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Preconditioning for Implicit Ocean Models

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Sandia National Laboratories

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Collaborators

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- LANL - Wilbert Weijer.
- U. Groningen - Fred Wubs, Jonas Thies.
- U. Utrecht - Henk Dijkstra, Erik Bernsen.

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Outline

- Ocean Modeling 101.
- The Coriolis Term.
- Pressure Coupling.
- Final Thoughts.

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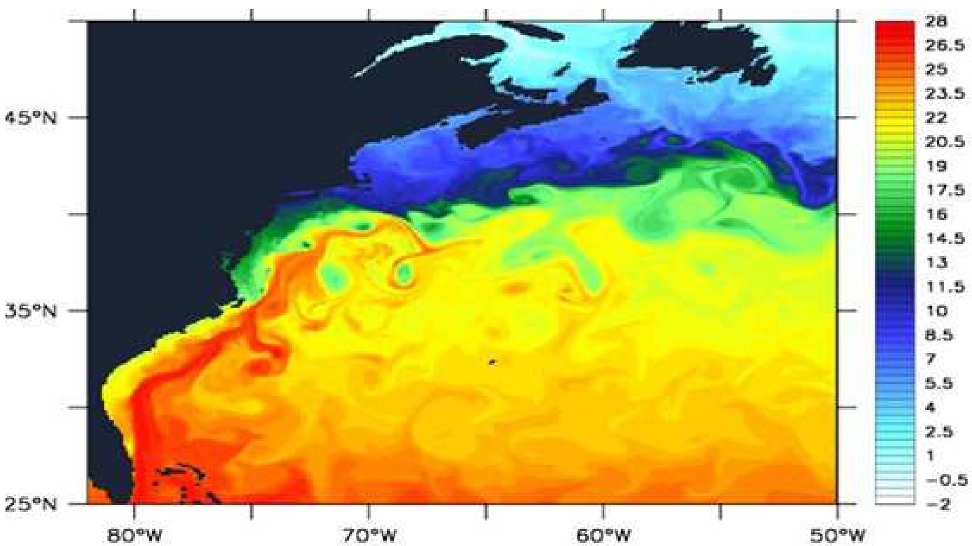
Ocean Modeling in POP

- Parallel Ocean Program (POP) is one of the models in the Community Climate System Model (CCSM).
- Physics of POP
  - Finite difference of thin stratified fluid equations w/ hydrostatic and Boussinesq approximations.
  - Coupled temperature & salinity advection-diffusion.



Left: THCM, J. Thies; Right: POP website

POP Equations

$$\frac{\partial u}{\partial t} + \mathcal{C}_1(u) - \alpha_1 uv - \Omega v + \alpha_2 \frac{\partial \eta}{\partial \lambda} + \alpha_3 \frac{\partial p_{bc}(S,T)}{\partial \lambda} - D_1(u,v) = 0$$

$$\frac{\partial v}{\partial t} + \mathcal{C}_1(v) + \alpha_1 u^2 + \Omega u + \alpha_4 \frac{\partial \eta}{\partial \phi} + \alpha_5 \frac{\partial p_{bc}(S,T)}{\partial \phi} - D_2(u,v) = 0$$

$$\frac{\partial S}{\partial t} + \mathcal{C}_2(S) + \mathcal{C}_3(u,v,S) - D_3(S,T) = 0$$

$$\frac{\partial T}{\partial t} + \mathcal{C}_2(T) + \mathcal{C}_3(u,v,T) - D_3(S,T) = 0$$

$$\frac{\partial \eta}{\partial t} + \int_{-H}^{0} \left(\alpha_6 \frac{\partial u}{\partial \lambda} + \alpha_7 \frac{\partial v}{\partial \phi} \right) dz = 0$$

Diffusion
Convection
Coriolis
Coupling #1
Coupling #2

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POP Equations (Take 2)

$$\boxed{\frac{\partial u}{\partial t} + \mathcal{C}_1(u) - \alpha_1 uv - \Omega v} + \alpha_2 \frac{\partial \eta}{\partial \lambda} + \alpha_3 \frac{\partial p_{bc}(S,T)}{\partial \lambda} \boxed{- D_1(u,v)} = 0$$

$$\boxed{\frac{\partial v}{\partial t} + \mathcal{C}_1(v) + \alpha_1 u^2 + \Omega u} + \alpha_4 \frac{\partial \eta}{\partial \phi} + \alpha_5 \frac{\partial p_{bc}(S,T)}{\partial \phi} \boxed{- D_2(u,v)} = 0$$

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~~Coupling #1~~
~~Coupling #2~~

The Coriolis Term (1)

- Consider the Coriolis-Diffusion equation:

$$\begin{bmatrix} -\Delta & \Omega \\ -\Omega & -\Delta \end{bmatrix} \begin{bmatrix} u \\ v \end{bmatrix} = \begin{bmatrix} f \\ g \end{bmatrix}.$$

- (Periodic) Fourier analysis gives spectral radius for Jacobi:

$$\rho_J = \left(\frac{16\cos^2(2\pi h) + \Omega^2 h^4}{16} \right)^{1/2}$$

which means Jacobi converges if:

$$\Omega < \frac{4}{h^2} |\sin(2\pi h)|.$$

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The Coriolis Term (2)

$$\begin{bmatrix} -\Delta & \Omega \\ -\Omega & -\Delta \end{bmatrix} \begin{bmatrix} u \\ v \end{bmatrix} = \begin{bmatrix} f \\ g \end{bmatrix}.$$

- Problem: Ocean models have kilometer scale h.
- Solution: Ω is diagonal, so it can be block inverted.
- (Periodic) Fourier analysis shows Block(2) Jacobi is stable for any Ω and spectral radius,

$$\rho_B = \left(\frac{16 \cos^2(2\pi h)}{16 + \Omega^2 h^4} \right)^{1/2}$$

- In fact, larger $\Omega \Rightarrow$ *faster* convergence.

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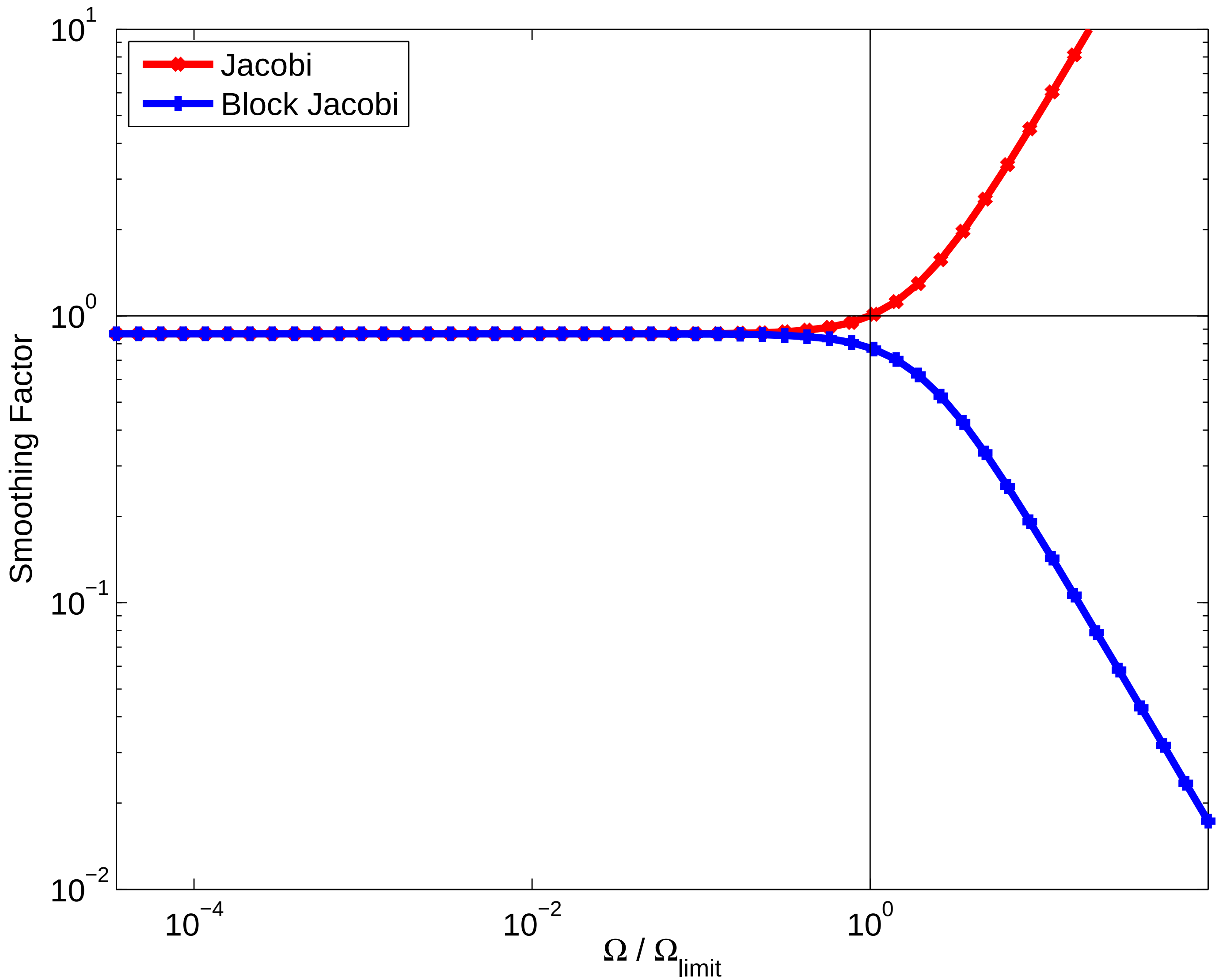
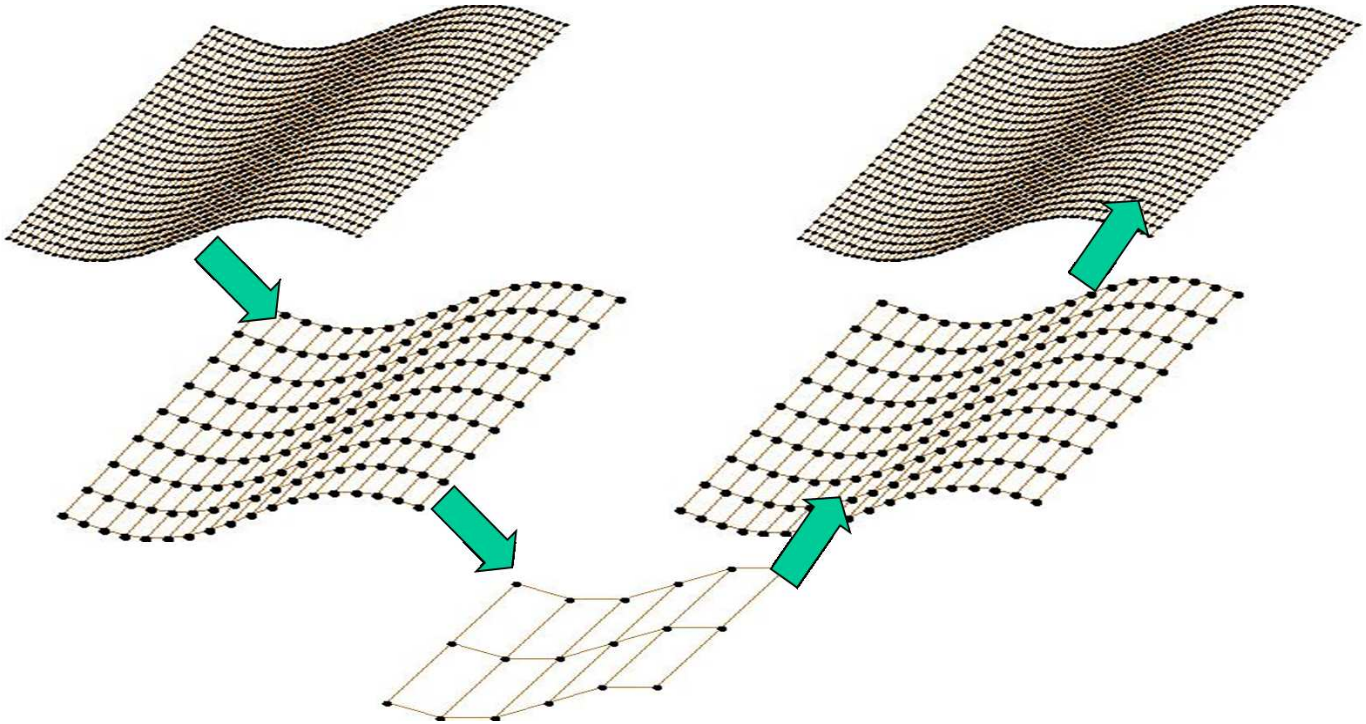
Smoothing Factor by Ω

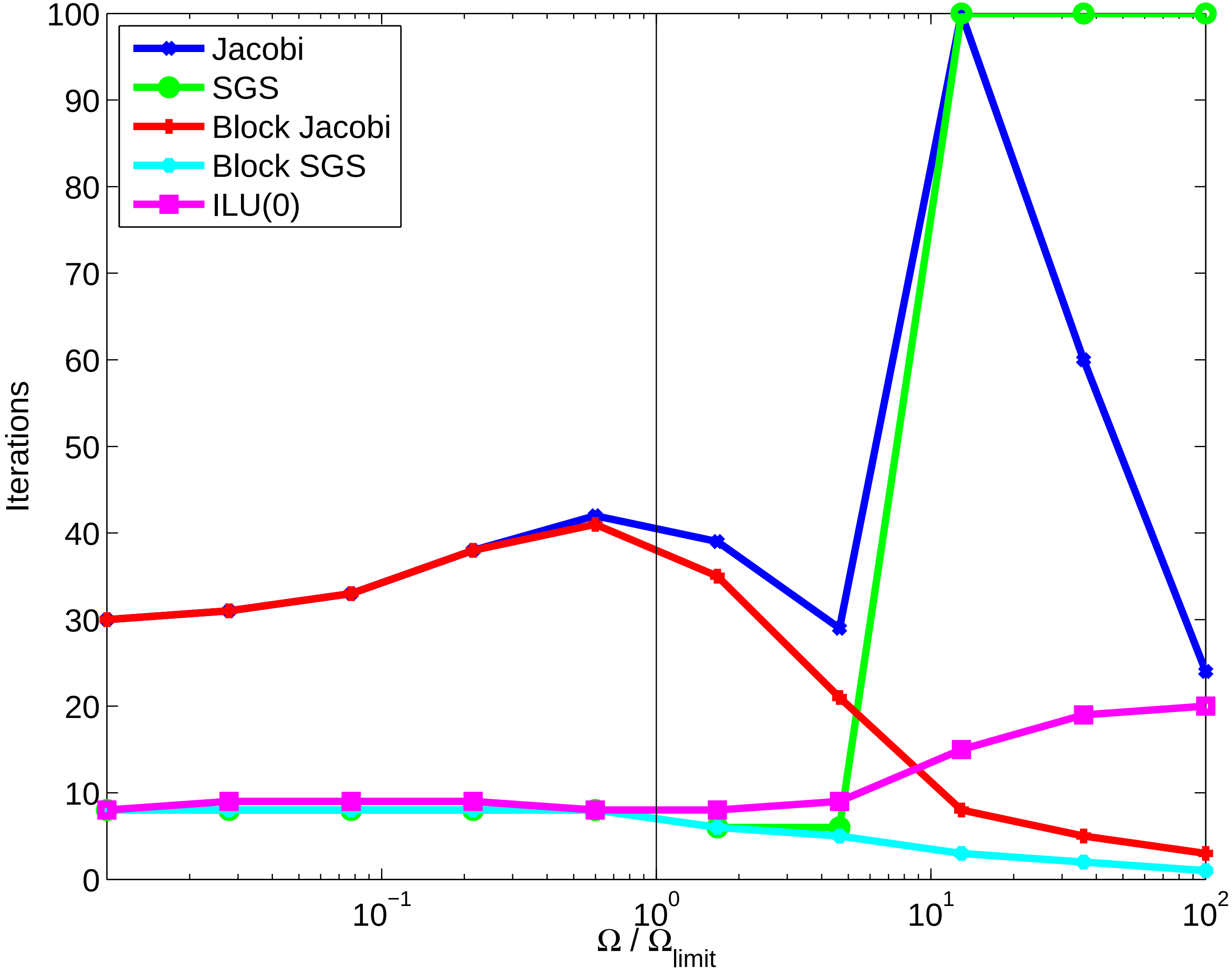


AMG for Coriolis-Diffusion



- Consider (Block) Jacobi, (Block) SGS and (Block) ILU(0) as smoothers to a 2-grid method w/ GMRES acceleration.
- Coarse grid: Non-smoothed aggregation w/ 2 nullspace vectors.

2-Level AMG Convergence



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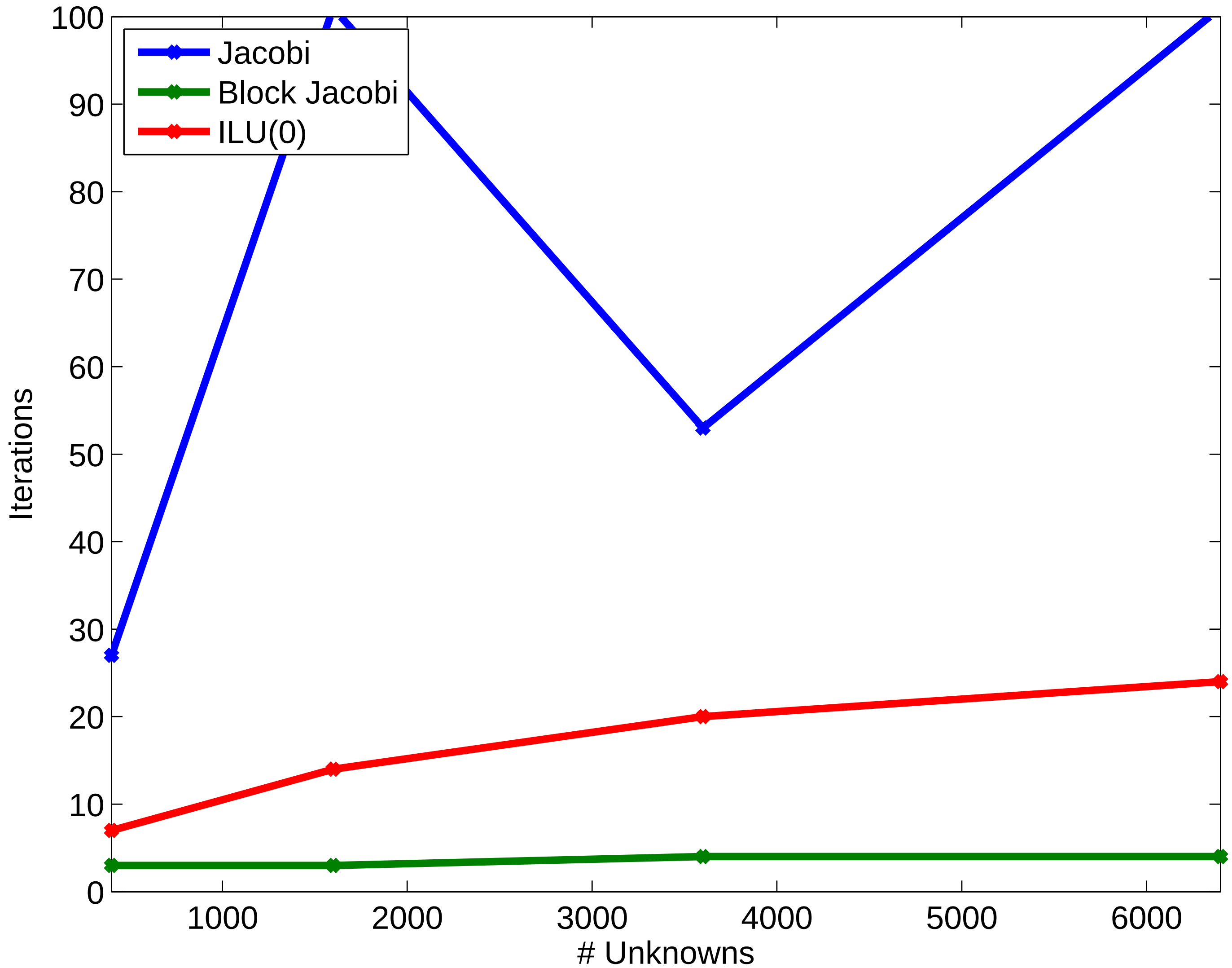
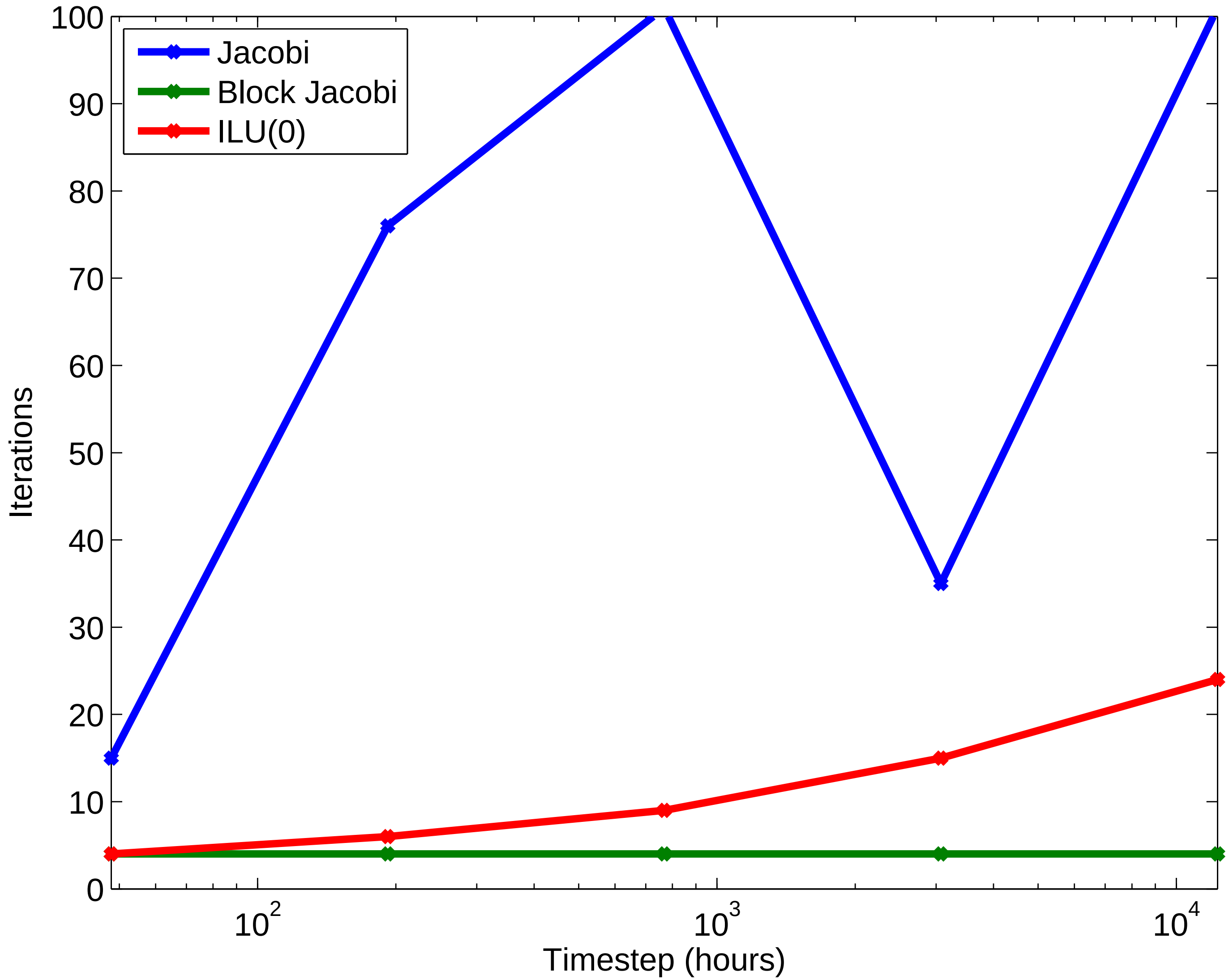
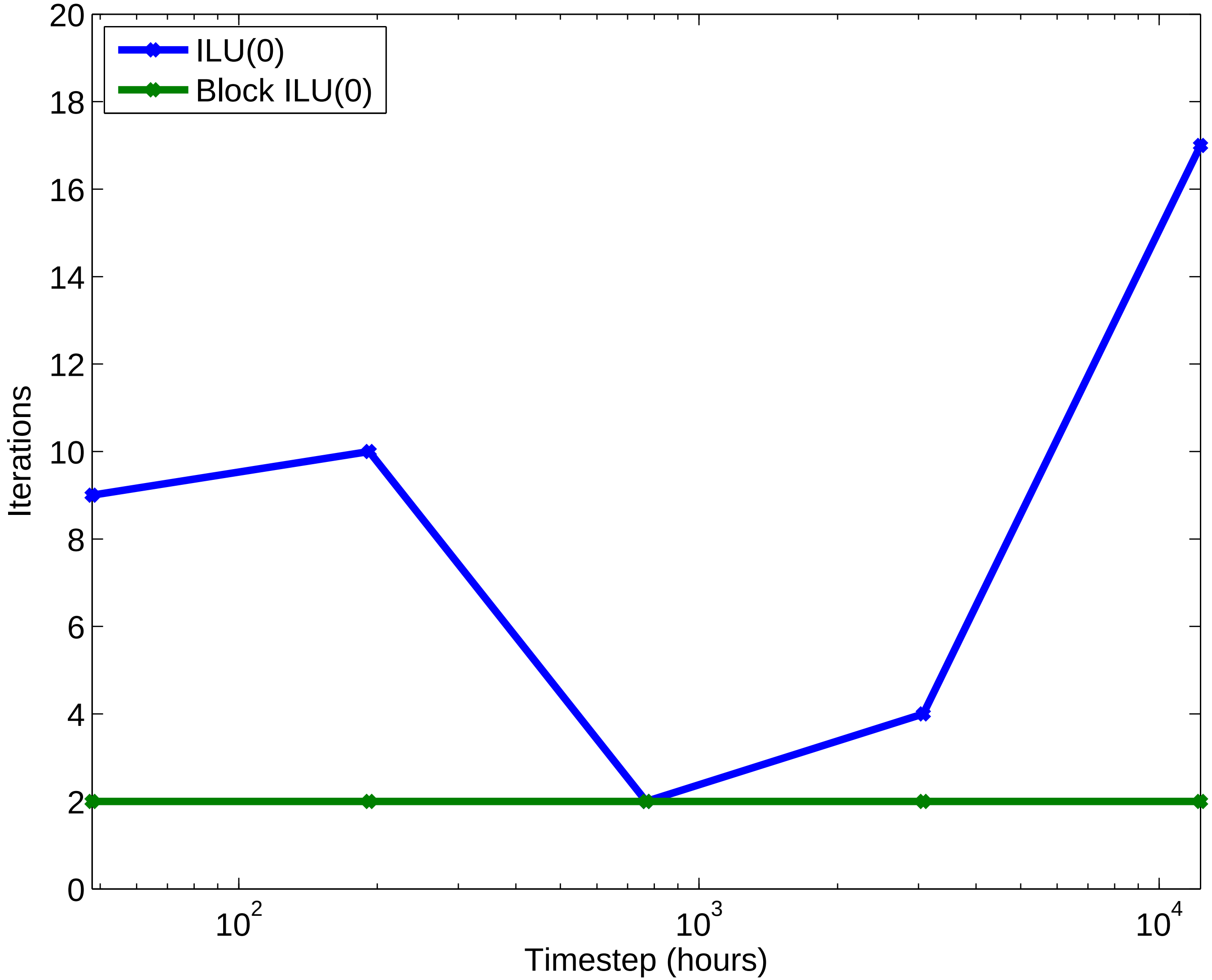
POP w/o wind: Δx Refinement



POP w/o wind: Large Δt



POP w/ wind: Large Δt



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POP Equations (Take 3)

$$\frac{\partial u}{\partial t} + \mathcal{C}_1(u) - \alpha_1 uv - \Omega v + \alpha_2 \frac{\partial \eta}{\partial \lambda} + \alpha_3 \frac{\partial p_{bc}(S,T)}{\partial \lambda} - D_1(u,v) = 0$$

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Schur Complements

- Treat Velocity-Pressure problem as a Schur Complement system.
- Factor a block 2×2 matrix:

$$\begin{bmatrix} A & B^T \\ C & D \end{bmatrix} = \begin{bmatrix} I & \\ CA^{-1} & I \end{bmatrix} \begin{bmatrix} A & \\ & S \end{bmatrix} \begin{bmatrix} I & A^{-1}B^T \\ & I \end{bmatrix}$$

  where $S = D - CA^{-1}B^T$.
- Approximate this factorization in order to precondition.
- Now we're left with two questions:
  - How to approximate A^{-1}? (See previous slides).
  - How to approximate S^{-1}?

SIMPLE & Block SIMPLE

- SIMPLE-like methods approximate $S = D - CA^{-1}B^T$, with

$$S = D - CF^{-1}B^T,$$

where $F = \text{diag}(A)$.

- Wait a second!
  - Isn't SIMPLE's F a lot like a step of point Jacobi on the convection-diffusion-Coriolis block?
  - Didn't we just show that this is unstable for large Ω?
- We'd best look at a Block SIMPLE as well.

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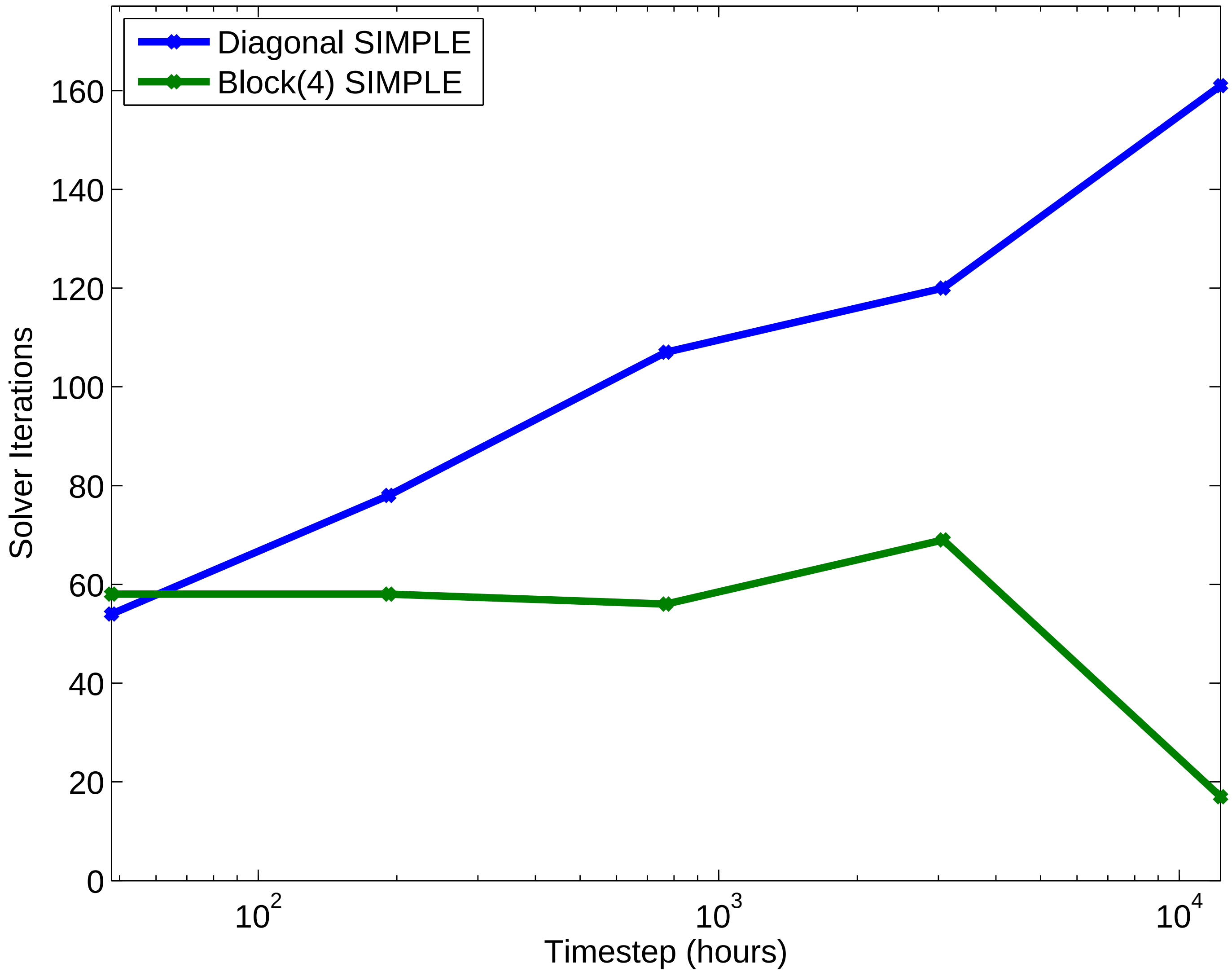
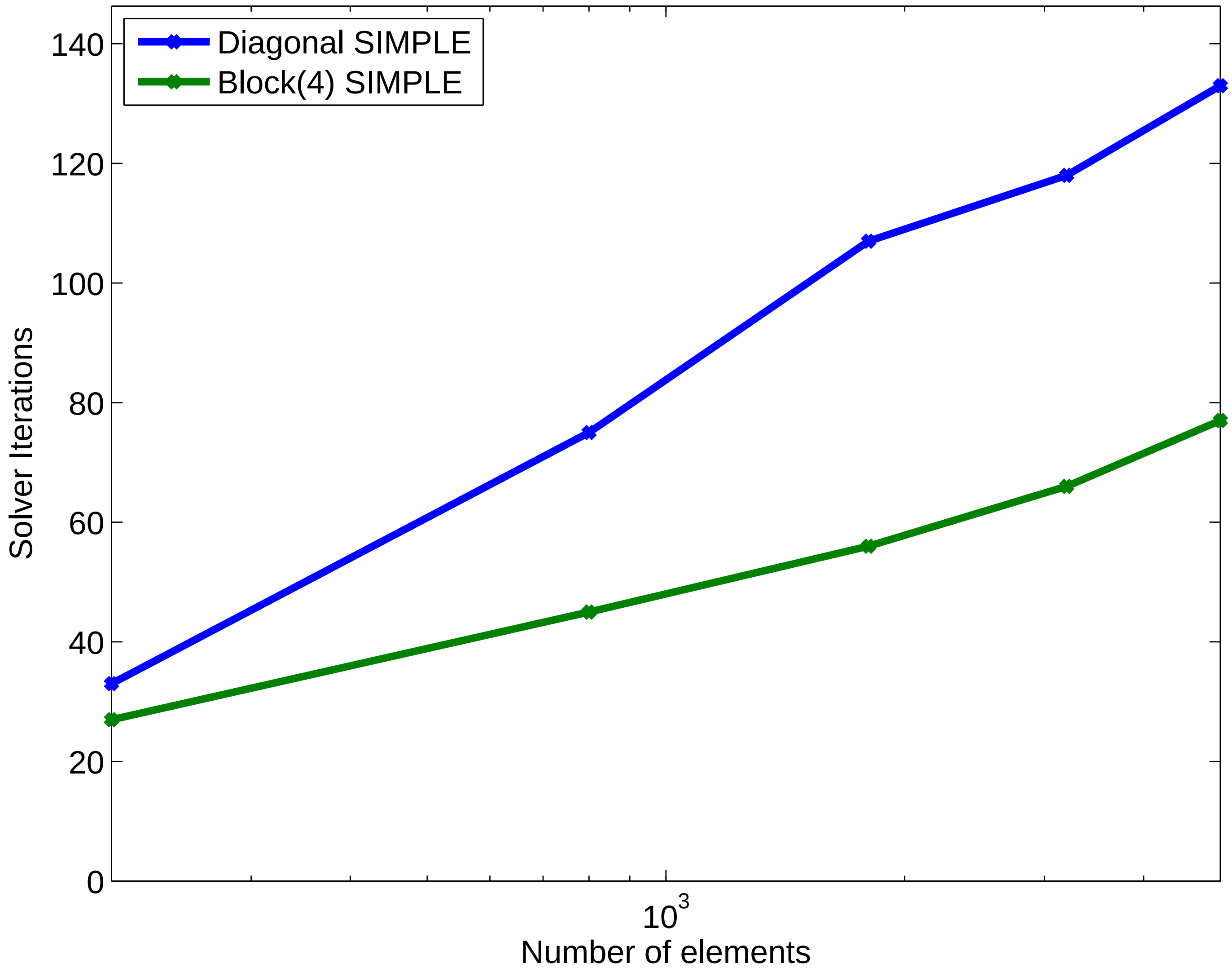
POP w/o wind: Large Δt



POP w/o wind: Δx Refinement



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Conclusions

- Block methods are important for Coriolis-Diffusion.
  - AMG + Block Jacobi / Block GS works well.
  - Convection may require more powerful block smoothers (Block ILU?).
- Schur complement must capture (1,1)'s block nature.
  - Block SIMPLE is OK even w/ large timesteps.
- Future work
  - Less mesh dependence.
  - Robust implementation in POP.
  - Harder & more realistic problems.