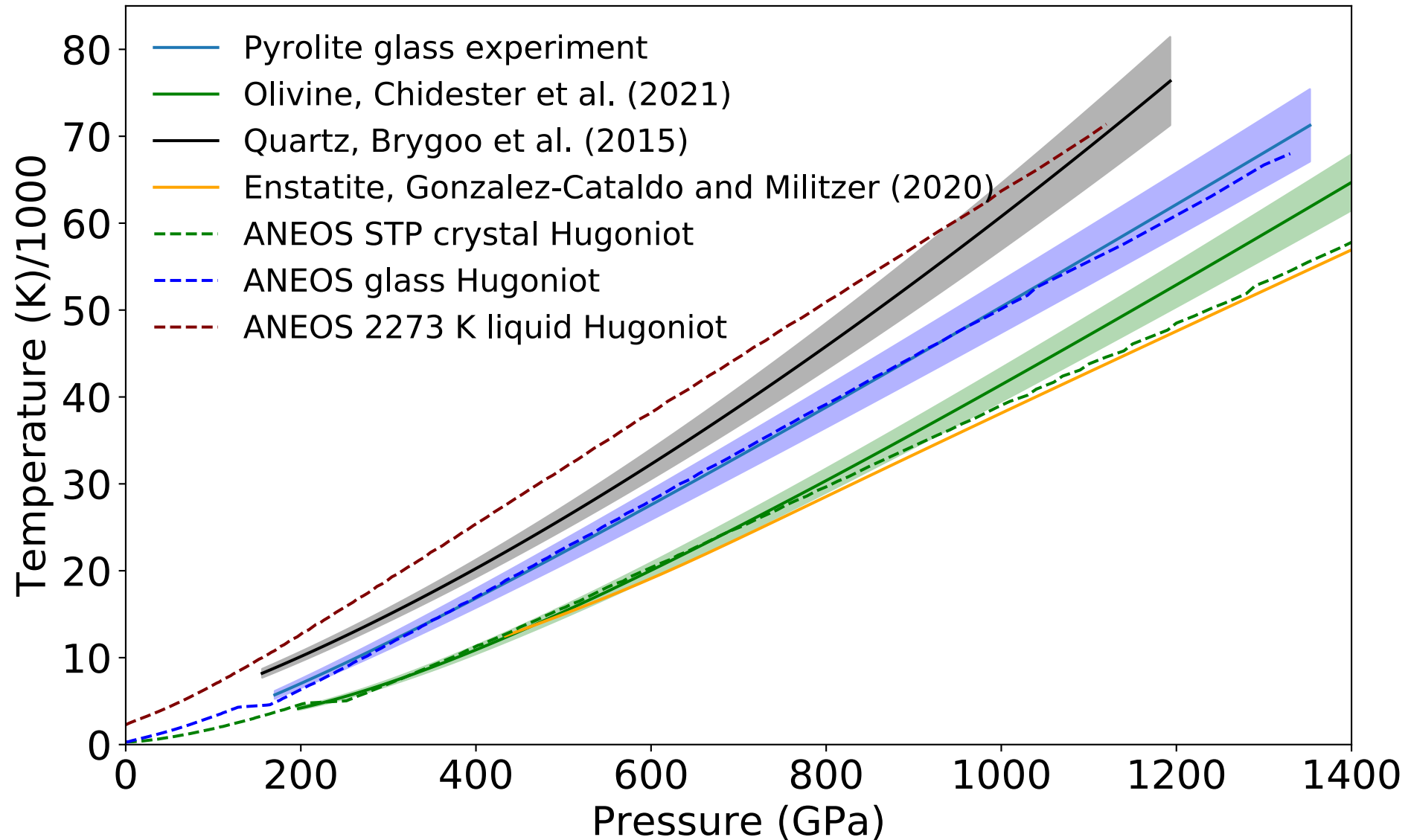
Analytical EOS constrained by experimental results



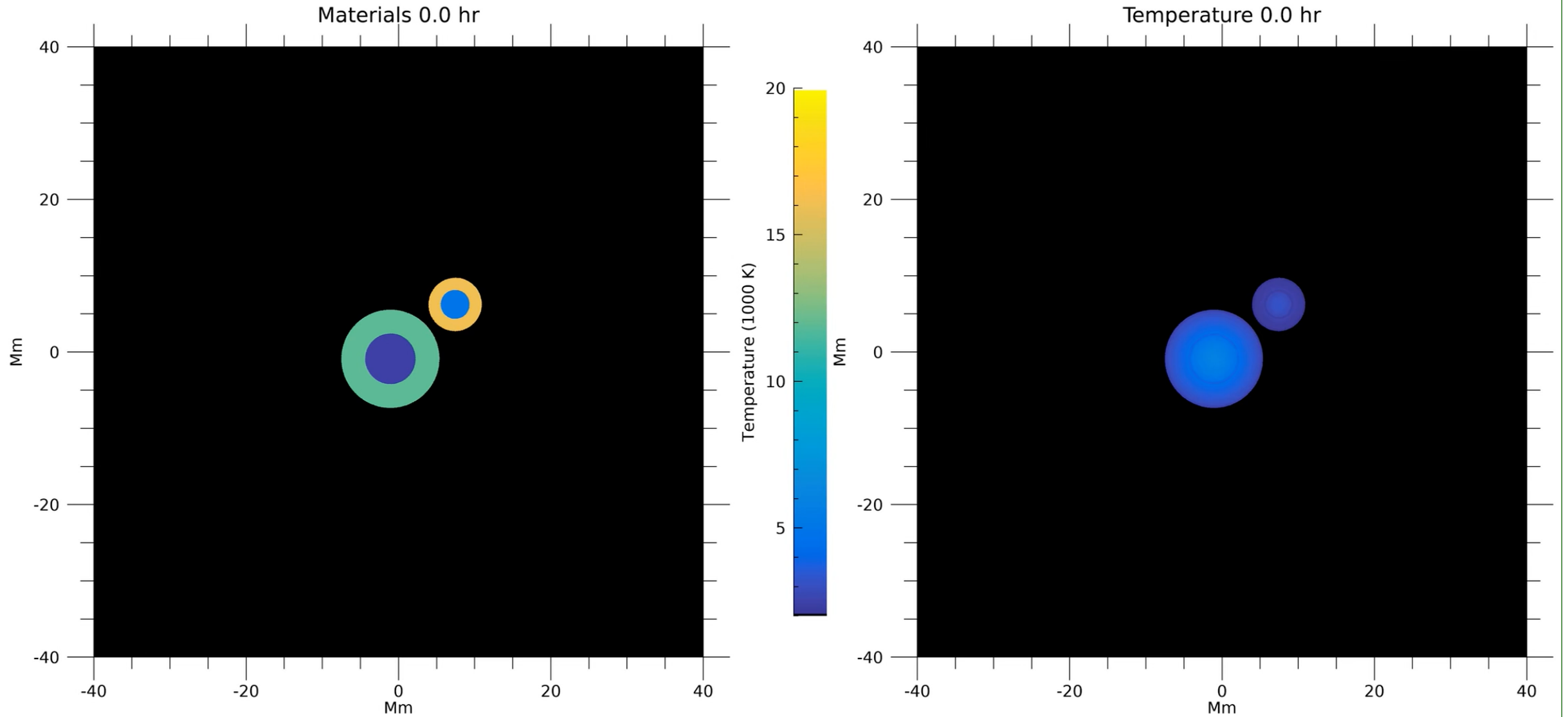
Analytical EOS constrained by experimental results



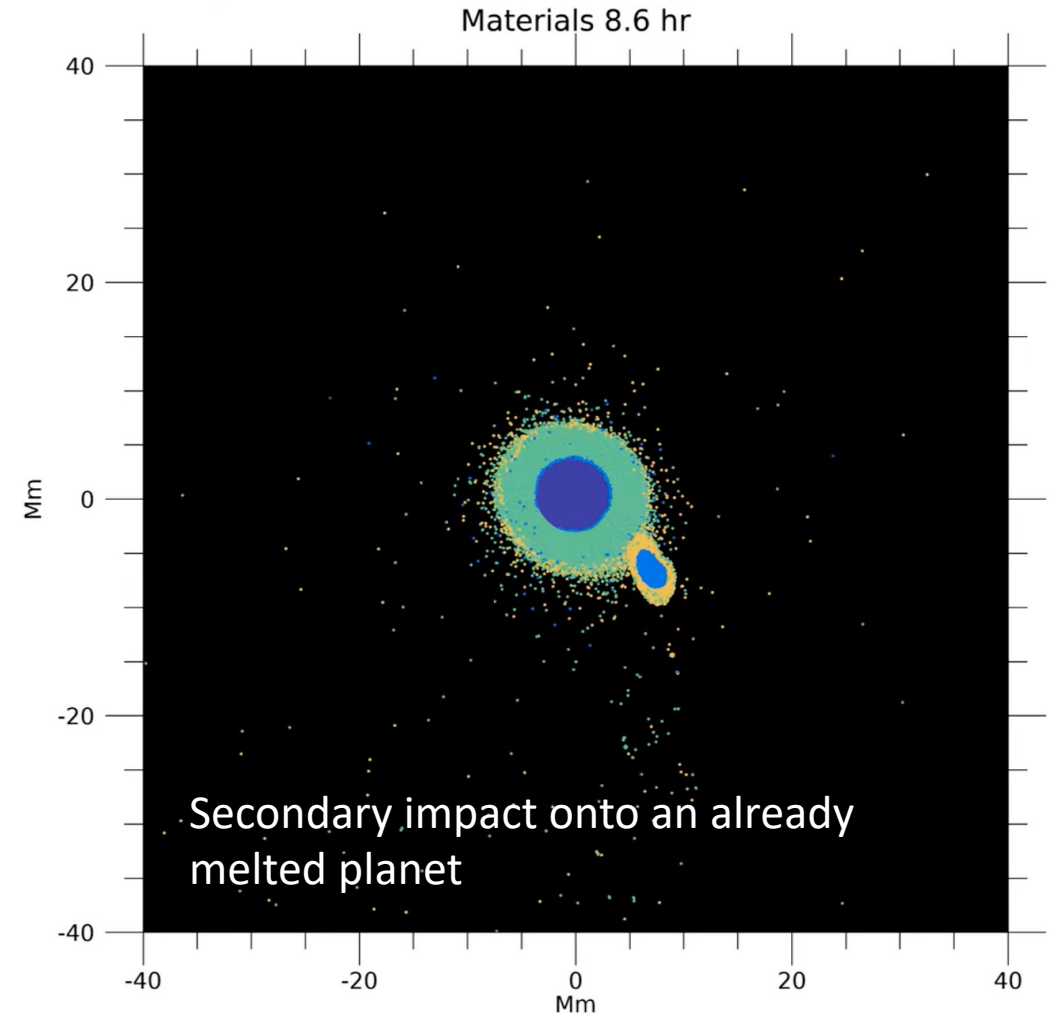
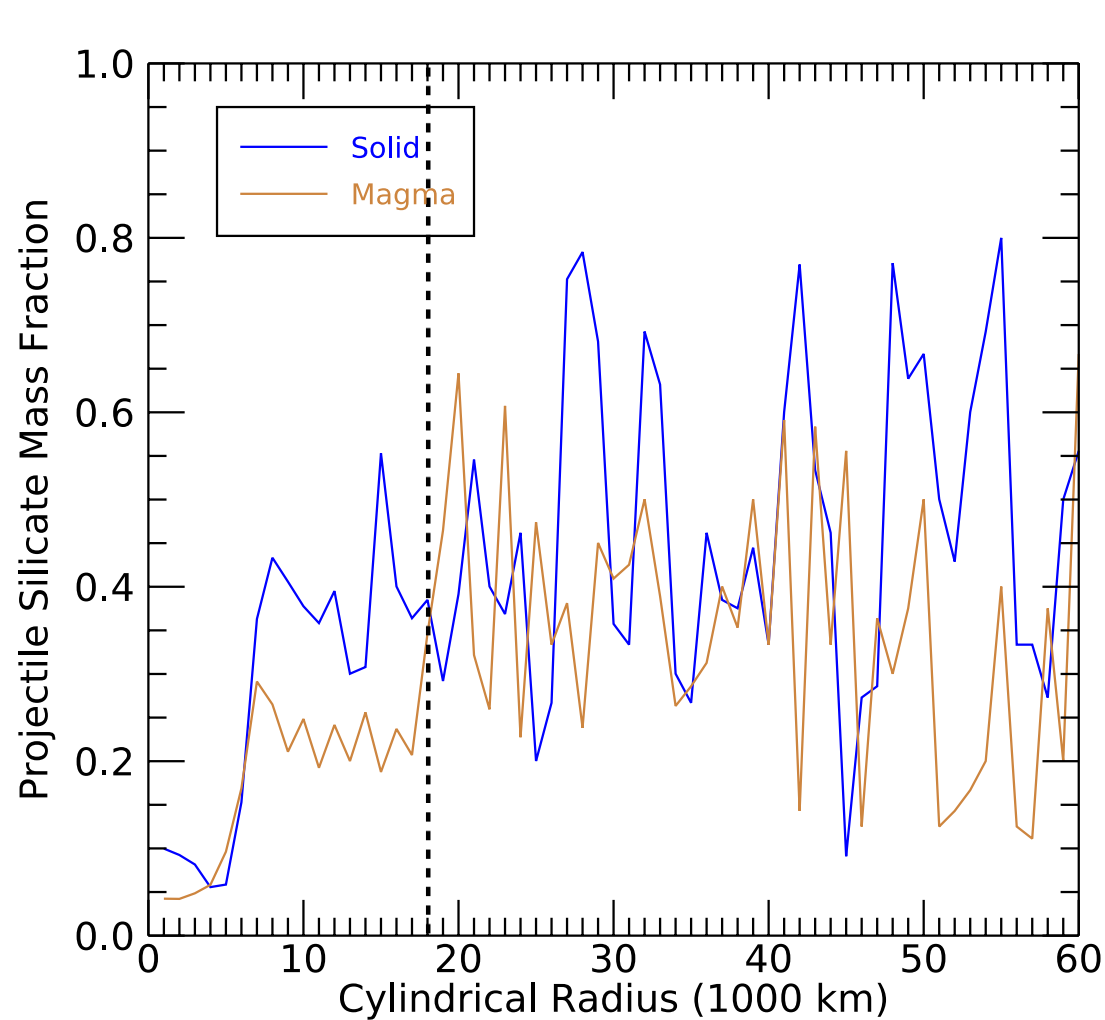
Analytical EOS constrained by experimental results



Magma ocean proto-Earth

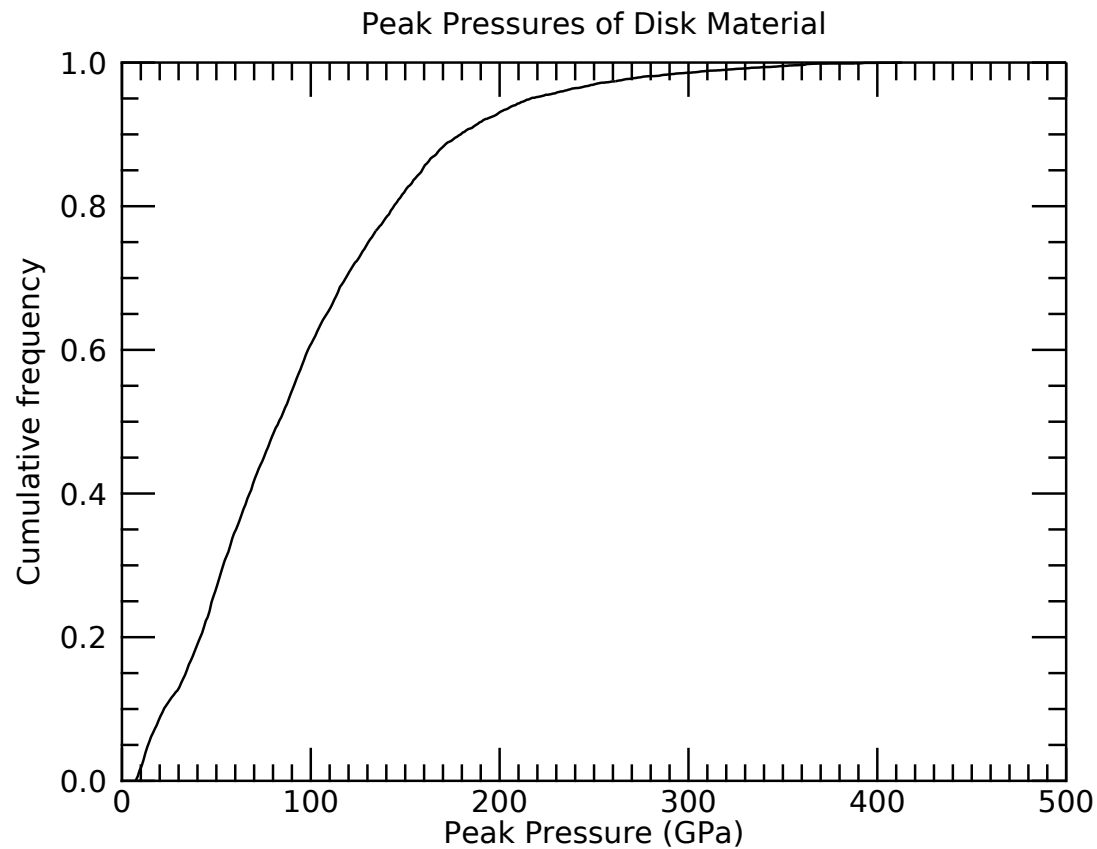


Result 1: Giant impacts onto a magma ocean are not different than giant impacts onto a solid planet

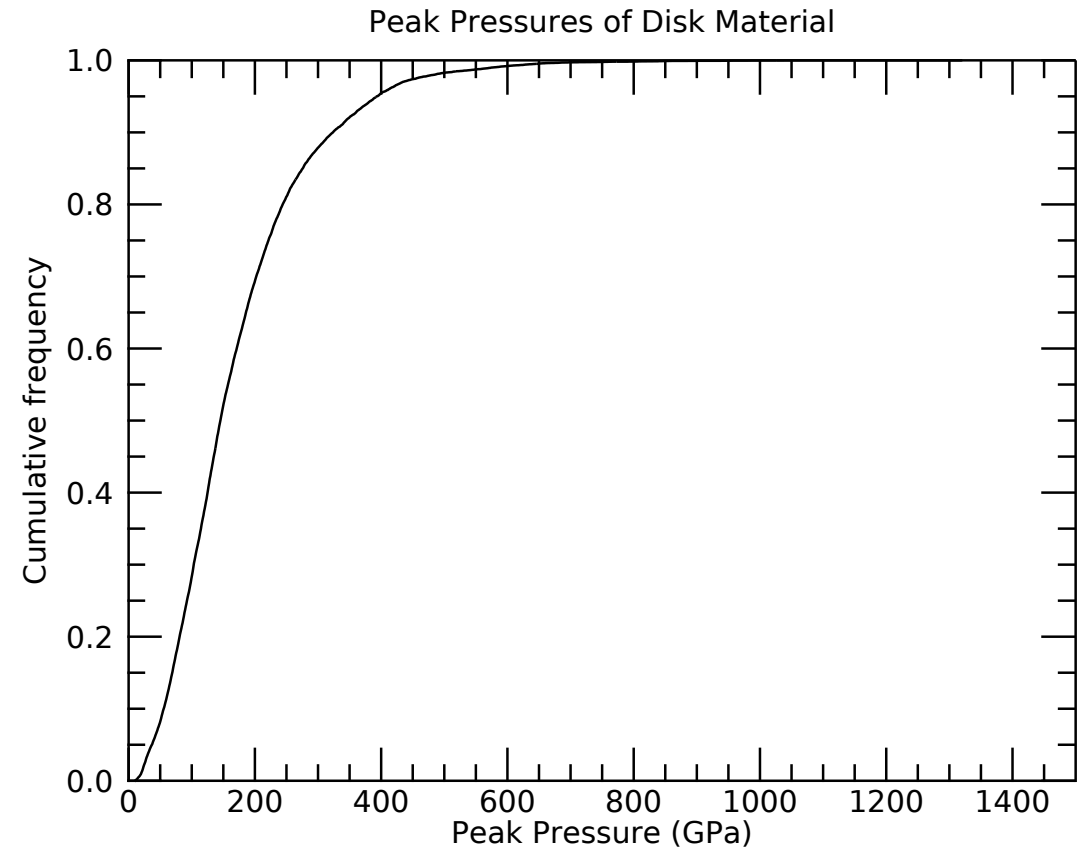


Result 2: The disk is heavily shocked

Mars-Mass Giant Impact



Two Half-Earths Colliding



Conclusions

- Experimentally-constrained EOS are needed for accurate planet formation models
- Shocked pyrolite glass is very similar to known crystalline silicates.
- The canonical giant impact model provides the same results if the planet is entirely solid or has a magma ocean
- These experiments provide information on the pressures and temperatures reached in giant impacts, which is important for chemistry