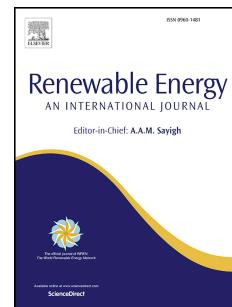


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# Wave climate and energy resources in the Mariana Islands from a 42-year high-resolution hindcast

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13 **Abstract:**

14 A high-resolution wave hindcast from 1979 to 2020 was developed for Guam and  
15 the Commonwealth of the Northern Mariana Islands, which are influenced by mesoscale  
16 and synoptic weather systems from the western Pacific Rim to the Central Pacific. The  
17 spectral wave modeling utilized structured global and unstructured regional grids with  
18 improved resolution in nearshore waters, where wave resources development is most  
19 feasible. The hindcast demonstrates its capability in capturing multi-modal seas through  
20 reproduction of bulk and partitioned wave parameters derived from altimetry and buoy  
21 measurements. The local wave climate is influenced by the year-round trade winds and  
22 tropical cyclones as well as seasonal monsoons, frontal systems, and mid-latitude high-  
23 and low-pressure systems. The strong intra-annual variation of the significant wave  
24 height and power flux reflects transitions between local and distant weather systems,

25 which account for the diverse wave conditions comprised of short-period wind waves,  
26 intermediate-period seas, and long-period swells around the islands. Among the Pacific  
27 island regions, spectral partitioning is particularly important for describing the dominant  
28 wave components by period and direction with pertinent information for energy  
29 resources characterization and applications related to coastal engineering, ecosystem,  
30 and management in the Mariana Islands.

31 **Keywords:** Wave hindcast, Wave climate, Wave energy resources, Partitioned wave  
32 spectrum, Mariana Islands

33 **1. Introduction**

34 Most remote Pacific Island communities depend heavily on imported fossil fuels  
35 as an energy source, which is vulnerable to natural disasters and market volatility. A  
36 diversified energy portfolio including ocean wave resources is favorable for promoting  
37 energy independence in these isolated locations with proximity to the ocean. The  
38 Mariana Islands, located in the northwestern Pacific Ocean, are 2000 km from Japan,  
39 the Philippines, and New Guinea (Figure 1a). The island group includes Guam, an  
40 unincorporated territory of the US, as well as the most populous islands of the  
41 Commonwealth of the Northern Mariana Islands (CNMI) - Rota, Tinian, and Saipan.  
42 Although the islands are relatively small, their Exclusive Economic Zone (EEZ) covers a  
43 large region (970,772 km<sup>2</sup>) with vast marine resources. On the inhabited islands, ocean  
44 waves play a significant role in coastal hazard management, coral reef ecology, and  
45 maritime activity all year round. To capture real-time wave conditions, a Waverider buoy  
46 was installed at 200 m depth off the east shore of Guam in 2003. Two additional wave  
47 buoys were deployed north of Guam at 500 m depth and north of Saipan at 487 m  
48 depth in 2012 (Figure 1b). The buoy records provide vital information for coastal  
49 activities but are insufficient to describe the complex wave climate due to sparse

50 coverage, island blocking, limited years of record, and irregular downtimes. Numerical  
51 modeling, when appropriately validated, serves to extend the observation record by  
52 creating a synthetic time history, which can support wave energy resource quantification  
53 and coastal management activities.

54 Third-generation spectral wave models, such as WAVEWATCH III (Tolman and  
55 Chalikov, 1996; Tolman 2006, 2008, 2009) and WAM (WAMDIG 1988, Komen et al.  
56 1994), have been utilized with coupled model systems including atmosphere, land,  
57 ocean, and sea ice components to produce wave hindcasts as part of global reanalysis  
58 products. The ERA-interim reanalysis produced by the European Centre for Medium-  
59 Range Weather Forecasts (ECMWF) includes wave data products from WAM for 1979  
60 to 2019 at 1 arc-deg grid resolution (Dee et al., 2011). It was superseded by the recently  
61 released ERA5 reanalysis, which provides a global wave hindcast from 1950 to the  
62 present with 0.36 arc-degree resolution (Hersbach et al., 2020). The US National  
63 Centers for Environmental Prediction (NCEP) produced comparable datasets such as  
64 the Climate Forecast Reanalysis (CFSR) at 0.5 arc-deg resolution from 1979 to 2009  
65 (Sara et al., 2010) and CFS version 2 (CFSv2) at 0.25 arc-deg resolution from 2010  
66 onward (Saha et al., 2014). This facilitates development of global and regional wave  
67 hindcast at 0.5 arc-deg to 4 arc-min resolution using WAVEWATCH III (Chawla et al.,  
68 2012). The global hindcast datasets are suitable for wave climate and resources  
69 assessment at the basin level (Stopa and Cheung, 2014a, b), but do not have sufficient  
70 spatial resolution and provisions for shoaling and refraction in coastal waters around  
71 many of the Pacific islands including the Mariana Islands. In addition, shadowing around  
72 islands is a major local factor influencing available ocean wave resources in a multi-  
73 modal sea state (Stopa et al., 2011).

74 The Centre for Australian Weather and Climate Research (CAWCR) developed a  
75 1979-2009 wave hindcast for the Central and South Pacific with resolution as fine as 4  
76 arc-min in coastal waters (Durrant et al., 2019). The Changing Waves and Coasts of the

77 Pacific project utilized an earlier version of the CAWCR hindcast to assess the wave  
78 climate and seasonal variability for Guam (Durrant et al., 2014). The study focused on  
79 bulk parameters such as significant wave height and peak period derived from the  
80 combined sea state to provide a baseline assessment. Waves generated by local and  
81 distant weather systems, which have surface winds of varying strength and fetch, can  
82 have distinct heights, periods, and directions in the waters surrounding Guam  
83 (Bosscherelle et al., 2014). To better understand the wave climate for Pacific islands such  
84 as Guam and the CNMI, an in-depth investigation of wave components within the mixed  
85 seas and their relationship with the near and far-field weather system is needed. The  
86 latest versions of spectral wave models such as WAVEWATCH III and SWAN, also  
87 known as Simulating WAves Nearshore (Booij et al., 1997; Booij et al., 1999), include a  
88 watershed algorithm (Hanson and Phillips, 2001) to partition the frequency-direction  
89 spectrum by dominant component. Wave parameters from the partitions can be used for  
90 separate analyses of wind waves and swells to better quantify global and regional wave  
91 resources (Arinaga and Cheung, 2012; Portilla-Yandún et al., 2015). The partitioned  
92 wave parameters have the advantage over the bulk parameters in providing detailed  
93 classification of energy resources by direction and frequency ranges (Ahn et al., 2022)  
94 and facilitate their statistical analyses for more accurate projection of extrema (Portilla-  
95 Yandún et al., 2020).

96 Under the US Department of Energy's nationwide Powering the Blue Economy  
97 initiative, the Pacific Northwest National Laboratory (PNNL) in collaboration with the  
98 University of Hawaii developed a high-resolution wave hindcast from 1979 to 2020 for  
99 characterization of energy resources around the US coastal regions and affiliated  
100 islands in the Pacific Ocean (García Medina et al., 2019, 2021; Yang et al., 2020). The  
101 modeled island regions include the EEZs around Hawaii, American Samoa, Baker and  
102 Howland Islands, the CNMI, Guam, Jarvis Island, Johnston Atoll, Palmyra Atoll,  
103 Kingman Reef, and Wake Island in the central and western Pacific as illustrated in

104 Figure 2. The use of unstructured computational grids in SWAN (Zijlema, 2010) is  
105 especially important for small islands by allowing a gradual, local refinement of the  
106 spatial representation for complex bathymetry and jagged shores. This provides a more  
107 accurate description of the wave shoaling, refraction, and dissipation in the nearshore  
108 region, where wave energy resources development is most feasible. The wave  
109 hindcasting follows the Technical Specification published by the International  
110 Electrotechnical Commission Technical Committee 114 (IEC-TC 114) to assure  
111 compatibility for resources assessment among oceanic and continental regions (García  
112 Medina et al., 2019, 2021; Yang et al 2018, 2019). Li et al. (2021) utilized the hindcast  
113 to study the wave climate in Hawaiian waters and complemented the IEC TS with the  
114 use of partitioned spectra to characterize wave resources in multimodal seas.

115 Located in the northwestern Pacific Ocean, the inhabited Mariana Islands are  
116 subject to waves influenced by ocean-surface winds originating from both Asia  
117 continental and Pacific oceanic weather systems with unique characteristics among  
118 other Pacific Island regions covered by the PNNL hindcast. Prior studies examined the  
119 local wave climate (Durrant et al., 2014; and Bosserelle et al., 2014), but additional work  
120 is still needed to explore its spatial-temporal relationship with weather systems and to  
121 quantify the multimodal seas by component for meaningful applications. In the present  
122 paper, we utilize the 42 years of high-resolution hindcast to investigate the weather  
123 patterns and wave climate characterizing the wave energy resources for Guam and the  
124 CNMI, while a companion paper in this special collection focuses on the PNNL hindcast  
125 around the Samoan Islands in the South Pacific Ocean (García Medina et al., 2023).  
126 Section 2 provides a summary of the hindcast model and data system from García  
127 Medina et al. (2021) followed by validation with satellite altimetry and buoy  
128 measurements around the Mariana Islands in Section 3. The validation includes an  
129 analysis of the spectral partitions at the buoys, delineating dominant components in the  
130 multimodal sea state for subsequent analyses. Section 4 links the representative wave

131 patterns and weather systems to investigate the sources of the multimodal seas. This  
132 provides a framework to understand the composition and intra-annual variations of the  
133 hindcast significant wave height and energy flux in Section 5. The concluding remarks in  
134 section 6 includes key features of the wave climate, potential for wave energy  
135 development, and applications of the data products in the Mariana region.

136 **2. Hindcast model and data system**

137 The wave hindcast system is comprised of the third-generation spectral models  
138 WAVEWATCH III and SWAN with a suite of nested computational grids from the globe  
139 to the Mariana Islands (Figure 2). WAVEWATCH III covers the entire globe at 0.5 arc-  
140 deg, the central and western Pacific at 10 arc-min resolution, and the regions around  
141 the US and affiliated Pacific Island territories at 4 arc-min resolution. The three levels of  
142 model grids are two-way nested with direct information exchange during run time. The 4  
143 arc-min regional grid, which extends about 1 arc-deg outside the EEZ of Guam and the  
144 CNMI, serves as a transition from the Pacific Island grid by producing two-dimensional  
145 spectral boundary conditions for the implementation of SWAN around the Mariana  
146 Islands. The SWAN grid is built on an unstructured triangular mesh with resolution  
147 refining from 5 km offshore to 100 m at the shore with less than 10% area change  
148 between adjacent elements for smooth transitions (Figure 3). The grid flexibility allows  
149 detailed local descriptions of the rugged shorelines, shallow shelves, harbors, and  
150 marinas extending the model validity into nearshore waters with optimal efficiency.  
151 Important coastal structures, not resolvable by the 100-m nearshore mesh, are  
152 manually incorporated for proper geospatial representation in the model.

153 Accurate wave modeling requires a digital elevation model (DEM) that can  
154 describe intricate seafloor features such as seamounts, shelves, channels, and  
155 nearshore reefs in shoaling waters. We compiled a DEM from high-resolution  
156 bathymetry and topography datasets for Guam and the CNMI with the Earth

157 Topography (ETOPO) Global Relief Model as the background. The datasets were  
158 obtained from multiple sources and rectified with the World Geodetic System 84  
159 spheroid and the local mean sea level as the horizontal and vertical datums. They are  
160 listed in the order of precedence for computational grid generation:

161 • 2007 U.S. Army Corps of Engineers (USACE) LiDAR topography at 0.5 m resolution  
162 for the entire island of Guam.

163 • 2008 US Navy & National Oceanic and Atmospheric Administration (NOAA)  
164 multibeam bathymetry of Apra Harbor at 1 m resolution.

165 • 2001 USACE Scanning Hydrographic Operational Airborne (SHOALS) LiDAR  
166 bathymetry to 40 m depth at 4 m resolution.

167 • 2007 USACE LiDAR bathymetry at 4 m resolution (limited coverage).

168 • 2007 University of Hawaii School of Ocean and Earth Science and Technology  
169 (SOEST) multibeam bathymetry to 400 m depth at 5 m resolution.

170 • 2011 University of Hawaii SOEST multibeam bathymetry to 3.5 km depth at 60-m  
171 resolution.

172 The multibeam datasets are available from the Pacific Islands Benthic Habitat Mapping  
173 Center at the University of Hawaii (<https://www.soest.hawaii.edu/pibhmc/cms/data-by-location/>). The LiDAR datasets have recently been integrated into the Continuously  
174 Updated Digital Elevation Model (CUDEM) - 1/3 Arc-Second Resolution Bathymetric-  
175 Topographic Tiles (<https://www.fisheries.noaa.gov/inport/item/48051>).

177 The input forcing for wave hindcasting includes hourly surface winds with spatial  
178 resolutions of 0.5 arc-deg from CFSR for 1979-2020 and 0.25 arc-deg from CFSV2 for  
179 2010- 2020. Orographic effects are not fully resolved in the wind field but are unlikely to  
180 influence the wave field due to the relatively small size of the Mariana Islands. The  
181 CFSR and CFSV2 winds are well reproduced over the ocean and proven accurate in  
182 wave modeling for other US regions (Allahdadi et al. 2019; Ahn et al. 2020; Garcia-  
183 Media et al. 2019, 2021a, 2021b; Yang et al. 2017, 2018, 2019, 2020; Wu et al. 2020).

184 The ST4 source term package developed by Arduin et al. (2010) is used in  
185 WAVEWATCH III and the spectrum is discretized into 24 direction bins following the  
186 setup of the NCEP global wave forecast (Chao et al., 2003a; Chao et al., 2003b) and 29  
187 frequency bins with a logarithmic increment factor of 1.1 from 0.035 to 0.505 Hz. In  
188 SWAN, the default options for the source term, bottom friction, depth-induced breaking,  
189 and triad interactions are used. The SWAN spectrum is discretized with 5-deg  
190 directional bins and 31 logarithmically-spaced frequency bins from 0.035 to 0.505 Hz to  
191 account for energy density distributions. The higher directional resolution is needed to  
192 mitigate the garden sprinkler effects, which are amplified by refraction over the steep  
193 volcanic island slopes. The model configuration is consistent with the hindcast for the  
194 US West Coast, Alaska, and the Hawaiian Islands by following the IEC TS for wave  
195 resources characterization.

196 We implement the model system for wave hindcasting around the Mariana  
197 Islands for the period from 1979 to 2020. The 42-year hindcast dataset includes the bulk  
198 and partitioned wave parameters such as the significant wave height, peak period, and  
199 peak direction for all grid points within the SWAN domain as well as the frequency-  
200 direction spectra along the 100-m water depth contour, at distances 2 and 5 km from  
201 the shore, and at the three buoys near the islands. The computational time step is 2 min  
202 for a balance of accuracy and efficiency, while the output interval of 3 hours is sufficient  
203 to capture wave events from mesoscale weather phenomena. The dataset for the  
204 Mariana Islands along with those for other US and affiliated Pacific islands will be  
205 publicly available at the Registry of Open Data on AWS (Amazon Web Services) for  
206 implementation in coastal environment studies, wave climate analysis, and energy  
207 resources assessment.

208 **3. Validation of wave hindcast**

209 Satellite-borne altimeters have provided nearly continuous measurements  
210 covering most of the ocean since the GEOSAT (GEOdetic SATellite) mission in 1985.  
211 Despite the long orbital repeat cycle, the altimetry data serve as a valuable resource for  
212 model validation and assessment (e.g., Li et al., 2016; Stopa and Cheung, 2014; Dutheil  
213 et al., 2021). The Australian Ocean Data Network portal (<https://portal.aodn.org.au/>)  
214 provides the significant wave height derived from the Ku band measurements on  
215 multiple altimetry platforms (Ribal and Young, 2020). The dataset includes more than  
216 four million observations within the SWAN model domain from 1985 to 2020. We binned  
217 the dataset over a 0.2 arc-deg grid to allow for an adequate description of the spatial  
218 variation and provide sufficient data in each bin for statistical analysis. The data density  
219 in Figure 4a shows sufficient coverage with distinct tracks having more measurements.  
220 The hindcast significant wave height is interpolated to match the time and location of  
221 each available altimetry observation and bins with less than 200 data pairs are omitted  
222 in the computation of error metrics. Figures 4b, c, and d show the bias, linear correlation,  
223 and the root-mean-square-error (RMSE) of the hindcast against the altimetry data.  
224 Good agreement is indicated by a bias of less than 0.1 m as well as a correlation  
225 coefficient of over 0.9 and its relatively uniform distribution across the model region. The  
226 negative bias of the model results near Guam is consistent with experiences from local  
227 weather forecasters with the operational data products. The RMSE of 0.2~0.5 m is  
228 typical among hindcast datasets for other Pacific Island regions (Li et al., 2016, Durrant  
229 et al., 2014) and is not concerning given the similar range of errors in the altimetry-  
230 derived significant wave height (Yang and Zhang, 2019). The slightly higher RMSE to  
231 the north and west of the Mariana Islands is a result of the more dynamic weather  
232 systems around the western Pacific rim.

233 Wave buoy measurements are commonly considered as ground truth for the  
234 validation of model results. There are three Waverider buoys deployed, respectively, off  
235 the north shore of Saipan, and the north and east shores of Guam with Coastal Data

236 Information Program (CDIP) indexes 197, 196, and 121 (Figure 1b). The wave buoys  
237 continuously record sea-surface movement, from which the frequency-direction  
238 spectrum and parameters such as significant wave height, peak period, and peak  
239 direction are automatically derived every 30 minutes. Buoy 196 immediately north of  
240 Guam is the most exposed to the dominant wave components. The time series of the  
241 hindcast and recorded wave parameters from 2013 to 2020 are shown in Figure 5a and  
242 a close-up view of the comparison for 2017 in Figure 5b. The buoy records suggest a  
243 wide range of conditions with significant wave heights of 0.3-6.9 m, peak periods of 3.3-  
244 18.2 s, and directions typically from the southwest through east clockwise. The diverse  
245 waves have strong seasonal variations with large events above 4 m in the winter  
246 months and waves of 1-2 m in the summer months. The moderate summer sea state is  
247 occasionally interrupted by large waves from tropical cyclones passing nearby as  
248 indicated by rapid changes in wave directions from south to north. The hindcast follows  
249 the buoy measurements to capture both the seasonal variations and the individual wave  
250 events, despite slight overestimation of the small events and a tendency to  
251 underestimate the peak of some large events. The latter is likely due to limited  
252 resolution of strong winds near the core of large storms (Stopa and Cheung, 2014).

253 We pair the hindcast and buoy measurements of the same hour for a direct  
254 comparison. Figures 6a and 6b provide the scatter and Q-Q plots for the significant  
255 wave height from 2013 through 2020. The scatter plot gives a small RMSE of 0.3 m and  
256 a high correlation coefficient of 0.91 with 90% of the data pairs within  $\pm 0.46$  m of a  
257 perfect match. This point comparison corroborates the spatial comparison with altimetry  
258 data with an improved agreement and reiterates the good correspondence between the  
259 hindcast and observations. The large spreading for waves above 4 m may be attributed  
260 to timing differences of events between the two datasets. The quantile-quantile (Q-Q)  
261 plot eliminates any time difference to compare the statistical distributions. Despite the  
262 slight underestimation of the significant wave height in the range of 2 to 4 m, the

263 hindcast is within 10% of the measurements even at the 99th and higher percentiles.  
264 Comparisons with available data at Buoys 121 and 197 show similar patterns (Figures  
265 A1-A4 in the appendix). The location of Buoy 121 on the east side of Guam is open to  
266 waves from the central Pacific that are partially blocked at Buoys 196 and 197. The  
267 recorded waves, which approach Guam from the east through the south in the summer,  
268 are complementary to the validation. Despite mismatching the peaks for some extreme  
269 events, the long-term hindcast provides a reliable dataset for statistical analysis and an  
270 accurate description of the overall sea state in the entire region.

271 The sea state comprises waves from multiple sources with distinct characteristics.  
272 Figure 7 illustrates the occurrence probability of the measured significant wave height  
273 and peak period at buoy #196 along with the distributions in the cardinal and ordinal  
274 directional sectors. Waves from the northeast are dominant 42% of the time, and these  
275 are followed by easterly and northerly waves with 24% and 18% occurrences. The  
276 remaining directional sectors account for less than 10% each. Waves from the  
277 southeast and south are rare and partially blocked by the island, and those that reach  
278 the buoy after wrapping around the headlands account for less than 0.2% of the time.  
279 Despite the disproportionate distribution among the directions, all show a dominant  
280 peak at 10 s and rapidly diminishing occurrence probability and significant wave height  
281 below 7 s. A second peak appears at 12 s from north and northwest with significant  
282 wave height ranges similar to those of the primary peak at 10 s. The results suggest a  
283 mix of dominant wave components in the Mariana Islands delineated by 7 and 11 s  
284 periods. Following the marine definitions of the US National Weather Service  
285 (<https://www.weather.gov/gum/MarineDefinitions>), we classify the multimodal seas into  
286 three main components identified as wind waves, seas, and swells for periods below 7s,  
287 of 7-11 s, and above 11 s, respectively. The classification of the dominant wave  
288 components is also consistent with local knowledge and the weather patterns in the  
289 Western and Central Pacific (Jordan, 1955; Lander 1994; Lander and Guard, 2003).

290 Since the wave climate analysis will be based on separate wave components from the  
 291 hindcast, it is necessary to independently verify their delineation and to understand the  
 292 model performance for each sea state.

293 For each component, we integrate the 2D spectrum from SWAN over the period  
 294 range for the zeroth moment ( $m_0$ ) and calculate the representative wave height  $h_s =$   
 295  $4\sqrt{m_0}$  that reflects the energy level. Figure 8 shows good correspondence of the  
 296 representative wave heights computed from the hindcast and records at #196 for 2017.  
 297 The predefined period ranges capture distinct peaks of wind waves, seas, and swells,  
 298 but are not absolute in delineating the wave components. There are some concurrent  
 299 peaks between adjacent period ranges as the atmospheric forcing from weather  
 300 systems is not always well distinguished from one another. The three components have  
 301 distinct seasonal characteristics adding confidence to their delineation by period. The  
 302 wind waves under 7 s are the most persistent with representative wave heights in the  
 303 range of 0.5-2 m throughout the year. Seas of 7 to 11 s also have a year-round  
 304 presence with higher energy in the winter than in summer. In particular, the  
 305 representative wave height reaches 4 m in February and typically stays within 0.5~1.5  
 306 m from June to August. There is also a stronger correlation between the wind waves  
 307 and seas in the winter months suggesting their generation from large weather systems  
 308 with long fetches extending to the Mariana Islands. Swells with periods above 11 s are  
 309 mostly independent of the other components and dominant in the winter months with  
 310 heights exceeding 3 m, while being sporadic in the summer. The hindcast and the buoy  
 311 measurements show good agreement in terms of the total energy for the mixed seas as  
 312 well as the energy distribution among the components for in-depth investigation of  
 313 weather and wave climate as well as its implications for wave resources.

314 **4. Weather and wave patterns**

315 The strong seasonal, multi-modal sea state around the Mariana Islands is a  
316 result of synoptic and mesoscale weather systems in different parts of the Western and  
317 Central Pacific. The National Weather Service reports the atmospheric and marine  
318 conditions daily in the Area Forecast Discussion (AFD) for Guam and the CNMI  
319 (<https://mesonet.agron.iastate.edu/wx/afos/list.phtml>). The historical AFD archive and  
320 the hindcast provide a basis for identifying typical weather and wave patterns in the  
321 region for understanding the sea state in terms of its components delineated by period.  
322 Figure 9 shows representative summer events with the CFSR surface wind and  
323 pressure as well as the hindcast significant wave height, peak period, and peak  
324 direction to demonstrate the met-ocean conditions. Among the synoptic weather  
325 phenomena in the region, trade winds are prevalent throughout the year (Eldredge,  
326 1983, Jordan, 1955). They originate from a subtropical high ridge extending across the  
327 North Pacific with anti-cyclonic flows producing easterly winds toward the Mariana  
328 Islands. The trade winds are relatively gentle and uniform with periodic fluctuations  
329 during most of the summer. Figure 9a shows typical trade-wind conditions with steady  
330 moderate winds of 8 m/s in early July 2020. The seas around the islands, which are  
331 generated along an open fetch, are a result of the upstream flow and local winds at  
332 earlier and present time, respectively. The significant wave height of 1.5 m is quite  
333 uniform east of the islands, but decreases to 0.5 m off the sheltered west shores.  
334 Outside the sheltered waters, the peak period is consistent at 8 s, associated with the  
335 open fetch from the North Central Pacific to the Mariana Islands.

336 The summer trades near the Mariana Islands are sometimes interrupted by the  
337 monsoons originating from Southeast Asia (Becerro et al., 2006, Lander, 1996, Beattie  
338 and Elsberry, 2012, Lander and Guard, 2003). On July 30, 2018, the westerly monsoon  
339 flow with speeds of 7 m/s extends from the Philippines to the Mariana Islands (Fig. 9b).  
340 The event lasted for a few days with fully developed seas of 1.2 m significant wave  
341 height around Guam and the CNMI on August 1, 2018. The peak period increases from

342 7 to 9 s across the islands due to the increased fetch toward the east. Tropical cyclones  
343 occur typically in the summer and transition months (Holliday, 1975, Rupp and Lander,  
344 1996) and influence the wave conditions around the Mariana Islands, even from a  
345 distance. As an example, Figure 9c shows the northwest migration of Typhoon Haisen  
346 away from the CNMI in early September 2020. It strengthens rapidly from September 3  
347 to 4 and becomes a super-typhoon reaching category 4 strength with maximum 1-min  
348 sustained wind speeds above 69 m/s, but did not have a direct impact on the islands  
349 due to its location 1600 km away. The typhoon swells generated earlier traverse the  
350 open ocean to Mariana waters reaching 2.2 m wave height and 13 s peak period,  
351 except for the sheltered areas to the southeast where background seas of about 1 m  
352 and 9 s from the trade winds are dominant. Tropical cyclones which make direct landfall  
353 or pass near the islands can generate hazardous waves to the coastal communities.

354 In addition to trade-wind flows, mid-latitude cyclones and anticyclones migrating  
355 from west to east are rather common during the winter months (Harr et al., 2000). The  
356 resulting northerly swells and seas produce a notable contribution to the wave climate in  
357 the islands. Figure 10a shows a high-pressure system in the East China Sea on  
358 January 21, 2020, moving eastward to the north of the Mariana Islands 1.5 days later.  
359 The anticyclonic system produces northeasterly winds with speeds reaching 12 m/s  
360 across the islands that augment the easterly trades to generate heightened seas.  
361 Meanwhile, the low-pressure system east of the Kuril Islands strengthens and moves  
362 northeast toward the Kamchatka Peninsula. The low-pressure system with strong  
363 northwesterly winds generates swells reaching the Mariana islands. The combined  
364 significant wave height from seas and swells reaches 3.5 m on January 22, 2020.  
365 Around each island or island group, prominent shadows of the seas develop to the  
366 southwest with near-shore wave heights of less than 2.0 m. More extended shadows of  
367 the swell develop to the south of the islands at the same time. In contrast to the gradual  
368 transition of the significant wave height, the peak period shows a patchy distribution

369 around the islands associated with shadowing of the multi-modal seas. The period of  
370 northeast seas from the nearby mid-latitude high increases from 8 s to 11 s along the  
371 fetch to the southwest. The background north swells with 14 s periods from the more  
372 distant mid-latitude low are visible off the western shores in the shadows of the  
373 moderate seas and to the east of the island chain.

374 The migrating mid-latitude systems can also generate northerly winds reaching  
375 the islands with influence on the local met-ocean conditions (Donets et al., 2018, Lander,  
376 1994). In late January 2020, a low-pressure system east of Japan is accompanied by a  
377 high-pressure system at the East China Sea during the passage of an occluded cold  
378 front (Figure 10b). South-southeasterly winds prevail in the prefrontal atmosphere north  
379 of the Mariana Islands. As the high and low-pressure systems strengthen and move  
380 eastward, strong northwesterly winds sweep across the East China Sea bringing cold  
381 air from the mid-latitudes toward the tropics. The cold front reaches the Mariana islands  
382 with post-frontal northerly winds of 10 m/s on February 1. The post-frontal winds  
383 converge with moderate easterly trade winds of 6 m/s creating a shearline southeast of  
384 the islands. These near and far-field weather systems have profound influence on the  
385 complex local wave conditions. The significant wave height of the combined seas  
386 reaches 2.5 m west of the islands primarily due to the strong post-frontal northerly winds  
387 and decreases to 1.5 m on the east side, where the sea state can be mostly attributed  
388 to the moderate easterly seas augmented by north swells generated earlier by the low-  
389 pressure system east of the Kuril Islands. The distinct characteristics of each wave  
390 system can be identified in the peak period distribution. The short 5-7 s and long 14 s  
391 peak period west and east of the island chain, respectively, indicate wind waves from  
392 the frontal system and swells from the mid-latitude low-pressure system, while the  
393 intermediate 8-10 s period around the islands represents the combined effect of the  
394 easterly trade winds and northerly postfrontal winds. These typical weather systems and

395 the corresponding wave fields provide a basis to elucidate the seasonal and spatial  
 396 variations of the wave energy resources around the Mariana Islands.

397 The characteristics of the dominant wave components are best assessed using  
 398 the partitioned hindcast. The IEC wave resources parameters, which are integrated  
 399 from the entire spectrum, can be extended through spectral partitioning to provide a  
 400 more complete description for multi-modal seas. The spectral partitions of the hindcast  
 401 have already been validated with measurements from buoy 196. Figure 11 shows the  
 402 six IEC parameters computed from the validated spectrum as well as its partitions for  
 403 the three dominant wave components in 2017. The wave power and significant wave  
 404 height time series have similar trends, but the former more clearly depicts the  
 405 dominance of seas between 7 and 11 s due to the year-round occurrence of trade winds  
 406 as well as their intensification by migrating mid-latitude high in the northwestern Pacific  
 407 during the winter. The energy is occasionally augmented by swells with periods greater  
 408 than 11 s from mid-latitude lows further north. The wind waves can have significant  
 409 wave heights in the 0.5~2 m range, but only minor contributions to the total energy  
 410 because of their short periods. The integrated energy period is close to the period for  
 411 seas due to their primary contribution.

412 The spectral width parameter clearly differentiates the narrowband  
 413 characteristics of the swells and the broadband wind waves. The seas have very  
 414 consistent spectral width slightly above the swells. The same parameter for the  
 415 combined sea state is noticeably larger than those of the individual components but is  
 416 closer to the wind waves despite their low energy level throughout the year. This is  
 417 complemented by the direction of the maximum directionally resolved wave power and  
 418 the directionality coefficient, which correspond well to the characteristics of each  
 419 component from a specific directional window. In the winter, eastward migration of mid-  
 420 latitude lows and highs produces periodic direction changes for swells from north to  
 421 northeast and seas from northeast to east. Wind waves throughout the year and seas in

422 the summer are mainly from the east due to persistent trades albeit with occasional  
423 disruptions from tropical cyclones and monsoons. The results show wider directional  
424 spread of wind waves and seas due to their multiple sources, but more focused swells  
425 primarily from more distant low-pressure systems in the winter. Computation of the IEC  
426 parameters by spectral partition provides more pertinent information that can assist  
427 designers and operators to tune their devices for optimizing energy production in the  
428 strongly multimodal sea state around the Mariana Islands.

429 **5. Seasonal wave resources**

430 The validated long-term hindcast provides a wealth of information for analysis of  
431 the mixed seas and the individual components to understand the wave climate and  
432 resources. Figure 12 shows the average significant wave height and wave power as  
433 well as their intraannual variability. Averaged over the 42-year hindcast, the significant  
434 wave height is quite uniform across the region with a slight variation from 2.0 m to 1.8 m  
435 along the island chain from north to south. The wave height decreases to under 1 m off  
436 the western shores due to sheltering of the prevailing seas from easterly trade winds.  
437 This is important for the development of wave energy resources in the Mariana Islands  
438 because the western shores host major harbors and infrastructure, which can provide  
439 logistic support and serve as transportation hubs for deployment and maintenance. The  
440 wave conditions have large seasonal variations. Calmer trade winds occur in the boreal  
441 summer months of June, July, and August with about 1.5 m average significant wave  
442 height in the open ocean. Waves from tropical cyclones and summer monsoons can  
443 interrupt the moderate sea state and have minor contributions to the seasonal average  
444 due to their relatively low occurrence, but are enough to introduce faint shadows off the  
445 east shores. The trade winds and the resulting seas are strengthened by mid-latitude  
446 high-pressure systems in the northwestern Pacific during the boreal winter months of  
447 December to February. Frontal systems and mid-latitude low-pressure systems to the

448 north also become active and produce severe wind waves and large swells impacting  
449 the Mariana Islands. Multi-modal seas become more frequent in winter with occasional  
450 treacherous conditions increasing the average significant wave height to above 2.4 m  
451 off the island shores. Such conditions are less favorable for installation and  
452 maintenance of wave energy devices and mooring systems.

453 The wave power is proportional to the square of wave height producing a larger  
454 contrast in the spatial distribution. The year-round average reaches 18 and 15 kW/m in  
455 the north and south with 10 kW/m in the shadows west of the islands. The Mariana  
456 Islands have steep insular slopes in the absence of shelves. The wave power in the  
457 deep water around islands increases linearly with the wave period to augment the  
458 seasonal variation. With mostly seas from westerly monsoons and east trade winds in  
459 the summer, the seasonal average is about 10 kW/m away from the islands and below  
460 6 kW/m in the shadows off the west and east shores. In the winter months, the wave  
461 power reaches 26 and 22 kW/m in the northern and southern Mariana Islands primarily  
462 due to stronger easterly trade wind seas. The slight northward increase is a result of  
463 intermittent occurrences of north swells from mid-latitude low-pressure systems off the  
464 Kamchatka peninsula as well as northerly seas and wind waves from mid-latitude high-  
465 pressure systems and northwest frontal systems originating in the East China Sea. The  
466 north-facing shore experiences the most diverse wave climate as well as the highest  
467 energy level of each island. The wave power at the east and west shores is under 10  
468 kW/m due to island sheltering. The overall increase of the significant wave height and  
469 wave power from the summer to winter in Guam and the CNMI is directly related to  
470 transition of the dominant weather systems in the Western and Central Pacific. Bulk  
471 parameters derived from the combined seas in Mariana waters do not always provide  
472 pertinent information for practical applications.

473 Performance of most wave energy conversion devices also depends on the wave  
474 period. The wave power can characterize the available wave resources but might not be

475 sufficient for engineering design and operation planning. The mixed seas around the  
476 islands are generated by a range of synoptic and mesoscale processes in the Western  
477 and Central Pacific. The hindcast partitions allow separate analysis of each wave  
478 component defined by its period range. There are up to 10 partitions from the watershed  
479 algorithm for each hindcast spectrum in the SWAN output. We consider the three  
480 primary partitions by energy level and classify the computed wave height under the wind  
481 wave, seas, and swell categories based on the period being under 7 s, between 7 and  
482 11 s, and above 11 s. When more than one partition is assigned to the same  
483 component at a time step, only the largest one is recorded to minimize noise from small  
484 and intermittent events and concatenated into a continuous sub-dataset. The three sub-  
485 datasets provide a basis for spatial and temporal analysis by dominant wave  
486 components. Figure 13 provides the average partitioned wave height for the  
487 occurrences in each category during the boreal summer and winter months. In contrast  
488 to the combined seas, the three components have distinct energy levels, patterns, and  
489 seasonality associated with climatology in the western and central Pacific.

490 The short-period wind waves are primarily generated by local winds with a  
491 relatively uniform distribution across the island chain. The summer has an average  
492 partitioned wave height of less than 0.6 m associated with the light to moderate easterly  
493 trades and southwesterly monsoon winds in the region. The partitioned wave height  
494 increases to 1.0 m from the strengthened trades and northerly frontal winds in the winter.  
495 In contrast, the intermediate-period seas are influenced by upstream trade wind flows  
496 with a longer fetch extending to the central Pacific. The average partitioned wave height  
497 reaches 1.1 m and 1.7 m on the windward side in the summer and winter, with a  
498 prominent reduction on the leeside to the west due to sheltering by the island chain. The  
499 long-period swells have distinct seasonal patterns due to the source location. Tropical  
500 cyclones are most active in the western Pacific during the summer months producing an  
501 average wave height of 1.3 m west of the island chain. The wave height from a tropical

502 cyclone varies with the intensity, track, and forward speed, and the peak value could be  
503 much larger than the average. Mid-latitude cyclones occur to the north of the islands  
504 during the winter months. The average partitioned wave height of the swell reaches 1.5  
505 m with a prominent shadow to the south. Similar to tropical cyclones, the peak swells  
506 could be much larger than the average. The seasonal distribution of individual wave  
507 components provides an additional level of detail for the description of wave climate and  
508 energy resources that assist selection and design of wave energy conversion devices  
509 for deployment near the Mariana Islands.

510 **6. Discussion and Conclusions**

511 The multimodal wave conditions in the Mariana Islands are influenced by local  
512 and distant weather systems in the Western and Central Pacific. An assessment of the  
513 recorded significant wave height, peak period, and peak direction identifies distinct  
514 components of wind waves, seas, and swells of below 7 sec, 7-11 sec, and above 11  
515 sec, respectively. The 42-year hindcast, which was generated from reanalysis winds  
516 following the IEC TC-114 Technical Specification for a Class 2 (Feasibility) study,  
517 provides an effective approach to quantifying their spatial and temporal distributions  
518 around the islands for in-depth analysis of the wave climate and energy resources. The  
519 wave parameters and spectral partitions are thoroughly validated with measurements  
520 around the islands. The comparison with the altimetry observations shows good  
521 reproduction of the overall wave height and spatial pattern. The slight increase of  
522 RMSEs toward the west and north is associated with transition to more dynamic  
523 weather systems from the Central to the Western Pacific. In comparison with buoy  
524 records, the wave hindcast shows slight overestimation for the small events and  
525 underestimation for the peaks of large events, while demonstrating good agreement for  
526 the energy distribution among the dominant wave components in the multi-modal sea  
527 state. Although the demarcation of wind waves, seas, and swells is not always well

528 defined as with the atmospheric forcing and is location or event dependent, the good  
529 correlation obtained in the spectral partitions ensures that the selected period ranges  
530 can capture the dominant energy in the subsequent analysis.

531 The validated hindcast and partitioned parameters provide details in the temporal  
532 and spatial distributions of the wind waves, seas, and swells. The year-round trade  
533 winds, driven by a subtropical ridge, account for the prevailing and background seas  
534 that play a significant role in the local wave climatology. The sea state is gentler with the  
535 prevailing direction from the east in the summer and can be augmented on the leeside  
536 of the islands by seas from monsoons as well as swells from tropical cyclones to the  
537 west even at some distance away. In the winter months, the trade wind waves are  
538 energized by northeasterly winds from mid-latitude high-pressure systems and  
539 augmented by swells from the mid-latitude low-pressure systems to the north. In  
540 addition, frontal systems can reach the Mariana Islands from the northwest generating  
541 short-period wind waves with large amplitude on the west side of the islands. These  
542 local and distant weather systems bring waves of various occurrence frequencies  
543 accounting for the seasonal variation of significant wave height, peak period, and  
544 energy level. The wave components from different directions are sheltered by the  
545 islands leading to complex spatial variations. Implementation of the IEC wave energy  
546 resource parameters by spectral partition is an effective approach to describe the multi-  
547 modal sea state with distinct energy characteristics for each component.

548 The 42-year wave hindcast provides valuable insights into the wave resources  
549 along the island chain in support of energy converter design and planning. The wave  
550 power has strong seasonal variation due to transition between near and far-field  
551 weather systems affecting both the wave height and period. The summer has an  
552 average around 10 kW/m associated with light to moderate trade winds, despite  
553 occasional large events from tropical cyclones and monsoons. The sea state becomes  
554 more treacherous in the winter months with average wave power up to 22 kW/m around

555 the islands. Despite having different energy levels in the two seasons, the northern  
556 portion of the island archipelago has slightly greater energy resources. Around each  
557 Island, the energy level decrease dramatically off the west shores because of sheltering  
558 of persistent trade wind waves and seasonal north swells. The north-facing shores of  
559 Saipan, Rota, and Guam have the most diverse energy resources comparable to those  
560 in the open ocean due to direct exposure to seas all year round as well as strong wind  
561 waves and swells in the winter. The available wave power, which meets the  
562 requirements of most energy converters, is at a short distance from shores due to the  
563 steep insular slope and the absence of shelves. However, candidate energy converters  
564 should be adaptable over the wide range of peak periods and directions to be viable for  
565 the Mariana Islands.

566 The PNNL hindcast has been proven to be a valuable dataset for wave climate  
567 research and energy resource assessment through the present study with the Mariana  
568 Islands as well as the corresponding studies for American Samoa (García Medina et al.,  
569 2023) and Hawaii (Li et al., 2021). The Mariana Islands experience the most diverse  
570 wave conditions among the studied regions due to their location in a typhoon spawning  
571 area as well as the presence of monsoons and cold fronts from the western Pacific Rim.  
572 The dataset covers all US and affiliated Pacific islands, in which the ocean has a  
573 significant role in the livelihood of local communities. The archived model output of bulk  
574 and partitioned parameters within the EEZs and frequency-direction spectra near the  
575 shore is meant to support a wide variety of scientific and engineering applications. The  
576 100-m resolution on an unstructured grid provides a long-term description of the  
577 nearshore wave environment that influences coastal circulation, sediment transport, and  
578 mixing processes. The spectral data can provide boundary conditions for phase-  
579 averaged or phase-resolving modeling of coastal waves and circulations at much higher  
580 resolution at the shore. The full hindcast dataset, which will be publicly available online,

581 can be implemented in studies related to coastal hazards, shoreline erosion, maritime  
582 operations, and coral reef ecosystems.

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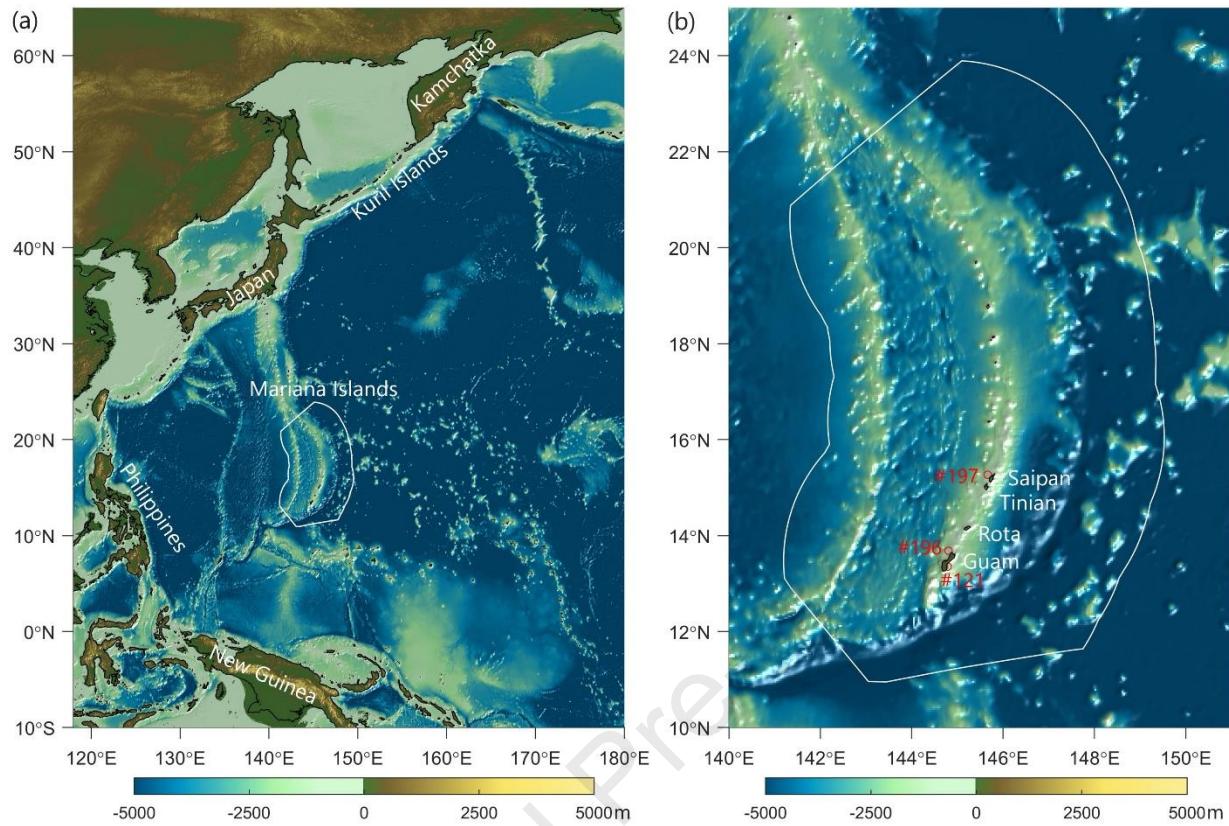


Figure 1. Location and bathymetric maps. (a) Western and central Pacific. (b) Mariana Islands. The red circles and labels indicate CDIP buoys. The white line delineates the Exclusive Economic Zone.

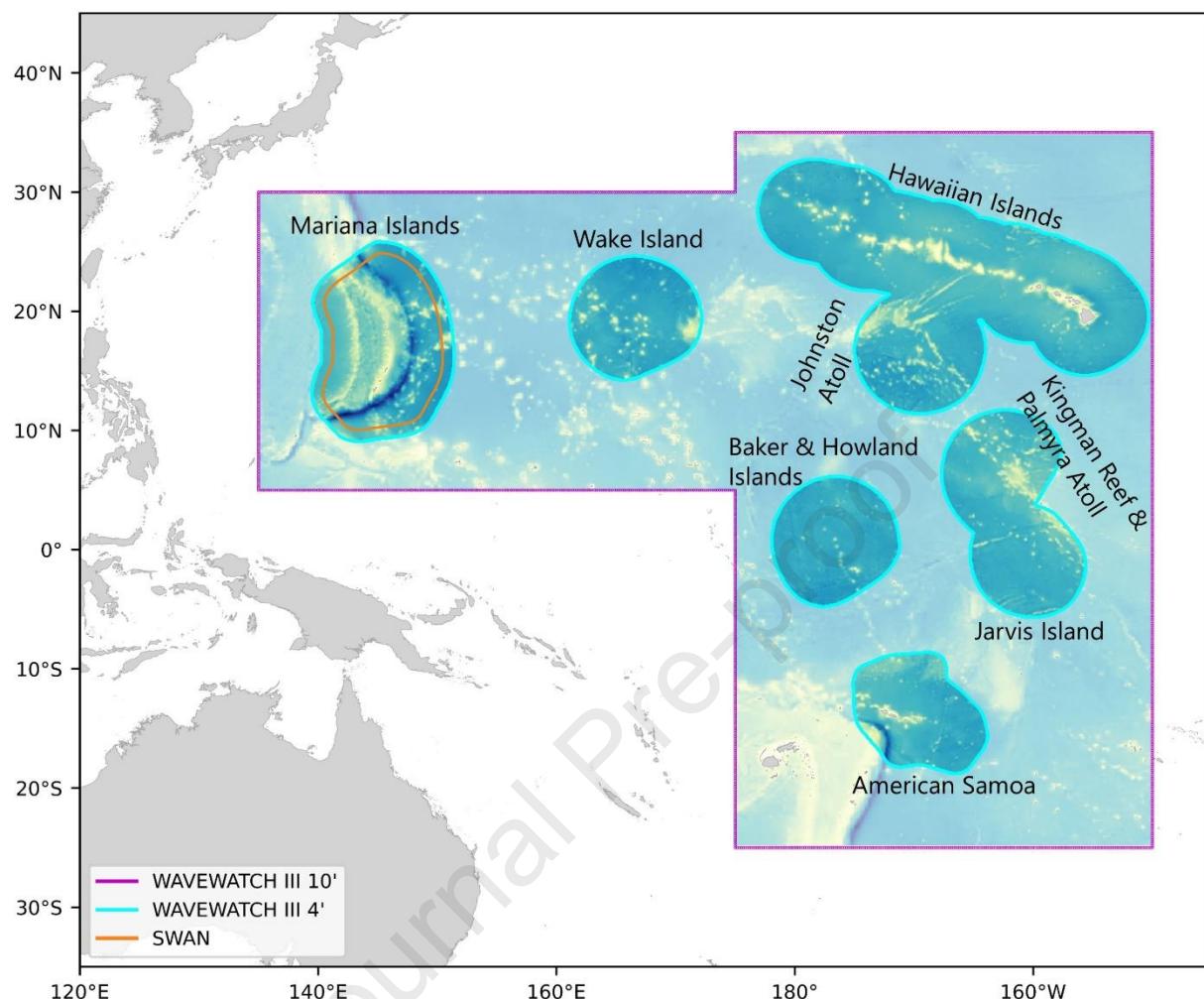


Figure 2. Regional WAVEWATCH III nested grids for the Pacific Islands and the SWAN unstructured grid for the Mariana Islands.

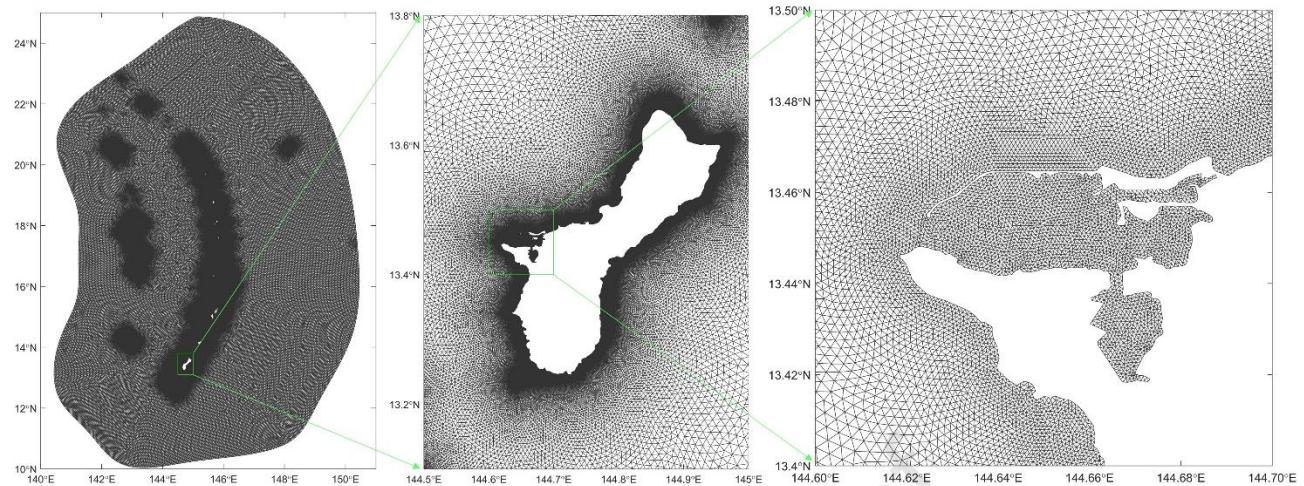


Figure 3. Telescopic views of the SWAN unstructured grid with resolution from 5 km at the boundary to 100 m at the shore.

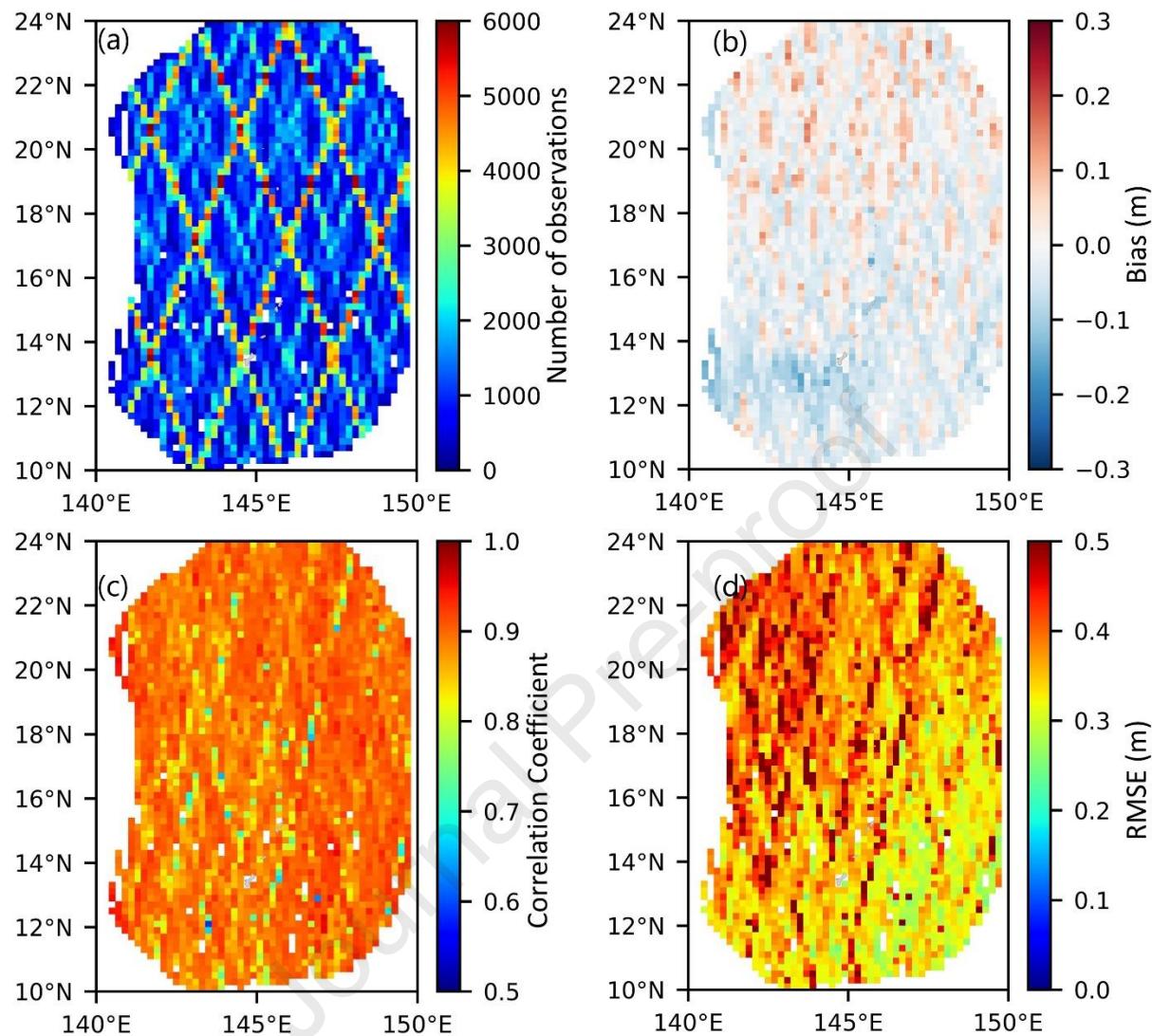


Figure 4. Comparison between hindcast and altimetry observations. (a) Bins and data density over the SWAN domain. (b) Bias. (c) Correlation Coefficient. (d) Root-mean-square error (RMSE).

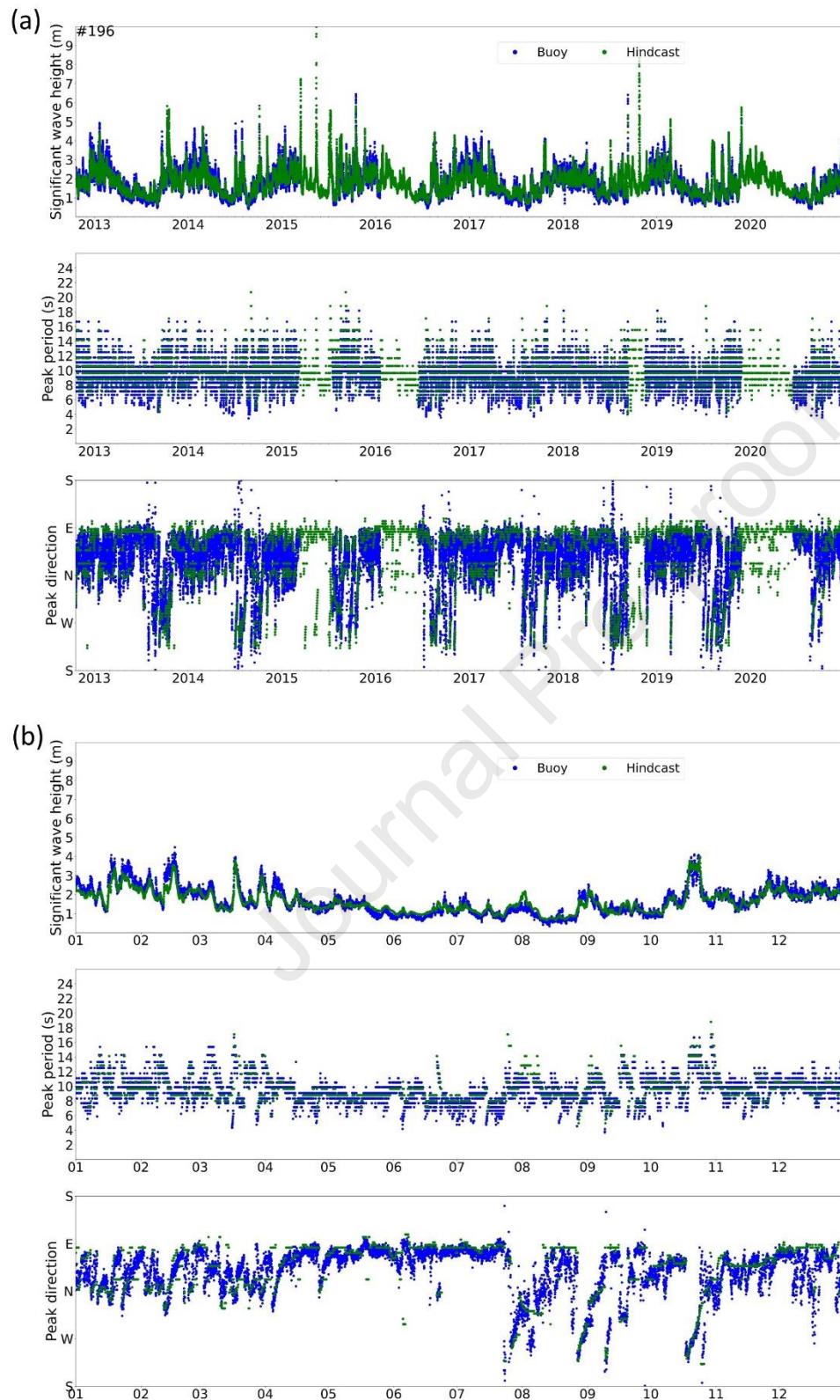


Figure 5. Time series of significant wave height, peak period, and direction from buoy #196 and the wave hindcast (a) from 2013 to 2020 and (b) in 2017.

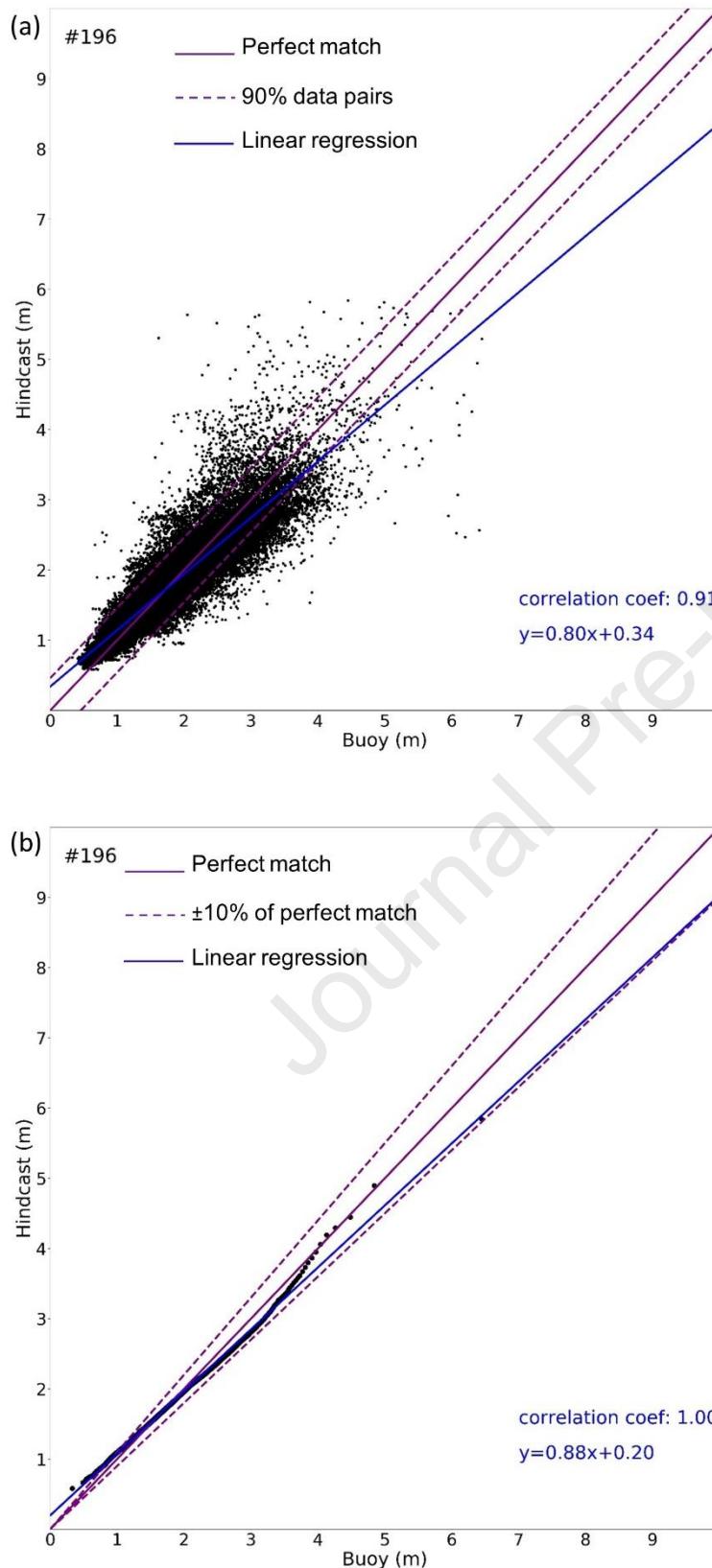


Figure 6. Scatter and Q-Q plots of measured and hindcast significant wave height at buoy #196.

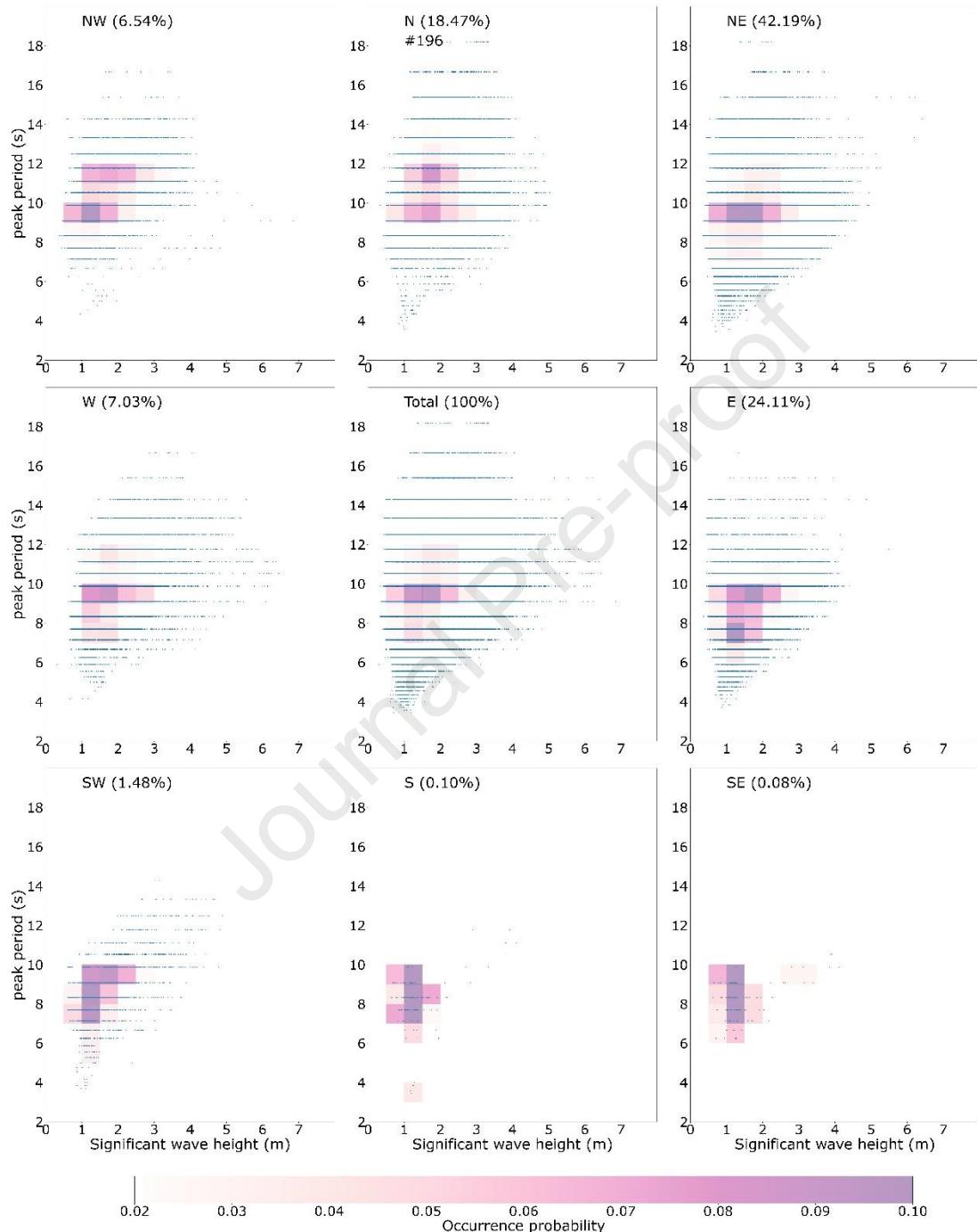


Figure 7. Scatter diagram for the measured significant wave height and peak period in eight directional windows from buoy #196.

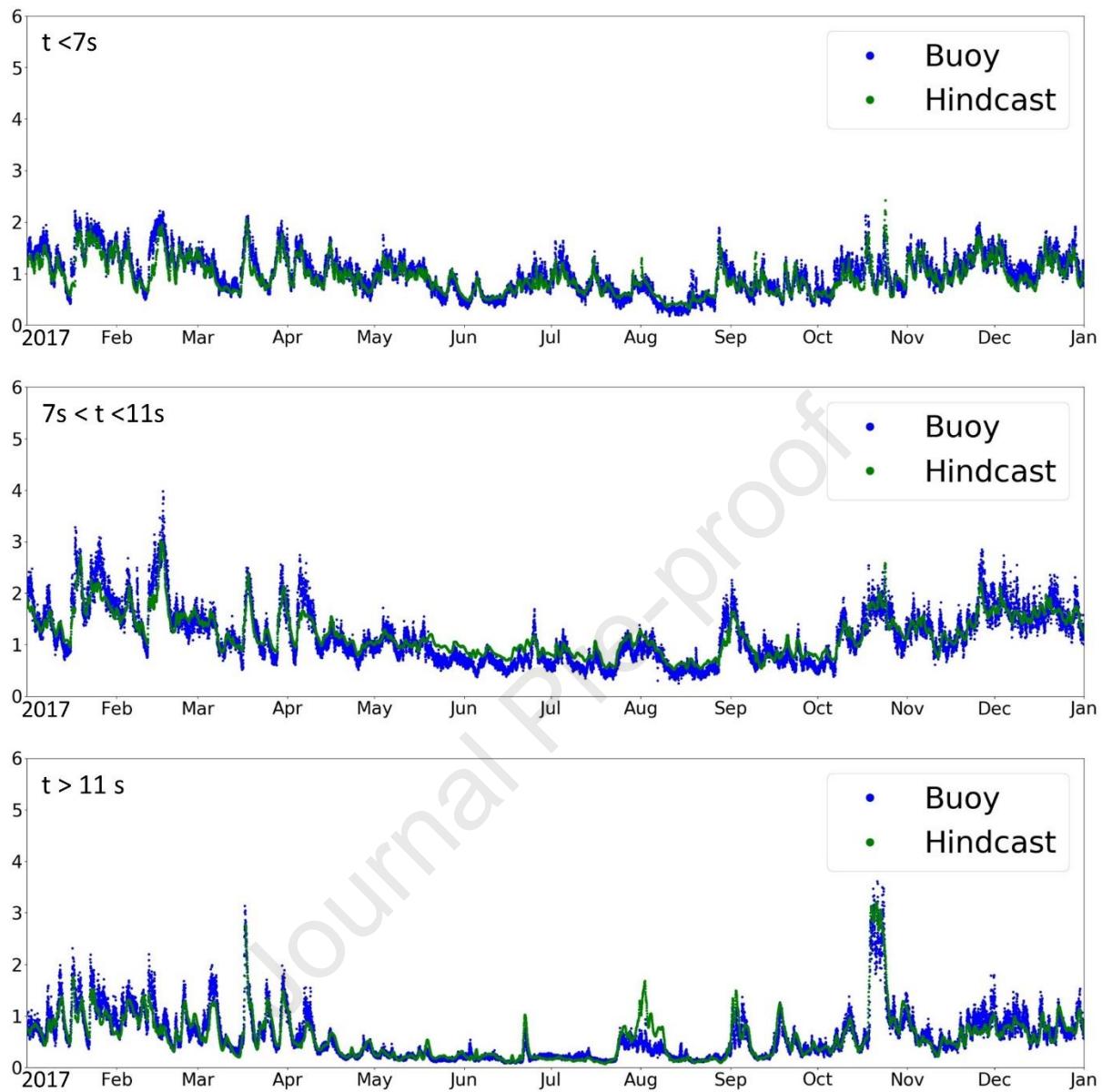


Figure 8. Time series of representative wave heights for periods less than 7s, 7-11 s, and above 11 s at buoy #196 in 2017.

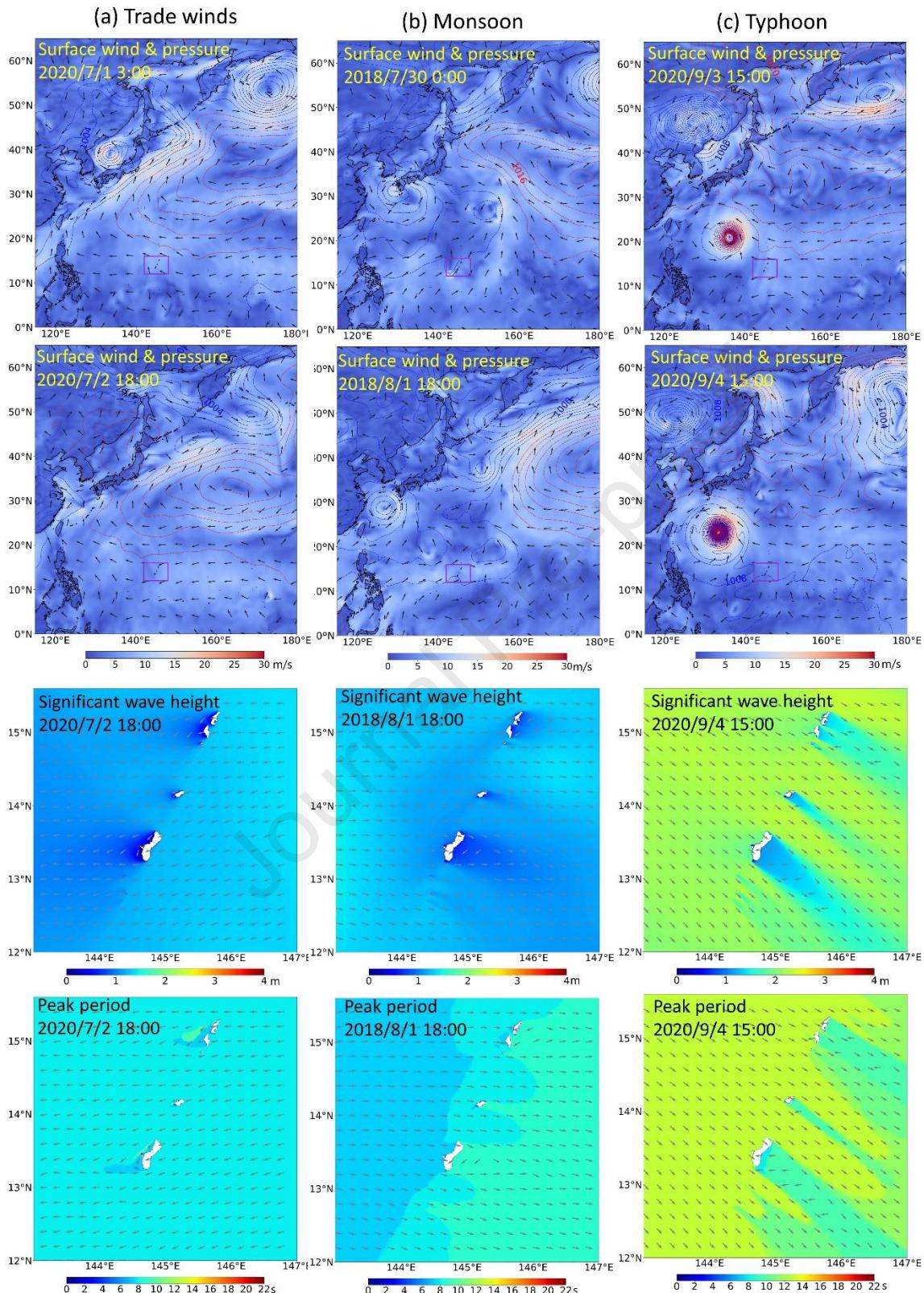


Figure 9. Representative summer weather systems in the Western and Central Pacific and wave patterns near Guam and the CNMI. (a) Trade winds. (b) Monsoon. (c) Tropical cyclone.

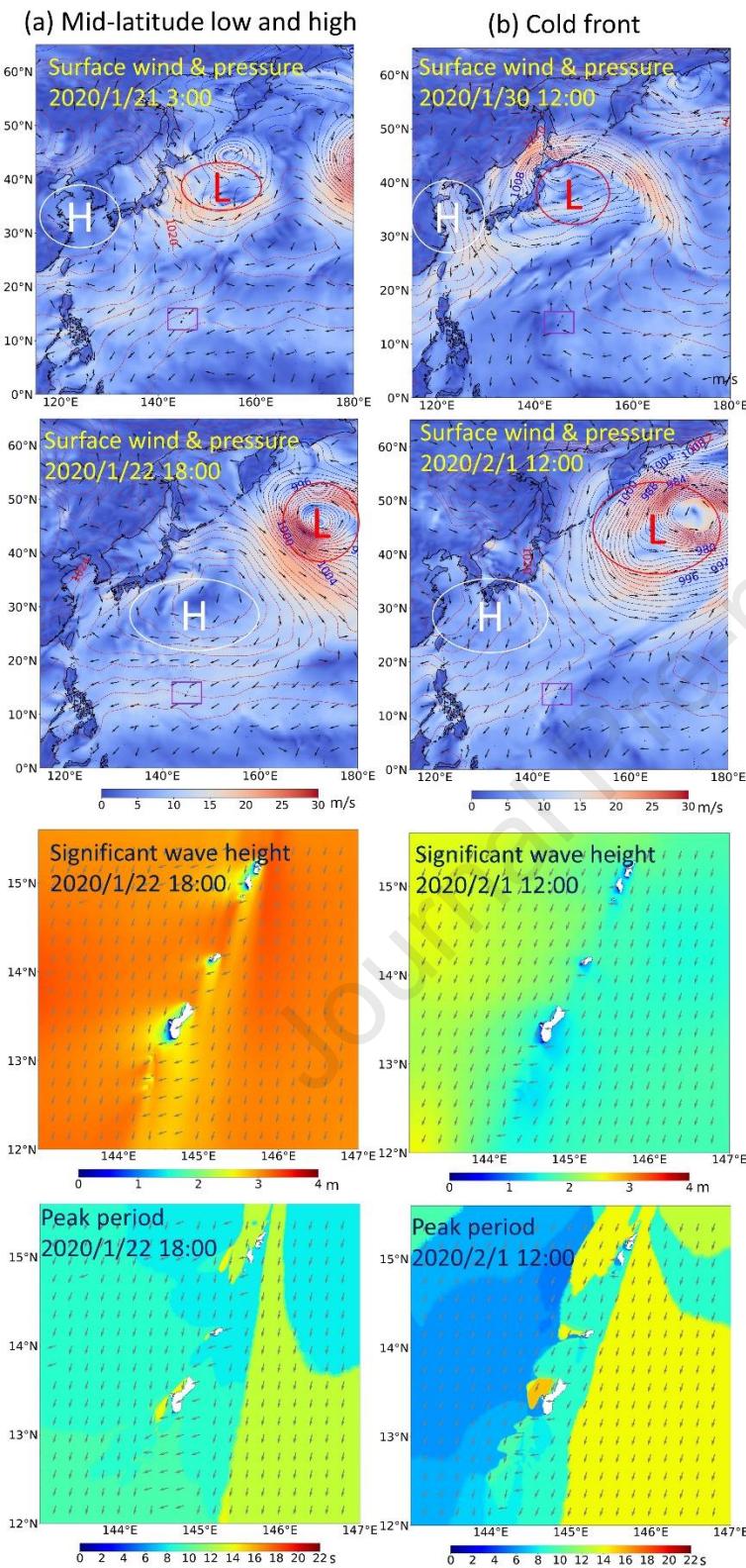


Figure 10. Representative winter weather systems in the Western and Central Pacific and wave patterns near Guam and the CNMI. (a) Mid-latitude low and high (b) Cold front

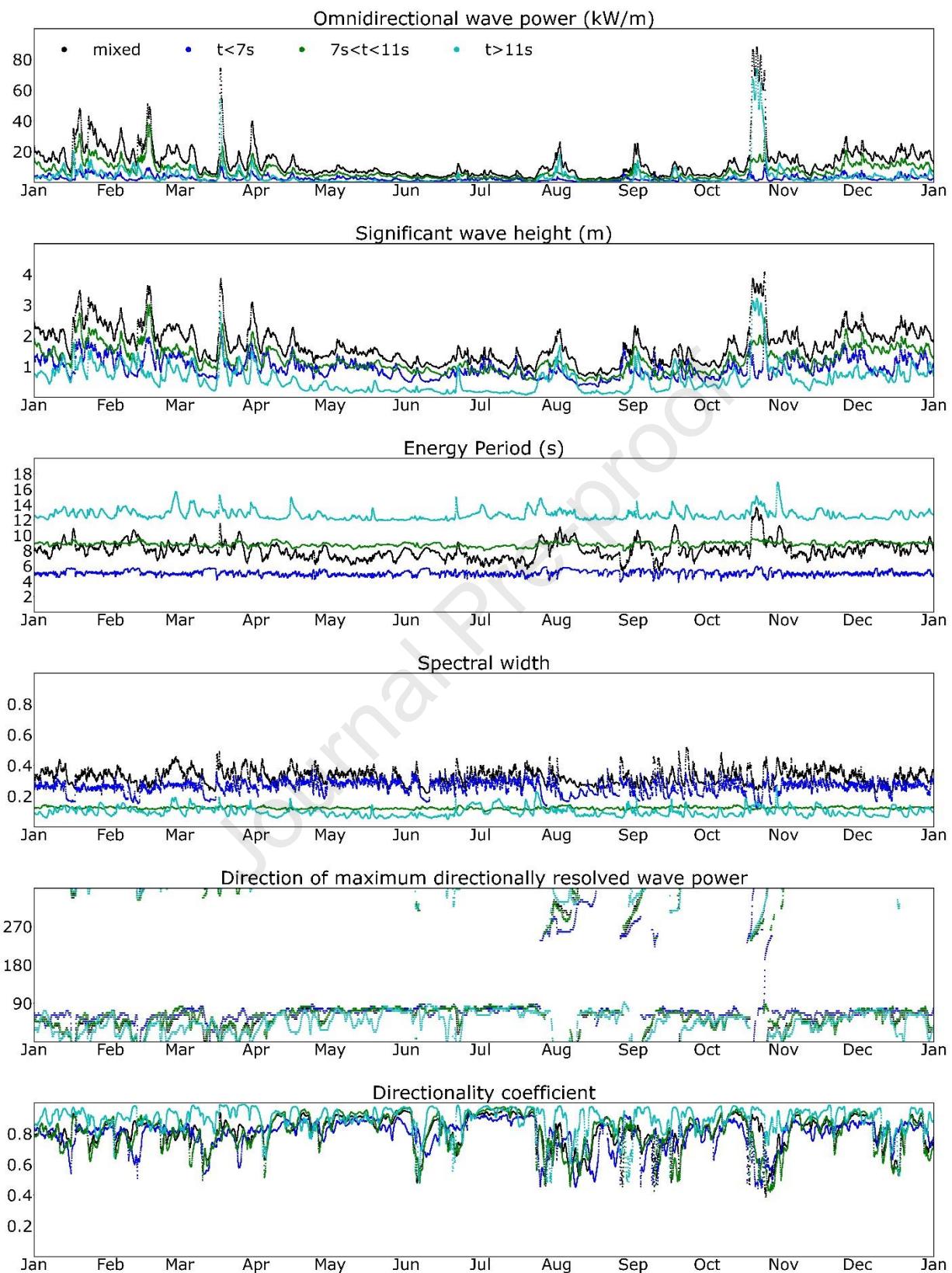


Figure 11. Time series of the six IEC wave energy resources parameters for the full and partitioned spectra from the hindcast at buoy #196 in 2017.

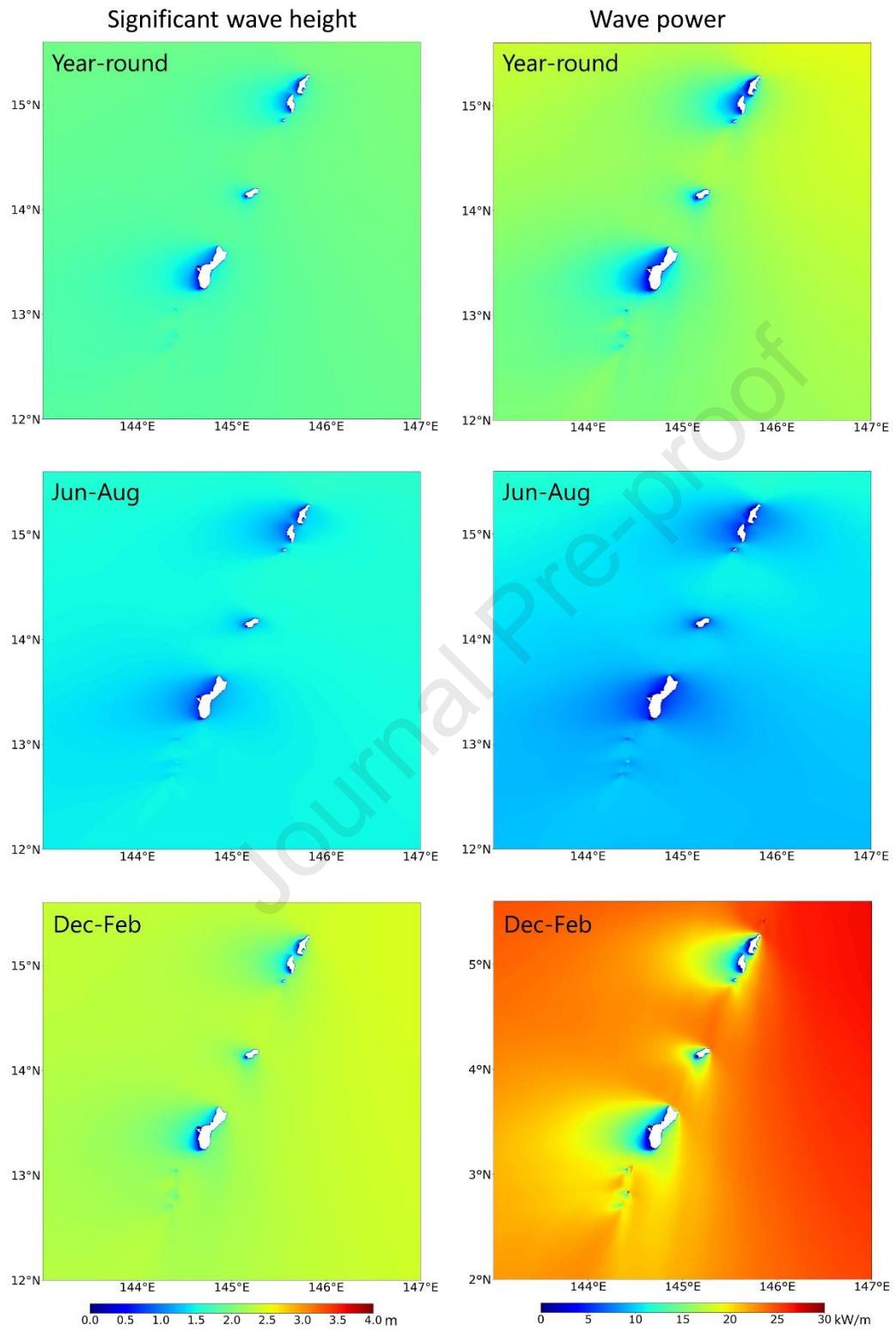


Figure 12. Year-round and seasonal averaged significant wave height and wave power from the 1979-2020 hindcast.

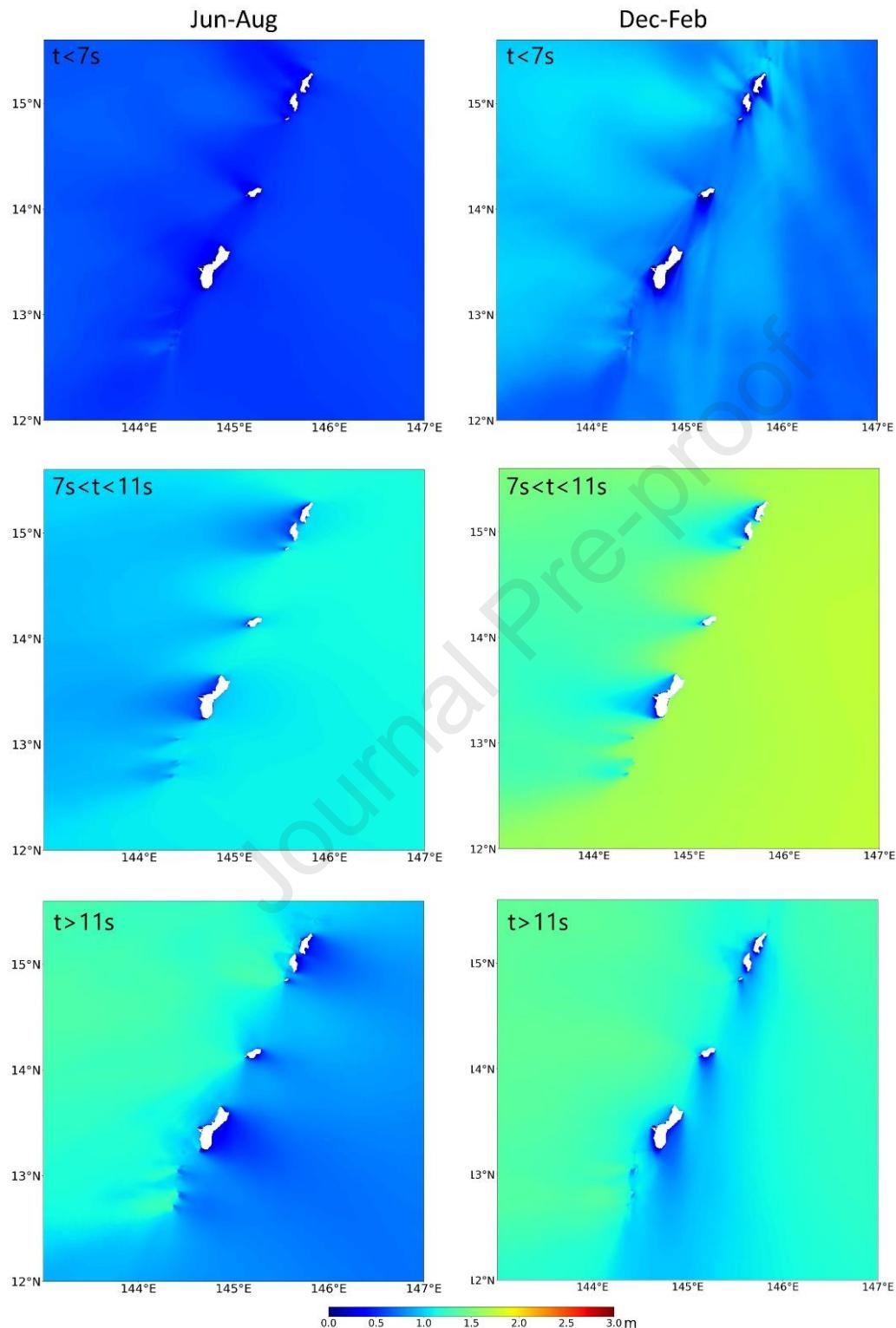


Figure 13. Seasonal average wave heights for spectral partitions of wind seas, wind waves, and swells from the 1979-2020 hindcast.

## Appendix

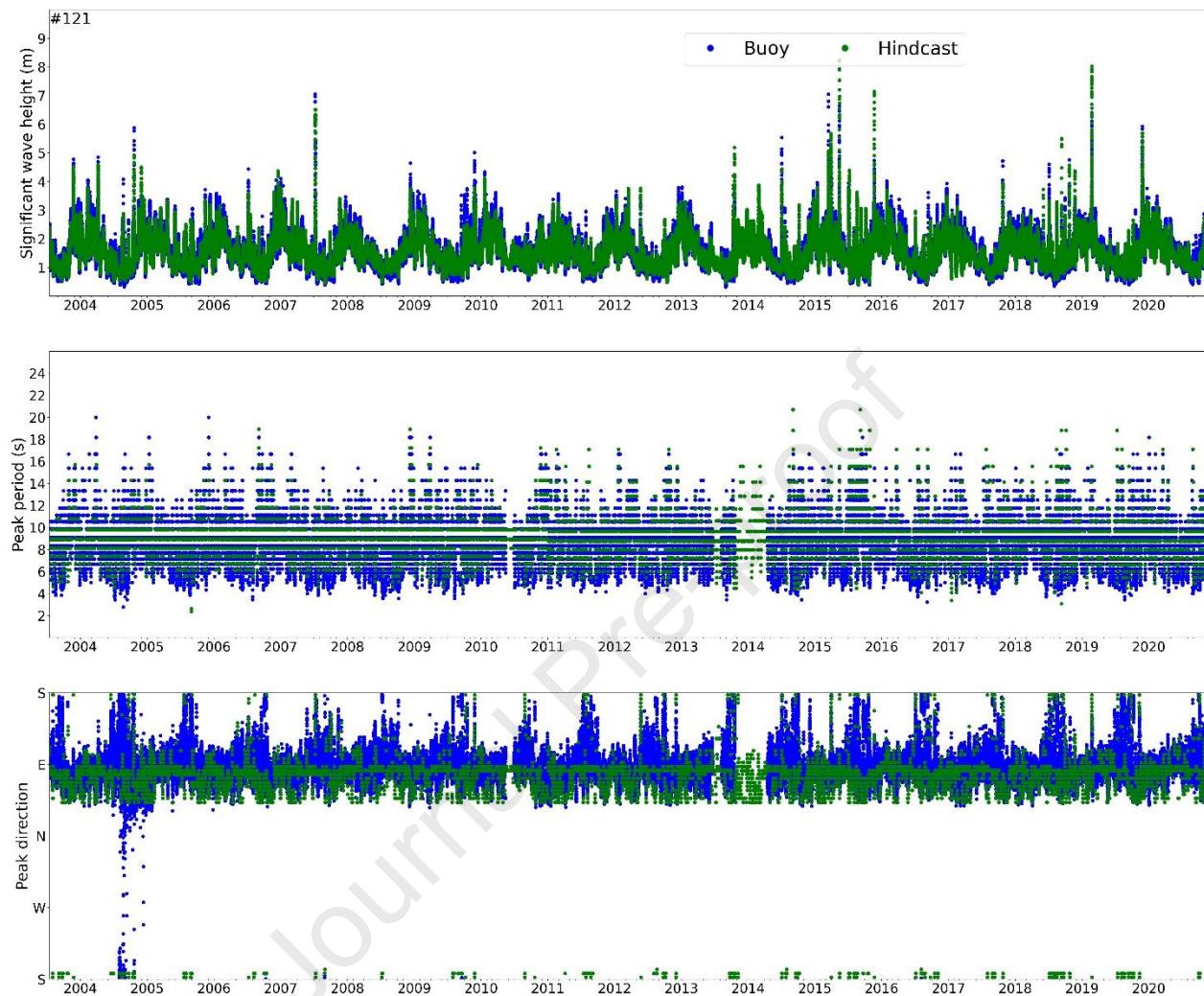


Figure A1. Time series of significant wave height, peak period, and direction from buoy #121 and the wave hindcast.

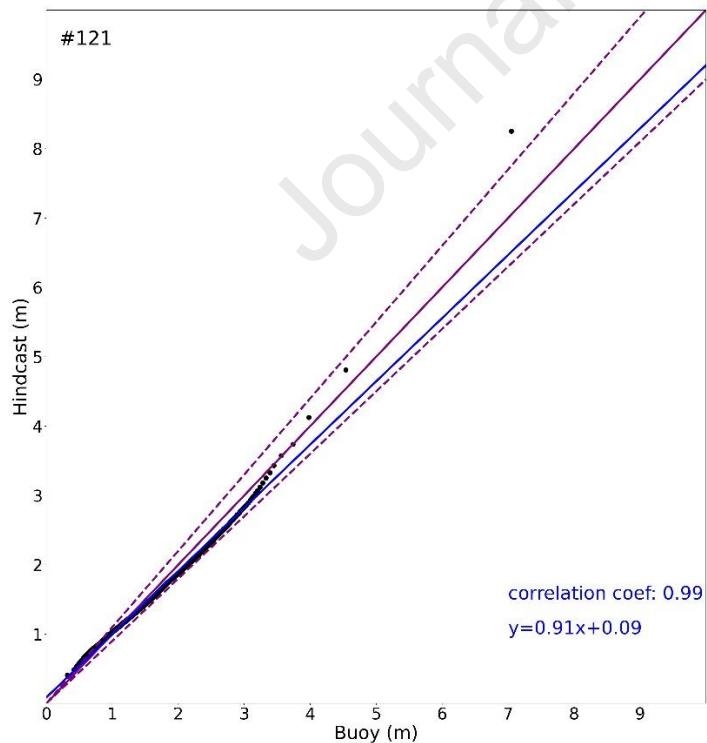
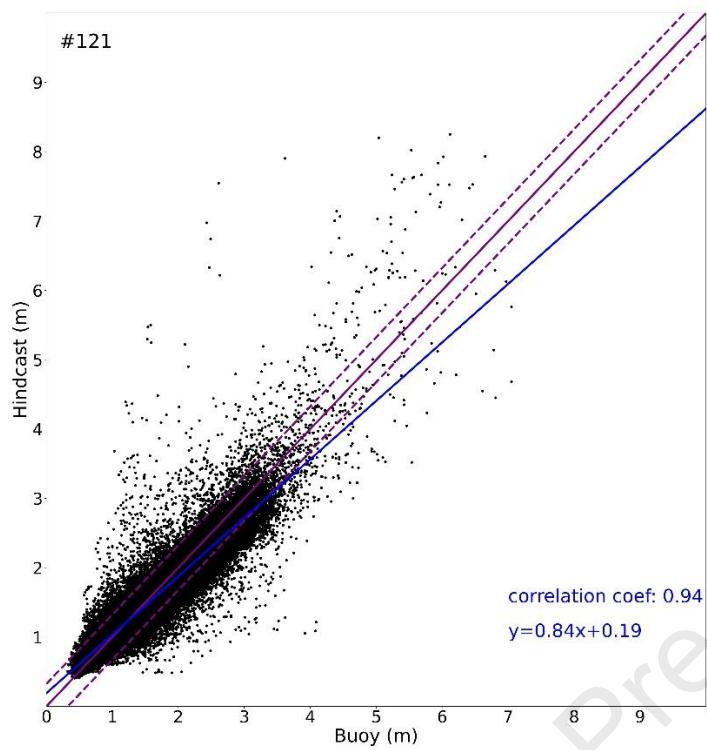


Figure A2. Scatter and Q-Q plots of measured and hindcast significant wave height at buoy #121.

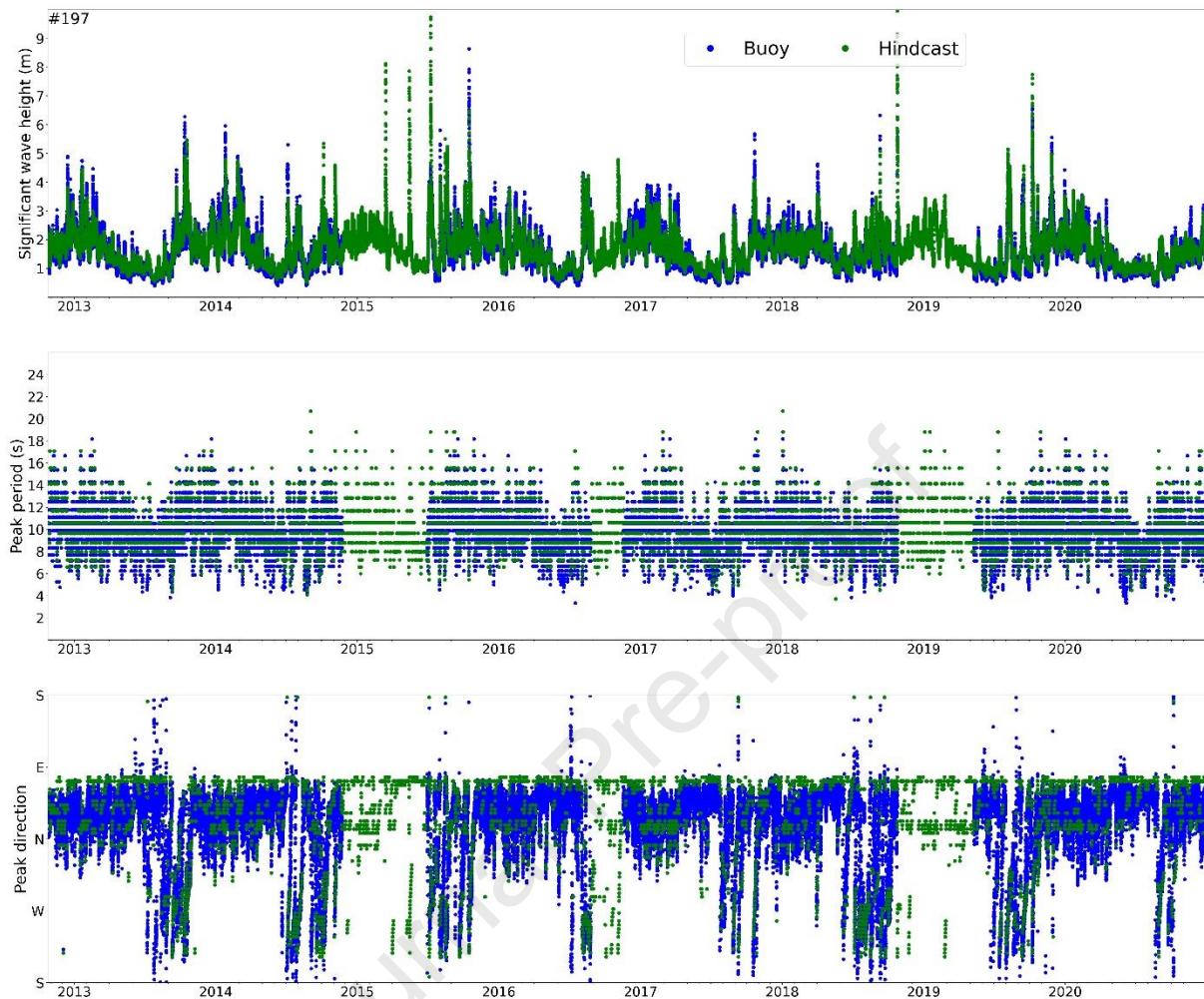


Figure A3. Time series of significant wave height, peak period, and direction from buoy #197 and the wave hindcast.

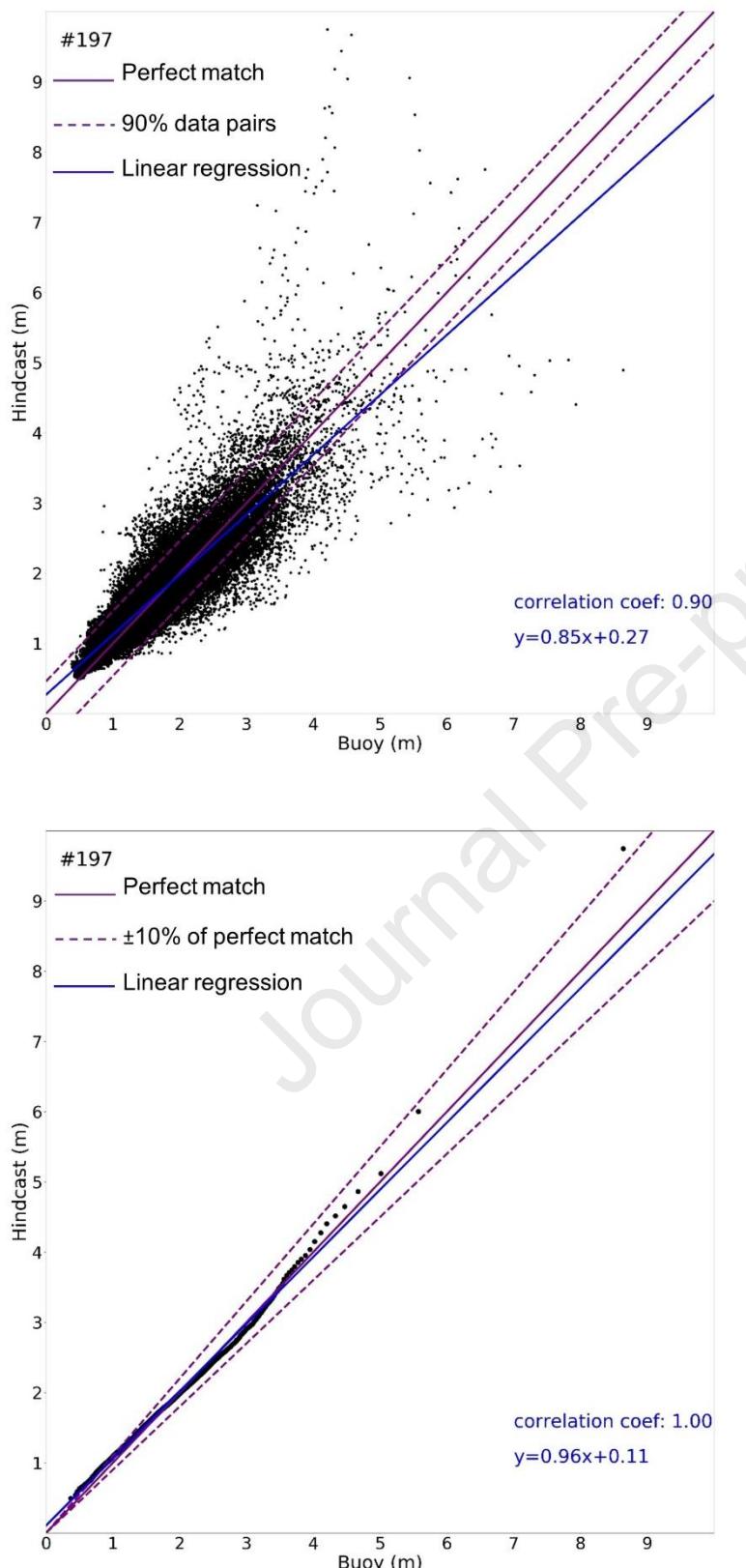


Figure A4. Scatter and Q-Q plots of measured and hindcast significant wave height at buoy #197.

**Highlights:**

- A 42-year wave hindcast by unstructured SWAN following IEC-TC 114 technical specifications.
- Hindcast validation of bulk and partitioned spectral energy for multimodal seas.
- Strong correlation between wave components to local and regional weather systems.
- Spatiotemporal descriptions of dominant wave components and energy levels.

**Declaration of interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests: