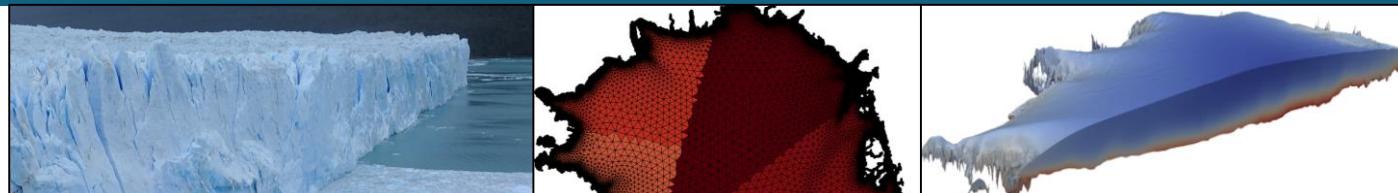




Sandia
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Improving Large-Scale PDE-Constrained Optimization for Ice Sheet Initialization



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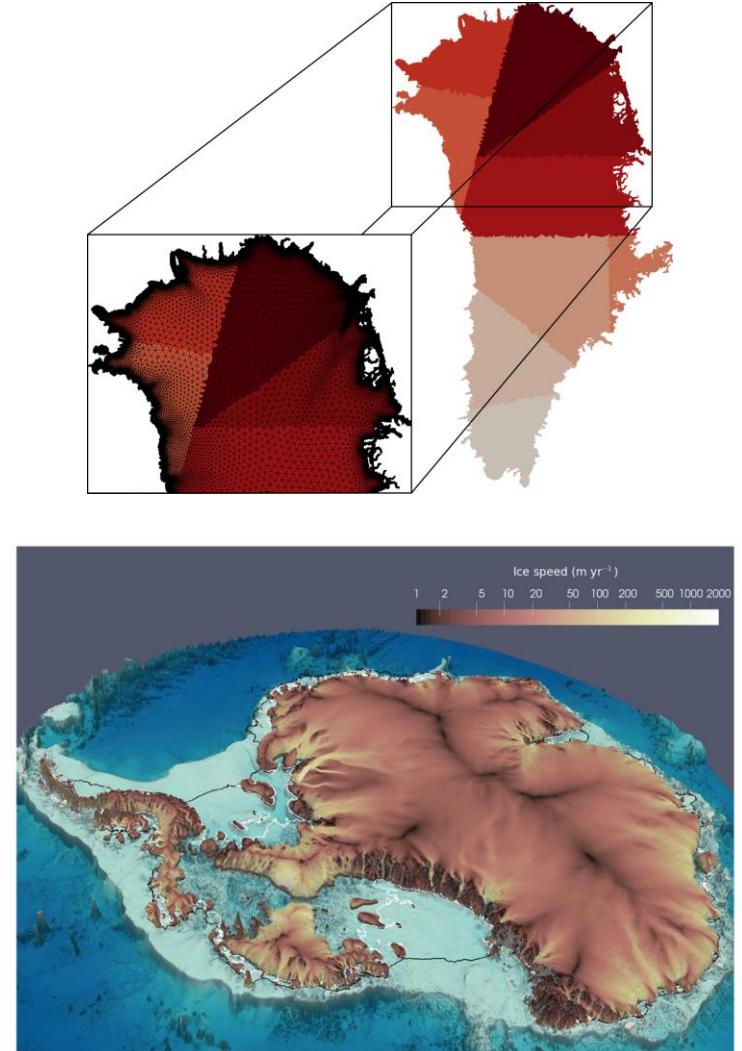
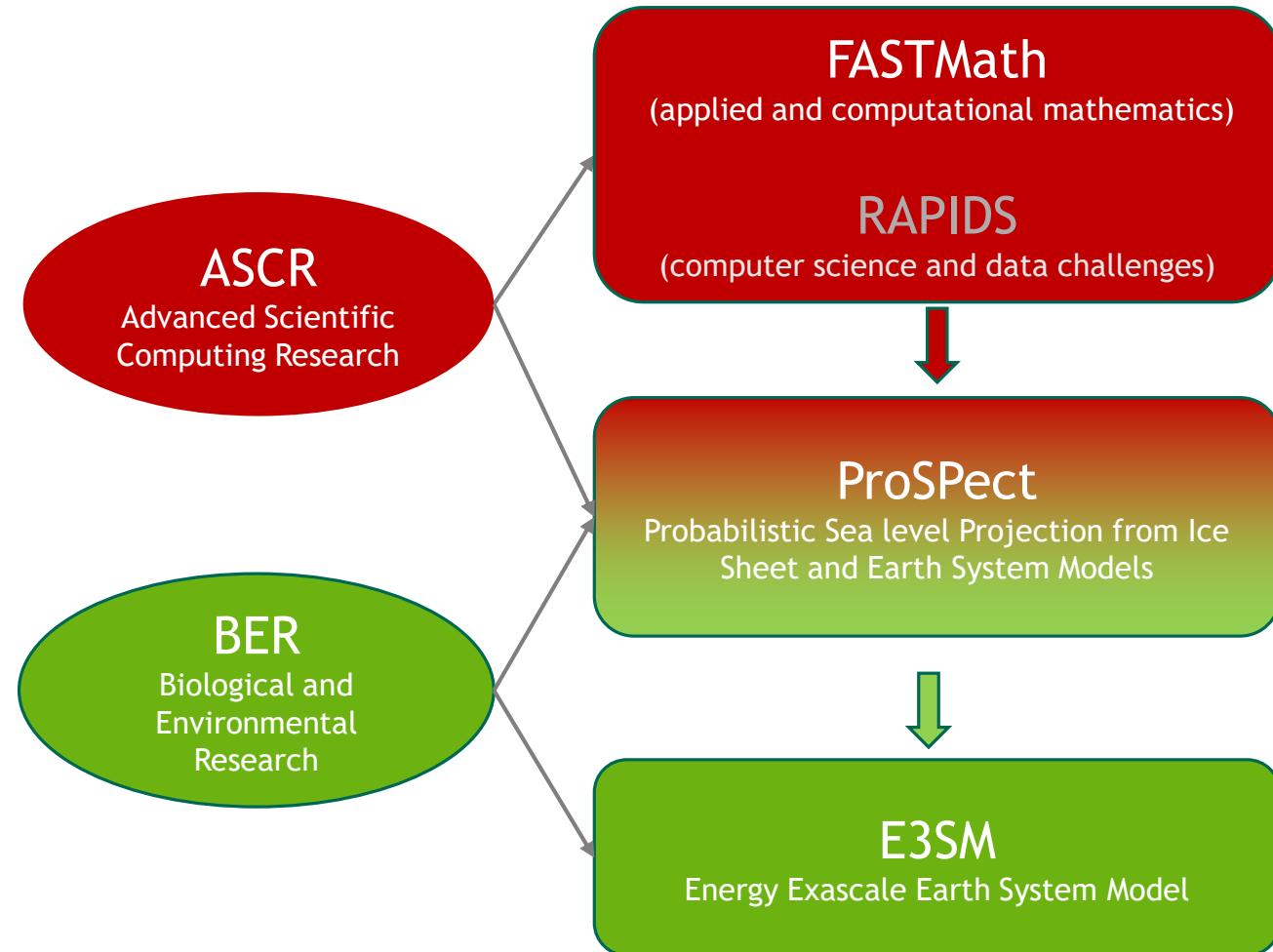


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- Overview of DOE programs impacting Land Ice modeling
- Brief motivation and introduction to ice sheet models
- Ice sheet initialization
- Improvements enabled by FASTMath institute:
 - Performance speed-up in assembly phase
 - Novel preconditioner for thermo-mechanical problem
 - Newton-Krylov approaches for PDE-constrained optimization

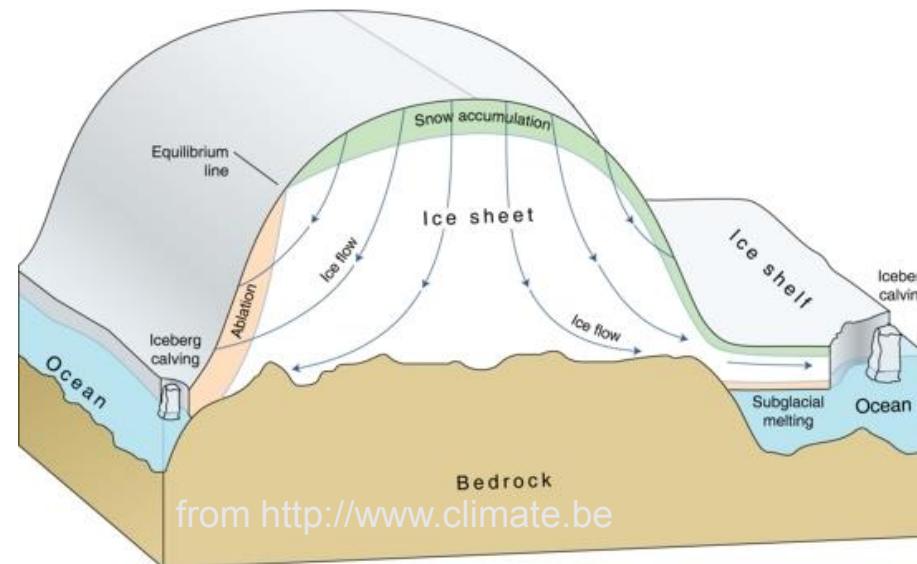
DOE Land Ice Modeling Efforts



Brief Motivation an basic physics



- Modeling ice sheets (Greenland and Antarctica) dynamics is essential to provide estimates for sea-level rise in next decades to centuries.
- Ice behaves like a very viscous shear-thinning fluid (similar to lava flow) driven by gravity.
- Several unknown or poorly known parameters (e.g. basal friction, bed topography) and processes (calving laws, basal hydrology)



from <http://www.climate.be>

Model: Ice velocity equations

Stokes equations:

$$\begin{cases} -\nabla \cdot \sigma = \rho \mathbf{g} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}$$

ice velocity gravit. acceleration



Stress tensor:

$$\sigma = 2\mu \mathbf{D} - pI, \quad \mathbf{D}_{ij}(\mathbf{u}) = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

Ice viscosity (dependent on temperature):

$$\mu = \frac{1}{2} A(T) |\mathbf{D}(\mathbf{u})|^{\frac{1}{n}-1}, \quad n \geq 1, \quad (\text{typically } n \simeq 3)$$

In this work we use a simplification of Stokes equations, called **First Order** equations, obtained by scaling arguments given the shallow nature of the ice sheets and using hydrostatic pressure.

Model: Ice velocity equations



Stokes equations:

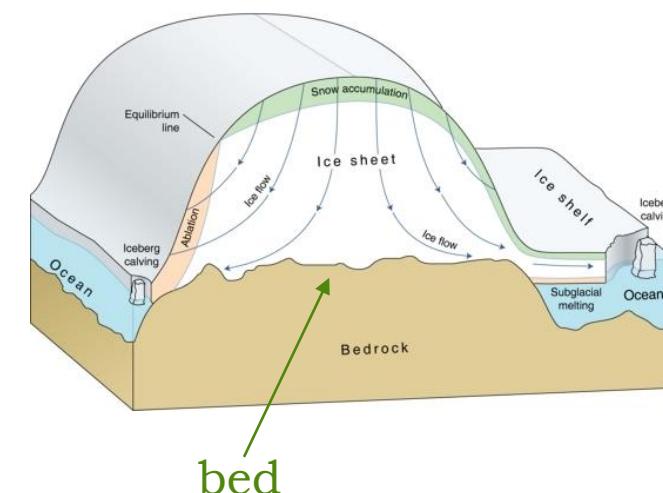
$$\begin{cases} -\nabla \cdot \sigma = \rho \mathbf{g} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}$$

Sliding boundary condition at ice bed:

$$\begin{cases} \mathbf{u} \cdot \mathbf{n} = 0, \quad (\text{impenetrability}) \\ (\sigma \mathbf{n})_{\parallel} = \beta \mathbf{u} \end{cases}$$

Free slip: $\beta = 0$

No slip: $\beta = \infty$



Model: Temperature equation



Heat equation (for cold ice):

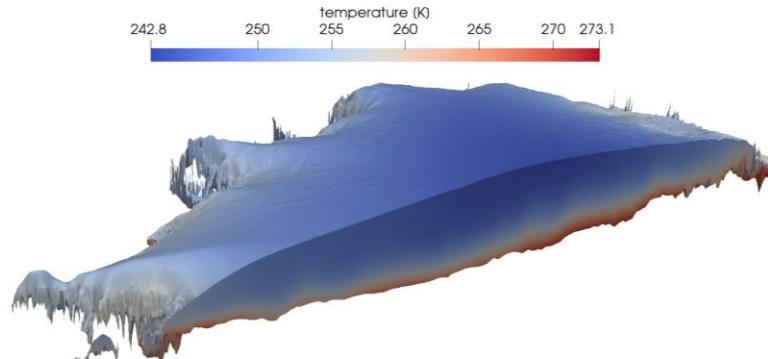
$$\rho c \partial_t T + \nabla \cdot (k \nabla T) + \rho c \mathbf{u} \cdot \nabla T = 4\mu |D(\mathbf{u})|^2$$

↑ conductivity ↑ heat capacity ↓ dissipation heating

Boundary condition at the ice bed
(includes melting and refreezing):

$$m = G + \beta |\mathbf{u}|^2 - k \nabla T \cdot \mathbf{n}$$

↑ melting rate ↑ geothermal heat flux ↑ frictional heating
 ↓ temperature flux



In this work we use a enthalpy formulation that accounts for temperate ice as well.

8 Ice sheet initialization



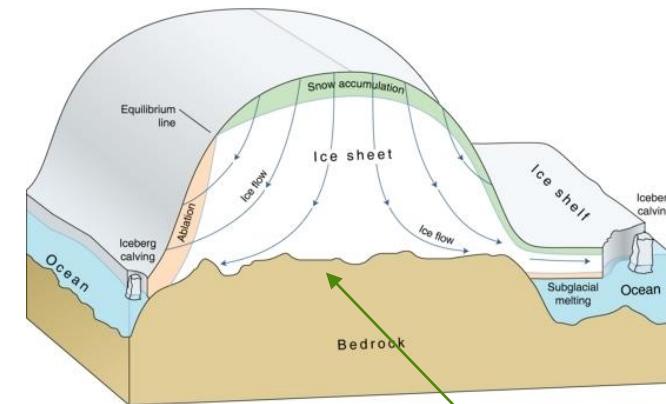
Goal: Find the initial/present-day thermo-mechanical state of the ice sheet and estimate the unknown/poorly known model parameters, by matching observations

Approach: **PDE-constrained optimization**

Find basal friction coefficient β that minimizes the mismatch with surface velocity:

$$\min_{\beta} J(\beta) = \int_{\Omega} \frac{|u - u_{obs}|^2}{\sigma^2} + R(\beta)$$

Subject to the coupled velocity/temperature problem



unknown sliding parameter β

Software Requirements

- Large Scale optimization library (ROL), featuring gradient-based methods (ROL)
- Computation of gradients of the PDE residual and the loss functional w.r.t. the solution and the parameters. **Automatic Differentiation** is crucial for complex physics
- Faster, more robust methods available using **Hessian** (second derivatives)

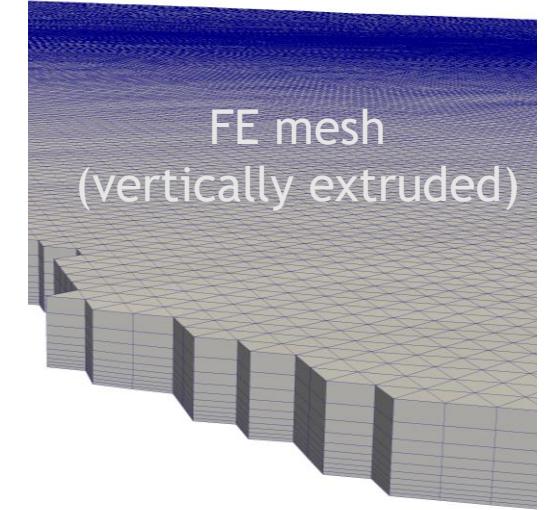


RAPID OPTIMIZATION LIBRARY

Software: MPAS-Albany Land Ice model (MALI)



ALGORITHM	SOFTWARE TOOLS
Linear Finite Elements on tets/prisms	Albany Land Ice
Optimization	ROL
Nonlinear solver (Newton method)	NOX
Krylov linear solvers/Prec	Belos/MueLu, Belos/FROSch
Automatic differentiation	Sacado



MPAS (Model for Prediction Across Scales): *Fortran, finite volumes* library, conservative Lagrangian schemes for advecting tracers (evolution of ice thickness)

Albany Land Ice: C++ finite element library built on top of **Trilinos** achieving performance portability through **Kokkos** programming model. Provides large scale PDE constrained optimization capabilities

References:

- Hoffman, et al. *GMD*, 2018
- Tuminaro, Perego, Tezaur, Salinger, Price, *SISC*, 2016.
- Tezaur, Perego, Salinger, Tuminaro, Price, Hoffman, *GMD*, 2015
- Perego, Price, Stadler, *JGR*, 2014

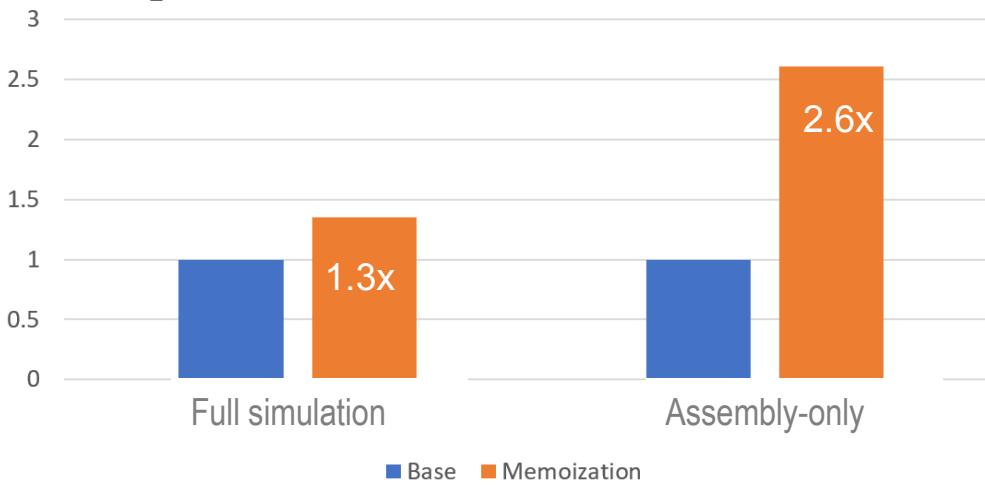


Performance improvements on CPUs

(Greenland glacier, velocity solver)

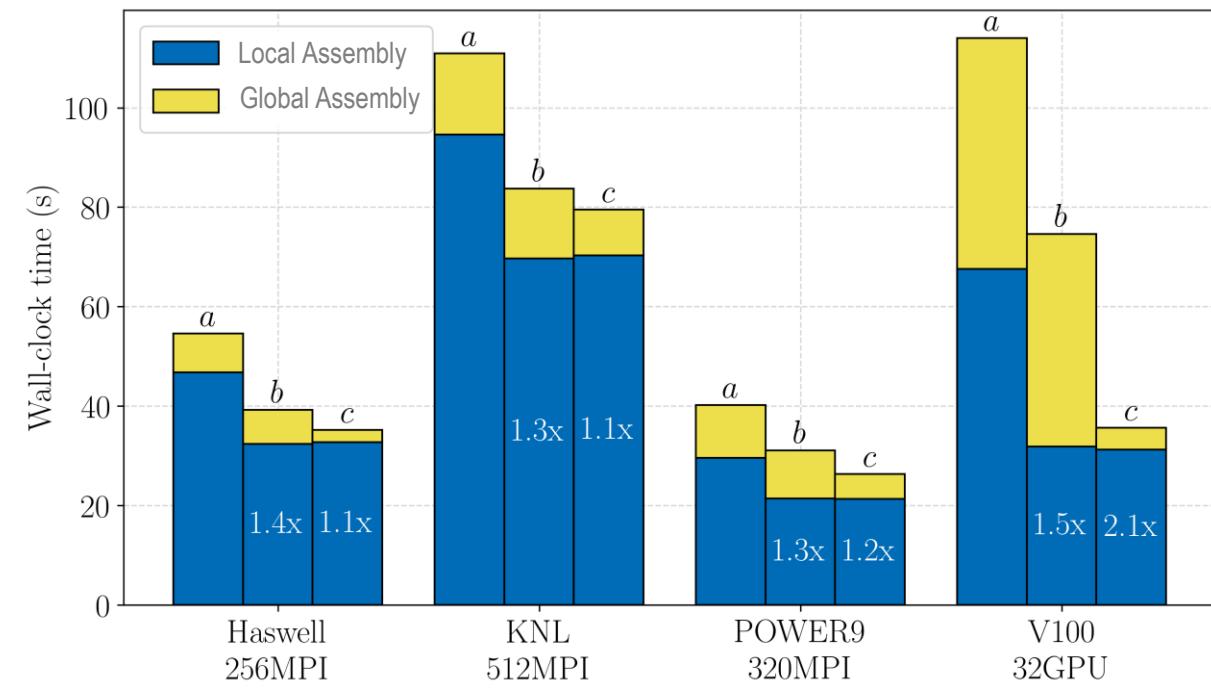


Improvements to initialization
problem due to **memoization**:



AIS initialization 10 iterations
(HSW, 128 nodes, 4096 procs)

Combined improvements due to **memoization** and
improved storage of FE matrix (Tpetra FE Crs Matrix)

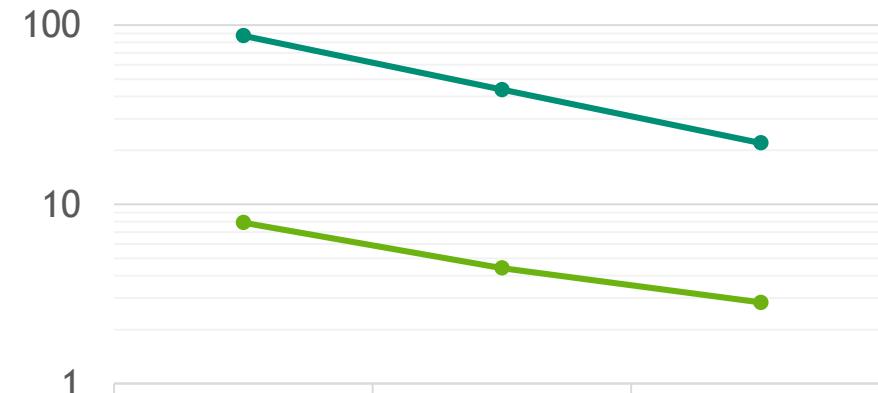


Assembly speed up on GPUs

(Thwaites glacier, velocity solver)

Haswell Node: 32 CPU cores/node
 V100 Node: 4 GPUs/node

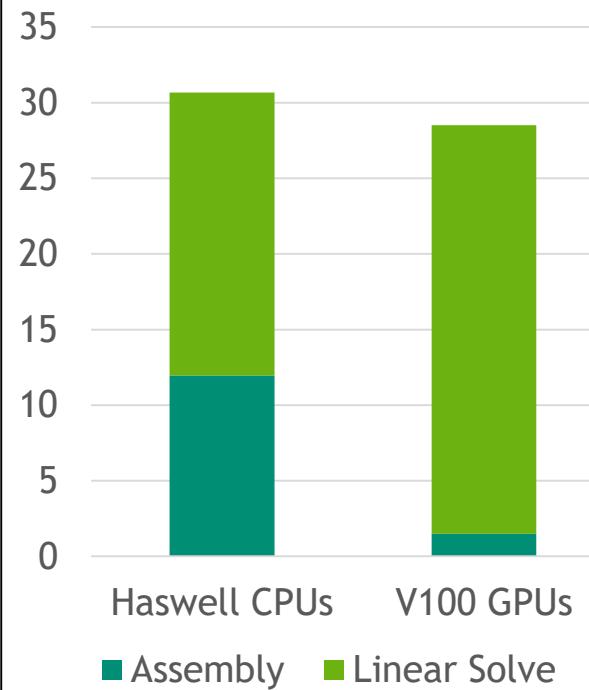
Finite Element Assembly Strong Scaling
 Wall-clock time (s) vs. Nodes



	4	8	16
Haswell CPUs	87.2051	43.6626	22.0101
V100 GPUs	7.89909	4.41412	2.8328
Speedup	11.03989194	9.891575218	7.769733126
DoFs/GPU	87230	43615	21807

Up to **11x** speedup for assembly on GPUS.
 Time to work on linear solvers!

Total Solve Time (s)



Two-level Schwarz preconditioner (FROSch) (Greenland Ice sheet, coupled velocity/temperature problem)

Challenge: our “workhorse” multigrid preconditioner does not work out of the box for the coupled problem.

Solution: FROSch (Fast and Robust Overlapping Schwarz)

Single-level Schwarz preconditioner vs Two-levels Schwarz preconditioner

MPI ranks	avg. #iters	avg. setup time	avg. solve time
512	48.7	45.0	11.3s
1024	61.9	54.3	5.29s
2048	89.9	59.1	2.52s
4096	116.1	78.7	1.17s
			10.53s
			5.41 s
			5.36s
			4.59s
			4.75s
			4.31s
			2.32s
			5.70s
			3.99
			1.37s
			3.68
			3.30

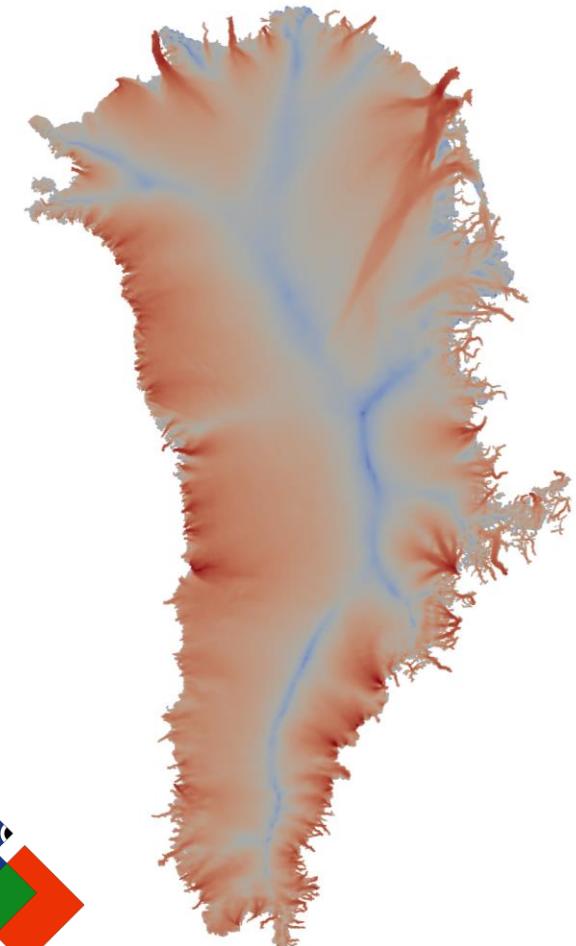
- The two-level preconditioner does better in terms of number of iterations but it is almost equivalent in term of CPU time to the one-level one.
- Time is dominated by direct solves of the sub-problems and the coarse solve.

Simulations by **A. Heinlein**

Support: CMDV-SM/FASTMath/ProSPect

Resources: Cori (NERSC)

Reference: Heinlein, Perego, Rajamanickam, arXiv, 2021



Ice sheet initialization

Hessian computation using automatic differentiation (using Sacado package)

Capabilities required for Newton-Krylov optimization methods:

Hessian of residual \mathbf{f} dotted with the Lagrange multiplier $\boldsymbol{\lambda}$ in the direction \boldsymbol{v} :

$$\begin{aligned} \partial_{\mathbf{u}\mathbf{u}}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \boldsymbol{v}, \quad \partial_{\mathbf{u}\mathbf{p}}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \boldsymbol{v}, \\ \partial_{\mathbf{p}\mathbf{u}}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \boldsymbol{v}, \quad \partial_{\mathbf{p}\mathbf{p}}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \boldsymbol{v} \end{aligned}$$

Hessian of loss functional J in the direction \boldsymbol{v} :

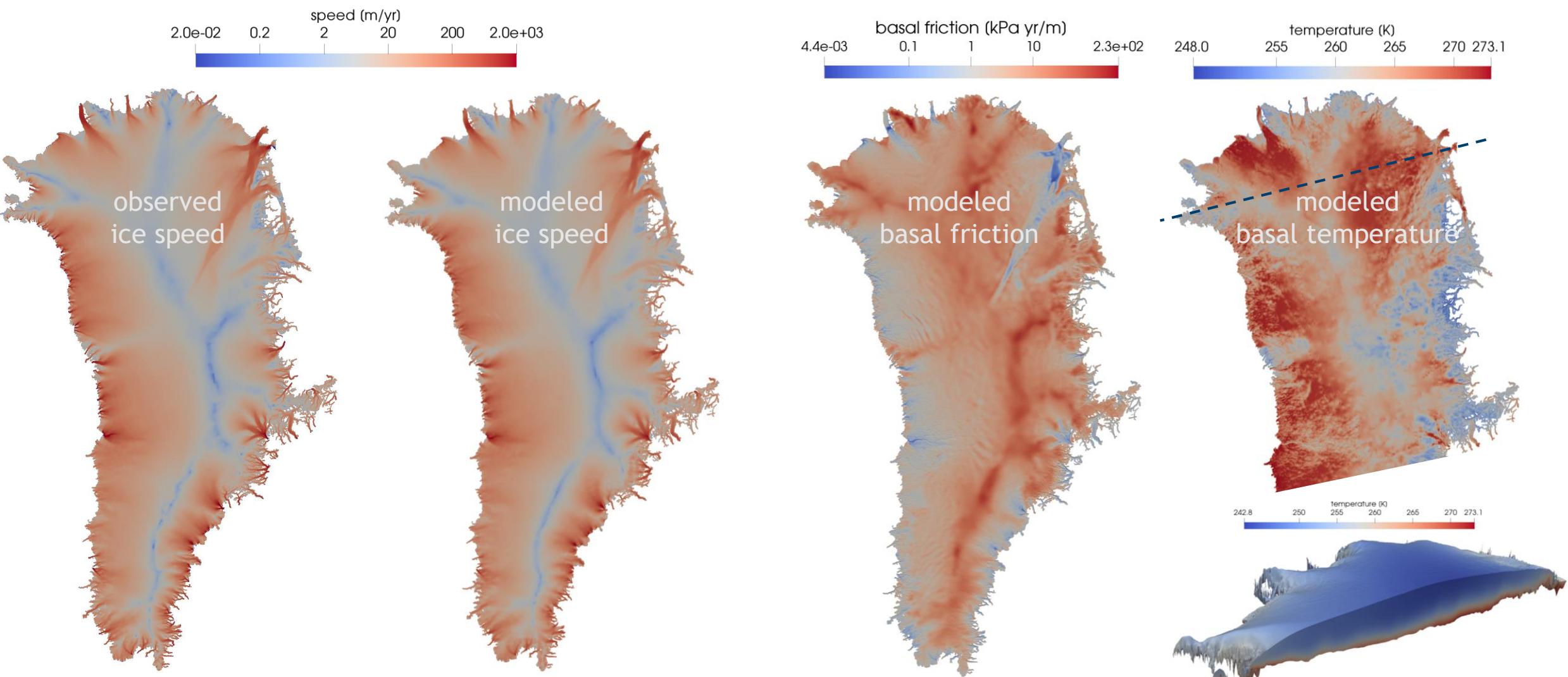
$$\begin{aligned} \partial_{\mathbf{u}\mathbf{u}} J(\mathbf{u}, \mathbf{p}) \boldsymbol{v}, \quad \partial_{\mathbf{u}\mathbf{p}} J(\mathbf{u}, \mathbf{p}) \boldsymbol{v}, \\ \partial_{\mathbf{p}\mathbf{u}} J(\mathbf{u}, \mathbf{p}) \boldsymbol{v}, \quad \partial_{\mathbf{p}\mathbf{p}} J(\mathbf{u}, \mathbf{p}) \boldsymbol{v}, \end{aligned}$$

Computed w/ **automatic differentiation**, based on the formula:

$$\partial_{\mathbf{p}\mathbf{p}} J(\mathbf{p}) \boldsymbol{v} = \partial_r \left(\partial_{\mathbf{p}} J(\mathbf{p} + r \boldsymbol{v}) \right) \Big|_{r=0}$$

We also provide the full matrix $\partial_{\mathbf{p}\mathbf{p}} J$, that can be used to initialize reduced-space optimization methods like BFGS, and the *reduced* Hessian-vector product.

Thermo-mechanical initialization of Greenland ice sheet



300K parameters, 14M unknowns. Initialization: ~10 hours on 8k nodes on NERSC Cori (Haswell)

Support: FASTMath/E3SM/ProSPect

MS270 on Thursday targets ice sheet modeling