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Improving Large-Scale PDE-Constrained Optimization for Ice Sheet Initialization



CCR
Center for Computing Research



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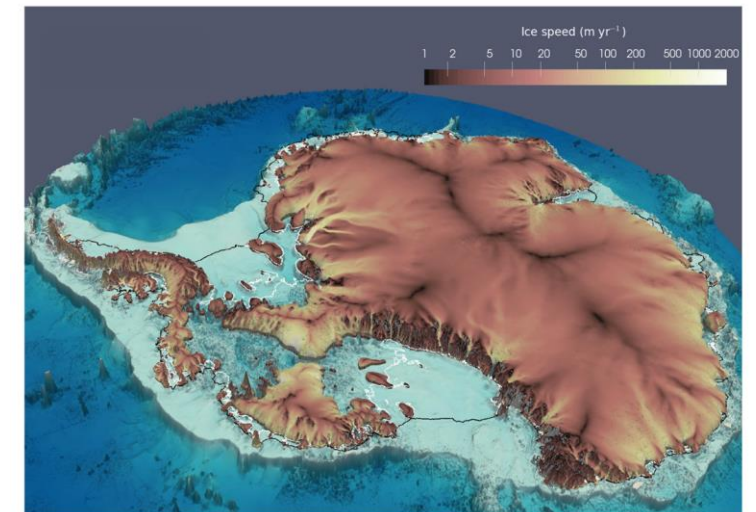
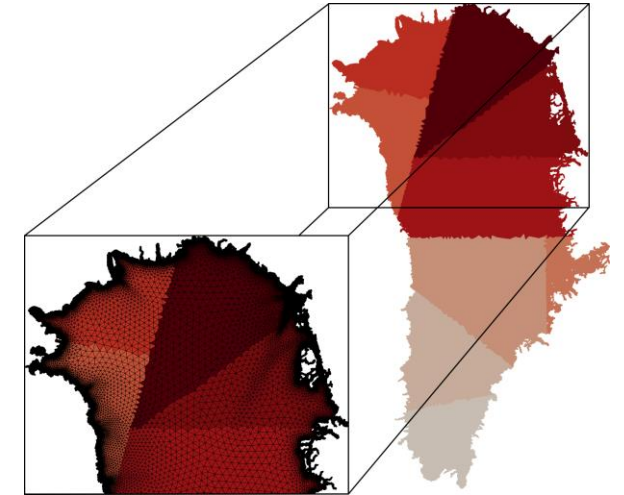
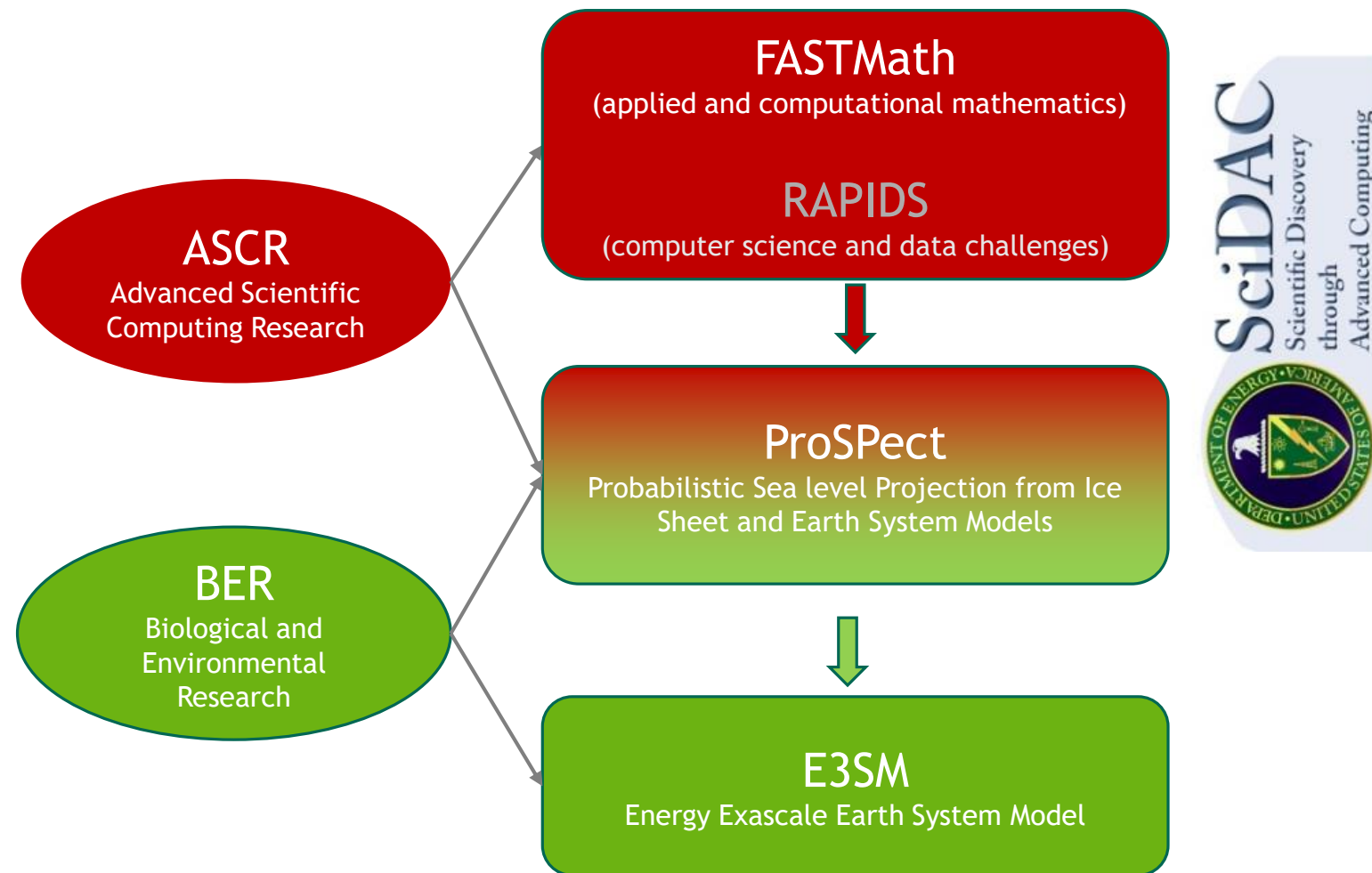


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- Overview of DOE programs impacting Land Ice modeling
- Brief motivation and introduction to ice sheet models
- Ice sheet initialization
- Improvements enabled by FASTMath institute:
 - Performance speed-up in assembly phase
 - Novel preconditioner for thermo-mechanical problem
 - Newton-Krylov approaches for PDE-constrained optimization

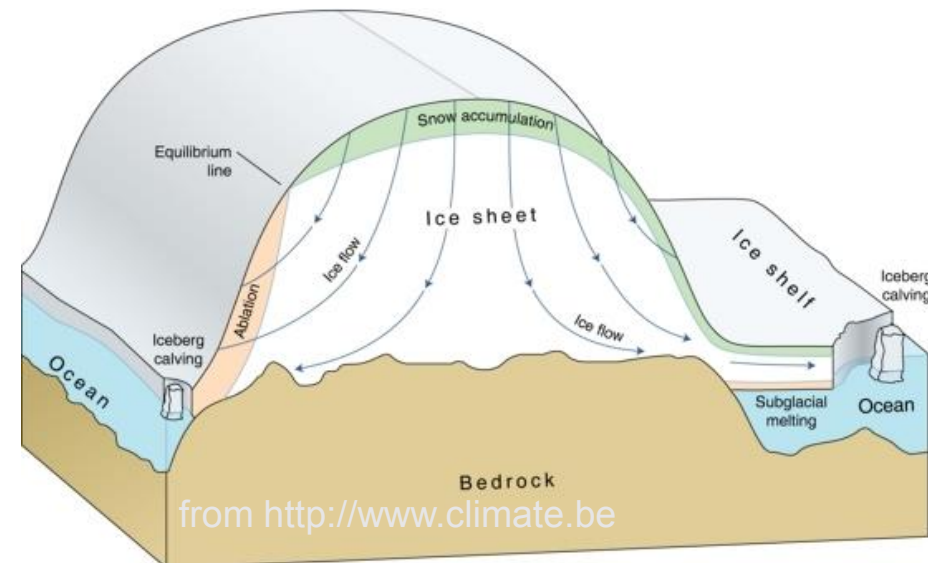
DOE Land Ice Modeling Efforts



Brief Motivation and basic physics



- Modeling ice sheets (Greenland and Antarctica) dynamics is essential to provide estimates for sea-level rise in next decades to centuries.
- Ice behaves like a very viscous shear-thinning fluid (similar to lava flow) driven by gravity.
- Several unknown or poorly known parameters (e.g. basal friction, bed topography) and processes (calving laws, basal hydrology)



Model: Ice velocity equations



Stokes equations:

$$\begin{cases} -\nabla \cdot \sigma = \rho \mathbf{g} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}$$

gravit. acceleration

ice velocity

Stress tensor:

$$\sigma = 2\mu \mathbf{D} - pI, \quad \mathbf{D}_{ij}(\mathbf{u}) = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

Ice viscosity (dependent on temperature):

$$\mu = \frac{1}{2} A(T) |\mathbf{D}(\mathbf{u})|^{\frac{1}{n}-1}, \quad n \geq 1, \quad (\text{typically } n \simeq 3)$$



In this work we use a simplification of Stokes equations, called **First Order** equations, obtained by scaling arguments given the shallow nature of the ice sheets and using hydrostatic pressure.

Model: Ice velocity equations

Stokes equations:

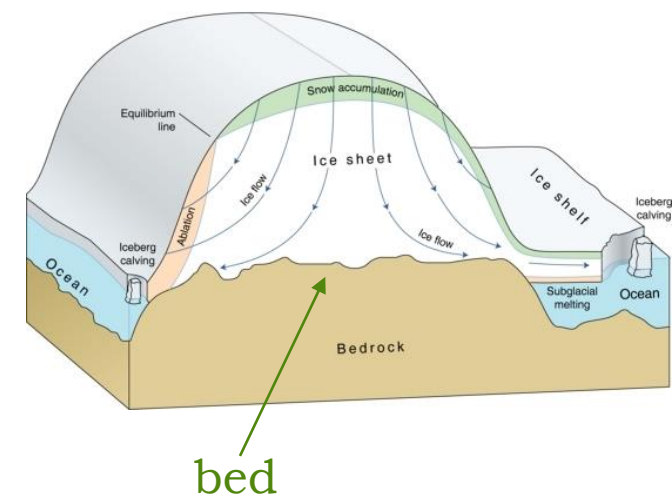
$$\begin{cases} -\nabla \cdot \sigma = \rho \mathbf{g} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}$$

Sliding boundary condition at ice bed:

$$\begin{cases} \mathbf{u} \cdot \mathbf{n} = 0, & (\text{impenetrability}) \\ (\sigma \mathbf{n})_{\parallel} = \beta \mathbf{u} \end{cases}$$

Free slip: $\beta = 0$

No slip: $\beta = \infty$



Model: Temperature equation



Heat equation (for cold ice):

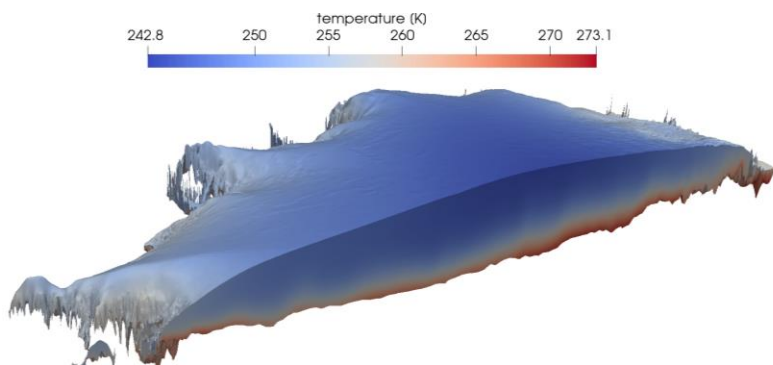
$$\rho c \partial_t T + \nabla \cdot (k \nabla T) + \rho c \mathbf{u} \cdot \nabla T = 4\mu |D(\mathbf{u})|^2$$

conductivity
heat capacity
dissipation heating

Boundary condition at the ice bed
(includes melting and refreezing):

$$m = G + \beta |\mathbf{u}|^2 - k \nabla T \cdot \mathbf{n}$$

melting rate
geothermal heat flux
frictional heating
temperature flux



In this work we use an enthalpy formulation that accounts for temperate ice as well.

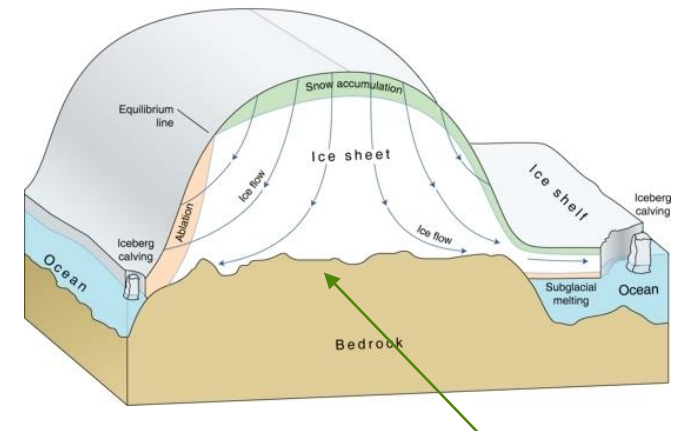
Goal: Find the initial/present-day thermo-mechanical state of the ice sheet and estimate the unknown/poorly known model parameters, by matching observations

Approach: **PDE-constrained optimization**

Find basal friction coefficient β that minimizes the mismatch with surface velocity:

$$\min_{\beta} J(\beta) = \int_{\Omega} \frac{|u - u_{obs}|^2}{\sigma^2} + R(\beta)$$

Subject to the coupled velocity/temperature problem



unknown sliding
parameter β

Software Requirements

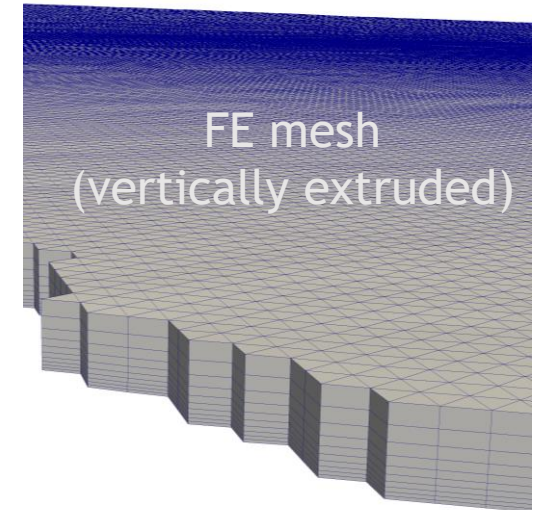
- Large Scale optimization library (ROL), featuring gradient-based methods (ROL)
- Computation of gradients of the PDE residual and the loss functional w.r.t. the solution and the parameters. **Automatic Differentiation** is crucial for complex physics
- Faster, more robust methods available using **Hessian** (second derivatives)



Software: MPAS-Albany Land Ice model (MALI)



ALGORITHM	SOFTWARE TOOLS
Linear Finite Elements on tets/prisms	Albany Land Ice
Optimization	ROL
Nonlinear solver (Newton method)	NOX
Krylov linear solvers/Prec	Belos/MueLu, Belos/FROSch
Automatic differentiation	Sacado



MPAS (Model for Prediction Across Scales): *Fortran*, **finite volumes** library, conservative Lagrangian schemes for advecting tracers (evolution of ice thickness)

Albany Land Ice: C++ finite element library built on top of **Trilinos** achieving performance portability through **Kokkos** programming model. Provides large scale PDE constrained optimization capabilities

References:

Hoffman, et al. GMD, 2018

Tuminaro, Perego, Tezaur, Salinger, Price, SISC, 2016.

Tezaur, Perego, Salinger, Tuminaro, Price, Hoffman, GMD, 2015

Perego, Price, Stadler, JGR, 2014

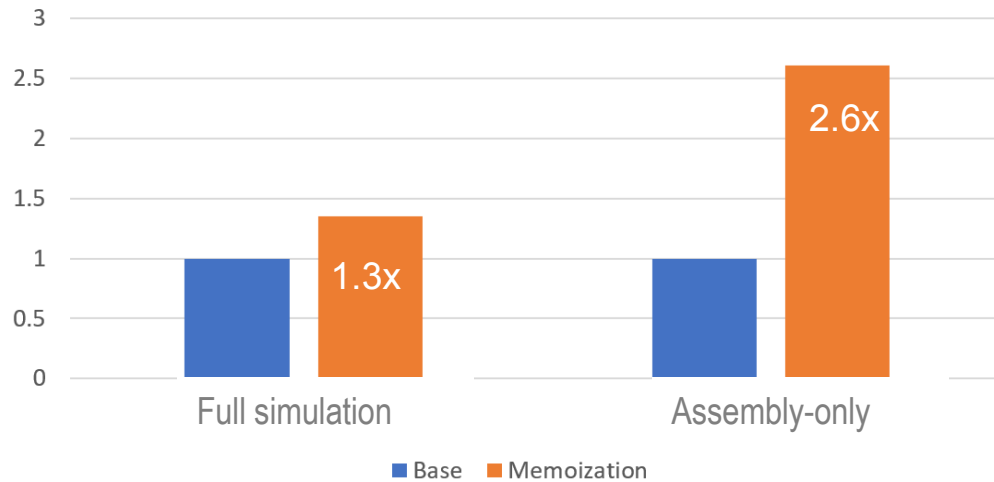


Performance improvements on CPUs

(Greenland glacier, velocity solver)

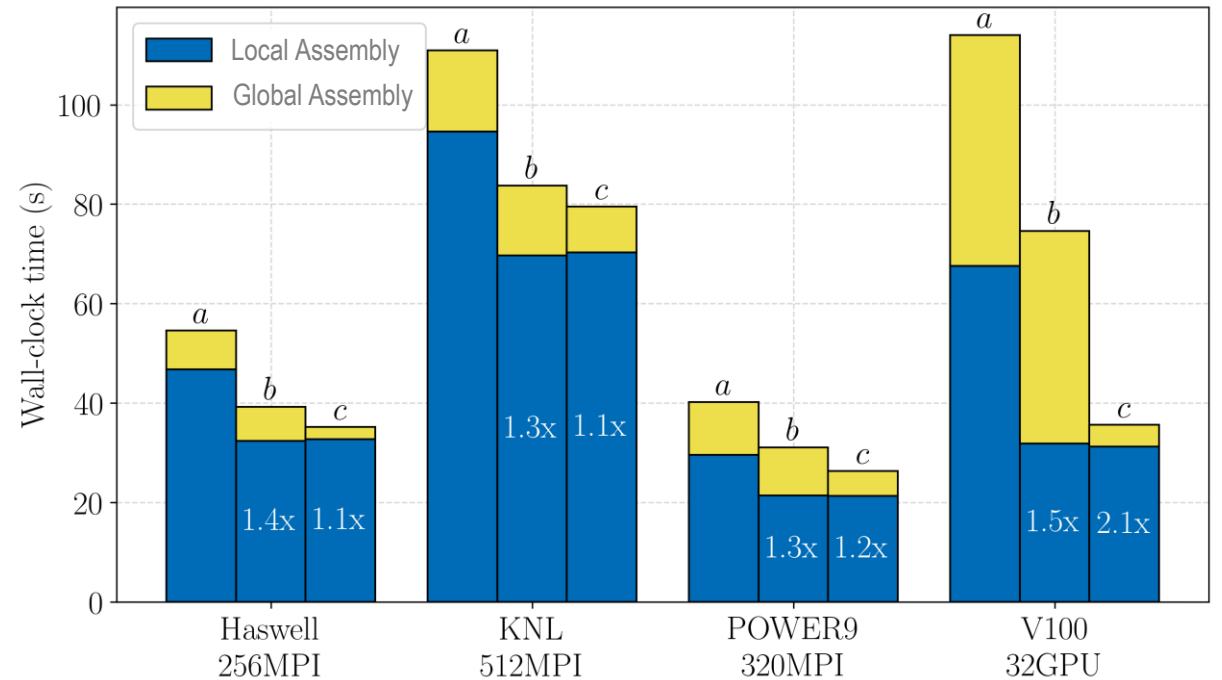


Improvements to initialization problem due to **memoization**:



AIS initialization 10 iterations
(HSW, 128 nodes, 4096 procs)

Combined improvements due to **memoization** and **improved storage of FE matrix** (Tpetra FE Crs Matrix)



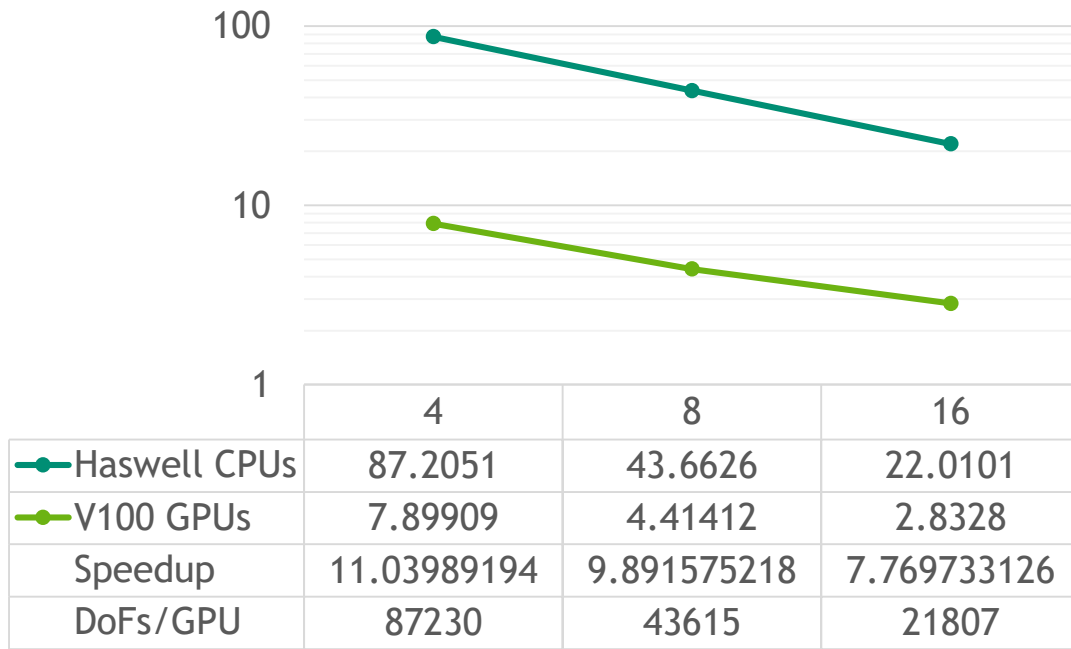
Assembly speed up on GPUs (Thwaites glacier, velocity solver)



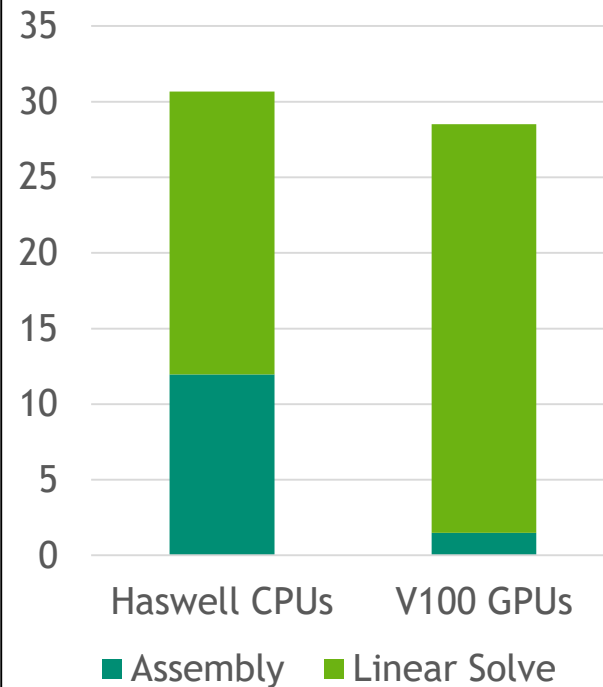
Haswell Node: 32 CPU cores/node
V100 Node: 4 GPUs/node

Up to **11x** speedup for assembly on GPUS.
Time to work on linear solvers!

Finite Element Assembly Strong Scaling
Wall-clock time (s) vs. Nodes



Total Solve Time (s)



Two-level Schwarz preconditioner (FROSch)

(Greenland Ice sheet, coupled velocity/temperature problem)



Challenge: our “workhorse” multigrid preconditioner does not work out of the box for the coupled problem.

Solution: FROSch (Fast and Robust Overlapping Schwarz)

Single-level Schwarz preconditioner vs Two-levels Schwarz preconditioner

MPI ranks	avg. #iters		avg. setup time		avg. solve time	
512	48.7	45.0	11.3s	10.53s	5.41 s	5.36s
1024	61.9	54.3	5.29s	4.59s	4.75s	4.31s
2048	89.9	59.1	2.52s	2.32s	5.70s	3.99
4096	116.1	78.7	1.17s	1.37s	3.68	3.30

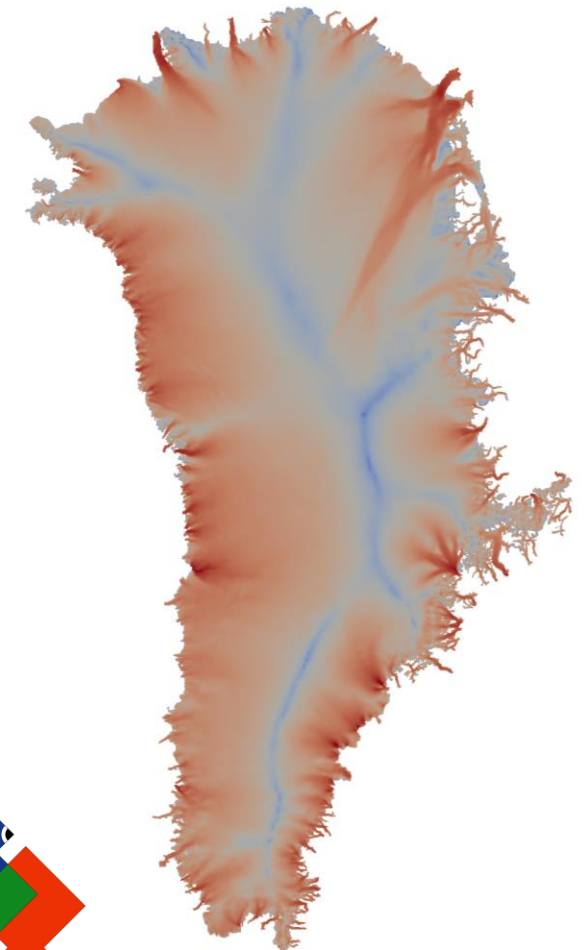
- The two-level preconditioner does better in terms of number of iterations but it is almost equivalent in term of CPU time to the one-level one.
- Time is dominated by direct solves of the sub-problems and the coarse solve.

Simulations by **A. Heinlein**

Support: CMDV-SM/FASTMath/ProSPect

Resources: Cori (NERSC)

Reference: Heinlein, Perego, Rajamanickam, arXiv, 2021



Ice sheet initialization



Hessian computation using automatic differentiation (using Sacado package)

Capabilities required for Newton-Krylov optimization methods:

Hessian of residual \mathbf{f} dotted with the Lagrange multiplier $\boldsymbol{\lambda}$ in the direction \mathbf{v} :

$$\begin{aligned} &\partial_{uu}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \mathbf{v}, & \partial_{up}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \mathbf{v}, \\ &\partial_{pu}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \mathbf{v}, & \partial_{pp}(\boldsymbol{\lambda}^T \mathbf{f}(\mathbf{u}, \mathbf{p})) \mathbf{v} \end{aligned}$$

Hessian of loss functional J in the direction \mathbf{v} :

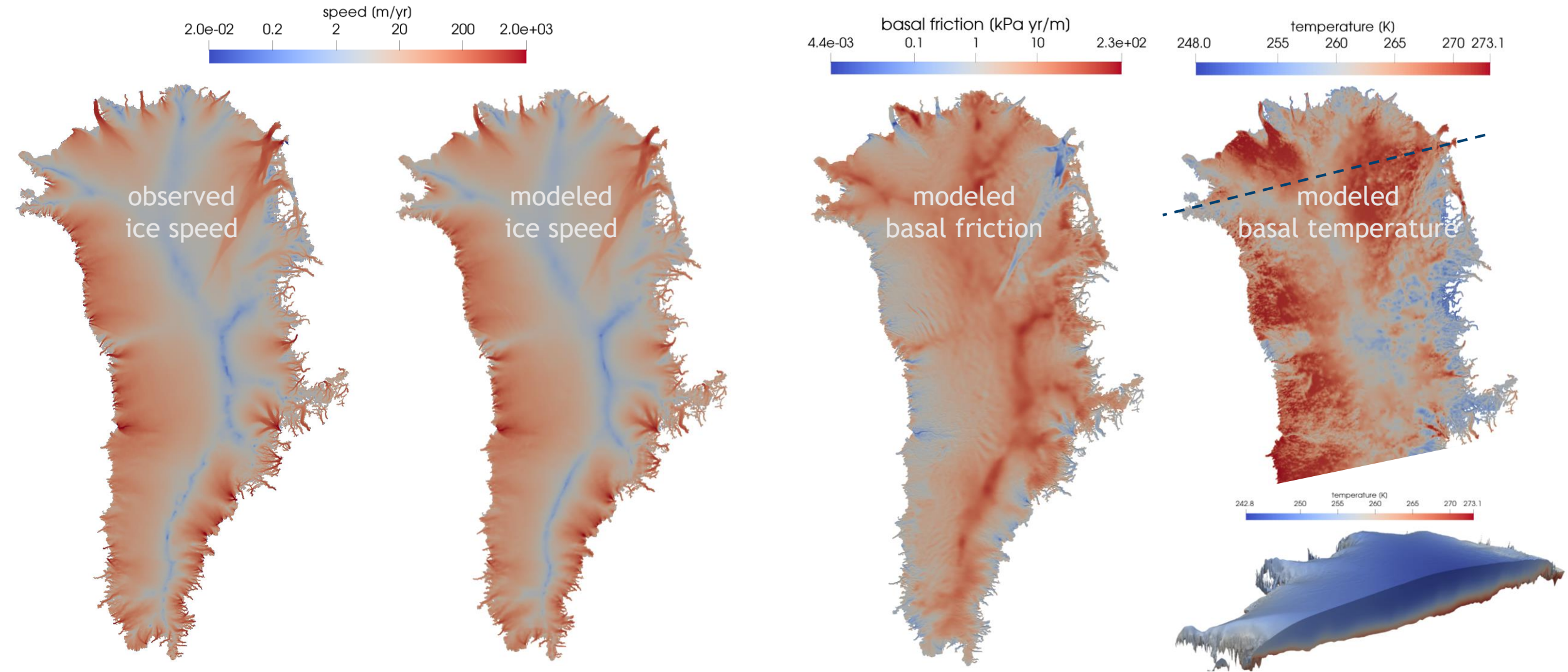
$$\begin{aligned} &\partial_{uu} J(\mathbf{u}, \mathbf{p}) \mathbf{v}, & \partial_{up} J(\mathbf{u}, \mathbf{p}) \mathbf{v}, \\ &\partial_{pu} J(\mathbf{u}, \mathbf{p}) \mathbf{v}, & \partial_{pp} J(\mathbf{u}, \mathbf{p}) \mathbf{v}, \end{aligned}$$

Computed w/ **automatic differentiation**, based on the formula:

$$\partial_{pp} J(\mathbf{p}) \mathbf{v} = \partial_r \left(\partial_p J(\mathbf{p} + r \mathbf{v}) \right) \Big|_{r=0}$$

We also provide the full matrix $\partial_{pp} J$, that can be used to initialize reduced-space optimization methods like BFGS, and the *reduced* Hessian-vector product.

Thermo-mechanical initialization of Greenland ice sheet



300K parameters, 14M unknowns. Initialization: ~ 10 hours on 8k nodes on NERSC Cori (Haswell)

Support: FASTMath/E3SM/ProSPect

MS270 on Thursday targets ice sheet modeling