

1 **Joint impedance and facies inversion of time-lapse seismic data for improving monitoring**  
2 **of CO<sub>2</sub> incidentally stored from CO<sub>2</sub> EOR**

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22 **ABSTRACT**

23 Time-lapse seismic monitoring is an effective and proven technology for mapping the  
24 distribution of CO<sub>2</sub> in a subsurface reservoir. When injected CO<sub>2</sub> displaces other reservoir fluids,  
25 porous-medium properties are changed and thus the seismic impedance changes, causing time-  
26 lapse seismic amplitude differences in the injection zones. The analysis and interpretation of  
27 images created from these amplitude differences can provide information about reservoir  
28 architecture and the CO<sub>2</sub> migration within the reservoir. Incorporating seismic inversion and rock  
29 physics into the interpretation of time-lapse seismic data can considerably improve the modeling  
30 and monitoring to detect and assess the location of CO<sub>2</sub> over time. The joint inversion method  
31 presented in this paper has an integral representation of the geology in the inversion algorithm  
32 using elastic facies, which provides information about the spatial distribution of the geologic  
33 heterogeneities controlling the movement of fluids in the reservoir. The method was successfully  
34 applied to time-lapse seismic data from a mature oil field undergoing CO<sub>2</sub> enhanced oil recovery.  
35 The estimated seismic acoustic impedances and facies reflect the characteristics of individual  
36 geologic facies and fluid conditions of the reservoir subject to CO<sub>2</sub> injection. The probabilities  
37 estimated by the joint impedance and facies inversion for the reservoir's litho-fluid facies can be  
38 used for forecasting CO<sub>2</sub> saturation and pressure changes within the target reservoir.

39

40 **KEYWORDS:** seismic inversion, time-lapse seismic, 4D seismic, CO<sub>2</sub> storage, enhanced oil  
41 recovery, seismic monitoring, facies, rock physics, reservoir simulation.

42

## 43 1. INTRODUCTION

44 The success of geophysical methods in carbon capture and storage projects has been  
45 illustrated in the literature (e.g., Davis et al., 2003; Arts et al., 2004; Alnes et al., 2008; White,  
46 2009, 2011; Chadwick et al., 2010; O'Brien et al., 2010; Bhuyian et al., 2012; Dodds et al., 2013;  
47 Park et al., 2013, Carrigan et al., 2013; Tucker et al., 2016; Schmidt-Hattenberger et al., 2016;  
48 Burnison et al., 2017; Dupuy et al., 2017; Furre et al., 2017; Zhao et al., 2017; Roach et al., 2017;  
49 Salako et al., 2017, 2018; Wan et al., 2018; Macquet et al., 2019; Barajas-Olalde et al., 2020;  
50 Trautz et al., 2020; White et al., 2021; Pevzner et al., 2021). The seismic method is the most  
51 popular geophysical method mainly because of its high horizontal resolution. When the seismic  
52 method is applied to monitoring time-lapse changes of a reservoir subject to CO<sub>2</sub> injection, at least  
53 two data sets are acquired. The baseline data set is acquired before the CO<sub>2</sub> injection, and the  
54 monitor data sets are acquired after the CO<sub>2</sub> injection has been initiated. Maps of seismic amplitude  
55 differences from baseline and monitor surveys can be used to interpret the time-lapse changes in  
56 the reservoir due to CO<sub>2</sub> injection. Incorporating seismic inversion and rock physics into the  
57 interpretation can considerably improve the modeling and monitoring to detect stored CO<sub>2</sub> and  
58 assess the location of the CO<sub>2</sub> distribution in the reservoir over time.

59 Seismic inversion transforms seismic reflection data into reservoir elastic properties (i.e.,  
60 P-wave velocity ( $V_p$ ), S-wave velocity ( $V_s$ ), density ( $\rho$ ), acoustic impedance (AI), shear  
61 impedance (SI) that produced the reflections) (e.g., Sen and Stoffa, 2013). When seismic  
62 inversion is integrated with rock physics, elastic properties can be quantitatively transformed into  
63 geologic properties (e.g., porosity, lithology, fluids) for reservoir characterization (e.g., Mavko et  
64 al., 2009).

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66           The main challenges in seismic inversion are the different sources of uncertainties such as  
67 observation errors, limited bandwidth of the seismic data, seismic, well logs, data processing, and  
68 rock physics modeling errors (Bosch et al., 2010; de Figueiredo, 2018). Therefore, a Bayesian  
69 framework is a natural option for a seismic inversion (Bosch et al., 2010, and literature therein; de  
70 Figueiredo, 2018, and literature therein).

71           The solution of the Bayesian inverse problem is defined by a posterior distribution  
72 (Tarantola, 2005), which is conditioned to all the available observation data, the errors, and prior  
73 knowledge of the subsurface, such as spatial correlations, depositional trends, and background  
74 model (de Figueiredo, 2018).

75           In the last decades, extensive activities have been conducted to develop Bayesian algorithms  
76 that incorporate uncertainty quantification, geology, rock physics, and geostatistics in seismic  
77 inversion (e.g., Buland and Omre, 2003; Gunning and Ginsky, 2007; Bosch et al., 2010; Ulvmoen  
78 and Omre, 2010; Kolbjørnsen et al., 2016; Fjeldstad and Grana, 2018; Kemper and Gunning, 2014;  
79 Grana et al., 2017; de Figueiredo et al., 2018; Grana, 2018; Gunning and Sams, 2018; Waters et  
80 al., 2019; Gunning and Huntbatch, 2020).

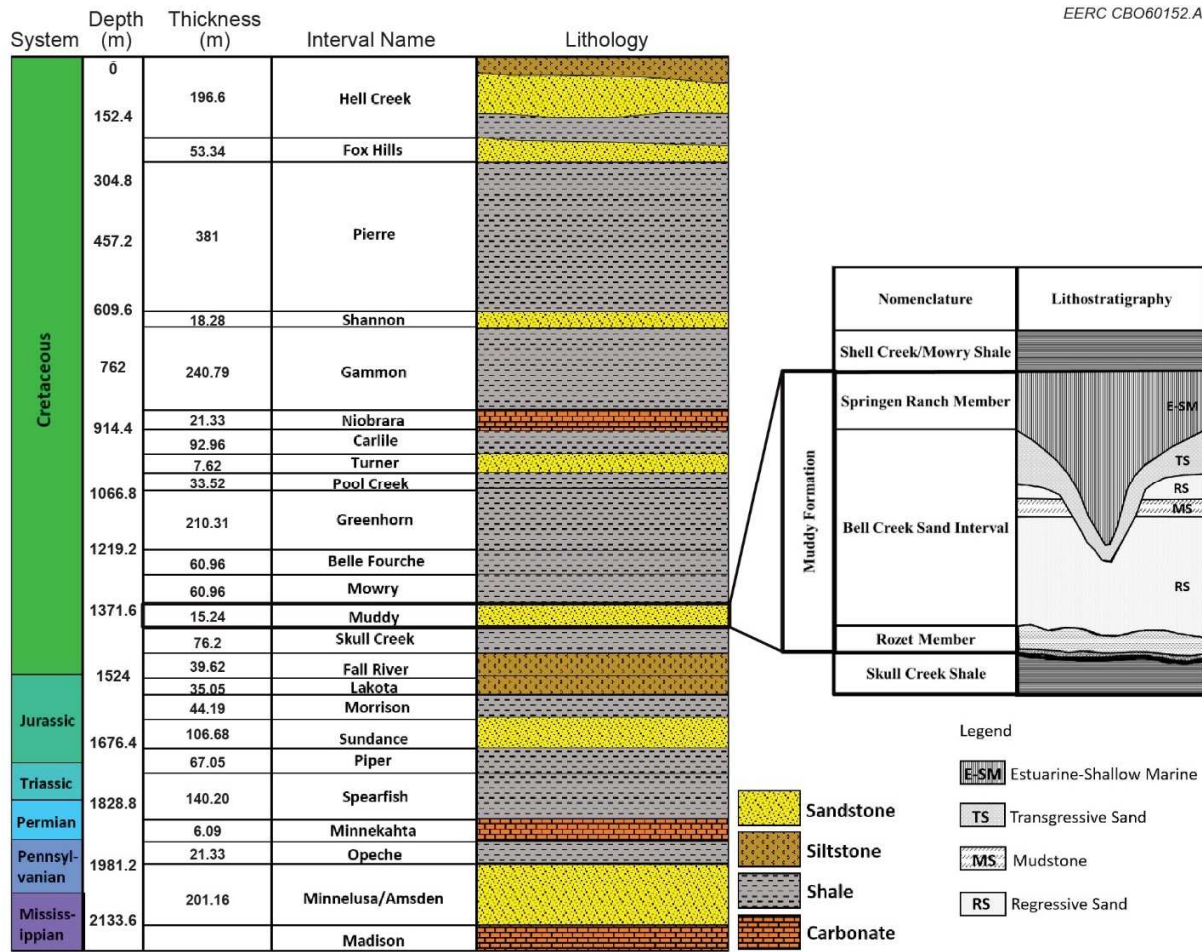
81           Instead of a two-step inversion in which elastic parameters are estimated first and then the  
82 elastic parameters are mapped to reservoir properties, one-step inversion (i.e., joint inversion) to  
83 reservoir properties is preferred (Connolly, 2017). Joint inversion algorithms allow the integration  
84 of a wide range of geological information and can predict facies directly from seismic reflection  
85 data (e.g., Rimstadt et al., 2012; Kemper and Gunning, 2014; Grana, 2018; Gunning and Sams,  
86 2018; Waters et al., 2019; Azevedo et al., 2020).

87           We apply a joint impedance and facies inversion (Ji-Fi) technology (Kemper and Gunning,  
88 2014) to two seismic data sets to monitor the time-lapse changes within the reservoir as a result of

89 CO<sub>2</sub> EOR (enhanced oil recovery) operations at a mature oil field. Ji-Fi estimates a three-  
90 dimensional (3D) facies classification scheme based on Bayesian principles, geological facies as  
91 priors, and powerful image-processing techniques. The oil field is in southeastern Montana near  
92 the northeastern edge of the Powder River Basin in the United States.

## 93 **2. GEOLOGICAL SETTING AND CO<sub>2</sub> EOR**

94 The oil field is a complex stratigraphic trap located structurally on a shallow monocline with  
95 a 1°–2° dip to the northwest and an axis trending southwest to northeast for approximately 32 km  
96 (Bosshart et al., 2015; Sharaf and Sheikha, 2021). The 6–9-m thick oil-bearing reservoir is part of  
97 the Lower Cretaceous Muddy Formation at a depth of approximately 1372 m (Fig. 1). This  
98 formation is dominated by high-porosity (15%–35%), high-permeability (150–1175-mD)  
99 sandstones (Saini et al., 2012). The Muddy Formation has been divided into three informal  
100 members based on Molnar and Porter’s study (1990), consisting of (from bottom to top) the Rozet  
101 Member, the Bell Creek Sand Interval, and Springen Ranch, (see Fig. 1). The Rozet Member  
102 overlies the Skull Creek Shale and consists of a 0.15–0.9-m sand marking the Skull Creek–Rozet  
103 contact and an overlying 1.5–3.0 m of shale/siltstone. The Bell Creek Sand Interval overlying the  
104 Rozet Member is described as having up to three different sand lobes (vertically). The Springen  
105 Ranch Member overlies the Bell Creek Sand Interval and is generally silty with some sand content.  
106 Overlying the Springen Ranch Member is the Shell Creek Shale. The overlying Lower Cretaceous  
107 Mowry Formation provides the primary seal, preventing migration to overlying aquifers and to the  
108 surface (Bosshart et al., 2015). The CO<sub>2</sub> injection target reservoir is a heterogeneous sand with  
109 fluvial channels, barrier bars, lagoonal, and nearshore facies leading to compartmentalized  
110 geobodies separated by permeability barriers and baffles, which lead to complex fluid  
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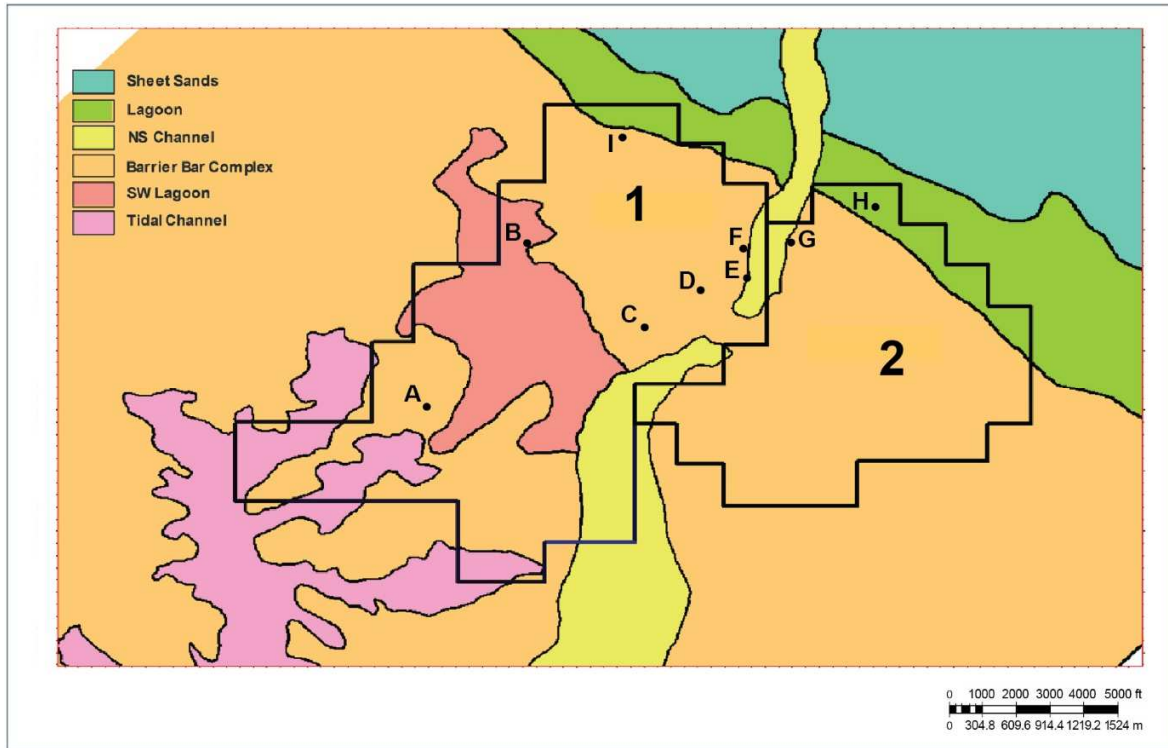
**Figure 1.** Stratigraphic column of the study area and simplified reservoir stratigraphy (modified after Burnison et al., 2017).

117 communication pathways between various compartmentalized geobodies (Bosshart et al., 2015).  
 118 The permeability barriers have been inferred from a) well penetration with little or no reservoir-  
 119 quality sandstone; b) wells with high water cut only a short distance updip from wells with notable  
 120 oil and/or gas production; c) significant variability in pressure, volume, temperature (PVT)  
 121 properties of produced hydrocarbons; d) considerable reservoir pressure differences, sometimes  
 122 over very short distances within the field; and e) 4D seismic difference maps and seismic attributes  
 123 (Spakiewicz et al., 1988; Honarpour et al., 1989; Molnar and Porter, 1990; Bosshart et al.; 2015,  
 124 Burnison et al., 2017; Salako et al., 2017).

125           Since the discovery of oil and gas in the 1960s, the oil field has undergone primary and  
126 secondary recovery that resulted in a gradual production decline leading to tertiary oil recovery  
127 implementation using CO<sub>2</sub> flooding (Hamling et al., 2013). CO<sub>2</sub> EOR is progressing through nine  
128 schematical development phases (regions within the oil field). Up to 1.6 million cubic meters of  
129 CO<sub>2</sub> per day is supplied by natural gas-processing plants located approximately 373 km from the  
130 study area (Burnison et al., 2017). This study focuses on Development Phases 1 and 2. Each phase  
131 was pressurized by water injection in preparation for CO<sub>2</sub> EOR. CO<sub>2</sub> injection began in Phase 1 in  
132 May 2013 and Phase 2 in December 2013.

133           Extensive work has been conducted at the Energy & Environmental Research Center  
134 (EERC) to build a geologic model incorporating 3D seismic data, well logs, and core data (Jin et  
135 al., 2016). The model includes depositional facies that can constrain petrophysical property  
136 distributions, enabling a better history match and more accurate predictive simulations of pressure  
137 response, injected/produced volumes, and saturation changes. Fig. 2 shows the facies interpreted  
138 in Development Phases 1 and 2 and an adjacent development phase within the Bell Creek Sand  
139 Interval Member.

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**Figure 2.** Geological facies interpreted in the study area: lagoon, barrier bar complex, North-South channel, and tidal channel. The polygons represent the study area (Development Phases 1 and 2). The boundaries of the map represent the area of the seismic surveys. The wells used in this study are represented by the letters A through I.

### 148 3. JOINT IMPEDANCE AND FACIES INVERSION

149 Rock physics and seismic inversion are fundamental tools of quantitative reservoir  
 150 characterization. In this process, the impedance and elastic parameters estimated by the inversion  
 151 of seismic data can be related to rock properties using rock physics models. Ji-Fi seamlessly  
 152 integrates rock physics and inversion using seismic facies (Kemper and Gunning, 2014).

153 Ji-Fi is model-based seismic inversion defined as a Bayesian approach, and it is expressed  
 154 as follows (Waters et al., 2019; Gunning and Huntbatch, 2020):

$$155 P(m, F | y) \sim L(y | m) * P(m | F) * P(F) \quad (1)$$

156 where  $m$  is the elastic model ( $V_p$ ,  $V_s$ ,  $\rho$ );  $F$  is the facies variable at discrete locations in the  
157 inversion volume;  $y$  is seismic data;  $L(y | m)$  is the likelihood of the synthetic seismic generated  
158 by the forward model;  $P(m | F)$  is the probability density function of the rock properties given the  
159 facies prior model; and  $P(F)$  is the facies distribution aligned to the stratigraphy.

160 As inferred from Equation 1, Ji-Fi solves for the distribution of elastic properties and facies.  
161 The low-frequency model (LFM) of elastic parameters required in seismic data inversion to  
162 account for the band limit and the lack of low-frequency signal in the seismic data is a product of  
163 the inversion.

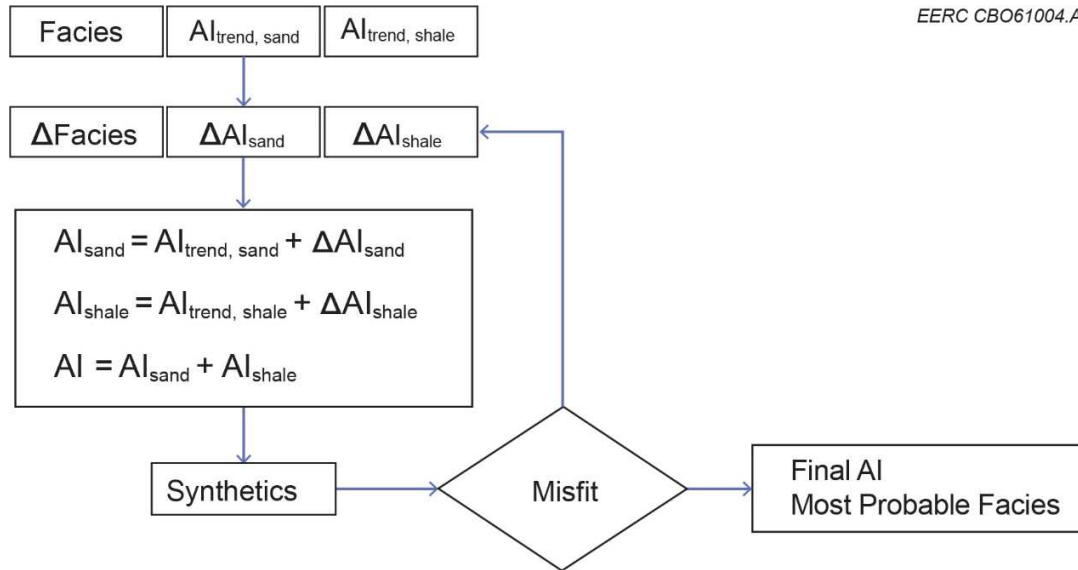
164 Contrary to standard inversion algorithms that build LFMs by interpolating well logs within  
165 a structural and stratigraphic framework, Ji-Fi uses a rock physics models defined by facies to  
166 build an LFM model per facies. The rock physics model for each of the facies is defined by a depth  
167 trend in terms of the elastic parameters  $V_p$ ,  $V_s$ ,  $\rho$ ,  $AI$ ,  $SI$ , and their relationships. Each  
168 component of the model has associated uncertainties. The rock physics models are assumed to  
169 apply over an entire study area and can be extrapolated to depths greater than the well control.

170 The inputs for the basic Ji-Fi workflow are the following: a set of estimated prior facies  
171 proportions within defined stratigraphic intervals, LFM models (depth-dependent rock physics  
172 models) for each expected facies and uncertainties for each facies, a set of seismic partial angle  
173 stacks, estimated noise in the partial-angle stacks, and seismic wavelets with estimated individual  
174 root mean square (RMS) error, which are used to solve for both the most likely facies model and  
175 its corresponding set of impedances (Waters et al., 2016). A detailed analysis of the requirements  
176 of the prior input information and its uncertainty for the Ji-Fi workflow is given by Waters and  
177 Kemper (2018) and Gunning and Sams (2018).

178           The selection of the (reservoir and nonreservoir) facies to invert for is based on the study  
179 area's lithologies and fluids, their relevance to a geologic interpretation, and their separation in the  
180 elastic domain. According to Equation 1, seismic data are used to constrain the distribution of the  
181 facies in the Ji-Fi scheme. Therefore, the facies must have some degree of separation in the elastic  
182 domain. Furthermore, a compromise must often be made on selecting the number of facies based  
183 on the degree of overlap in the elastic domain. Since these requirements might not be compatible  
184 with the idealized geologic facies definition, the facies should be called elastic facies (Sams et al.,  
185 2016).

186           The LFM models for each expected facies can be estimated from a depth trend analysis  
187 (DTA) of P-wave velocity ( $V_p$ ), shear-wave velocity ( $V_s$ ), density, and impedances derived from  
188 well logs, one-dimensional (1D) basin models, or analog fields. The LFM trends are referenced to  
189 a single datum, which reduces the need for detailed seismic interpretation. The crossplotted depth  
190 trends ( $V_p$  vs. density,  $V_p$  vs.  $V_s$ , AI vs. SI, AI vs.  $V_p/V_s$ ), and an assessment of their uncertainty  
191 are used for quality control. An essential feature of the Ji-Fi approach is that the LFM trends are  
192 not interpolated between wells. Therefore, the depth trends are not compromised by bias in  
193 interpreting seismic horizons or interpolation artifacts. Since the facies represented in the depth  
194 trends depend on rock physics models, more plausible elastic properties estimates can be obtained.

195           Fig. 3 depicts a diagram of a simplified Ji-Fi scheme with only the AI component and sand  
196 and shale facies. The Ji-Fi algorithm iteratively inverts seismic (trace by trace) for impedances  
197 given facies and then inverts for facies given an impedance model until convergence is reached.  
198 The convergence criterion is founded on matching synthetic seismic models with real seismic trace  
199 data until the misfit between the two is reduced to an acceptable level. It should be noted that  
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202 **Figure 3.** Simplified diagram of the Ji-Fi scheme for an acoustic inversion (modified after  
 203 Waters et al., 2016).

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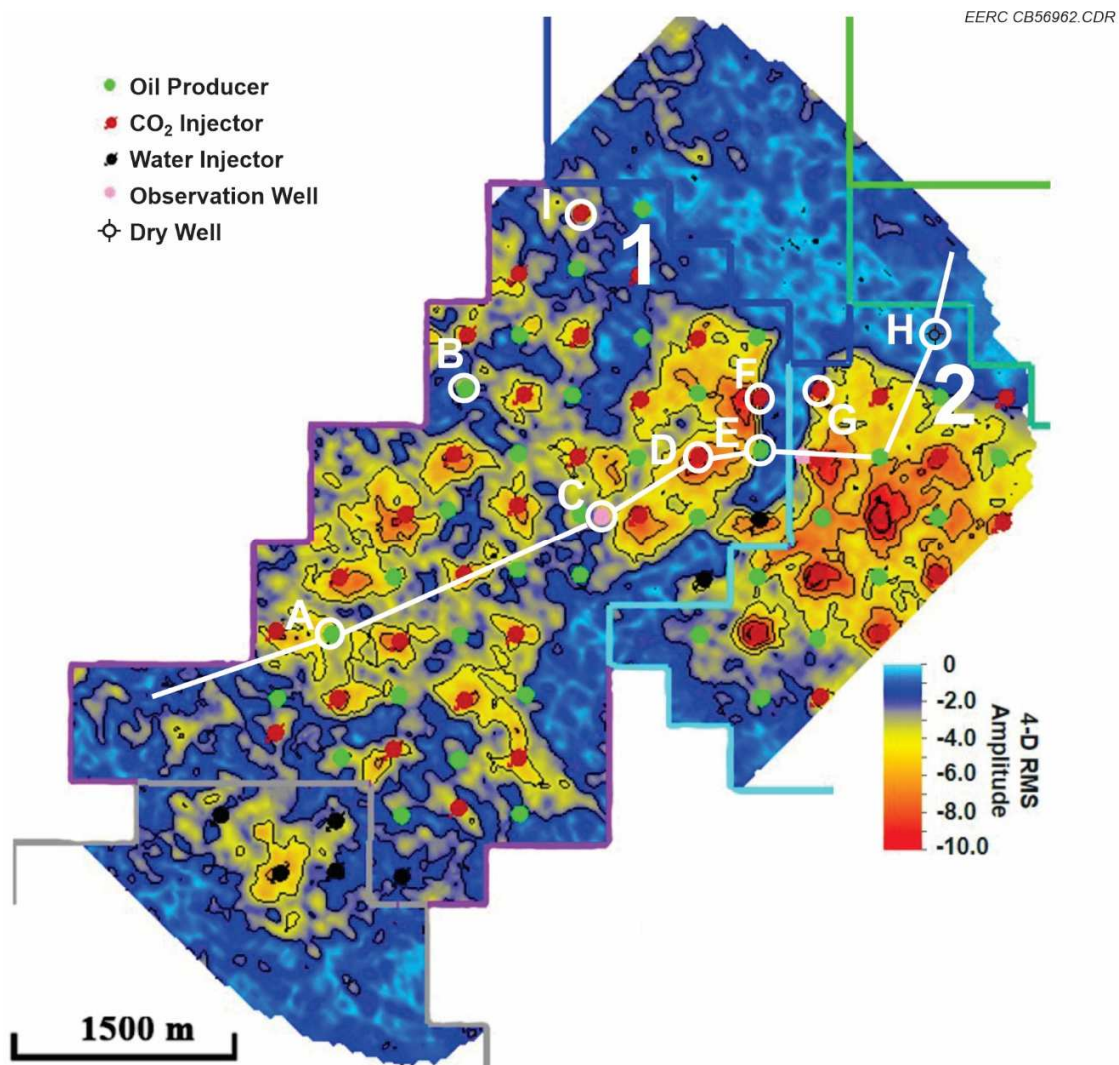
205 the facies result depends on the last set of impedance results and vice versa. This scheme implies  
 206 that better impedances lead to better facies estimates and vice versa at each workflow iteration.  
 207 The inversion can be constrained to promote facies continuity (lateral or vertical) and to honor  
 208 realistic geologic and hydrodynamic facies conditions (Kemper et al., 2017). The Ji-Fi workflow  
 209 outputs are  $V_p$ ,  $V_s$ , density, and the most probable facies, with standard deviations and  
 210 probabilities. Kemper and Gunning (2014) and Gunning and Huntbatch (2020) report details of  
 211 the Ji-Fi algorithm.

## 212 3.1 Input data

### 213 3.1.1 Seismic data and horizons

214 The Ji-Fi approach requires seismic data for inversion and seismic horizons to constrain the  
 215 time and depth model relationships. The 3D seismic data sets to monitor the time-lapse changes  
 216 within the reservoir due to the CO<sub>2</sub> EOR operations correspond to the September 2012 baseline  
 217 survey acquired prior to CO<sub>2</sub> injection activities and the October 2014 monitor survey acquired

218 after the start of commercial CO<sub>2</sub> EOR operations. The 2014 monitor survey overlaps the 2012  
219 baseline survey at Development Phase 1 and part of Development Phase 2. At the time of  
220 acquisition of the monitor survey, the Phase 1 and 2 areas had undergone CO<sub>2</sub> injection for 18 and  
221 10 months, respectively. A summary of repeat 3D seismic data acquisition and the data processing  
222 workflow is included in Appendix A of Salako et al. (2018). The seismic reflection of the 6–9-m-  
223 thick Bell Creek sand reservoir is that of a thin-bed response, composed of entering and exiting  
224 reflections from the Springen Ranch and Skull Creek intervals, respectively, near the 11-ms  
225 resolution limit of the embedded wavelet derived from an amplitude spectrum with a bandwidth  
226 of approximately 10–48 Hz estimated from a 500-ms window centered in the zone of interest  
227 (Burnison et al., 2014, 2017). Sets of baseline (2012) and monitor (2012) angle stacks (reverse  
228 polarity, common midpoint, and angle gathers stacked over 10° angle ranges to 5°, 15°, 25°, 35°  
229 midpoints) were used as input for our study. A seismic cube obtained from the stacking of the  
230 prestack time-migrated (PSTM) offset gathers was used to interpret time and depth horizons. The  
231 interpreted horizons used to build the model for the seismic inversion are the following: Horizon  
232 1: a strong event above the reservoir (this horizon was used as a datum for the DTA); Springen  
233 Ranch: top of the reservoir; Skull Creek—bottom of the reservoir and low boundary of Bell Creek  
234 (LBBC)—a strong event below the reservoir. Fig. 4 shows a time-lapse amplitude RMS difference  
235 map, calculated between the Springen Ranch and Skull Creek. The difference is computed as the  
236 monitor (2014) minus baseline (2012) survey. The high sensitivity of the seismic



237

238 **Figure 4.** Time-lapse amplitude difference map of monitor (2014) minus baseline (2012)  
 239 calculated as a summation of RMS amplitudes at the reservoir interval between the  
 240 Springen Ranch and the Skull Creek horizons. The wells in Development Phases 1 and 2  
 241 are shown with the following symbols” Green dot – oil producer, red dot – CO<sub>2</sub> injection  
 242 well, black dot – water injector, pink dot – observation well, black circle with lines – dry  
 243 well, white circle – wells used in this study, white line – cross section along some of the  
 244 wells used in this study (modified after Salako et al., 2017).  
 245

246 signal to the changing fluid properties within the reservoir induced by CO<sub>2</sub> injection is clearly  
 247 observed in the 4D seismic amplitude response (warm colors). The interpretation of the map  
 248 (Burnison et al., 2017; Salako et al., 2017) indicates the permeability barriers in Phases 1 and 2  
 249 associated with the little time-lapse response to low-permeability deposits resisting fluid flow and

250 pressure dissipation. Burnison et al. (2017) and Salako et al. (2017) interpreted the apparent high  
251 seismic amplitude response to the CO<sub>2</sub> injection in Phase 2 with a shorter period of CO<sub>2</sub> injection  
252 (10 months) than in Phase 1 (18 months) to greater change in sustained pore pressure that occurred  
253 in Phase 2 before CO<sub>2</sub> injection.

### 254 **3.1.2 Well data**

255 After data conditioning, eight wells with complete log suites for rock physics modeling and  
256 quantitative well tie were used in the Ji-Fi workflow. The location (and type) of these A (producer),  
257 B (producer), C (observation), D (CO<sub>2</sub> injector), E (producer), F (CO<sub>2</sub> injector), G (CO<sub>2</sub> injector),  
258 and H (dry hole) wells in the study area is shown in Figs. 1 and 4 Three of these wells (A, C, and  
259 G) were extensively geologically sampled with sidewall and whole core.

### 260 **3.1.3 Geologic data**

261 Geologic analysis of the well data included a description of whole cores. Sidewall cores lab  
262 work was done, including x-ray diffractometry (XRD), scanning electron microscopy (SEM),  
263 porosity and permeability measurements, and petrology. The analysis of thin sections and SEM  
264 mineral maps reported by Eylands et al. (2013) indicates that most of the sidewall cores from the  
265 C well are grain-supported and that the grain cement is dominantly clay. A small volume of quartz  
266 cement is also present. Braunberger et al. (2013) studied core from the Bell Creek reservoir and  
267 reported that carbonate may also be a locally important cement. The significant differences in the  
268 porosity and permeability of the Muddy Formation sand intervals of the Bell Creek reservoir can  
269 be attributed to clay concentration and distribution within the samples (acting as a contact cement).  
270 Decreased porosity and permeability can be attributed to higher clay concentrations, especially  
271 kaolinite, filling the pore spaces between sample grains (Eylands et al., 2013).

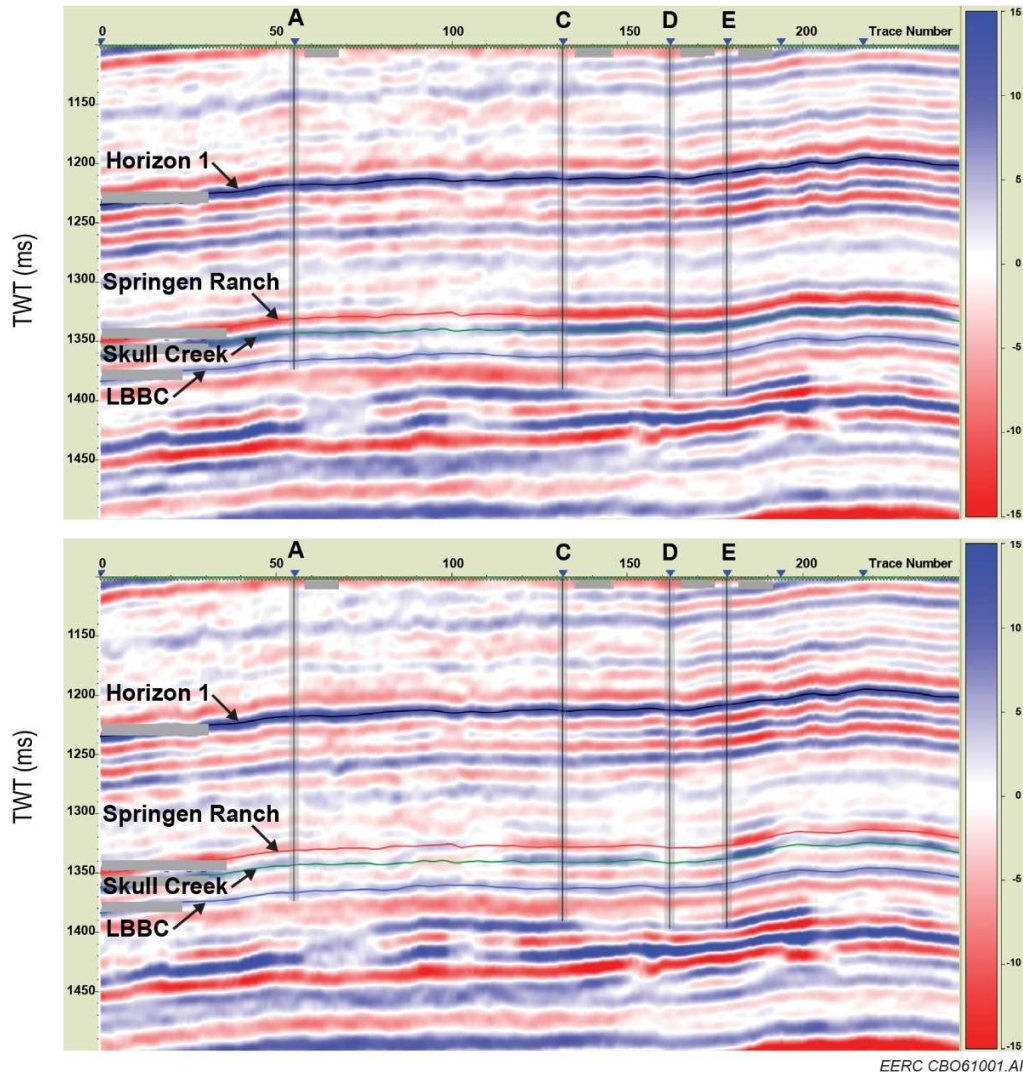
272 **3.2 Data conditioning**

273 **3.2.1 Seismic data**

274 PSTM offset gathers were converted to the partial-angle stacks required for the Ji-Fi prestack  
275 approach. Four partial-angle stacks were created: near (0–10 degrees), mid (10–20 degrees), far  
276 (20–30 degrees), and ultrafar (30–40). Although the original time-lapse conditioning scheme was  
277 satisfactory, additional data conditioning was performed. First, a 2–4–65–80-Hz finite impulse  
278 bandpass filter was applied to each partial-angle stack to remove high-frequency noise. After that,  
279 a robust time-shift estimation between near, mid, and far angles of the baseline and monitor data  
280 sets was applied to remove normal moveout correction (NMO correction) residual time-shifts  
281 included in the data processing. The time-shift estimation parameters were a 175-ms correlation  
282 window with a time-shift averaging over a three-trace radius followed by a 90-ms time-shift  
283 smoothing window. Fig. 5 depicts the results of the data conditioning on baseline (top) and monitor  
284 (bottom) stacked cross sections shown as a white line in Fig. 4. The A, C, D, and E wells and  
285 horizon 1, Springen Ranch, Skull Creek, and LBBC horizons are also depicted in the figure as  
286 reference.

287 **3.2.2 Well log data**

288 To assess any issues in the data acquisition and processing of the logs that can affect the  
289 inversion results, quality control of the available well logs was conducted. Any spikes and noise  
290 bursts were replaced, and gaps were filled using either spline functions away from the zone of  
291 interest or more robust rock-physics-based methods within the zone of interest. Histograms were  
292 used to assess the log distribution between the three wells A, C, and G with the complete sets of  
293 elastic logs (P-wave sonic, S-wave sonic, and density) in the Bell Creek interval from field data  
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295  
 296 **Figure 5.** Cross section of baseline (top) and monitor (bottom) stacks after data  
 297 conditioning. The cross section intersects the A (producer), C (observation), D (CO<sub>2</sub>  
 298 injector), and E (producer) wells. Horizon 1, Springen Ranch and Skull Creek, and the  
 299 LBBC horizons are shown as a reference. The cross sections do not include the part of the  
 300 line where the H (dry) well is located.  
 301

302 acquisition. Consistent results were obtained between wells, except for the density in the A well,  
 303 which had a wider distribution than expected. However, the values are still in the range for  
 304 sandstone reservoirs based on Appendix A1 of Mavko et al. (2009).

305 Since the other available wells do not include fundamental logs with elastic properties for  
 306 the application of Ji-Fi, a combination of rock physics methods with empirical relationships were

307 used to estimate  $V_s$  for B, D, E, F wells, and density for D, E, F wells. A hybrid method was used  
308 to estimate  $V_s$  for the defined working intervals: a modified Gassmann approach (Mavko et  
309 al., 2009) for the interval representing the reservoir. This approach is based on the linear form of  
310 Gassmann's equation to estimate  $V_s$  when the rock is hydrocarbon-saturated. The linear form of  
311 Gassmann's equation expresses the change in bulk modulus due to fluid substitution as a function  
312 of the intercept porosity parameter (PoR) (Mavko and Mukerji, 1995). As the fluid modulus is  
313 typically much less than the mineral modulus and the Gassmann's equation predicts no change in  
314 shear modulus during fluid substitution, the linear form of Gassmann's equation can be expressed  
315 as a function of the P-wave modulus and used to estimate the sonic velocity ( $V_p$ ) for a brine-  
316 saturated sample from the measured log data in a hydrocarbon interval. The Greenberg–Castagna  
317 approach (1992) can then be used to calculate the associated saturated  $V_s$ . The final step is to use  
318 the conventional Gassmann's equation to perform a fluid substitution of the brine-saturated  
319 properties to reestimate the  $V_p$ ,  $V_s$ , and density of the in situ saturation. The generated  $V_s$  product  
320 is the output of the process. Input data for this approach are porosities, in situ fluids, and  $V_p$  to  
321 predict water-saturated  $V_p$ – $V_s$ –density and then substituted for in situ fluids. Empirically  
322 calibrated  $V_p$ – $V_s$  relationships were used to predict the overburden (from Horizon 1 to Shell Creek  
323 markers) and the underburden (from Rozet to Skull Creek markers) working intervals. These  
324 relationships are given by the linear relationship:  $V_s = V_p * x + c$ . The constant values  $x$  and  $c$  can  
325 be derived from a linear fit using crossplots. In the case of density prediction, while linear  $V_p$ –  
326 density relationships calibrated [1)  $(\text{Rho} = V_p + \text{constant}) / \text{constant}$ ; 2)  $\text{Rho} = V_p * \text{constant} +$   
327  $\text{constant}$ ; the constants are derived from crossplots] per working interval were used for working  
328 intervals other than the Bell Creek interval, a mix of two methods scaled by porosity was used for  
329 the Bell Creek interval. The first method uses mineral volumes, calibrated mineral densities,

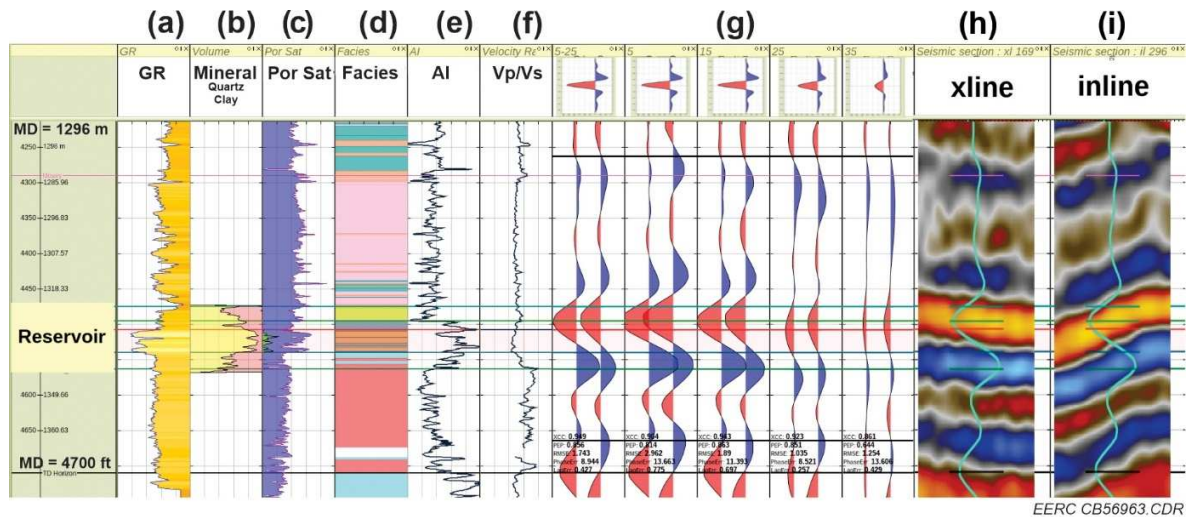
330 porosity, fluid logs, and fluid densities to predict log densities. In the second method, shale density  
331 is locally calibrated between the Springen Ranch and Skull Creek markers for fluid substitution  
332 (density of shale = 2.338) from a fluid subshale average set previously calculated. Attempts were  
333 made to predict Vp for some wells. However, it was impossible to define a method to predict all  
334 the working intervals for all the wells correctly. After data conditioning, the A, B, C, D, E, F, G,  
335 and H wells were considered in the Ji-Fi workflow.

### 336 **3.3 Well tie and wavelet extraction**

337 Well to seismic data (full stacks and partial stacks) ties were performed to estimate the  
338 wavelets for the inversion. The A, B, C, G, and H wells were selected for this process as the original  
339 and predicted logs were adequate in the depths from Horizon 1 to the LBBC horizons, which were  
340 considered as the inversion model limits. Since D and E wells did not fulfill this requirement, they  
341 were used for calibration purposes. The effective wavelet length was between 80–128 ms, and the  
342 analysis window was approximately 400 ms. An initial bulk shift was made to maximize  
343 correlation over the analysis window with a statistical 180-degree phase rotation wavelet or with  
344 a group average Bayesian wavelet (simultaneous multiwell wavelet estimation method). Next, the  
345 Roy White method (White, 1997) was used on the selected wells and surrounding seismic data.  
346 After estimating the deterministic wavelets at each well, an average wavelet was created and used  
347 to make a final bulk shift at all the well locations.

348 Comparable results to the well-to-full stack seismic data ties were obtained with partial  
349 stacks following the same methodology for full stacks mentioned above. However, in this case,  
350 one wavelet is estimated for each partial stack. While an averaged (Roy White) wavelet was  
351 estimated for the near angles, averaged Bayesian wavelets were estimated for the mid and far

352 angles. An example of the synthetic-to-seismic tie for the C well using baseline seismic data is  
 353 shown in Fig. 6. An excellent-quality tie was observed for the eight wells used in this study.  
 354



355  
 356  
 357 **Figure 6.** Well-tie and AVO QC at the C well: a) gamma ray with light fill identifying  
 358 sands; b) volume fraction set of quartz (yellow) and clay (pink); c) porosity with fluid  
 359 saturation (trace oil); d) petroelastic facies classification; e) AI; f) Vp/Vs ratio; g) synthetic-  
 360 to-seismic ties (full stack, 5°, 15°, 25°, and 35° angle stacks). The range of angle stacks is  
 361  $\pm 5^\circ$  from indicated angle with wavelets in the header (equal amplitude ranges for wavelets);  
 362 h) crossline; and i) in-line section QC with synthetic overlay. It should be noted that the  
 363 reverse-polarity wavelet convention (an increase in impedance results in a negative  
 364 reflection) was used in this study.  
 365

#### 366 4. ROCK PHYSICS

367 Rock physics analysis can estimate changes in the reservoir elastic properties due to changes  
 368 in stress, pore pressure, fluid type, saturation, and mineralogy variations.

##### 369 4.1 Rock properties

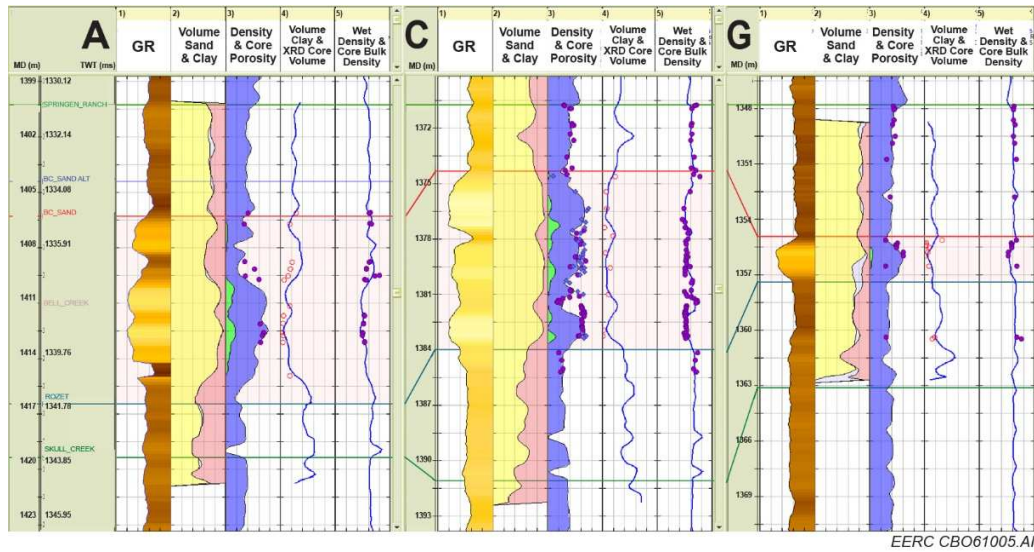
370 As the A, C, and G wells originally had shear sonic logs, they were used for rock physics  
 371 analysis. Sidewall cores in the A, C, and G wells were used to link wireline measurements with  
 372 the study area's geology. A close agreement of bulk density, porosity, and lithology volumes with  
 373 corresponding values from wireline logs was expected after allowing for differences due to depth

374 uncertainty, resolution, and sampling bias. In general, a  $\pm 0.61$ -m depth uncertainty for sidewall  
375 cores in the three wells was observed. The following comparisons were conducted, as shown in  
376 Fig. 7:

- 377 • Bulk densities derived from core grain density and core porosity. The results were  
378 compared with density log, good agreement between both types of density measurement  
379 was observed.
- 380 • Measured core porosities vs. various porosity logs (Schlumberger RST Advisor, neutron,  
381 sonic, density). The best agreement was found with density porosity (DPHI) on a  
382 sandstone matrix. This finding is consistent with the results reported by Sutherland (2011)  
383 and Pasternack (2009). A density of 2.65 g/cm<sup>3</sup> was used to calculate the sandstone  
384 density porosity curve. This value is close to 2.60–2.63 g/cm<sup>3</sup> for the Bell Creek clay  
385 grain density reported by Pasternack (2009).
- 386 • Mineral types and volumes estimated from XRD analysis of sidewall core samples vs.  
387 results from Schlumberger's RST SpectroLith estimates WCAR (carbonate), WQFM  
388 (quartz, feldspar, mica), and WCLA (clay) derived from wireline logs. The log data are  
389 consistent with the XRD data. The comparison suggests that the WCLA log provides a  
390 reasonable estimate of the weight fraction of clay. It can also be inferred from the XRD  
391 results that kaolinite is the dominant clay in the Bell Creek interval. The three high-quality  
392 logs used to compare with sidewall cores were considered the initial reference for rock  
393 physics analysis and inversion tests.

394

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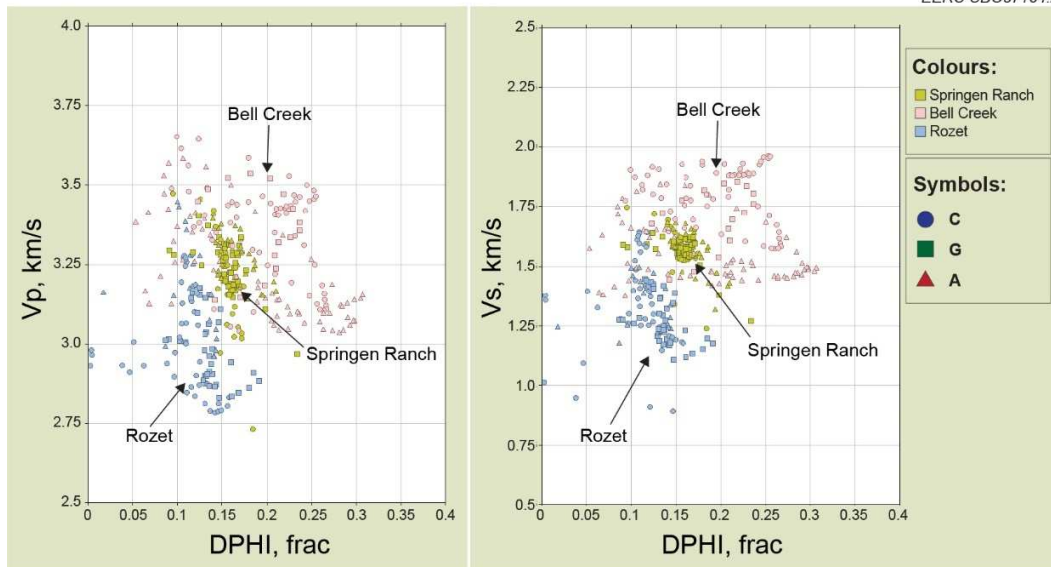


396

397 **Figure 7.** Comparison of core measurements of porosity, volume clay, volume carbonate, and bulk  
 398 density with log data from the A, C, and G wells. The tracks from left to right in each well are  
 399 gamma ray (GR), volume (yellow sand, pink clay), density porosity (DPHI) with core porosity,  
 400 volume clay (orange), or carbonate (blue) with core volume from XRD, wet density with core bulk  
 401 density. Core bulk density was calculated from grain density and porosity, assuming the fluid  
 402 density of 1.0 g/cm<sup>3</sup>.  
 403

#### 404 4.2 Muddy Formation

405 The Muddy Formation comprises three members: Springen Ranch, Bell Creek, and Rozet.  
 406 All three members are seismically thin and well below tuning. In combination, they reach a  
 407 sufficient thickness that seismic reflections are associated with the approximate top and base of  
 408 the Muddy Formation. Hence, all the members' elastic properties contribute to the apparent  
 409 impedance contrasts sensed by the reflections. After fluid substitution to replace any oil saturation  
 410 with brine, Fig. 8 shows that the elastic properties of all three units overlap to some extent. The  
 411 Bell Creek at any porosity has, in general, a faster V<sub>p</sub> and V<sub>s</sub> than Rozet. The Rozet has a smaller  
 412 range of porosities and is generally slower than the Bell Creek. Springen Ranch properties are near  
 413 the transition between Bell Creek and Rozet and are entirely contained within the data cloud  
 414 formed by the Bell Creek and Rozet.

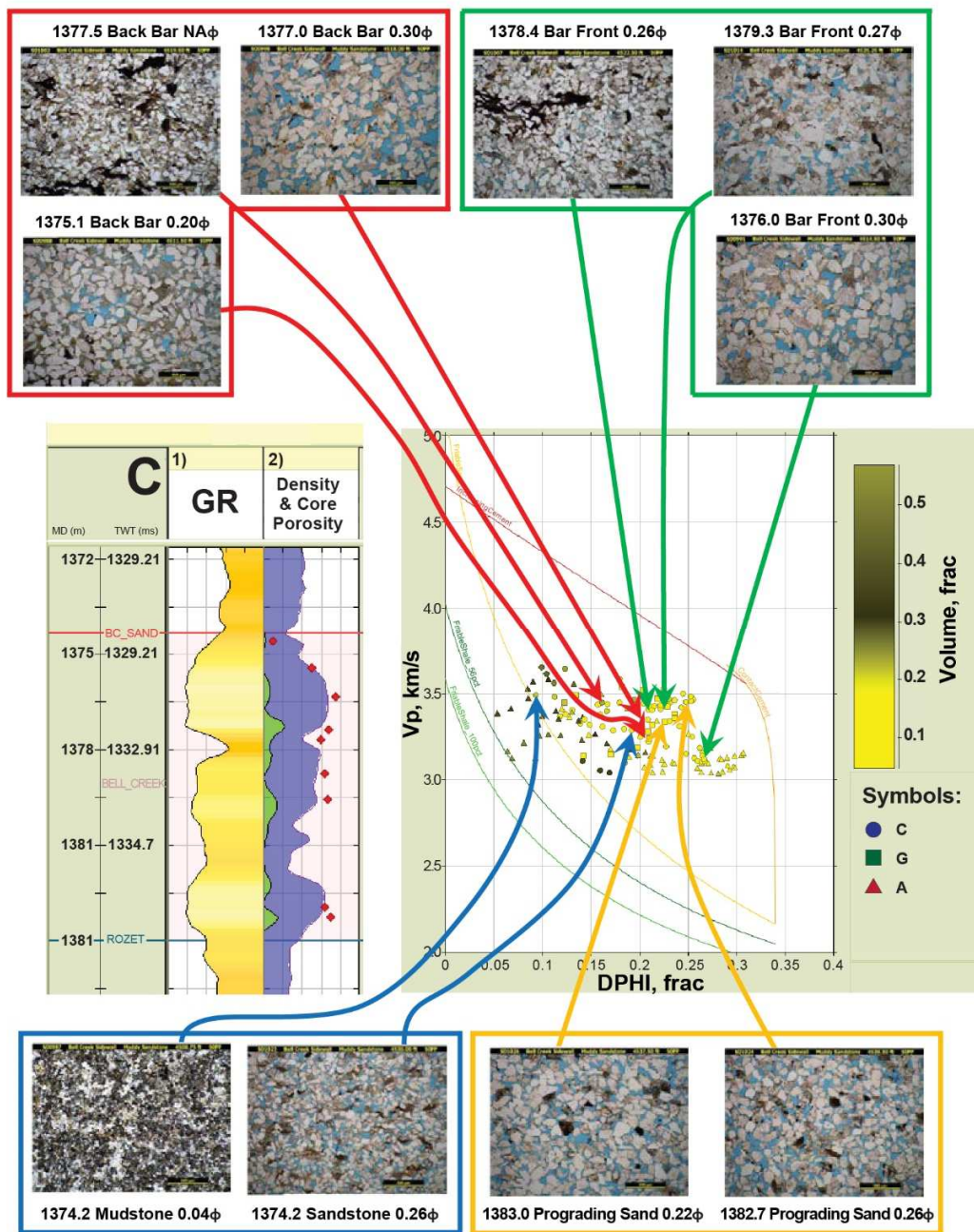


415  
 416  
 417 **Figure 8.** Muddy Formation. Springen Ranch (yellow) plots within the Bell Creek (pink). Rozet  
 418 (blue) has different properties that cause slower and lower porosity than the rest of the Muddy.  
 419 Octagon: C well; square: G well; triangle: A well.  $V_p$  and  $V_s$  represent the velocity conditions  
 420 after replacing any oil saturation with brine by fluid saturation.  
 421

422 The relationship between density and velocity in the Bell Creek reservoir is complex. In  
 423 general,  $V_p$  increases, and  $V_s$  decreases as clay volume increases. Moreover, high clay volume  
 424 samples have higher densities and lower velocities. The increase of bulk modulus and decrease in  
 425 shear modulus with an increase in volume of clay ( $V_{clay} < 0.3$ ) are much less apparent than it is  
 426 with velocity. Therefore, it can be assumed that the change in velocity is mostly density-driven.

### 427 4.3 Rock physics analysis and diagnostics

428 Rock physics analysis and diagnostics are used to relate seismic characteristics to geology.  
 429 The connection can be observed by plotting thin-section images from cores relative to their log  
 430 properties. The results can provide a useful reference for the rock physics analysis. Fig. 9 shows  
 431 thin sections from C well on a  $V_p$  vs. density porosity (DPHI) plot of the A (triangles), C (circles),  
 432 and G (squares) wells from the reservoir section colored by clay volume. The thin sections'  
 433



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435 **Figure 9.** Core data from Bell Creek reservoir at the C well are placed on a  $V_p$  (wet) vs. density  
 436 porosity (DPHI) plot colored by clay volume with data from A (triangles), C (circles), and G  
 437 (squares) wells. The curves represent the effective medium models: contact cement (dark orange),  
 438 friable sand (yellow orange), increasing cement (red brown), friable shale 56% (clay) (dark green),  
 439 and friable shale 100% (clay) (bright green). The depth (m), the interpreted facies labels of the  
 440 core plug, and core porosity are located below and at the side of the thin section.  
 441

442 porosity measurements, core measurements, and  $V_p$  measurements were obtained from the  
443 correlation with the log measurements as discussed above. Notice that the bar front and prograding  
444 sand facies have the highest porosities for their  $V_p$ , followed by the back bar facies and the  
445 nonreservoir facies. The loss of porosity is consistent with decaying sorting, where the bar front  
446 and prograding sand facies are best sorted, followed by the back bar facies. As expected,  $V_p$   
447 increases as porosity decreases.

448 Diagnostic rock physics models are used to relate seismic and elastic properties of the  
449 reservoir and the surrounding formations to geologic observations. They can facilitate predictions  
450 of seismic properties away from the wells that are created by stratigraphic conditions and fluid  
451 saturations that have not yet been encountered by drilling or in outcrops. A well-calibrated model  
452 may be used to predict geologic properties, away from well control.

453 When rock physics diagnostic models fit  $V_p$  and porosity, they often predict  $V_s$ , bulk, and  
454 shear moduli poorly. Granular rocks often have lower shear modulus than expected from the  
455 models, possibly as a result of some grain contacts that do not slip under compression, do slip  
456 under shear. Since  $V_p$  involves both moduli, fitting a model to  $V_p$  alone often results in an  
457 underprediction of bulk modulus and an overprediction of shear modulus and associated  $V_s$ . An  
458 ad hoc shear reduction factor was used in this study to represent grain slip and to improve the  
459 model fitting. One of the problems with the shear reduction factor is that it changes the entire  
460 model's shear components, including the well-defined values at the mineral points. The models  
461 used in this study were fitted in the moduli space to better constrain the shear reduction factor and  
462 to obtain an improved estimate of the bulk modulus. Following this approach, it was possible to  
463 fit the bulk modulus and then adjust the shear reduction factor to fit the shear modulus. Then,

464 fitting  $V_p$  and  $V_s$  only required adjusting the density model. Additionally, the model fitting was  
465 tested in  $V_p/V_s$  vs. AI space.

466 Several rock physics models were used in the rock physics diagnostic analysis of the Bell  
467 Creek reservoir. The following is a brief description of the models; more detail can be found in  
468 Avseth et al. (2005):

- 469 • Friable sand: models the elastic properties of uncemented sand as sorting changes and  
470 porosity is lost. The high-porosity well-sorted endpoint is defined by the Hertz–Mindlin  
471 theory, and the poorly sorted endpoint is defined by the mineral properties. The endpoints  
472 are connected by a modified Hashin–Shtrikman lower bound (Dvorkin and Nur, 1996).
- 473 • Friable shale: same as friable sand except that it is used to model changes in elastic  
474 properties of a constant mixture of sand and clay as sorting changes and porosity is lost.
- 475 • Contact cement: models the effect of initial cementation at grain contacts of a high-  
476 porosity grain pack. The model is only valid at high porosity and for small volumes of  
477 cement. The model is numerical, and the details are given in Dvorkin et al. (1994).
- 478 • Increasing cement: models the effect of increasing cement volumes above those modeled  
479 by the contact cement model. It is valid for intermediate cement volumes. The porosity  
480 limits are interpretive. The high-porosity endpoint is the intersection with the contact  
481 cement model, and the low-porosity endpoint is defined by a effective elastic properties  
482 of a mixture of sand grains and porosity-filling cement. The endpoints are connected by  
483 a modified Hashin–Shtrikman upper bound (Avseth et al., 2005).
- 484 • Constant cement: like the friable sand model, except here, the high-porosity well-sorted  
485 endpoint is defined by the contact cement model. The volume of cement is held constant  
486 as porosity decreases and sorting decays (Avseth et al., 2000).

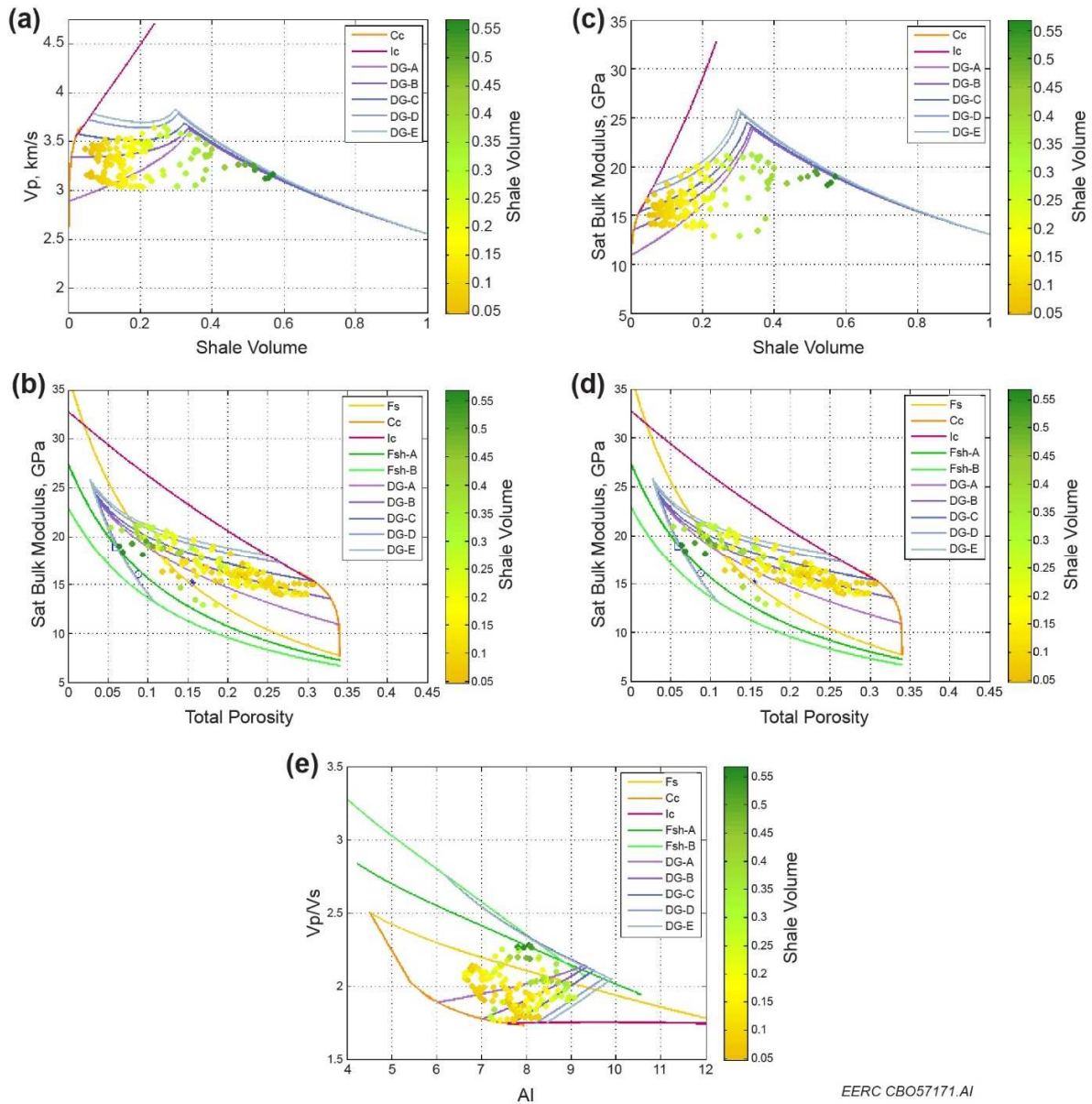
487 A composite framework model resulting from the combination of the friable sand, friable  
488 shale, contact cement, and increasing cement models was used to assess the volume of cement and  
489 shale properties of the Bell Creek reservoir, as shown in Fig. 10.

490 The Bell Creek reservoir is a mixture of shaley sand and sandy shale deposited in a barrier  
491 bar–strand plain sedimentary environment. The reservoir facies are relatively thin and complicated  
492 by the vertical and lateral variability of the depositional environment. This combination of factors  
493 creates inconsistency in porosity, sorting, and clay content, all of which influence the rocks' elastic  
494 properties.

495 Marion et al. (1992) developed a model to explain changes of elastic properties of  
496 unconsolidated mixtures of sand and dispersed clay as the mixtures are varied. The model was  
497 extended to consolidated mixtures measured with wireline logs by Dvorkin and Gutierrez (2002).  
498 This model starts from clean sand and adds porous clay to the pore space until filled without  
499 disturbing the grain contacts. This point is called the critical clay volume. If more clay is added,  
500 then sand must be removed, and the grains begin to separate. More clay is added, and sand removed  
501 until the sample is 100% clay. The critical clay volume marks the transition from grain-supported  
502 shaley sand to matrix-supported sandy shale.

503 From Marion et al. (1992) and Dvorkin and Gutierrez (2002) models, it can be inferred that  
504  $V_p$  increases, and porosity decreases as clay fills the pore space, replacing fluid until critical clay  
505 volume is reached. When clay volumes are greater than critical clay volume, porosity increases,  
506 and  $V_p$  decreases. The porosity increase is due to nonporous quartz being replaced with porous  
507 clay.  $V_p$  decreases as clay volumes become greater than the clay full point because of the  
508 replacement of high-velocity quartz with lower-velocity clay.

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**Figure 10.** Rock physics diagnostic Dvorkin–Gutierrez and framework models a)  $V_p$  vs. shale volume ( $V_{shale}$ ); b)  $V_p$  vs. total porosity ( $\Phi_{it}$ ); c) bulk modulus vs.  $V_{shale}$ ; d) bulk modulus vs.  $\Phi_{it}$ ; e)  $V_p/V_s$  vs. AI, where the models are Fs – friable sand, Cc – contact cement, Ic – increasing cement, Fsh-A – friable shale with 0.56 clay and 0.44 quartz, Fsh-B – friable shale with 1.0 clay, Dvorkin–Gutierrez shaley sand models DG-A – cement volume 0.001, DG-B – cement volume 0.011, DG-C – cement volume 0.035, DG-D – cement volume 0.075, DG-E – cement volume 0.092. Estimated shale properties from Table 1: blue square – Bell Creek, blue circle – Rozet, and blue asterisk – Sprigen Ranch.

	<b>V<sub>p</sub>, km/s</b>	<b>V<sub>s</sub>, km/s</b>	<b>Density, g/cm<sup>3</sup></b>	<b>K, GPa</b>	<b>Mu, GPa</b>	<b>DPHI, frac</b>
Bell Creek	3.18	1.39	2.52	18.9	4.9	0.06
Springen Ranch	3.07	1.53	2.40	15.2	5.59	0.155
Rozet	2.93	1.27	2.51	16.1	4.07	0.09

**Table 1.** Shale properties in the C well for Springen Ranch, Bell Creek, and Rozet based on well log data. K: bulk modulus and Mu: shear modulus.

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523  
524

525 The Dvorkin and Gutierrez (2002) model was applied to the Bell Creek and Springen Ranch  
526 to find the critical clay volume, which is identified as the point in a sand–clay mixture where V<sub>p</sub>  
527 is maximum. (Fig. 10). Maximum V<sub>p</sub> (3.652 km/s) occurs at a volume of 0.267 dry clay. This is  
528 in the range expected for critical clay volume (the critical clay volume is generally between  
529 0.2 and 0.4 volume clay). Density porosity is 0.1 at the critical clay volume. When converted to a  
530 total volume, the critical clay volume including clay porosity is 0.34. This value is close to the  
531 maximum density porosity in the Bell Creek, which is 0.31 in the A well with 0.05 volume clay.  
532 When converted to total volume with the clay removed, a clean sand porosity of 0.34 is estimated,  
533 which agrees with the critical clay volume. The critical porosity for clean sand is likely to be close  
534 to 0.34, which supports the hypothetical clean sand critical porosity used in the rock physics  
535 models. In general, the data show much variability in V<sub>p</sub> for any given volume of clay. This  
536 variability could be explained by a changing depositional environment causing changes in sorting  
537 and clay’s contribution to the elastic properties. Some of the clays are likely in laminated sands,  
538 and others may be in the contact between sand grains. In both cases, V<sub>p</sub> is low relative to a grain  
539 pack with dispersed clay. Variation in the volume of contact cement may also contribute to the  
540 variability. Changes in clay physical properties between Rozet and Springen Ranch could also add  
541 to the variability.

542 The following parameters were used to produce the Dvorkin–Gutierrez models shown in  
543 Fig. 10: Dvorkin–Gutierrez shaley sand, sand endpoints, DG-A – intersects contact cement model

544 at 0.001 volume cement, DG-B – intersects contact cement model at 0.011 volume cement, DG-C  
545 – intersects increasing cement model at 0.035 volume cement, DG-D – intersects increasing  
546 cement model at 0.075 volume cement, and DG-E – intersects increasing cement model at  
547 0.092 volume cement.

548 Although the standard practice is to parameterize the model based on the cleanest sand and  
549 shaliest shale, poor results were obtained with this parameterization. Therefore, the contact,  
550 cement, or increasing cement model for properties of the sand and the friable shale model at 100%  
551 clay for the shale properties were used for the model parameterization and iterated to find the shale  
552 porosity that fit the data. The models derived with these properties in Fig. 10 are consistent with  
553 the observations. The sandy shale model produces a trend that passes close to markers for the Bell  
554 Creek and Rozet shales' average properties. Although individual models in Fig. 10 are challenging  
555 to identify in the observations, the upper limit of the models appears to conform to the upper limit  
556 of the data when plotted on bulk modulus vs. clay and Vp vs. clay volume plots. The sandy shale  
557 portion of the model fits best when clay porosity is 0.11.

558 The depositional complexity reflected in the wide variation of sand porosity may create  
559 variation in the elastic properties, leading to uncertainty in the critical clay volume. Variation in  
560 clay grain properties and porosity may also contribute. The depositional environment changes from  
561 nearshore to barrier bar, to lagoon, to onshore; in some cases, the deposition style (e.g., laminated  
562 deposits) is inconsistent with the model's assumptions. Given this complexity, it is likely that there  
563 is more than one critical clay volume. Finally, a potential second shale trend is observed in the  
564 bulk modulus vs. clay volume plot in Fig. 10 but not in the Vp vs. clay volume; thus, a second  
565 higher-porosity shale trend potentially related to the Springen Ranch is suspected.

566 **4.4 Pressure sensitivity**

567 Time-lapse seismic measurements are primarily affected by changes in reservoir pore  
568 pressure and fluid saturation (e.g., Landrø, 2001; Lumley, D., 2010). Rock physics modeling was  
569 conducted to estimate the pressure and saturation scenarios of the Bell Creek reservoir. Reservoir  
570 and fluid properties obtained from petrophysics, geology, and reservoir simulation were used for  
571 the pressure and fluid substitution modeling.

572 Since neither laboratory measurements of  $V_p$  and  $V_s$  on dry rocks of the Bell Creek nor an  
573 analogous set of measurements from an area with similar geology and structural history were  
574 available, the MacBeth (2004) empirical model was used to predict the pressure sensitivity for the  
575 Bell Creek reservoir. Pressure response data are typically limited to the reservoir's permeable  
576 parts; therefore, the model's application was limited to the Bell Creek's sandy facies.

577 The stress sensitivity relations from the MacBeth model are based on laboratory  
578 measurements on rocks over a wide range of conditions. The model is given for  $K_{Dry}$  and  $\mu_{Dry}$  for  
579 effective pressure ( $P_{eff}$ ) as follows:

580 
$$K_{Dry} = \frac{K_{\infty}}{1 + \left( \frac{S_k}{(1-S_k)} \right) e^{\left( \frac{-P_{eff}}{P_k} \right)}} \quad (2)$$

581

582 
$$\mu_{Dry} = \frac{\mu_{\infty}}{1 + \left( \frac{S_{\mu}}{(1-S_{\mu})} \right) e^{\left( \frac{-P_{eff}}{P_{\mu}} \right)}} \quad (3)$$

583 This model requires knowledge of the average porosity, initial and final pore pressures, and  
584 overburden pressure. Average porosity is the only parameter required to determine model  
585 coefficients from MacBeth (2004) pressure sensitivity crossplots. The Bell Creek average porosity  
586 was derived from histograms of density porosity in the reservoir's sandy facies at the A, C, and G  
587 wells. The estimated average porosity of 23% was used to derive the following MacBeth model

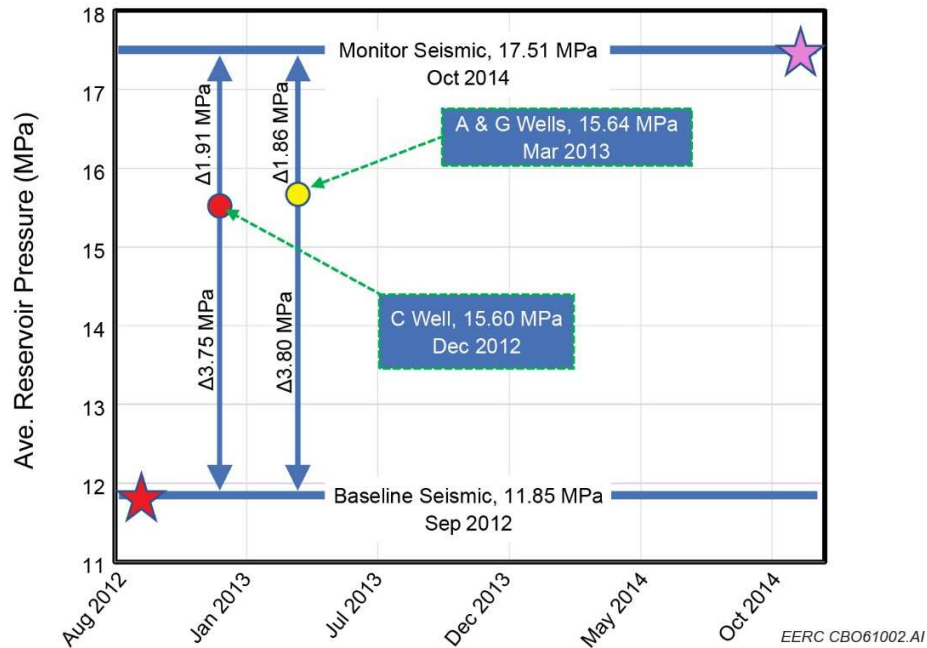
588 parameters: total stress for bulk modulus ( $S_{\kappa}$ ) = 0.59; pressure constant for bulk modulus ( $P_{\kappa}$ ) =  
589 6 MPa; highest-pressure asymptote for bulk modulus ( $\kappa_{\infty}$ ) = 12.4 MPa; total stress sensitivity for  
590 shear modulus ( $S_{\mu}$ ) = 0.59; pressure constant for shear modulus ( $P_{\mu}$ ) = 7 MPa, and highest-pressure  
591 asymptote for shear modulus ( $\mu_{\infty}$ ) = 10.9 GPa.

592 An overburden pressure of 30.19–30.42 MPa was calculated by integrating a density log  
593 from the C well between the surface and reservoir, assuming a density of 2.07 g/cm<sup>3</sup> between the  
594 top of the log and the surface. A value of 30.33 MPa was adopted for the overburden pressure of  
595 the Bell Creek reservoir.

596 Pore pressure estimation is complicated because pressure values were required when the  
597 well was drilled and when each seismic survey was acquired. Pore pressure can be derived from  
598 wireline formation testers, drill stem tests, production data, or reservoir modeling. As pore pressure  
599 data for individual wells at Bell Creek were not available, average values from the reservoir  
600 simulation were used.

601 The available data are discussed below:

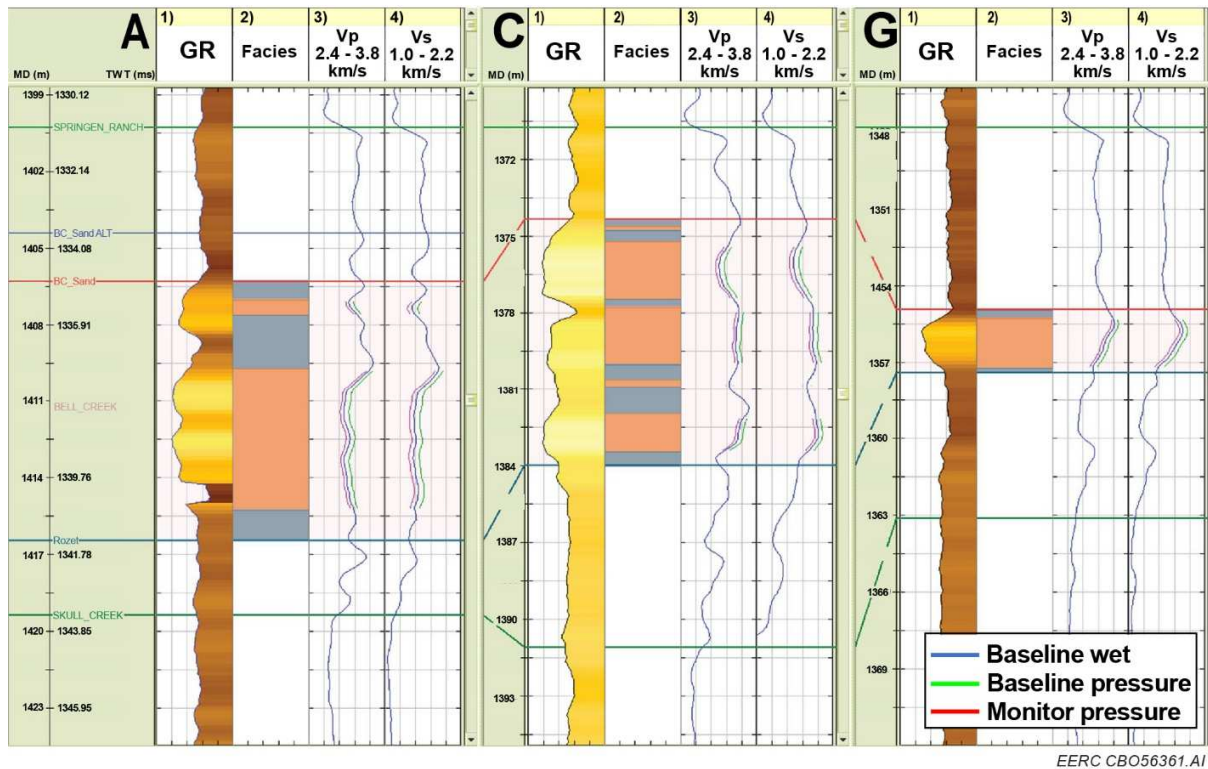
- 602 • As described in Fig. 11, the C well was drilled and logged in December 2012, which is  
603 between acquisitions of baseline and monitor seismic surveys. Pore pressure at the time  
604 of the baseline survey was 3.75 MPa lower than when C well was logged. Pore pressure  
605 at the time of the monitor survey was 1.91 MPa higher than when C well was logged.
- 606 • The G and A wells were drilled and logged in March 2013, which is between the  
607 acquisition of the baseline and monitor seismic surveys. Pore pressure at the time of the  
608 monitor survey was 3.89 MPa lower than when G and A wells were logged. Pore pressure  
609 at the time of the monitor survey was 1.86 MPa higher than when the wells were logged.



610

611 **Figure 11.** Average reservoir pressure conditions during the time-lapse seismic surface  
 612 survey. Red star: baseline seismic surface survey (11.85MPa in September 012); pink star:  
 613 monitor seismic surface survey (17.51 MPa in October 2014); red circle: C well (15.60 MPa  
 614 in December 2012); A and G wells (15.64 MPa in March 2013).  
 615

616 The pressure substitution based on MacBeth’s relationships used here involves a  
 617 combination of fluid substitution to remove fluid effects and pressure substitution applied to the  
 618 dry matrix and resubstituting of the fluids modeled at the new pore pressure back into the rock.  
 619 Therefore, the rocks in the model were fluid-substituted to wet before pressure substitution. In  
 620 general, increasing pore pressure should cause  $V_p$  and  $V_s$  to decrease. The application of the model  
 621 to the Bell Creek reservoir produces the results shown in Fig. 12. The velocities in all three wells  
 622 are influenced by pore pressure changes. Whereas the pore pressure at the time of the baseline  
 623 seismic survey was lower (green curves) and  $V_p$  and  $V_s$  were higher than when the wells were  
 624 drilled, the pore pressure at the time of the seismic monitor survey was higher (pink curves), and  
 625 velocities were lower.



626

627 **Figure 12.** Results of pressure sensitivity modeling for A, C, and G wells. Left to right  
 628 tracks at each well: GR, facies (orange: sandy, gray: shaley), Vp, and Vs. Blue curves:  
 629 original water-saturated Vp and Vs.; green curves: baseline survey; pink curves: monitor  
 630 survey.  
 631

632 **4.5 Fluid properties and substitution**

633 Prasad et al. (2021) provide an excellent discussion on fluid the challenges associated with  
 634 the standard fluid substitution using the Gassmann approach for CO<sub>2</sub> storage projects as part of the  
 635 seismic inversion workflow. Prasad et al. (2021) report the significant changes in rock properties  
 636 due to CO<sub>2</sub> injection and the time-lapse seismic signatures associated with those changes. Prasad  
 637 and coauthors emphasize the importance of developing rock physics models to assess second-and  
 638 third-order changes in seismic properties throughout laboratory studies, small-scale field pilots,  
 639 and monitoring large-scale geologic storage of CO<sub>2</sub>.

640 Diagnostic analysis of rock properties and the effects of fluid saturation changes require that  
 641 the reservoir in all the wells has the same initial fluid saturation. Gassmann's equations (1951)  
 642 were used for fluid substitution from in situ to brine-filled conditions. The equations require  
 643 information on initial fluid saturation, fluid elastic properties, rock elastic properties, and total  
 644 porosity.

$$645 \quad \frac{K_{Sat}}{K_{Mineral}-K_{Sat}} = \frac{K_{Dry}}{K_{Mineral}-K_{Dry}} + \frac{K_{Fluid}}{\varphi(K_{Mineral}-K_{Fluid})} \quad (4)$$

$$646 \quad \mu_{Sat} = \mu_{Dry} \quad (5)$$

648 where  $K_{Sat}$  is the effective bulk modulus of the rock with pore fluid,  $K_{Mineral}$  is the bulk modulus  
 649 of mineral material making up the rock,  $K_{Dry}$  is the effective bulk modulus of dry rock,  $K_{Fluid}$  is  
 650 the effective bulk modulus of pore fluid,  $\varphi$  is the porosity,  $\mu_{Sat}$  is the effective shear modulus of  
 651 rock with pore fluid, and  $\mu_{Dry}$  is the effective shear modulus of dry rock.

652 Initial fluid saturation was obtained from the Schlumberger RST Advisor analysis of the  
 653 wireline and baseline pulsed-neutron logs (PNL). Fluid properties for in situ oil and brine were  
 654 obtained from measurements on produced fluids. Rock composition was obtained from  
 655 Schlumberger RST SpectroLith analysis. Grain bulk modulus was calculated using standard  
 656 properties and Voigt–Reuss–Hill average (Mavko et al., 2009). Porosity is based on density  
 657 porosity on a sandstone matrix.

#### 658 **4.6 Elastic properties of fluids and their mixtures**

659 Modeling of elastic fluid properties is critical to the quantitative time-lapse analysis of the  
 660 Bell Creek reservoir. Table 2 provides an overview of reservoir and fluid properties used for fluid  
 661 substitution modeling. The reservoir temperature and pore pressure indicate that CO<sub>2</sub> will be  
 662

<b>Muddy Reservoir</b>	<b>Baseline Seismic Sep 2012</b>	<b>Well C Dec 2012</b>	<b>Wells A and G Mar 2013</b>	<b>Monitor Seismic Oct 2014</b>
Temperature, °C	42.2	42.2	42.2	42.2
Ave Pore Pressure, MPa	11.84 @ Well C	15.60 @ Well C	15.64 @ Well C	17.51 @ Well C
Pore Pressure Range (Inj/Prod), MPa	18.62/12.41	18.62/12.41	18.62/12.41	22.75/16.55
Brine Salinity, mg/L	6013	6013	6013	6013
Oil GOR, m <sup>3</sup> /m <sup>3</sup>	2	2	2	31
Oil API	37.8	37.8	37.6	33.1
Gas Gravity	0.83	0.84	0.84	2.18

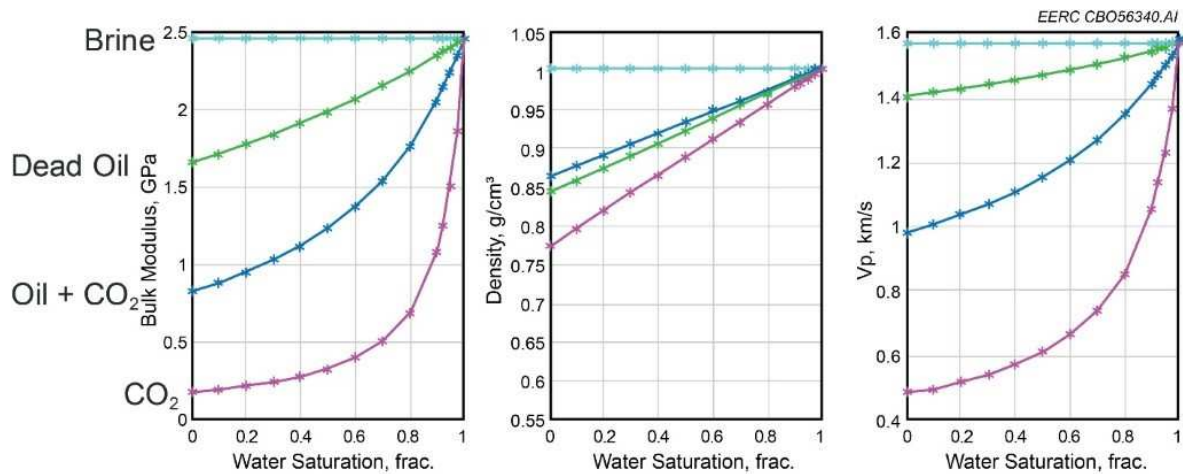
663 **Table 2.** Reservoir and fluid properties for the Bell Creek, Muddy reservoir. The changes in oil  
664 properties in October 2014 are due to the presence of CO<sub>2</sub>.

665

666 supercritical. The Fluid Acoustics for Geophysics (FLAG) modeling algorithm (Han and Batzle,  
667 2014) was used to model the elastic properties of brine, dead oil, and supercritical CO<sub>2</sub> (sCO<sub>2</sub>).  
668 The FLAG model is designed to work with mixtures of hydrocarbons and water, and mixtures  
669 containing sCO<sub>2</sub> act differently, thus alternative models are required for mixtures containing sCO<sub>2</sub>.  
670 Models of sCO<sub>2</sub>-saturated oil and water were built from relationships from the literature; mixtures  
671 of oil saturated with sCO<sub>2</sub> were modeled using Han et al. (2012) equations; the water saturated  
672 with sCO<sub>2</sub> was modeled using Han and Sun (2013) equations. All models above are empirical and  
673 based on fitting laboratory measurements of V<sub>p</sub> and density of defined fluid mixtures under  
674 controlled temperature and pressure conditions.

675 Fig. 13 shows the model results. Notice the following at the zero-water saturation endpoint:  
676 Water and water saturated with CO<sub>2</sub> (gas-water ratio [GWR] = 26.9) have very similar properties;  
677 dead oil is the second highest V<sub>p</sub> fluid. Saturation with CO<sub>2</sub> lowers its velocity significantly and  
678 slightly increases its density, and CO<sub>2</sub> at reservoir conditions has the lowest velocity and density.

679



680

681 **Figure 13.** Modeling results of elastic properties of fluids and their mixtures. October 2014  
 682 fluids mixed with water using Wood's equation. Initial water saturation before EOR is 0.5–  
 683 0.6 and decreases to an average of 0.35 after EOR. Approximate irreducible water  
 684 saturation for Bell Creek is  $S_{wirr} = 0.25$ .

685

686 Any fluid other than water is unlikely to exist in the reservoir at 100% saturation; other fluids  
 687 exist as mixtures primarily with water. Using Wood's (1941) equation to describe uniform mixing  
 688 of fluids in Fig. 13, it is observed that as water saturation increases, the fluid properties converge  
 689 toward water. Expected water saturations are between an irreducible water saturation of 0.25 and  
 690 approximately 0.5–0.6, which is the approximate saturation after waterflooding and before CO<sub>2</sub>  
 691 EOR. The water saturation after CO<sub>2</sub> EOR is expected to be near 0.35.

692 Uniform fluid mixing (Wood's equation) was selected to calculate the bulk modulus of the  
 693 fluid mixtures used in the inversion. The mixture of CO<sub>2</sub>-saturated brine with free CO<sub>2</sub> is likely to  
 694 exhibit partial saturation. The amount of partial saturation is expected to vary within the reservoir  
 695 due to varying geologic properties and reservoir conditions imposed by CO<sub>2</sub> injection and oil  
 696 production processes. It is worth mentioning that using uniform mixing for free CO<sub>2</sub> with CO<sub>2</sub>-  
 697 saturated brine may lead Ji-Fi to misestimate the free CO<sub>2</sub> fluid facies distribution.

698 Domenico (1976) performed lab experiments on unconsolidated sand samples saturated with  
699 gas–brine mixtures to determine the effects of changing fluid saturations on P-wave velocity. He  
700 determined that uniform fluid mixing best described his results. Most immiscible fluid mixtures  
701 are not uniformly mixed. Fluids distribution is controlled by the reservoir rock’s microtexture and  
702 the relative permeability of the fluids, among many other factors. The separation of a binary  
703 mixture into distinct components will cause the equivalent fluid to approximate a Voigt average.  
704 Most fluid mixtures have behaviors between these bounds. Brie et al. (1995) proposed an empirical  
705 relationship for the behavior of gas–brine mixtures observed in well log data.

$$706 \quad K_f = S_w^e (K_w - K_g) + K_g \quad (6)$$

707 where  $K_f$  = fluid bulk modulus,  $K_w$  = brine bulk modulus, and  $K_g$  = gas bulk modulus. The Brie  
708 exponent  $e$  is usually between 2 and 5, and when  $e = 1$ , Equation 7 is a Voigt equation, and when  
709  $e = 40$ , Equation 7 approximates Wood’s equation. The value of the Brie exponent is also expected  
710 to be sensitive to frequency of the seismic waves.

711 As an alternative to Wood’s equation, a value of Brie exponent in the ranges mentioned  
712 could have been selected. The choice of an appropriate value is problematic as this project does  
713 not have core flood measurements or time-lapse elastic well logs to help guide the selection.

#### 714 **4.7 Fluid substitution and Ji-Fi fluid facies**

715 The fluid properties were used to create end-member fluid models for Ji-Fi inversion  
716 (Table 3). The real fluid mixtures are expected to form a continuum; however, Ji-Fi was limited to  
717 four fluid facies models. Two fluid mixtures were selected to describe baseline (September 2012)  
718 reservoir conditions and four fluid mixtures to describe monitor (October 2014) reservoir  
719 conditions. Water saturations of 1.0, 0.55, and 0.35 were selected for the wet reservoir, oil reservoir  
720 after waterflood, and reservoir after CO<sub>2</sub> EOR, respectively.

- 721 The fluid models represent:
- 722 • All brine model – areas outside the field or below the oil–water contact (OWC).
  - 723 • Dead oil and brine – reservoir after water flooding and before CO<sub>2</sub> effect (S<sub>o</sub> = 0.45,
  - 724 S<sub>w</sub> = 0.55).
  - 725 • (oil + CO<sub>2</sub>) and brine – reservoir where EOR is active (S<sub>(oil + CO<sub>2</sub>)</sub> = 0.65, S<sub>w</sub> = 0.35).
  - 726 • (brine + CO<sub>2</sub>) and CO<sub>2</sub> – areas relatively close to the injectors (S<sub>(brine + CO<sub>2</sub>)</sub> = 0.35, S<sub>CO<sub>2</sub></sub>
  - 727 = 0.65).

• Baseline (September 2012)	Monitor (October 2014)
Brine	Brine
Dead oil (S <sub>o</sub> = 0.45) & brine (S <sub>w</sub> = 0.55)	Dead oil (S <sub>o</sub> = 0.45) & brine (S <sub>w</sub> = 0.55)
	Oil + CO <sub>2</sub> (S = 0.63) & brine (S <sub>w</sub> = 0.35)
	Brine + CO <sub>2</sub> (S = 0.35) & CO <sub>2</sub> (S <sub>CO<sub>2</sub></sub> = 0.65)

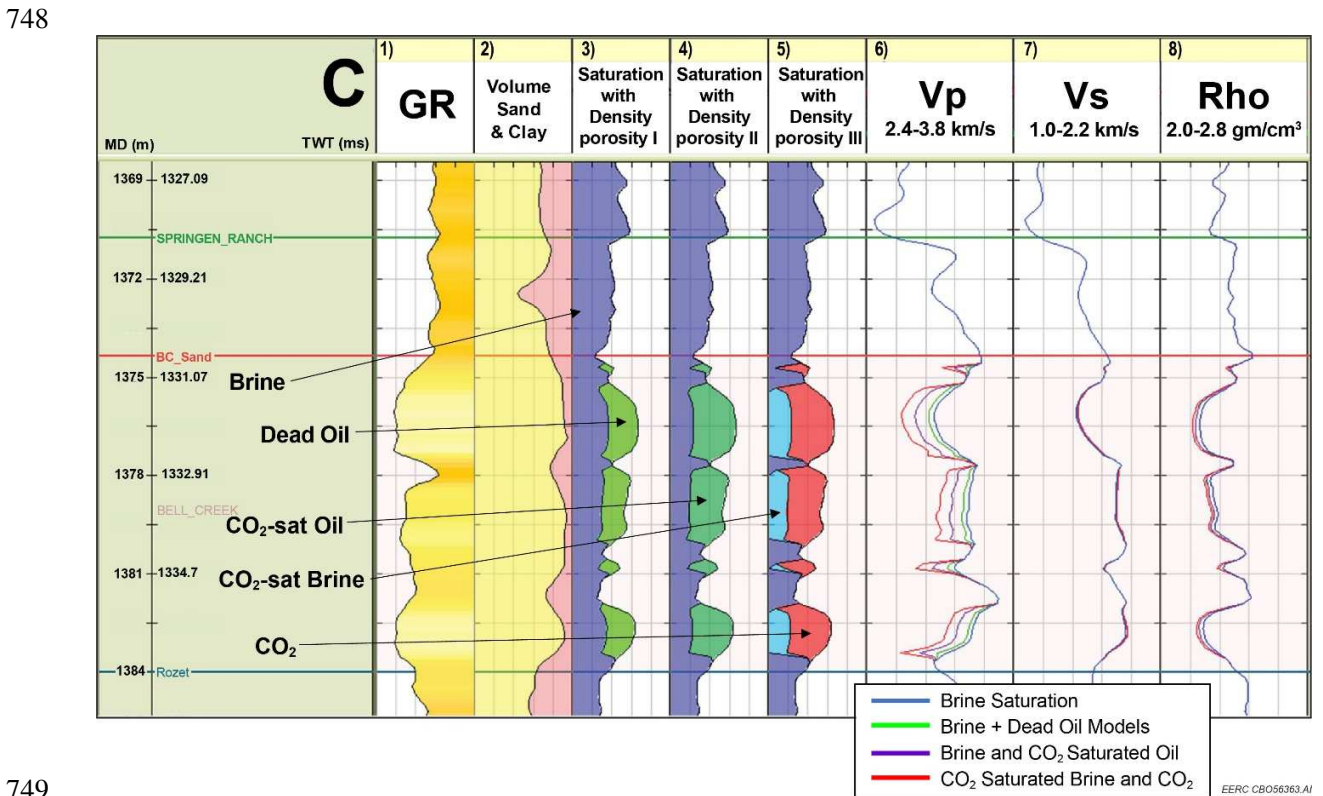
Note: The symbol “&” indicates fluid mixing with Wood’s equation while the symbol “+” indicates modeling of CO<sub>2</sub> saturated mixtures of oil and CO<sub>2</sub> or water and CO<sub>2</sub> using the models mentioned in the main text.

728 **Table 3.** Fluid mixture models for fluid substitution. The (&) symbol indicates fluid mixing with  
 729 Wood’s equation while the (+) symbol indicates modeling with either the oil and CO<sub>2</sub> or water  
 730 and CO<sub>2</sub> empirical fluid property models discussed above.  
 731

732 Gassmann (1951)-based fluid substitution was conducted. It consisted of replacing the brine  
 733 in the Bell Creek sandy facies of the pressure-adjusted baseline and monitor models with fluid  
 734 properties from the models in Table 3. Gassmann’s equations were used with total porosity from  
 735 density porosity and a rock model consisting of quartz, carbonate, and clay volume fractions. The  
 736 baseline fluid model and the monitor model are from September 2012 and October 2014,  
 737 respectively. Fluid substitution is limited to the sandy facies, which was also used in pressure  
 738 substitution.

739 Fluid substitution for the baseline case consisted of replacing the water in the water-saturated  
 740 baseline pressure-substituted model with a mixture of dead oil and water at four wells: A, B, C,  
 741 and G. A small reduction in V<sub>p</sub> and density relative to the brine-saturated model is observed for  
 742 the mixture of dead oil and brine.

743 In the fluid substitution for the monitor case, mixtures of dead oil and water, oil + CO<sub>2</sub> and  
 744 water, and CO<sub>2</sub> + water and water replaced the water in the water-saturated monitor pressure-  
 745 substituted model at four wells: A, B, C, and G. The results for C are shown in Fig. 14. Vp and  
 746 density decrease as brine is replaced with dead oil and brine, oil-CO<sub>2</sub> and water, and CO<sub>2</sub> and  
 747 water, respectively.



749  
 750  
 751 **Figure 14.** Monitor (October 2014) fluid substitution into sandy facies for C well. Left to right  
 752 tracks: GR, lithology (volume), and three saturations with density porosity (DPHI)  
 753 scenarios: 1) brine (dark blue) and dead oil (green), 2) brine (blue) and CO<sub>2</sub>-saturated oil (dark  
 754 green), 3) brine (dark blue), CO<sub>2</sub>-saturated brine (light blue) and supercritical CO<sub>2</sub> (red), P-wave  
 755 velocity (Vp), S-wave velocity (Vs), and density (Rho). Blue curves: 1.0 brine saturation, green  
 756 curves: brine + dead oil models, purple curve: brine and CO<sub>2</sub>-saturated oil, red: CO<sub>2</sub>-saturated  
 757 brine and CO<sub>2</sub>.

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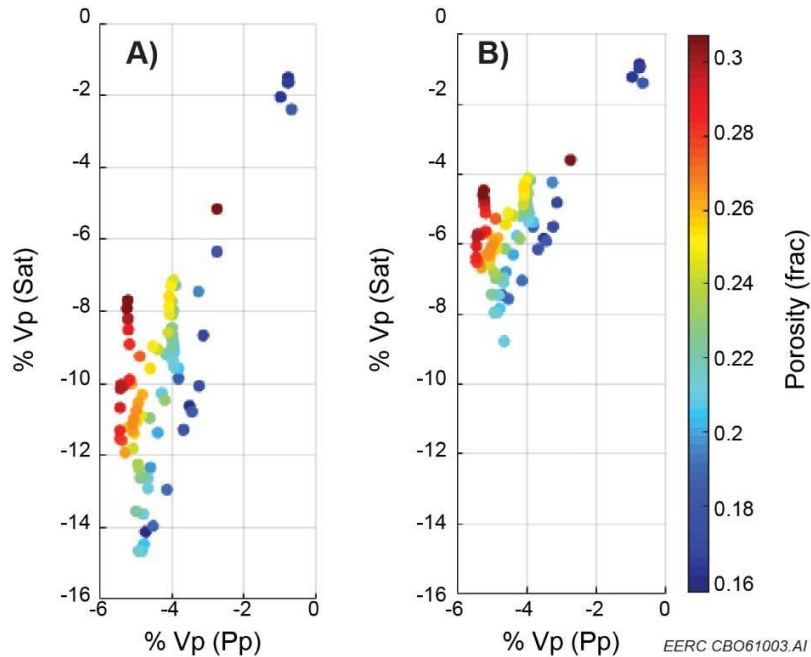
760 Baseline and monitor models show a small separation between brine-saturated and dead oil  
761 + brine Vp curves. A more significant separation is seen between CO<sub>2</sub> + brine, oil + CO<sub>2</sub>, and dead  
762 oil + brine Vp curves, suggesting that the monitor CO<sub>2</sub> + brine, CO<sub>2</sub> + oil, and unchanged fluid  
763 may be easier to separate than the pre-CO<sub>2</sub> baseline fluids. Minimal separation is observed for the  
764 density curves, suggesting a minimal contribution to amplitude variation with offset (AVO) at very  
765 far offsets.

#### 766 **4.8 Rock response to change in pore pressure and saturation**

767 The pressure and fluid substitution models can be summarized to show how pressure and  
768 saturation changes affect rock properties. When pore pressure is increased from baseline to  
769 monitor models, Vp and Vs decrease. When the wet reservoir is replaced with either CO<sub>2</sub> + oil or  
770 CO<sub>2</sub> + water, Vp decreases significantly, and Vs increases slightly because of the fluid's change  
771 in density. Hence, saturation change affects Vp, but pore pressure change affects Vp and Vs. The  
772 expected percentage of change in Vp for given porosities due to pore pressure and saturation  
773 conditions after the CO<sub>2</sub> injection into the reservoir in the period of the time-lapse seismic survey  
774 are shown in Fig. 15. Inversion results for saturation can have a pore pressure overprint. The wide  
775 range of Vp and Vs for such a small porosity range is probably associated with the reservoir's  
776 complexity.

#### 777 **5. ELASTIC FACIES**

778 Ji-Fi foregoes the traditional method of creating a low-frequency background model separate  
779 from a relative impedance inversion. Instead, Ji-Fi combines the two processes into an absolute  
780 inversion by constructing the background model iteratively from a set of defined elastic facies and  
781 related depth trends. Geologic facies are used to classify rocks according to their appearance or  
782



783

784 **Figure 15.** Effect of pressure and saturation changes on Vp due to injection of CO<sub>2</sub> into the  
 785 reservoir in the study area: a) percentage of change in Vp between brine saturated and the mixture  
 786 of brine and CO<sub>2</sub> at monitor conditions described in Table 2 (vertical axis) and b) percent change  
 787 in Vp between brine saturated and the mixture of oil and CO<sub>2</sub> at monitor conditions described in  
 788 Table 2 (vertical axis). The horizontal axis for both figures: percentage of change in Vp due to  
 789 increasing pore pressure in the period between the baseline and monitor surveys. The dots are  
 790 colored by porosity.

791

792 composition. In Ji-Fi implementation, facies classify rocks based on their shale volume, porosity,  
 793 and saturation. After creating facies logs from petrophysical logs, an automatic approach that  
 794 consistently interprets the elastic logs (Vp, Vs, and density) for the facies classification was  
 795 preferred to minimize the geologic interpretation bias. An unsupervised clustering approach was  
 796 applied to the elastic logs of the A, B, C, D, E, F, G, and H wells to estimate facies logs without  
 797 fluids. In this approach, the only initial interpreter intervention is definition of the number of  
 798 expected facies. The clustering algorithm uses this “hard” constraint in the multidimensional  
 799 analysis of log data to estimate groups with maximum homogeneity between the groups’ elements  
 800 and the maximum differences between groups. Several iterations were needed to estimate the dry

801 facies according to the study area geologic conditions (represented by the petrophysical facies  
 802 logs). In this process, the initial step consisted of analyzing the facies clusters obtained with no  
 803 prior information on the geology and then relating them to geologic/petrophysical facies.  
 804 Clustering success is reflected in the separation of the cluster in the seismic (elastic) domain and,  
 805 therefore, in the quality of the Ji-Fi results.

806 Since only the reservoir interval corresponding to the Bell Creek sandstone was considered  
 807 with fluid fill, the litho-fluid facies (facies reflecting not only rock type but also fluid fill) were  
 808 created according to the fluid substitution modeling cases mentioned in Table 3 (monitor). All of  
 809 the facies representing the geologic and fluid conditions in the period of the two seismic surveys  
 810 (September 2012 – October 2014) are shown in Table 4. These facies are integrated into the DTA  
 811 already explained in a previous section.

812

Color	Facies	Type	Comment
	Shallow Marine 1	Siltstone	
	Marine 1	Shale	
	Marine 2	Shale	
	Shell creek	Shale (lagoon, tidal flat)	
	Barrier bar	Sandstone (bar)	
	Shallow marine intrareservoir	Siltstone (transition)	
	Extra shale	Shale	
	Shallow Marine 2	Siltstone	
	Marine 3	Shale (deep sea)	
	Barrier bar – dead oil and water	Sandstone (bar)	Fluid substituting
	Barrier bar – oil	Sandstone (bar)	Fluid substituting
	Barrier bar – CO <sub>2</sub>	Sandstone (bar)	Fluid substituting
	Barrier bar – water	Sandstone (bar)	Fluid substituting

813

Note: The estimated geologic/petrophysical (dry) facies representing the Bell Creek reservoir at different wells are shallow marine intrareservoir, barrier bar, Shallow Marine 2, Marine 3, and Shell Creek. The fluid facies are assigned at the interval representing the Bell Creek sandstone (barrier bar facies). The prefix barrier bar is used in the name of the fluid facies to facilitate the reservoir interval identification.

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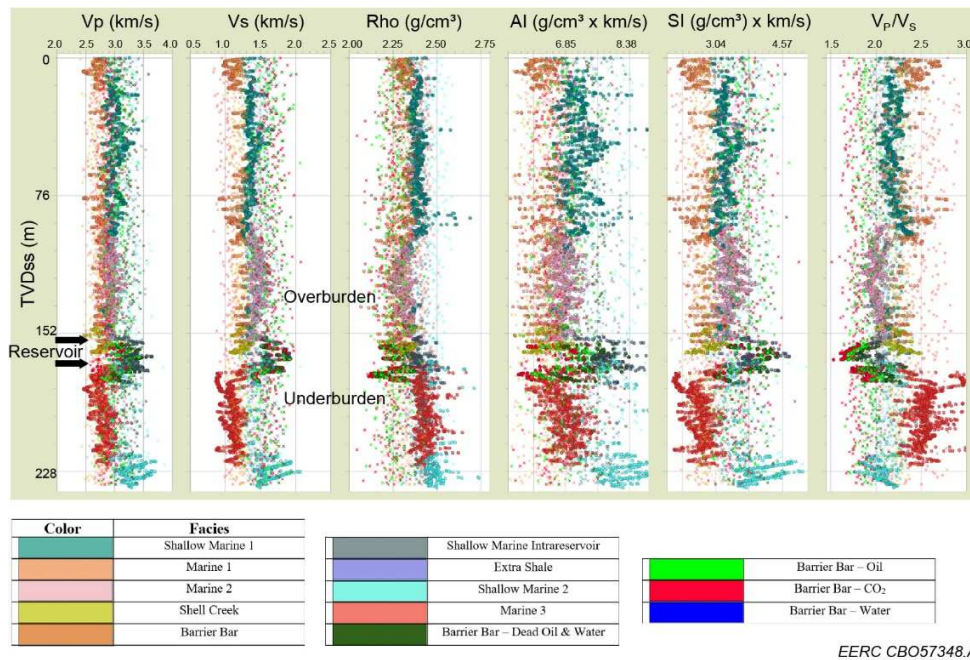
**Table 4. Facies estimated at the A, B, C, D, E, F, G, and H wells.**

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816 **6. JI-FI APPLICATION**

817 A DTA was performed to derive per facies relationships between  $V_p$  and depth and between  
 818  $V_p$ ,  $V_s$ , and density ( $\rho$ ) using the facies derived from the eight wells considered in this project.  
 819 The per facies background models and the rock physics relationships used in Ji-Fi are derived from  
 820 the DTA relationships. An example of the DTAs used for the Bell Creek data inversion is shown  
 821 in Fig. 16. The colored circles represent the trends of each facies derived from well logs at an  
 822 interval with the highest probability of finding specific facies. The colored crosses (the probability  
 823 distribution or prior data) complement the rest of the background associated with a specific facies.  
 824 Notice that in terms of input data for seismic inversion, one facies background model is the  
 825 equivalent to a low-frequency model in standard model-based inversions.

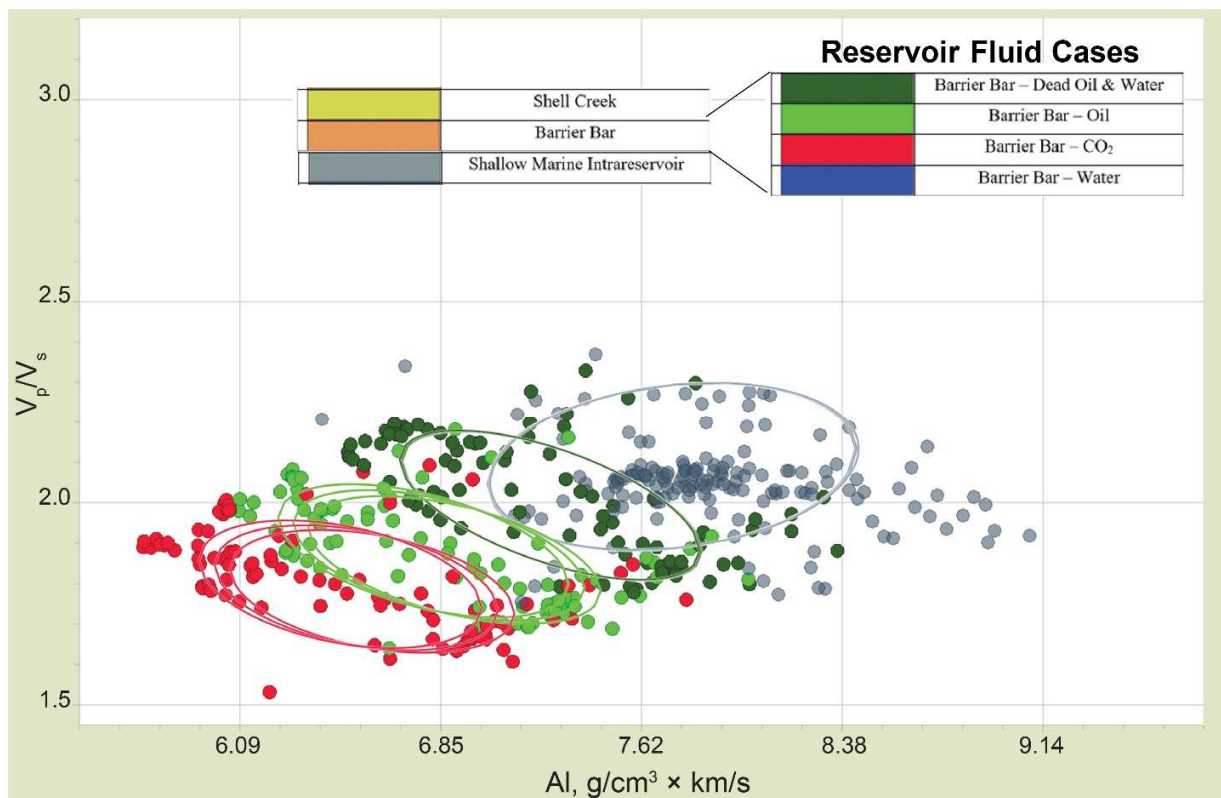
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829 **Figure 16.** Depth trend analysis, including fluid cases used in the Ji-Fi application to the seismic  
 830 data. Left to right tracks: P-wave velocity ( $V_p$ ), shear-wave velocity ( $V_s$ ), density ( $\rho$ ), acoustic  
 831 impedance (AI), shear impedance (SI),  $V_p/V_s$  ratio. Circles: the trend of specific facies at the  
 832 highest probability interval; cross: prior data (probability distribution) of the specific trend  
 833 covering the complete inversion interval. The colors of the symbols represent the facies.

834 The assessment of facies in the depth trends can be simplified when the facies are  
 835 crossplotted. The trends in Fig. 16 can be studied in crossplots of  $V_p$  vs.  $\rho$ ,  $V_p$  vs.  $V_s$ , AI vs. SI,  
 836 and AI vs.  $V_p/V_s$  ratio with their respective probability distributions shown as ellipses for given  
 837 standard deviations. Fig. 17 depicts the AI vs. velocity ratio of the reservoir. The circles and  
 838 ellipses represent the facies' distribution and the probability distribution (one standard deviation),  
 839 respectively. Note the separation in AI between the Shell Creek (yellow) and shallow marine  
 840 intrareservoir facies (gray). The separation between the three litho-fluid facies, barrier bar with  
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842 EERC CBO57180.AI

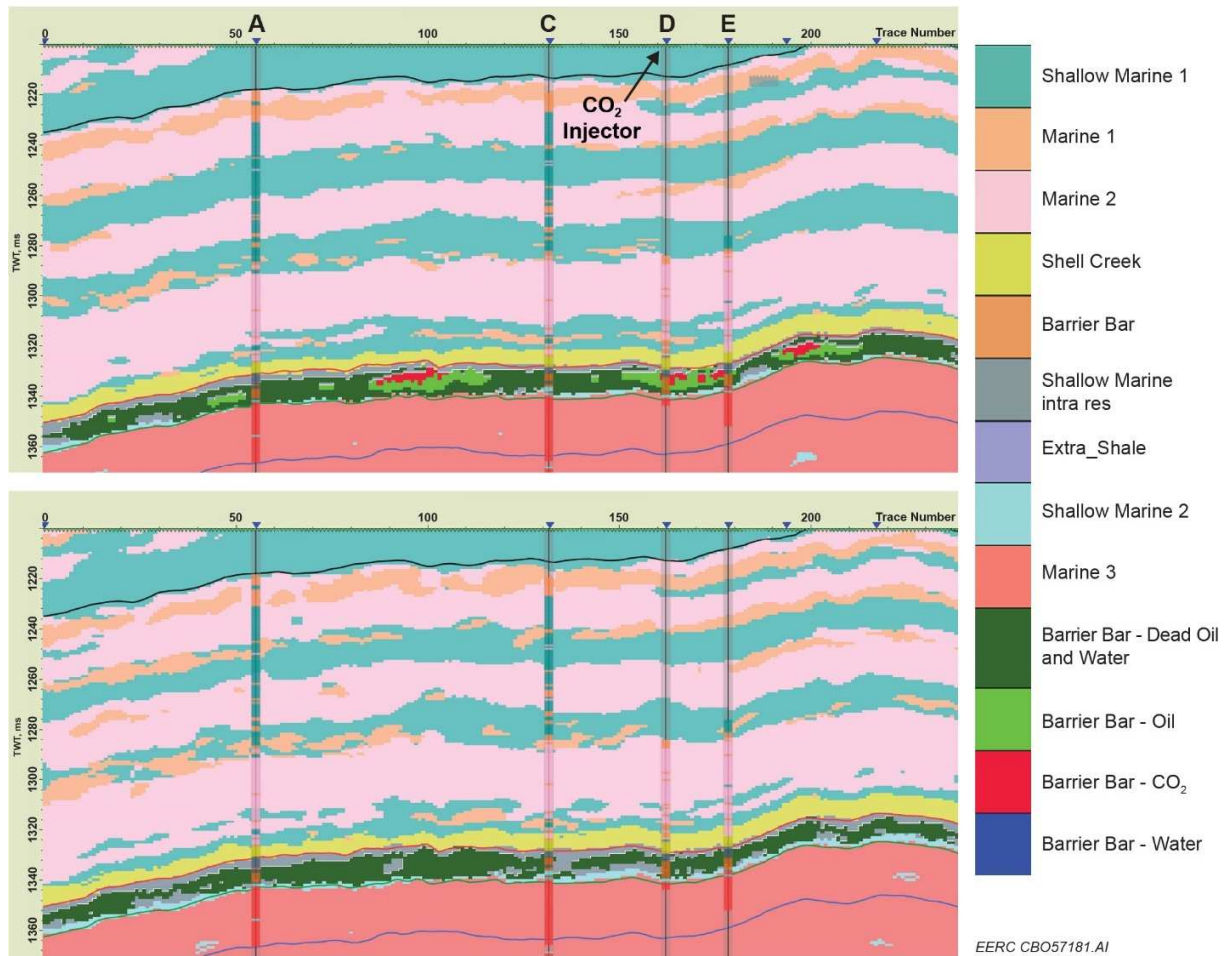
843 **Figure 17.** AI vs.  $V_p/V_s$  crossplot of the DTA shown in Fig. 16(9). Circles: individual facies;  
 844 ellipses: one standard deviation of prior data (probability distribution). Facies: shallow marine  
 845 intrareservoir (gray), barrier bar with dead oil and water (dark green), barrier bar with oil (bright  
 846 green), and barrier bar with CO<sub>2</sub> (red).

847

848 dead oil and water (dark green), barrier bar with oil (bright green), and barrier bar with CO<sub>2</sub> (red),  
849 is also evident from low AI and low-velocity ratio to high AI and velocity ratio. This type of  
850 separation reflects not only the geologic/petrophysical/fluid/seismic characteristics of the reservoir  
851 but also the quality of the initial facies estimation. In addition to high-quality input facies, a target  
852 zone and an estimation of the proportion of each facies in the target zone are fundamental for  
853 running Ji-Fi. The estimation of proportions can be derived from empirical observations generated  
854 from previous inversions or reservoir simulations.

855         When the Ji-Fi algorithm is applied to baseline and monitor seismic data sets, each survey  
856 is inverted separately. Hence, the time-lapse effect analysis due to changes in fluid conditions in  
857 the reservoir is conducted with two sets of impedances and (most probable) facies. This analysis  
858 can be simplified when Ji-Fi is applied to arbitrary lines defined through specific interest areas in  
859 the seismic cubes. Several arbitrary lines were defined as part of the work developed here. The  
860 arbitrary line presented here joins the A, C, D, E, and H wells. Several time-lapse anomalies  
861 observed along this line were crucial for refining the Ji-Fi parameters and the fast validation of  
862 results.

863         Fig. 18 shows the most probable facies estimated by Ji-Fi for the arbitrary line of the baseline  
864 and monitor data sets and the in situ facies (no fluids) at the well locations. The inversion target  
865 zone includes the interval between Horizon 1 (top of the target zone) and the LBBC horizon  
866 (bottom of target zone). Although the wells are not used as constraints in the Ji-Fi algorithm (as is  
867 the case with standard inversions), a good match is observed for the estimated Ji-Fi facies with the  
868 facies at wells outside the reservoir. The comparison of in situ (no fluids) from the wells with the  
869 litho-fluid facies reflects the fluid facies' distribution in the in situ facies. The changes between  
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873 **Figure 18.** Most probable facies estimated by the Ji-Fi algorithm at the arbitrary line for the  
874 baseline (bottom) and monitor (top) seismic data sets. The A (producer), C (observation), D (CO<sub>2</sub>  
875 injector), and E (producer) wells show the estimated in situ facies (no fluids). The seismic  
876 inversion shows the results using litho-fluid facies at the reservoir as input. The cross sections do  
877 not include the part of the line where the H (dry) well is located.

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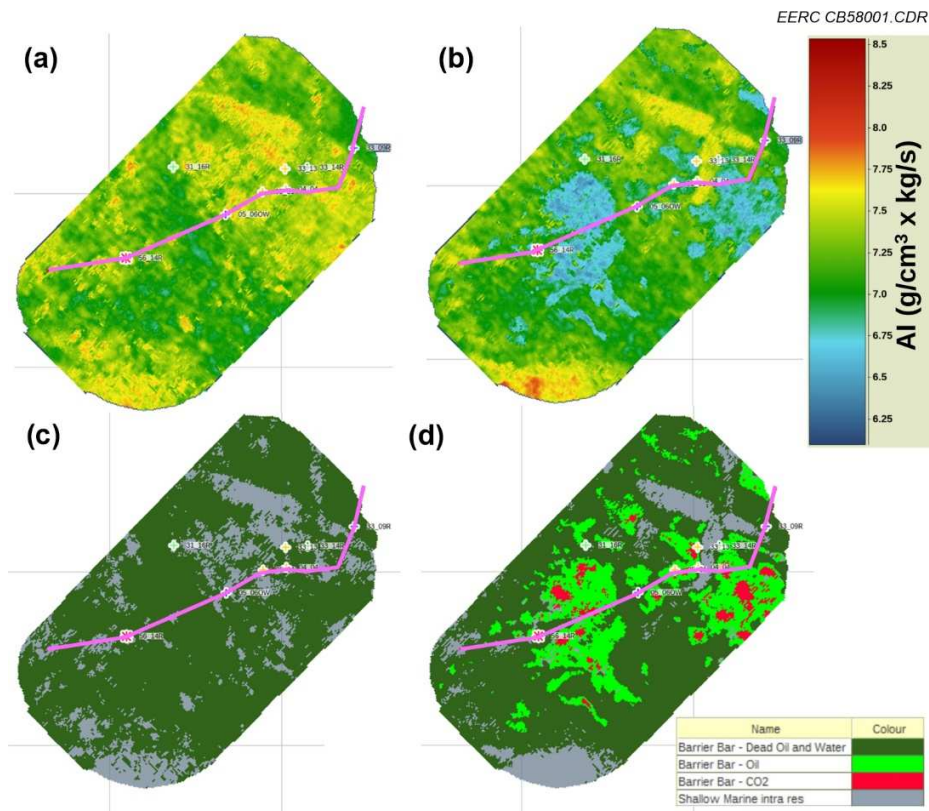
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881 the reservoir litho-fluid facies from baseline to monitor correctly predict the location of CO<sub>2</sub> and

882 oil saturated with CO<sub>2</sub> in the profile. These changes are easily seen on the estimated AI and litho-

883 fluid facies derived from the 3D Ji-Fi on a horizon at 6 ms below the top of the reservoir

884 (Fig. 19).



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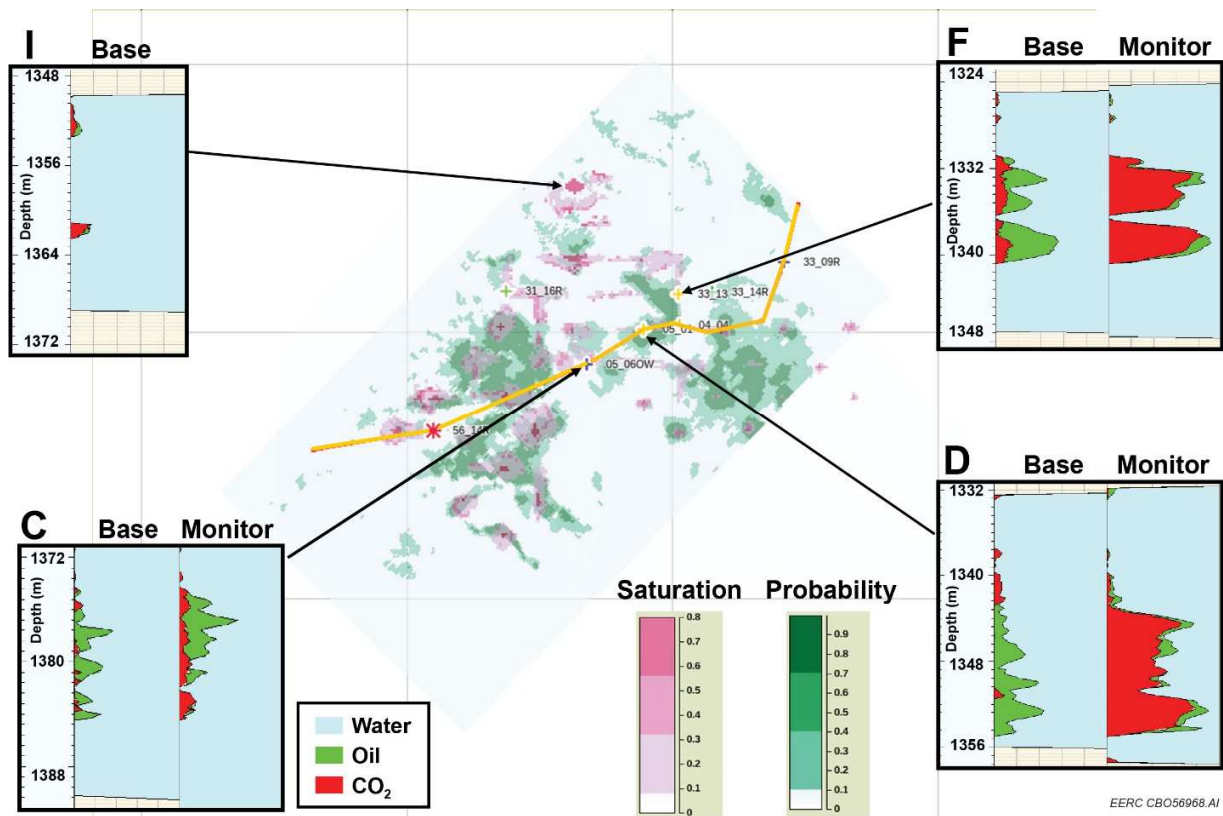
**Figure 19.** Top. AI estimated by Ji-Fi at the reservoir level (6 ms down from the reservoir) for the baseline (left) and monitor (right) seismic data sets. Bottom: Most probable facies estimated by Ji-Fi at the reservoir level (6 ms down from the top of the reservoir) for the baseline (left) and monitor (right) seismic data sets. The symbols (\* and +) represent the location of the wells used in this study. The pink line on all the maps represents the location of the arbitrary line shown in Figure 18.

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The good performance of the Ji-Fi algorithm should be noted. The estimated AI decreases (Fig. 19 top) because of reservoir fluids changes in the monitor survey. Several parameterizations of Ji-Fi were tested for the baseline and monitor (not shown here). Combinations of parameters generated some variations in the estimated facies and acoustic impedances. Eight Ji-Fi realizations were combined in a multirealization scheme to estimate the most probable facies and the probabilities of the barrier bar with different fluids.

901 **7. FLUID DISTRIBUTION FROM Ji-Fi VS. DISTRIBUTION FROM RESERVOIR**  
 902 **SIMULATION**

903 The probabilities of facies fluid distribution in the Bell Creek Reservoir estimated by Ji-Fi  
 904 are compared here with gas saturation distribution derived from reservoir simulations in  
 905 Fig. 20. The gas distribution from reservoir simulation corresponds to Layer 11 from October  
 906 2014, the same month when the seismic monitor survey was acquired. This layer was selected  
 907  
 908



909 **Figure 20.** Comparison of gas saturation from reservoir simulation results at Layer K11 (October  
 910 2014) with estimated probabilities for free CO<sub>2</sub> and oil saturated with CO<sub>2</sub> using a Ji-Fi  
 911 multirealization approach. The Ji-Fi results overlay the reservoir simulation results. Green bar: Ji-  
 912 Fi, pink bar: reservoir simulation. Time-lapse PNL of the C, D, and F wells from the same time of  
 913 acquisition as the seismic survey. The PNL from the I (CO<sub>2</sub> injector) well is only available after  
 914 the CO<sub>2</sub> injection.  
 915

916

917 because of its location approximately at the middle of the reservoir interval. The Ji-Fi probability  
918 fluid distribution is shown by a horizon with the same geometry as the Springen Ranch horizon,  
919 representing the reservoir top. The Ji-Fi horizon was placed 6 ms down from the top of the  
920 reservoir. It should be noted that although the geometry of Layer 11 and the Ji-Fi horizon are not  
921 the same, it was expected that both surfaces cross similar features of the reservoir. Overlaying the  
922 Ji-Fi facies probabilities horizon on the gas distribution map for free CO<sub>2</sub>, oil saturated with CO<sub>2</sub>,  
923 and the combination of free CO<sub>2</sub> with oil saturated with CO<sub>2</sub> shows a good correlation with the  
924 spatial distribution of the gas saturation from reservoir simulation, which confirms the ability of  
925 Ji-Fi to detect the changes in fluid saturation conditions in the reservoir due to CO<sub>2</sub> injection.

926         However, some discrepancies in the correlation are observed in the northwestern part of  
927 Phase 1. PNLs of some wells in the study area were used to review the geologic and fluid  
928 conditions in the reservoir at selected locations and assess the discrepancies. Fig. 20 shows the  
929 available PNLs of the C, D, F, and I wells and their location in the Ji-Fi reservoir simulation  
930 comparison map. Whereas the PNLs of the D and F-wells show a well-defined saturation in the  
931 reservoir, and therefore, a high-probability of reservoir fluids estimated by Ji-Fi, the C and I PNLs  
932 show a small saturation area. Only one PNL was available from the I well, where the main  
933 discrepancy in the correlation of Ji-Fi versus reservoir simulation was observed. If the small  
934 saturation area observed in the PNL is considered a reference for the reservoir's thickness, it is  
935 expected that such thickness is below the resolution of the seismic method.

936         In addition to the saturation effects observed in the reservoir simulation and Ji-Fi  
937 comparison, it is expected that some pressure effects are also in the seismic signal. Mur et al.  
938 (2020) have reported encouraging initial results of the Ji-Fi application to decouple pressure and

939 fluid effects in the study area. Extending the Ji-Fi method for quantitative pressure detection  
940 requires further detailed studies, analysis, and prior conditioning (Mur et al., 2020).

## 941 **8. CONCLUSIONS**

942 The Ji-Fi method has an integral representation of the geology in the inversion algorithm  
943 using elastic facies, which provides information about the spatial distribution of the geologic  
944 heterogeneities controlling the distribution of fluids in the reservoir.

945 In this paper, a reservoir properties-focused Ji-Fi workflow was introduced to derive the  
946 optimum input data for applying Ji-Fi to the baseline and monitor surveys from the study area.  
947 Comprehensive rock physics analyses using core data and well logs were conducted to establish  
948 the background geologic information and estimate the Bell Creek reservoir litho-fluid facies. This  
949 information was used to create a DTA. Facies model uncertainties were assessed using the facies-  
950 dependent probability density functions to perform Bayesian classification along with existing  
951 upscaled well elastic data. Good agreement was found between the upscaled petrophysical facies  
952 log and the Bayesian classification log with associated probabilities for each facies. The facies and  
953 their probability distributions corresponding to the highly heterogeneous target reservoir and its  
954 fluid conditions due to the CO<sub>2</sub> EOR activities were successfully separated in the seismic elastic  
955 space. This separation provided favorable conditions for the application of the Ji-Fi method.

956 Separate Ji-Fi of the baseline and monitor surveys and the differencing analysis to capture  
957 changes due to CO<sub>2</sub> were performed. Single and multirealization inversion schemes were also run  
958 and analyzed. Reliable acoustic impedances and most probably litho-fluid facies of the target  
959 reservoir were obtained. Information about the probability of the presence of litho-fluid facies is  
960 paramount to reducing uncertainty in forecasting CO<sub>2</sub> saturation changes within the target  
961 reservoir.

962           The successful application of the Ji-Fi method to time-lapse seismic data from a mature oil  
963 field undergoing CO<sub>2</sub> EOR shown in this paper suggests the critical role this type of technology  
964 can play in carbon capture, utilization, and storage (CCUS) projects to reduce the uncertainty in  
965 the detection and estimation of CO<sub>2</sub> distribution.

966           Since the most common technique to build an LFM (background model) for seismic  
967 inversion to estimate absolute elastic rock properties is interpolating well data with a structural  
968 and stratigraphic framework (Sams and Carter, 2017), when few wells or no wells are available in  
969 CCUS study areas, LFM building can be more problematic.

970           Estimating relative rock properties without building an LFM is an alternative in such  
971 scenarios. A Ji-Fi scheme can also be considered in such well scenarios as the LFM is a product  
972 of the inversion, and it is constrained by the elastic property depth trends associated with each  
973 expected facies and the inverted distribution of the facies.

974           The ability to derive the input LFM trends for Ji-Fi from few logs, 1D basin models, or  
975 analog fields and incorporate additional information in a Bayesian framework as it is available  
976 result in a good seismic inversion alternative.

977           Future enhancements of the Ji-Fi technology include the simultaneous inversion of the  
978 baseline and monitor surveys to relax the time-lapse data acquisition and processing constraints  
979 and reduce cost and turnaround time to deliver results to operators and deeper integration of time-  
980 lapse seismic inversion with reservoir simulation as suggested by Mur et al. (2020).

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