



# Constraining subglacial hydrology with ice surface velocity observations by solving an inverse ice dynamics-subglacial hydrology problem

L.Bertagna<sup>1</sup>, M.F.Hoffmann<sup>2</sup>, M.Perego<sup>1</sup>

<sup>1</sup>Sandia National Laboratories, Albuquerque, NM,

<sup>2</sup>Los Alamos National Laboratories, Albuquerque, NM

AGU21, Dec 13th, 2021

SAND SAND2021-00000



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# The First Order (FO) model

For the ice dynamics, we consider the usual First-Order model (Blatter 1995, Pattyn 2003) for the horizontal velocity  $\mathbf{u} = (u_x, u_y)$

$$-\nabla \cdot \underline{\underline{\sigma}} = -\rho_i g \nabla s$$

where  $s$  is the surface elevation, and

$$\underline{\underline{\sigma}} = 2\mu(\mathbf{u})\underline{\underline{\varepsilon}}, \quad \underline{\underline{\varepsilon}} = \begin{bmatrix} 2\varepsilon_{xx} + \varepsilon_{yy} & \varepsilon_{xy} & \varepsilon_{xz} \\ \varepsilon_{xy} & \varepsilon_{xx} + 2\varepsilon_{yy} & \varepsilon_{xz} \\ \varepsilon_{xz} & \varepsilon_{xz} & \varepsilon_{zz} \end{bmatrix}, \quad \mu(\mathbf{u}) = \frac{1}{2} A^{-\frac{1}{n}}(T) \|\underline{\underline{\varepsilon}}\|^{\frac{1}{n}-1}$$

At the ice-bedrock interface, we consider the sliding condition

$$\underline{\underline{\sigma}}\mathbf{n} + \beta\mathbf{u} = \mathbf{0}, \quad \beta > 0.$$

Constant field  $\beta = \beta(\mathbf{x})$

- Simple mathematical model, but pay price of high-dimension parameter space

Functional form  $\beta = \beta(\mathbf{u}, \mathbf{p})$

- Complex physically-informed model, but (hopefully) get a low-dimension parameter space

We consider a steady distributed model (Hewitt 2013) for subglacial hydrology.

$$\begin{cases} \cancel{\frac{\partial h}{\partial t}} + \nabla \cdot \mathbf{q} = \frac{m}{\rho_w} & \text{mass conservation} \\ \cancel{\frac{\partial h}{\partial t}} = \frac{m}{\rho_i} + \frac{h_r - h}{l_r} |u| - c_c A h N^3 & \text{cavities evolution} \end{cases}$$

with melting  $m = (G + \beta |u|^2)/L$ , water discharge  $\mathbf{q} = -k h^{\alpha_1} |\nabla \Phi|^{\alpha_2} \nabla \Phi$ , ice thickness  $H$ , transmissivity  $k$ , source geothermal flux  $G$ , ice sliding velocity  $u$ , bed bumps height/length  $h_r, l_r$  are given, effective pressure  $N = \rho_i g H - \Phi$ .

The unknowns are water hydropotential  $\Phi$ , and water thickness  $h$ .

**Idea:** perform data assimilation of surface velocity measures to identify uncertain parameters in *both* hydrology and sliding law, by solving a *coupled* FO-Hydrology problem.

**Method:** solve a deterministic inverse problem, by minimizing mismatch between computed and observed surface velocity. Use  $\mu_f$  and  $k$  as control variables.

$$[\mu, k] = \arg \min \int_{surf} |\mathbf{u} - \mathbf{u}_{obs}|^2 dx + R(\mu, k)$$
$$\text{s.t.} \begin{cases} -\nabla \cdot \underline{\underline{\sigma}} = -\rho_i g \nabla s \\ \nabla \cdot \mathbf{q} = \frac{m}{\rho_i} \\ \frac{m}{\rho_i} + \frac{h_r - h}{l_r} |u| - c_c A h N^3 = 0 \\ \underline{\underline{\sigma}} \mathbf{n} + \beta \mathbf{u} = \mathbf{0} \end{cases}$$

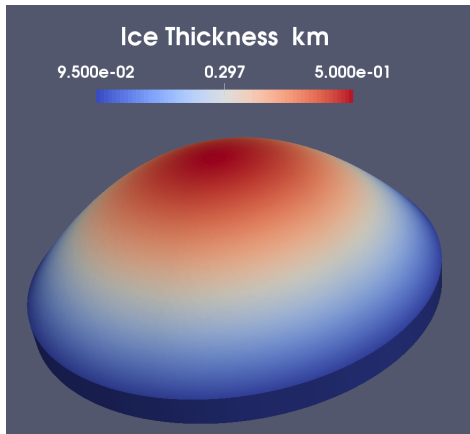
- Pre-processing: solve FO+Hydro for given  $(\mu_f, k)$  to generate surface velocity measures. Add fudge factor to measures, to emulate for measurements errors.
- Algorithm: Quasi-Newton (BFGS) with backtrack line search.

## Software

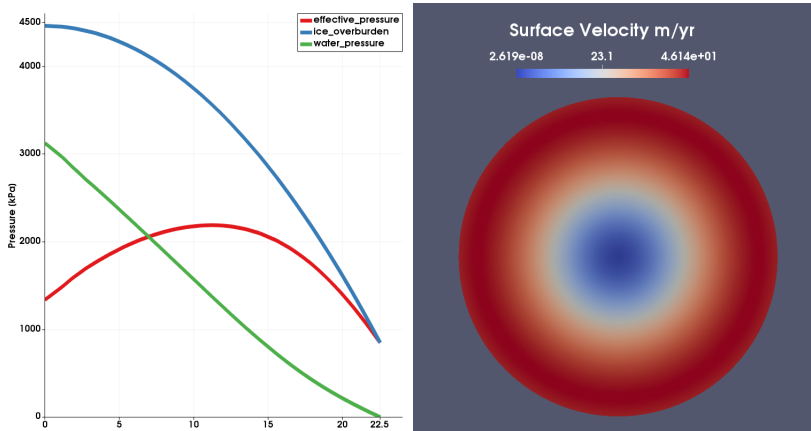
- **Trilinos**: general-purpose scientific library for high-performance distributed computing. Provides linear/nonlinear algebra, solvers, discretization tools.
- **Albany**: parallel finite element library (built on Trilinos), with analysis/sensitivities capabilities, timestepping, continuation, optimization, and more.
- **LandIce**: a package in Albany which implements ice-sheet and subglacial hydrology models.

# Idealized Dome: problem specs

- radially symmetric geometry:  
 $H(r) = s(r) = 0.5 (1 - r^2) / R^2$  km
- horizontal resolution: 1km
- prescribed uniform melting in mass eqn:  
 $m/\rho_w = 5.4$  mm/day
- no melting in cavities eqn
- power-law sliding:  $\beta = \mu_f Nu^{q-1}$ ,  $q = 4/3$
- linear water flux:  $\mathbf{q} = -kh\nabla\Phi$ .



# Idealized Dome: forward solve



**Figure:** Water/effective pressures (left) and surface velocity (right) obtained solving the coupled FO+Hydrology problem for  $k = 1e - 5$  and  $\mu_f = 0.01$ .

# Idealized dome: inverse problem

	Fwd	$\eta = 0$	$\eta = 5\%$	$\eta = 10\%$	$\eta = 20\%$
$\mu_f$	1e-2	0.994e-2	0.994e-2	0.995e-2	0.993e-2
$k$	1e-5	0.994e-5	0.994e-5	0.995e-5	0.993e-5
# BFGS it	-	39	39	40	35

Table: Estimated parameters for different choices of noise level.

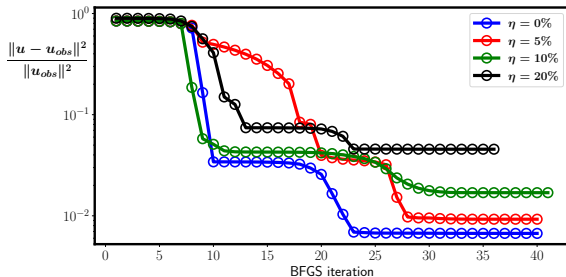
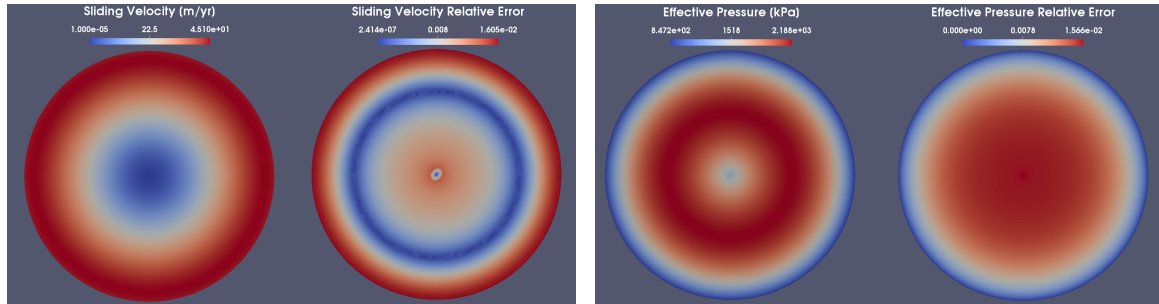


Figure: Convergence history of BFGS for different choices of noise level.



# Idealized dome: inverse problem



**Figure:** Exact solution and relative error plots for sliding velocity (left) and effective pressure (right), with 20% relative noise in the surface velocity measures.

# Considerations and future directions

- FO+Hydro solver in LandIce is robust on idealized geometry.
- Small parameter space makes inversion feasible.
- Next: synthetic problem on a real geometry (Humboldt).
- Next: assimilate real measures on a real geometry.