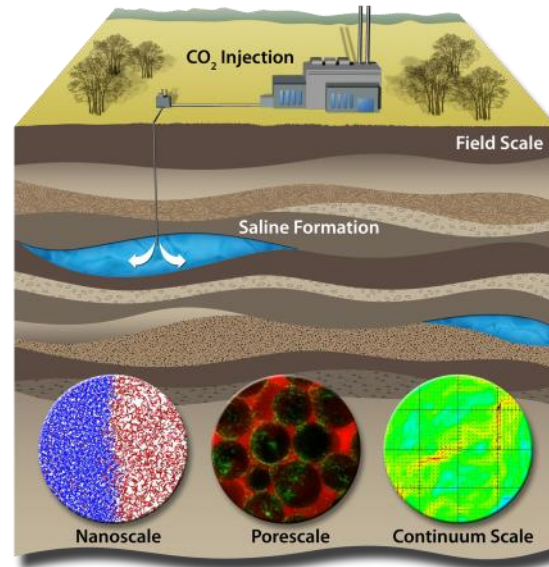


# Mixing-controlled Reactive Transport at the Pore Scale and Its impact on Flow Field and Upscaling of Reactive Transport



Hongkyu Yoon  
September 4, 2013



*Exceptional  
service  
in the  
national  
interest*



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# Outline

- **Motivations**
- Research Efforts
- Research Highlights
- Future Works

# Reactive Transport Processes during Geological Carbon Storage

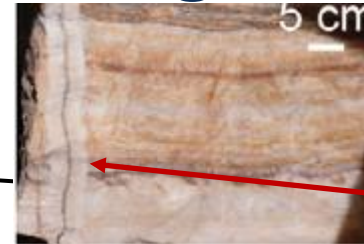
Injection well

Caprock

Caprock jointing

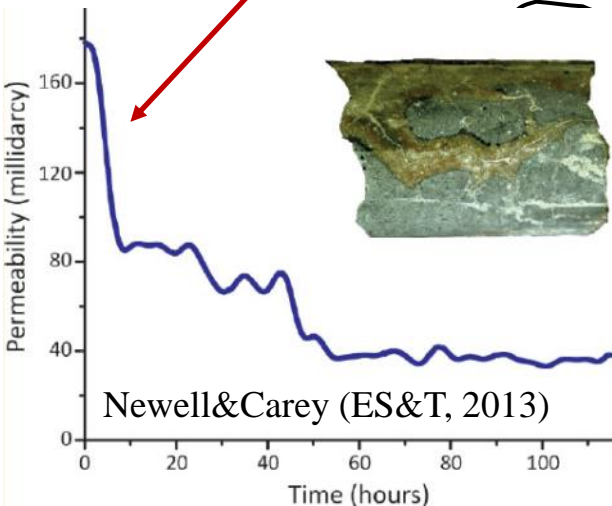
Storage zone

Gratier et al.,  
Geology, 2012;  
Crystal Geyser,  
Utah

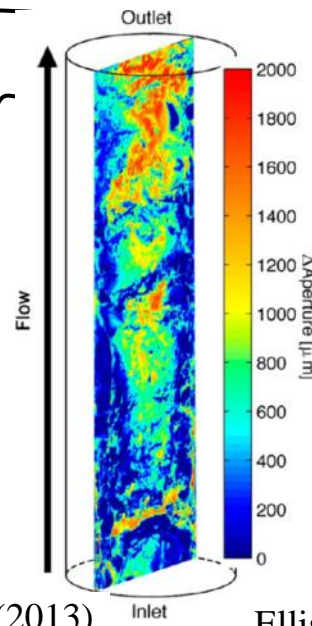


fault

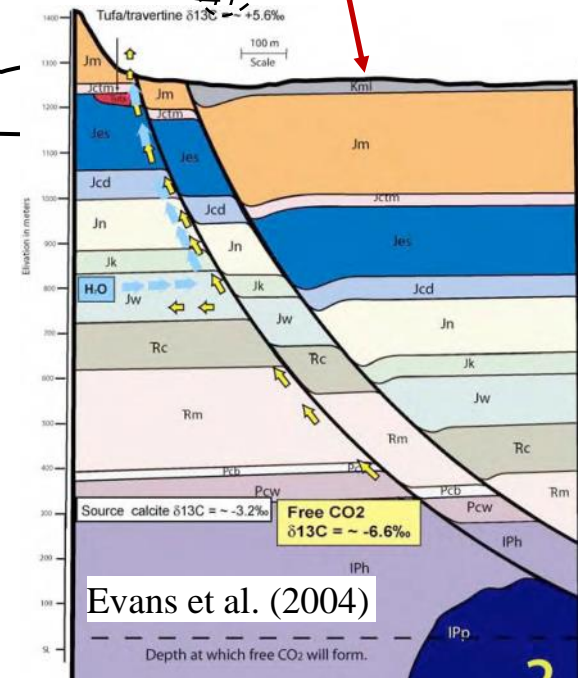
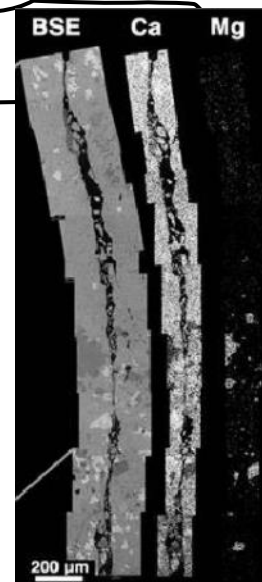
Permeability at  
the interface of  
cement and  
siltstone



Deng et al. (2013)

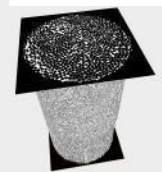


Ellis et al. (2013)

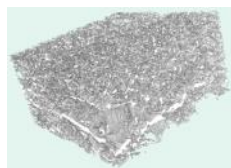


# Multi-Scale Problems

## Pore Scale



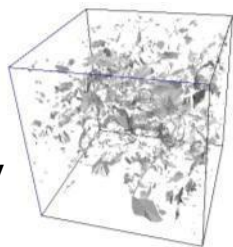
Micro-CT



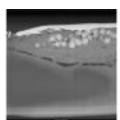
Statistical REV



Confocal  
Microscopy

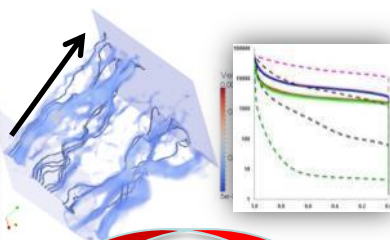


Pore throats

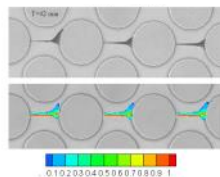
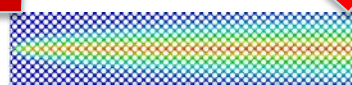


FIB-SEM

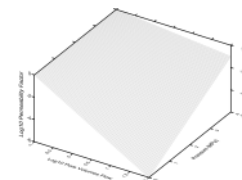
Core  
sample



UPSCALING



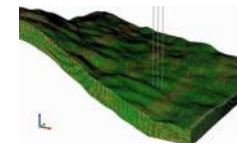
## Grid Block Scale



Reaction rate



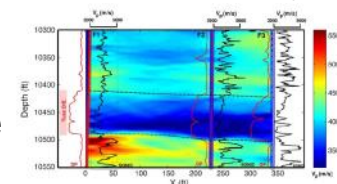
Reactive flow



3D permeability field

Observed Data

Different types  
Spatial scale  
Temporal scale



Seismic data

## Predictive

Fundamental (first-principles)

## Descriptive

Phenomenological (empirical)

Complex

Simple

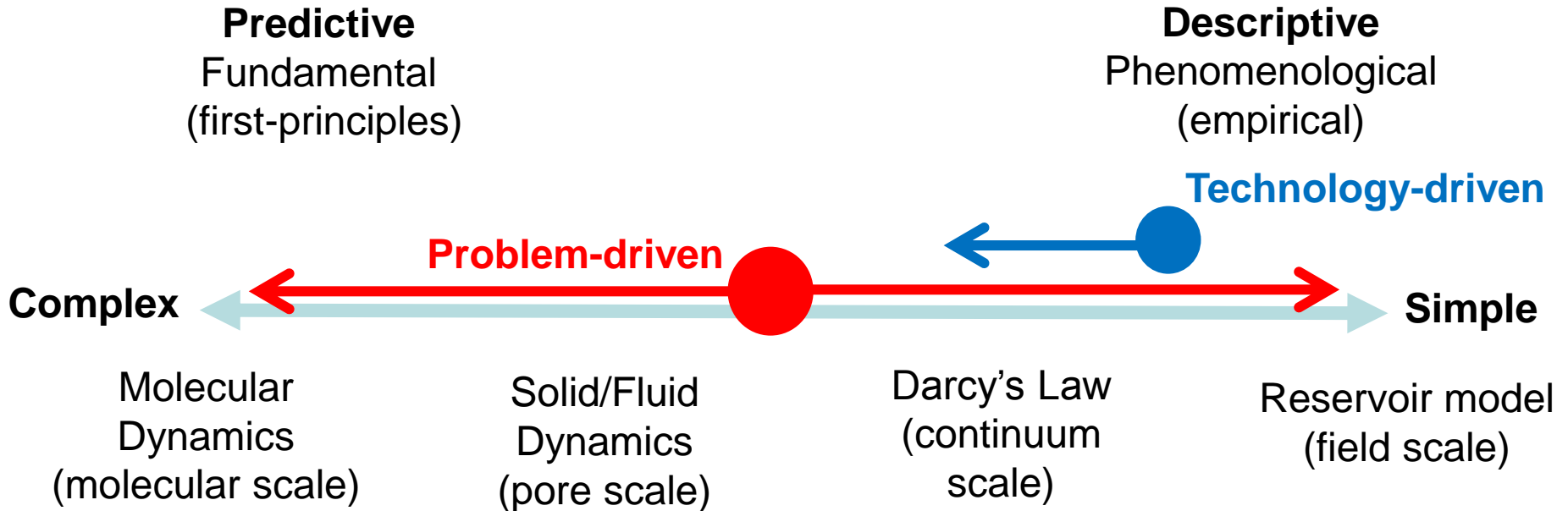
Molecular Dynamics  
(molecular scale)

Solid/Fluid Dynamics  
(pore scale)

Darcy's Law  
(continuum scale)

Reservoir model  
(field scale)

# What is the approximate level of complexity?



Where do we want (or need) to be?

Where can we be in practice?

# Research Direction

- Develop experimental and numerical tools to study a comprehensive understanding of multi-physics processes over a range of scales (nano to field scales)
- Apply recent advances in tool development to bridge the gap between two questions

## Example:

### Pore Scale Reactive Transport

- Developed a novel pore scale reactive transport model of coupled fluid flow, reactive transport, and precipitation and dissolution using pore scale experiments in a microfluidic pore-network (i.e., micromodel)

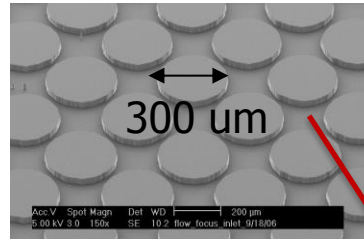
# Micromodel Experiment

## Micromodel Description

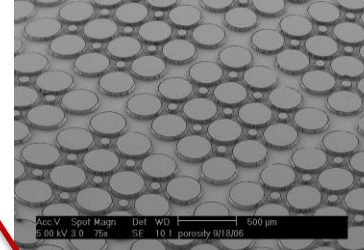
Depth: ~20 mm

Porosity: ~0.39

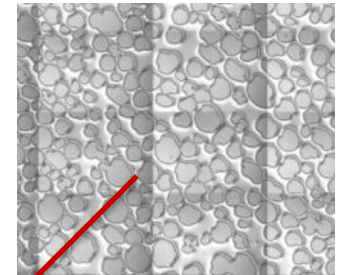
Flowrate: ~2 cm/min



Base Case  
Small Cylinder

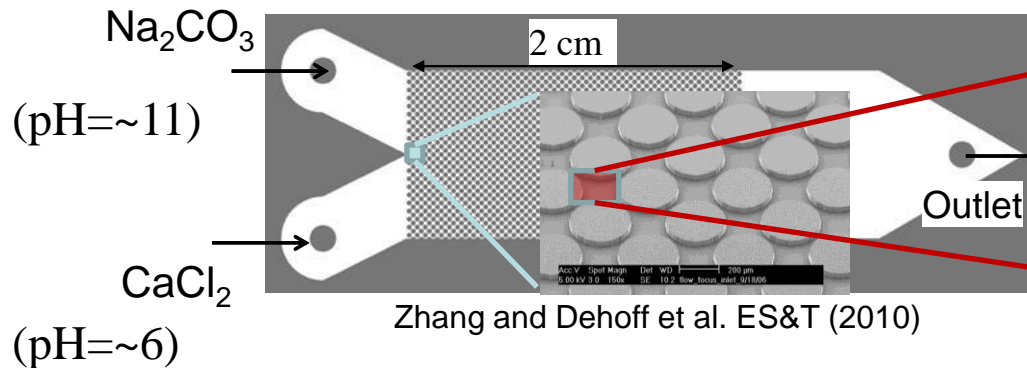


Aggregates

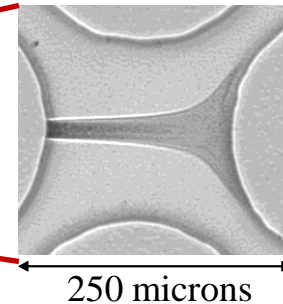


Irregular

Thermal oxidation:  
~ 100 micron thick oxide layer



Zhang and Dehoff et al. ES&T (2010)



Microscopic image of  
calcium carbonate  
(CaCO<sub>3</sub>) precipitates

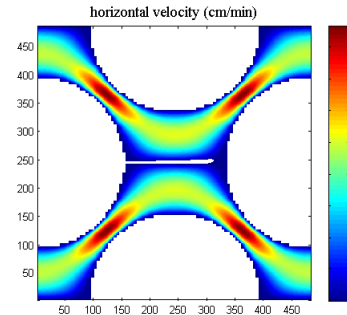
- Two solutions are mixing along the centerline and CaCO<sub>3</sub> precipitates
- Range of concentrations and solution chemistry vary
- Microscopic images are taken over time

# Pore Scale Model Framework

Lattice Boltzmann Method:  
Velocity field ( $u$ ) at pore scale



Finite Volume Method: Reactive transport at pore scale



Velocity

$\Delta t$

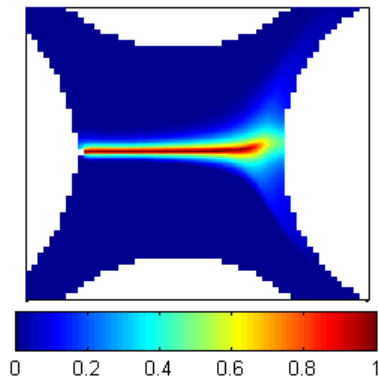
$\Psi_j = C_j + \sum_{i=1}^{N_{eq}} v_{ji} C_i$  Chemical equilibrium in bulk fluid (e.g.,  $H^+$ ,  $HCO_3^-$ , ...)  
Extended Debye-Hückel Equation for activity coefficients

$$D \frac{\partial \Psi_j}{\partial \mathbf{n}} = -I_m \quad \text{on reactive surface}$$

$$I_m = k_{cc} \left( [\Omega]^n - 1 \right)^m \quad \Omega = \frac{Q_{CC}}{K_{sp}} \text{ or } \ln \left( \frac{Q_{CC}}{K_{sp}} \right)$$

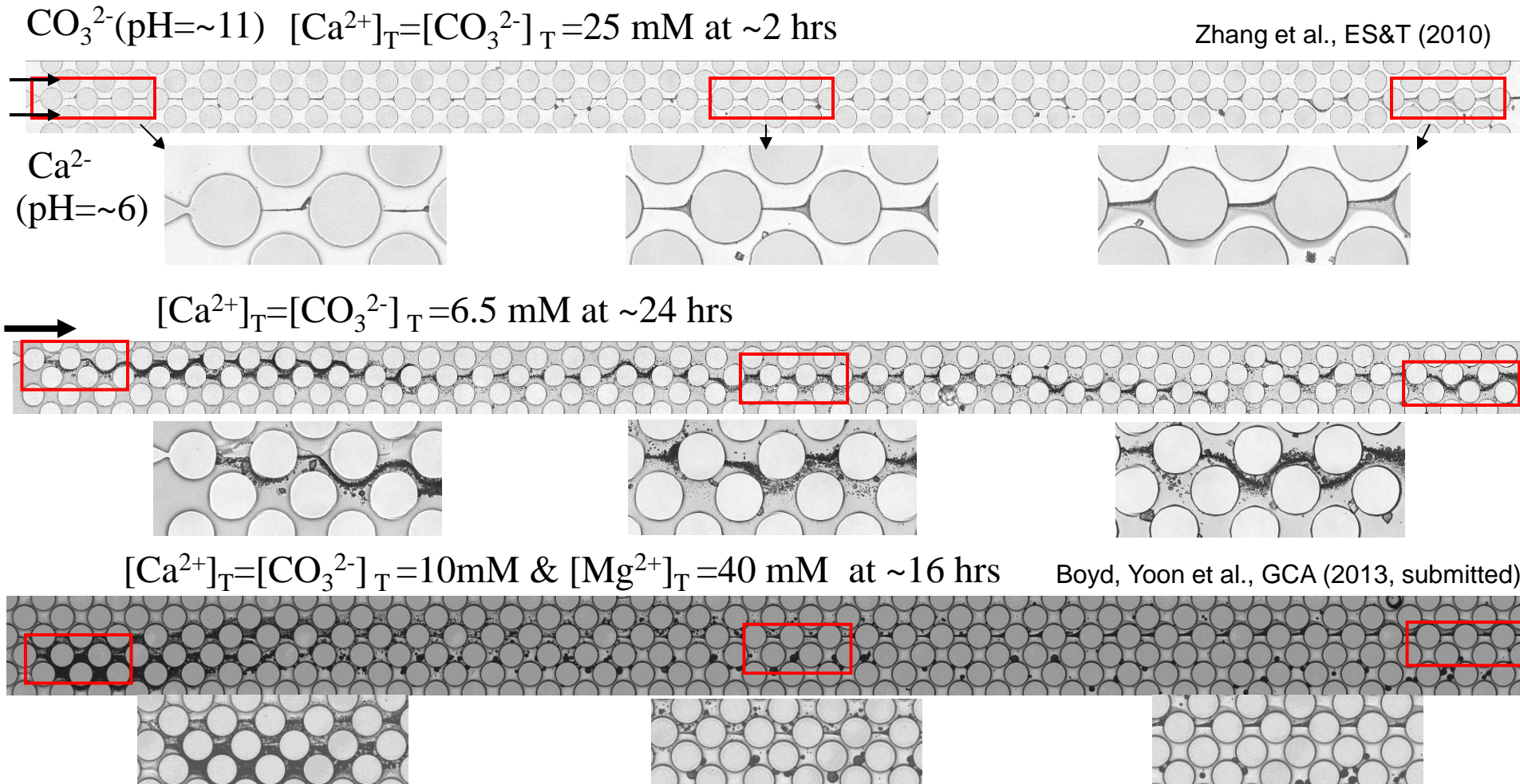
Update of  $CaCO_3$  volumetric content ( $V_m$ )

$$\frac{\partial V_m}{\partial t} = \overline{V}_m s_m k_{cc} \left( \left[ \frac{a_{Ca^{2+}} a_{CO_3^{2-}}}{K_{sp}} \right]^n - 1 \right)^m$$



Mineral phase  
volumetric content

# Experimental Results

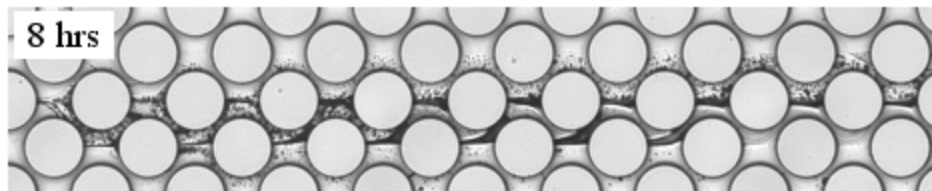
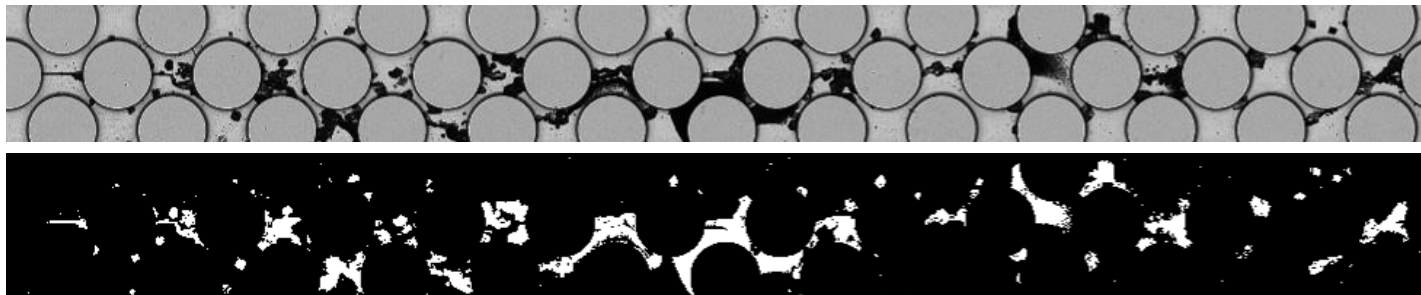


- Precipitation  $\sim$  along the centerline within 1-2 pore spaces in the transverse direction
- Width of the precipitate line  $\sim$  increase with distance from the inlet
- Rate of precipitation is concentration and species dependent

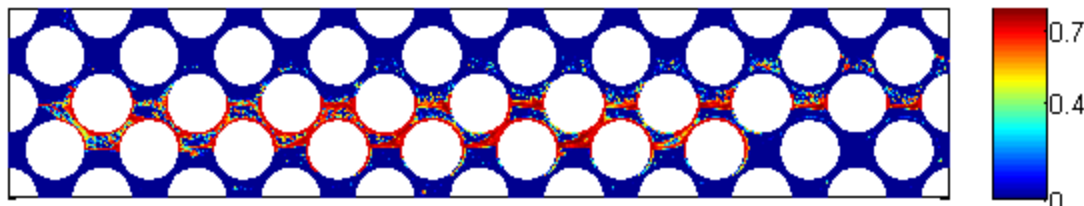
# Image Analysis

## Otsu thresholding

- Pixels segmented into foreground and background
- Uses threshold values that result in minimum interclass variance between foreground and background



Volumetric Fraction of Precipitate



# Results: Precipitation only at grain boundary

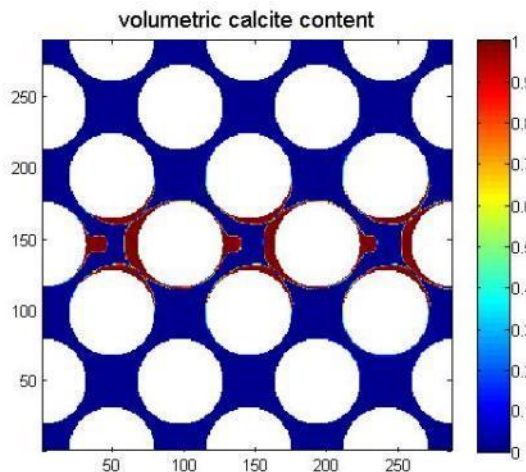
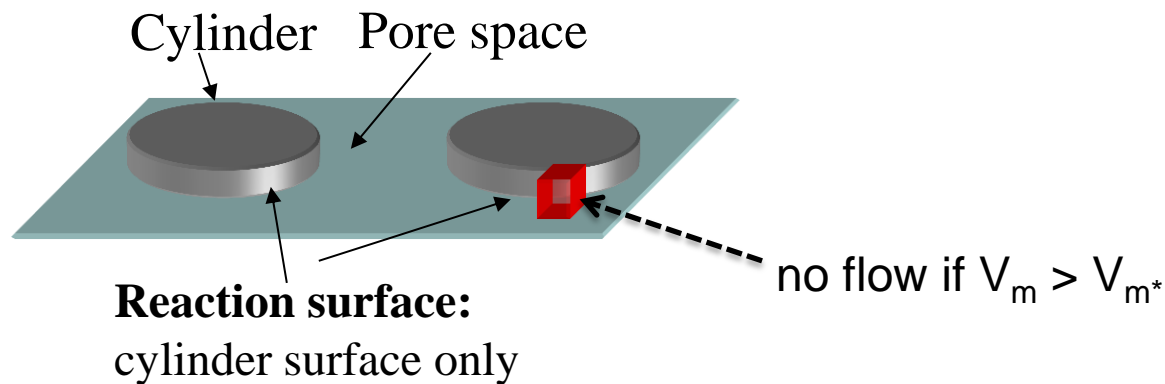
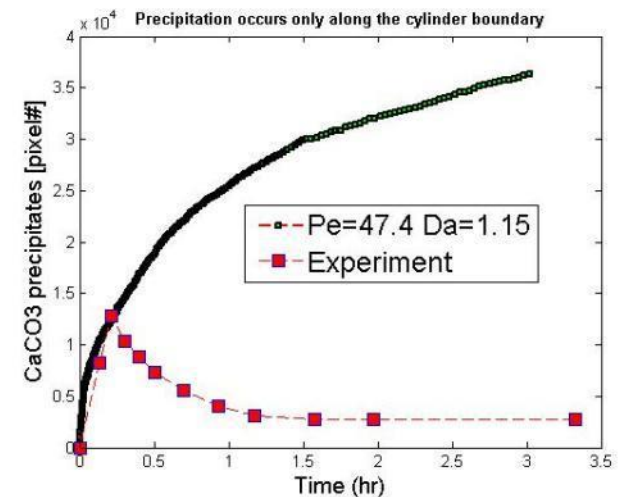


Image of precipitates  
at 180 min



25 mM  
Experiment



# Precipitation patterns

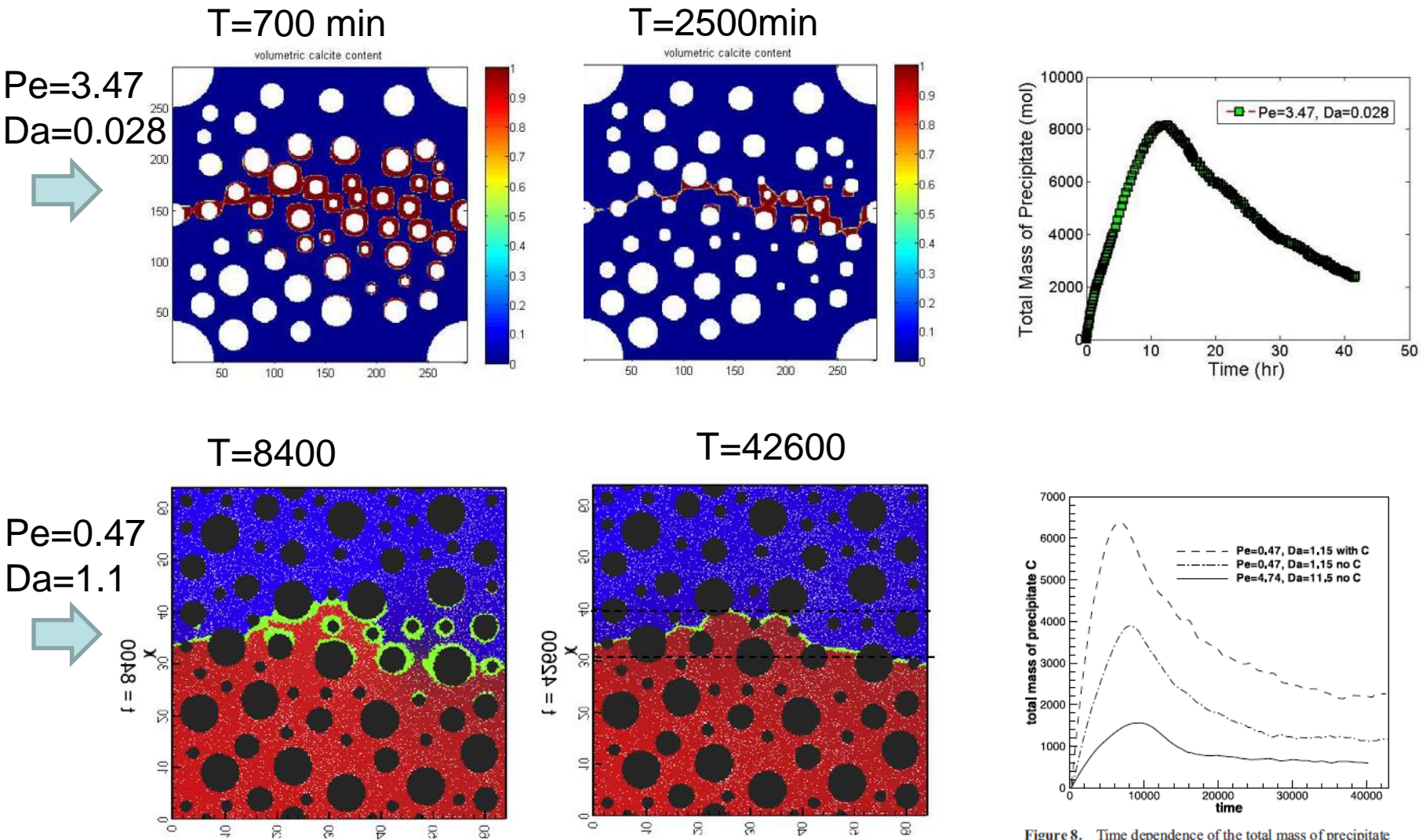
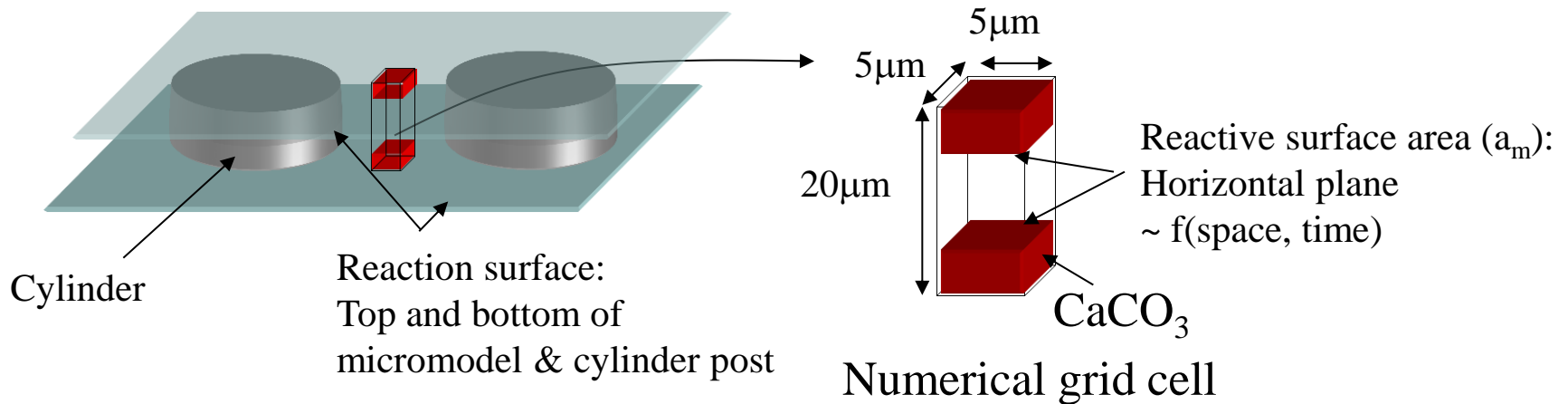


Figure 8. Time dependence of the total mass of precipitate  $C$  obtained from several simulations with and without the intermediate reaction produce,  $C$ , for different  $Pe$  and  $Da$ .

# Reaction in a micromodel system

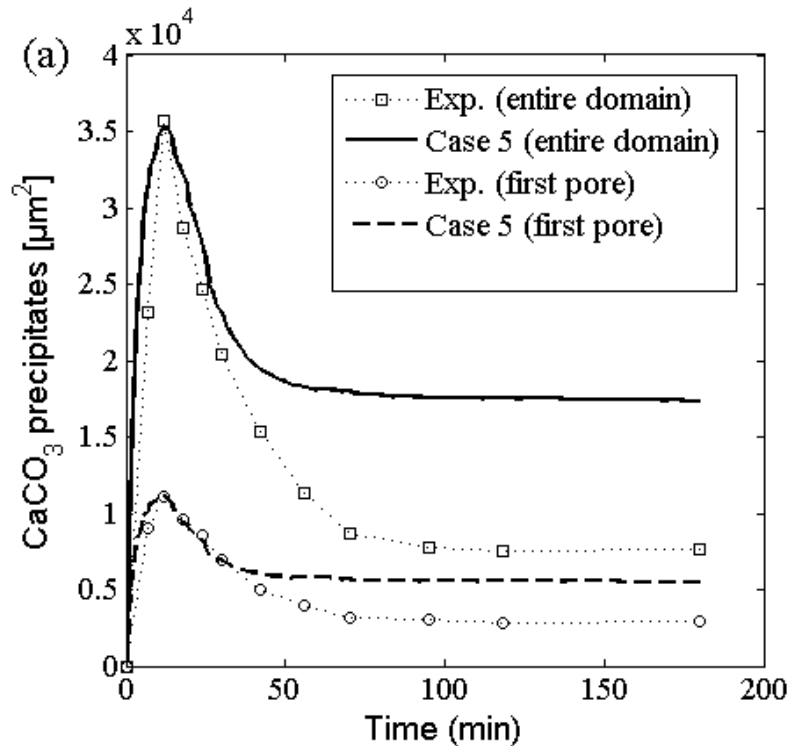
## 1. Quasi 3D grid cell for reactive surface



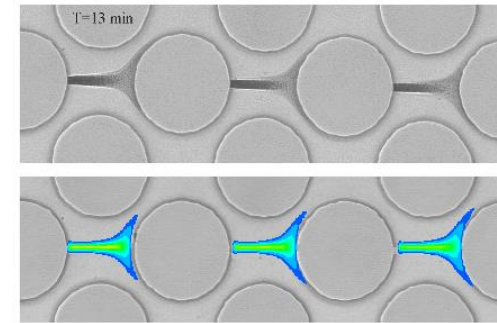
## 2. Effective diffusion coefficient = $D_m$ \* tortuosity ( $\tau$ )

- $\tau(V_m) = (1 - V_m)^n$  where  $n \sim 0$  to  $3$
- Diffusion is allowed until the grid cell is fully occupied by CaCO<sub>3</sub>

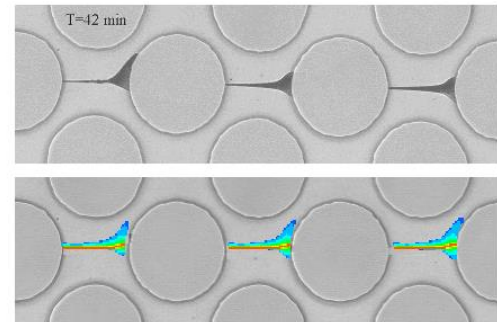
# Simulation results: Increase surface area during dissolution by 300



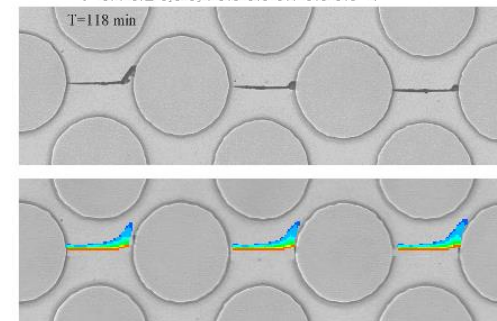
- Model results match thickness and area of precipitate until 30 min
- Model predicts dissolution below the centerline well, but not above the centerline



13 min

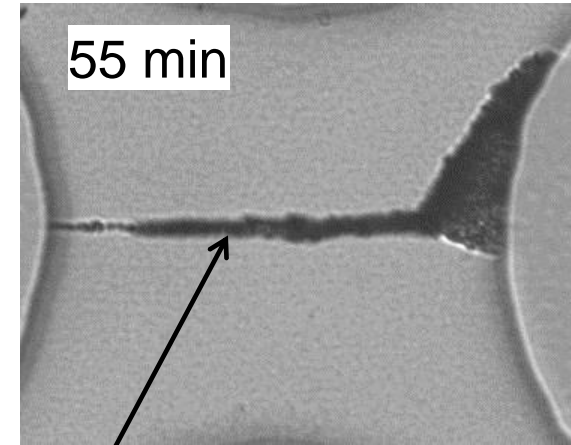
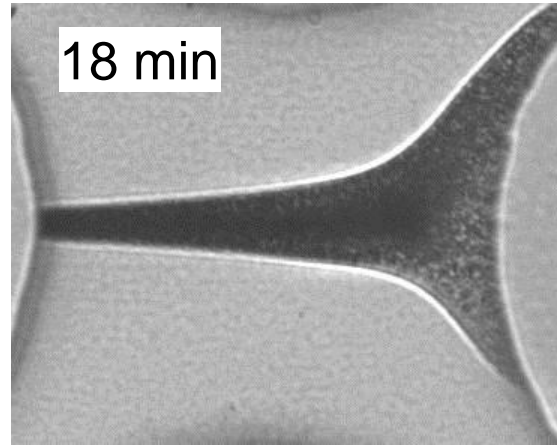
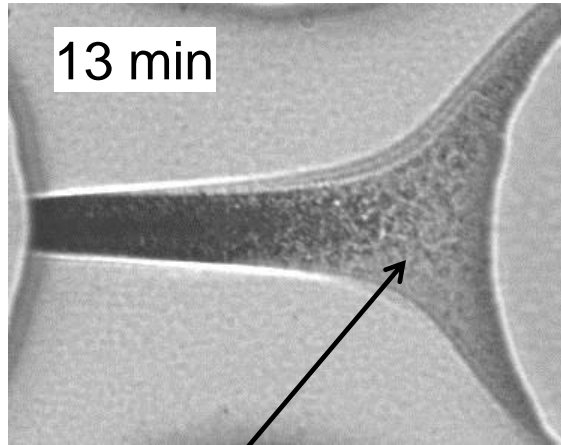


42 min

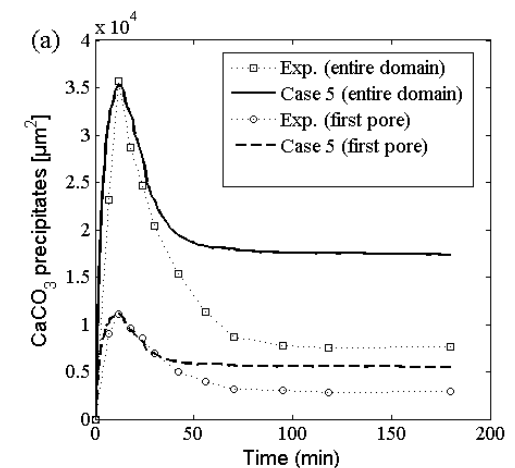


118 min

# Matching simulation to late-time dissolution



- Increase in surface area over time
- Transformation to different forms of  $\text{CaCO}_3$
- Stability of nano-particles after pore blocking (or reduced mixing along the centerline)
- Effect of nano-crystal size on solubility (Emmanuel and Ague, Chem. Geo. 2011)



# Rapid precipitation and transformation

## 6.5 mM Case

Amorphous Calcium Carbonate

1 min

6 min

Transformation

15 min

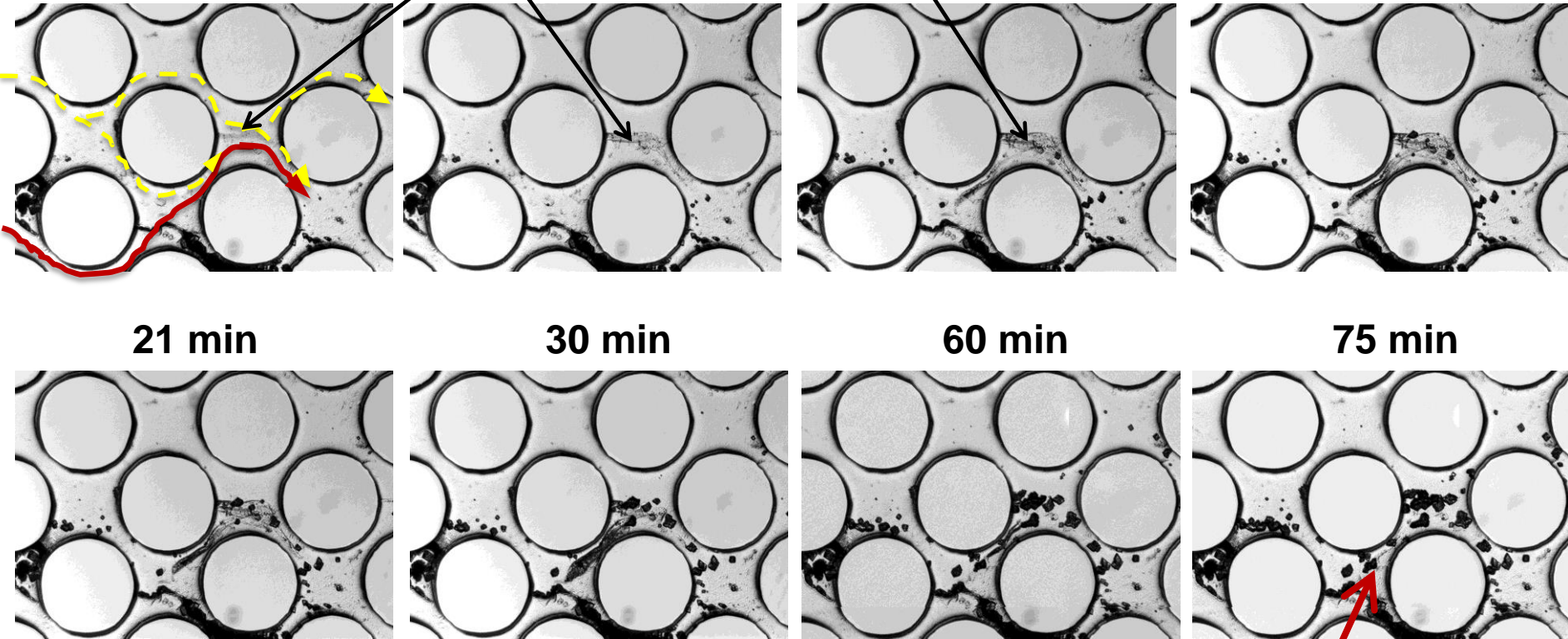
18 min

21 min

30 min

60 min

75 min



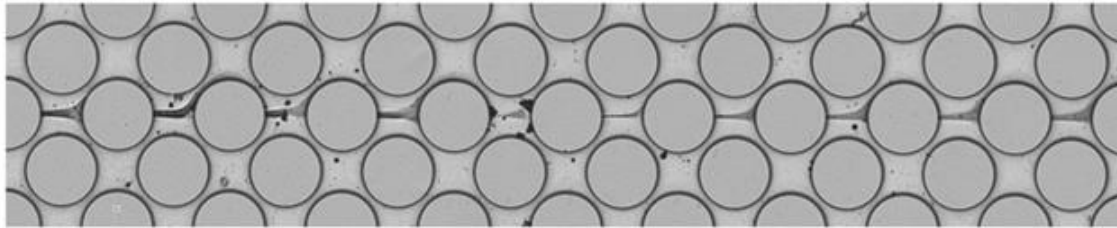
Calcite crystals

# Calcium Carbonate: Polymorph

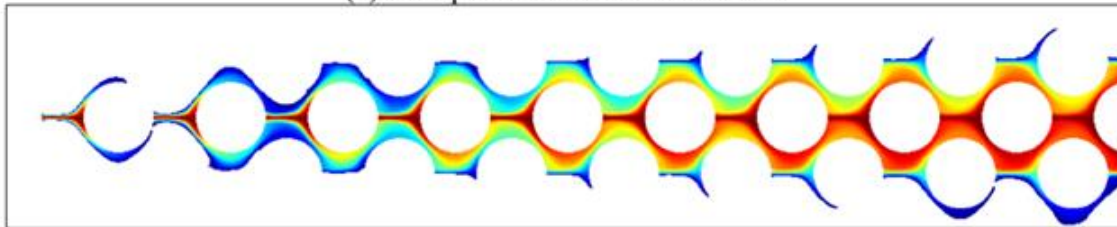
Saturation Ratio (SR) = Ion Activity product /  $K_{sp}$   
SR > 1: thermodynamically favorable to form

$$[Ca^{2+}]_T =$$
$$[CO_3^{2-}]_T$$
$$= 10 \text{ mM}$$

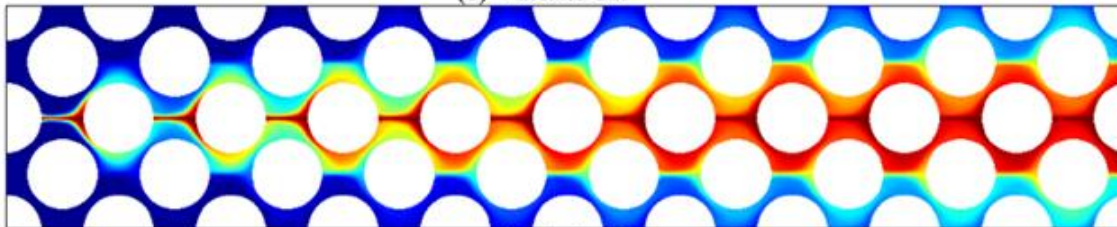
(a) 10 mM  $CaCl_2$  + 0 mM  $MgCl_2$  (0.67 hr)



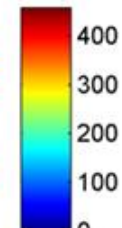
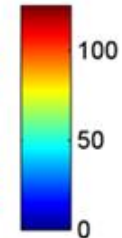
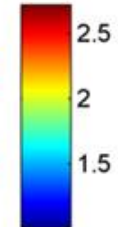
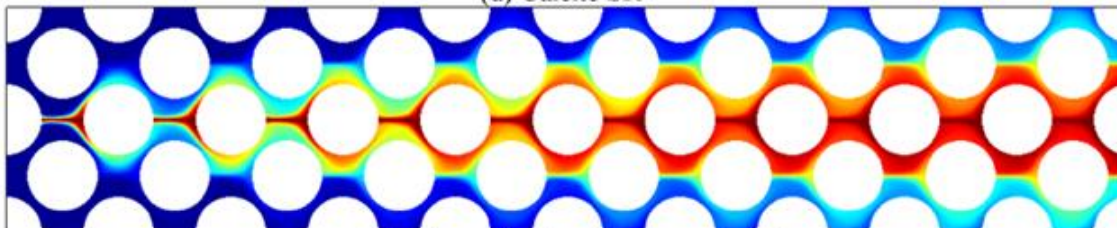
(b) Amorphous Calcium Carbonate SR



(c) Vaterite SR



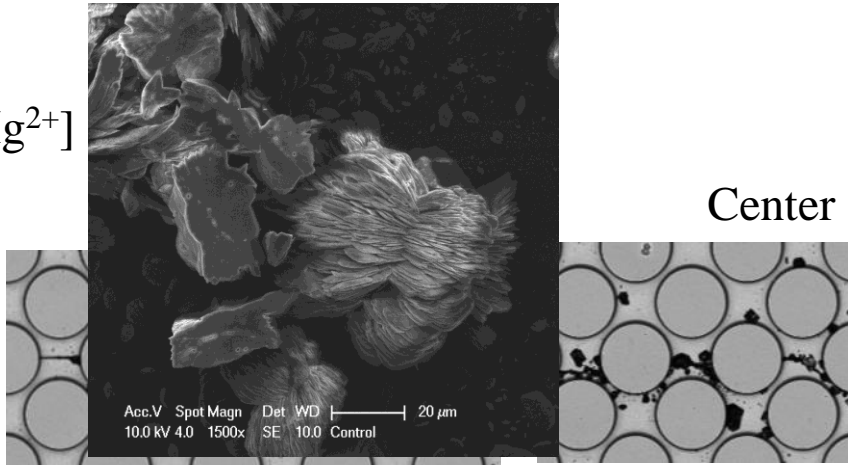
(d) Calcite SR



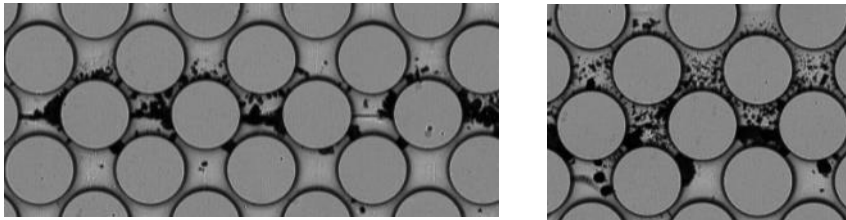
# Adding Magnesium

Adding  $[Mg^{2+}]$

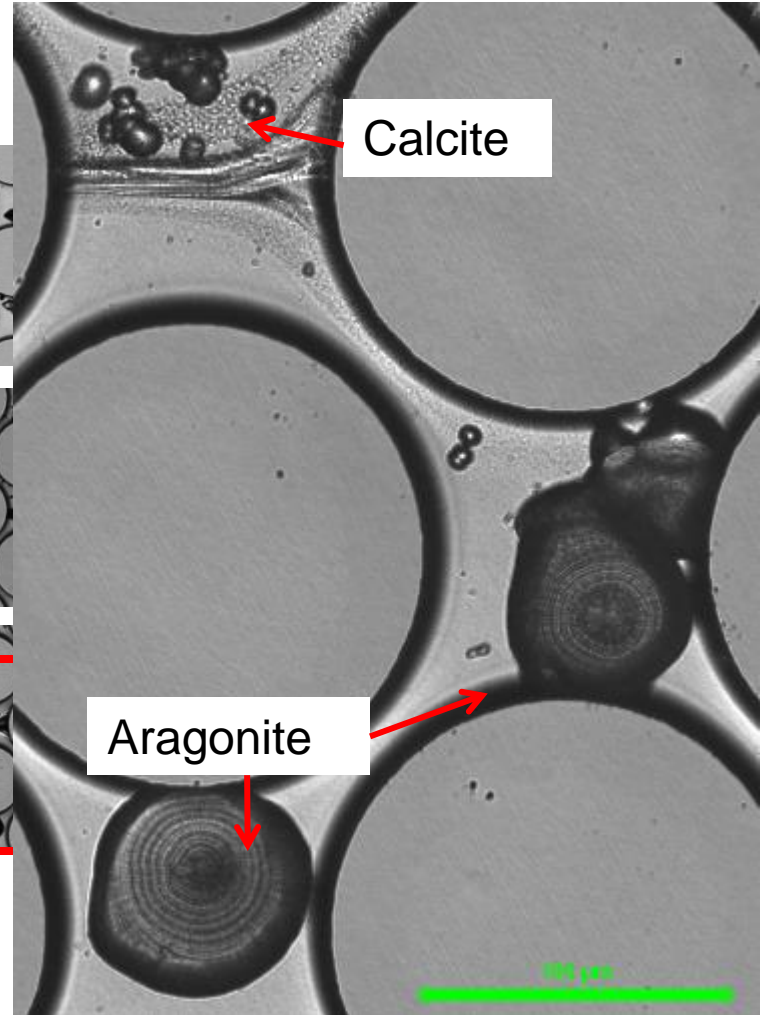
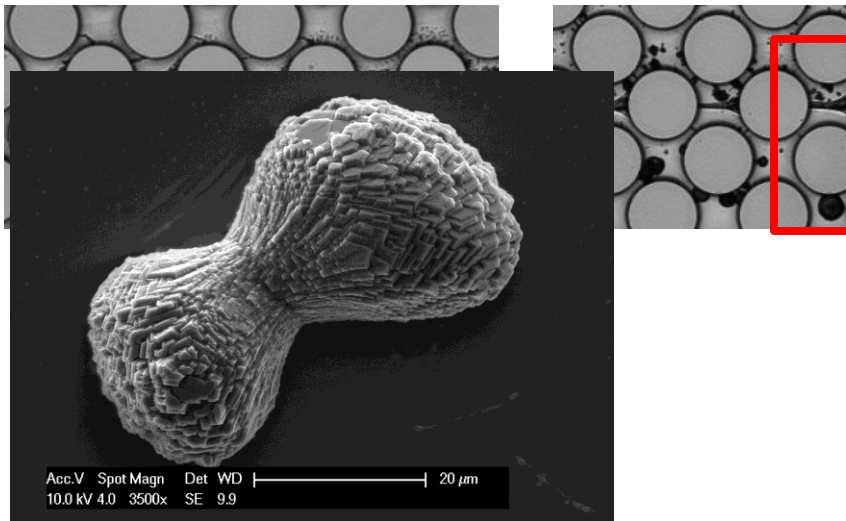
$[Mg^{2+}]$   
=0mM



$[Mg^{2+}]$   
=10mM



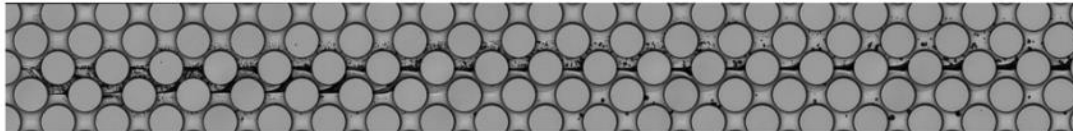
$[Mg^{2+}]$   
=40mM



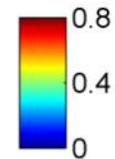
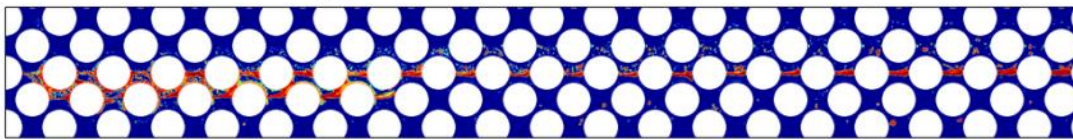
# Impact of precipitation on flow pattern and reaction kinetics

$$[\text{Ca}^{2+}]_{\text{T}} = [\text{CO}_3^{2-}]_{\text{T}} = 10 \text{ mM} \text{ \& } [\text{Mg}^{2+}]_{\text{T}} = 40 \text{ mM}$$

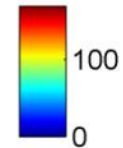
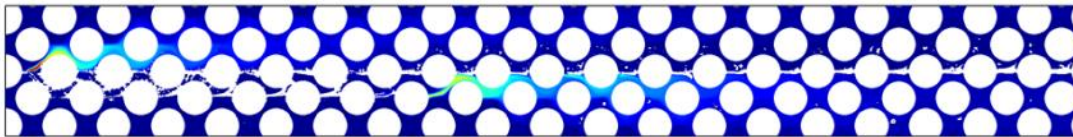
(a) Microscopy Image at 8hrs



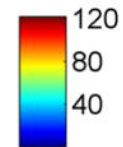
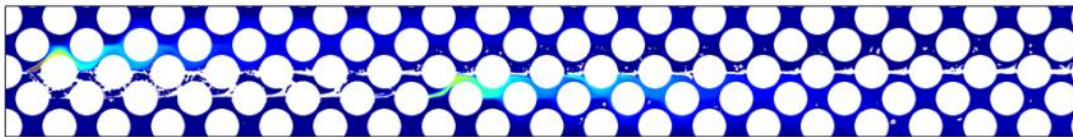
(b) Volumetric Fraction of Precipitate



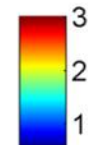
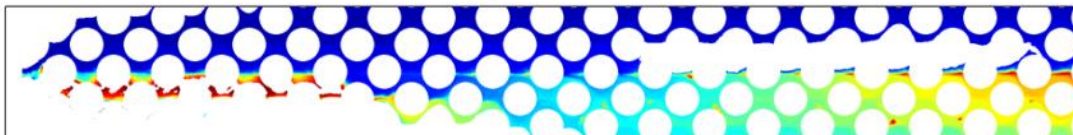
(c) Calcite SR



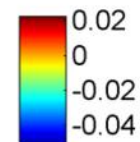
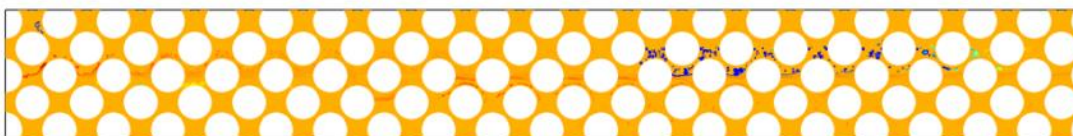
(d) Aragonite SR



(e)  $\text{Log}_{10}\{\text{Mg}^{2+}\}/\{\text{Ca}^{2+}\}$  Ratio

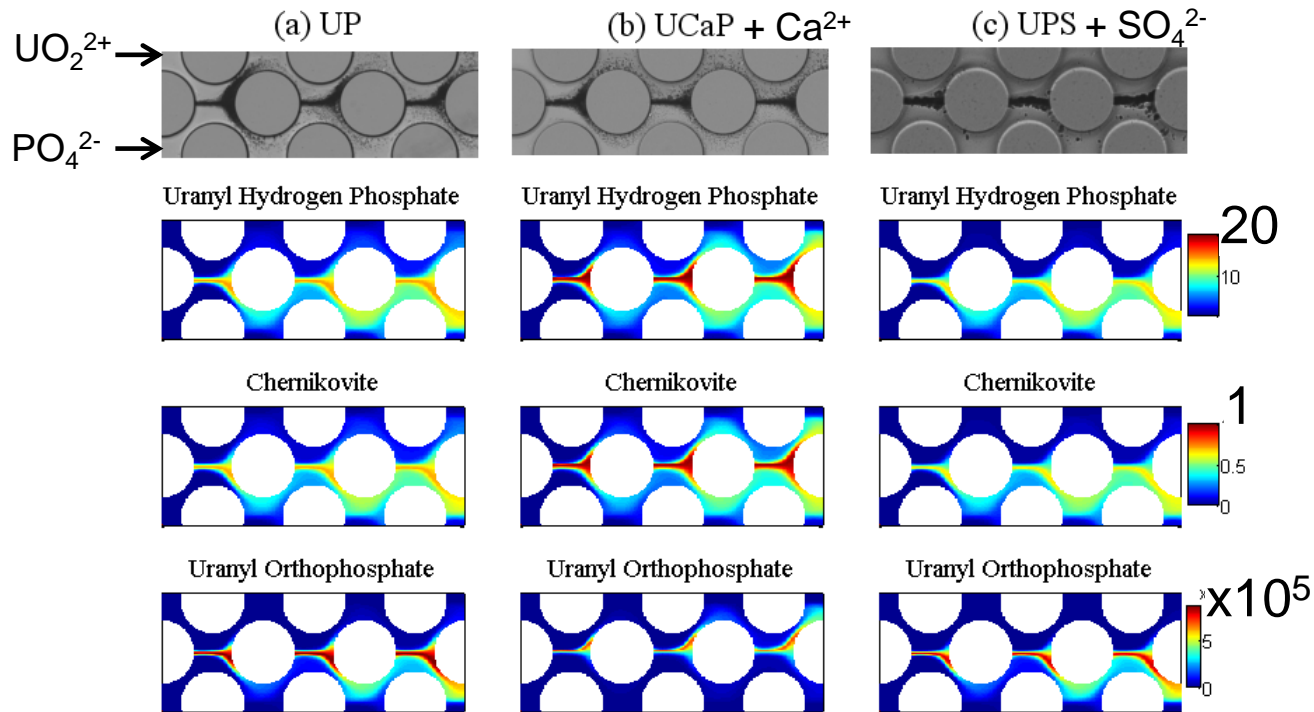


(f) Volumetric Reaction Rate (1/min)



# Impact of Solution Chemistry

Saturation Ratio (SR) = Ion Activity product /  $K_{sp}$   
 $SR > 1$ : thermodynamically favorable to form



Due to a kinetically controlled reaction,  
 only uranyl hydrogen phosphate (or chernikovite) is formed

# Summary and Implications

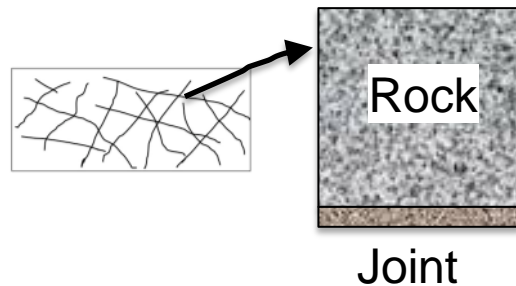
- Mineral precipitation rate along flow direction is concentration dependent and limited by transverse mixing
- $\text{CaCO}_3$  mineral phases are concentration dependent
- Overall, reaction kinetics, crystal growth and morphology are spatially and temporally affected by solution chemistry and hydrodynamics at pore scale
- Pore-scale model can be used to test if pore-scale processes observed in micromodels is predicted, and to develop an upscaled reaction model

Upscaling and  
Ongoing effort

# Upscaling/Hybrid modeling

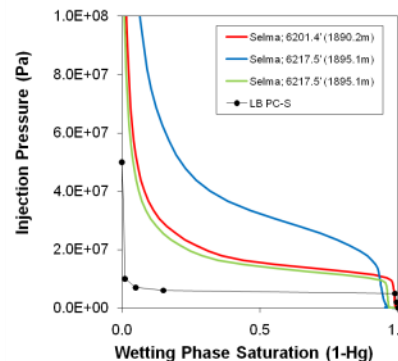
No general method/ framework for upscaling:

- Homogenization (multiple scale expansions)
- Volume averaging
- Pore network models (approx. physics)
- Mass balance principles based on pore scale models.
- Constitutive equation with closure based on detailed pore scale solutions (e.g., response function model)
- Hybrid pore-continuum approach



Geomodel (Kayanta):

Equivalent continuum joint model



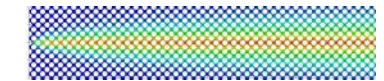
Pc-Sw Curve

Upscaled reaction rate

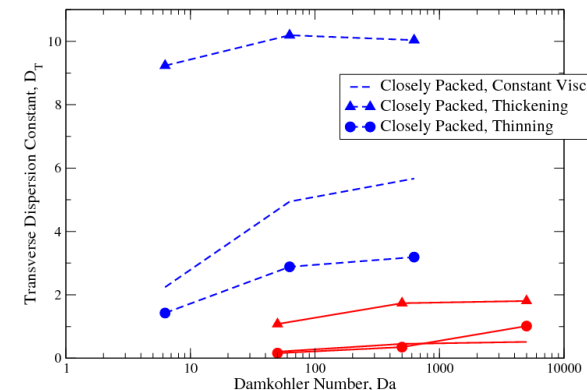
$$\dot{r}_{eff} = \frac{(\bar{C}_{in} - \bar{C}_{out})Q}{\nu A} \quad \bar{C}Q = \int_A \mathbf{u}c \cdot \mathbf{n} dS$$

(b)

Molins et al. WRR 2012



$$D_{T,R} = \frac{1}{X^3} \frac{9\pi\nu}{16C_0^2\phi^2} \left[ \int_0^X m(x) dx \right]^2$$

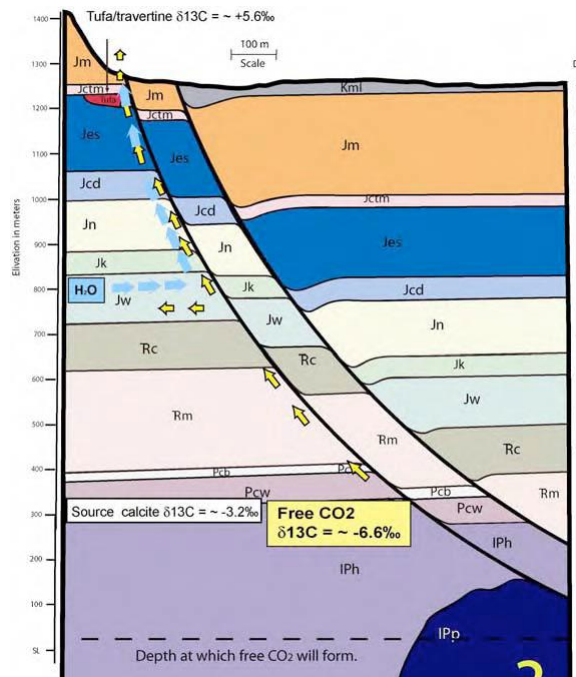


Upscaled dispersion

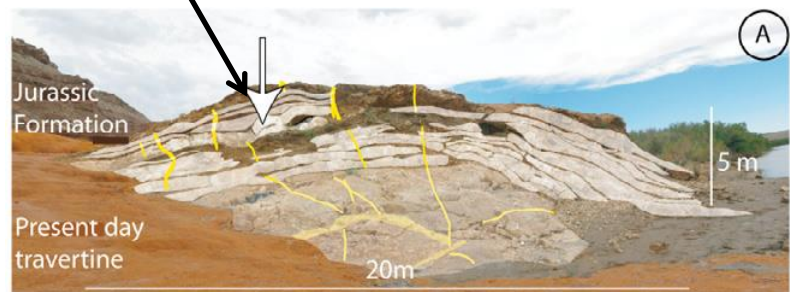
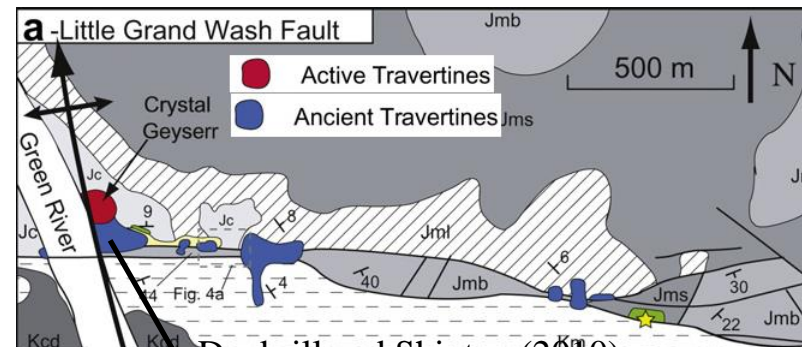
Davison, Yoon et al. (2012)

# Simulation of Little Grand Wash Fault

- Two hypotheses for CO<sub>2</sub> migration in the Crystal Geyser area:
  - CO<sub>2</sub> gas release from mantle along the fault
  - CO<sub>2</sub> dissolved in groundwater leaking up along the fault
- Develop scheme for selecting appropriate model for CO<sub>2</sub> leakage based on surface observation of travertine mounds



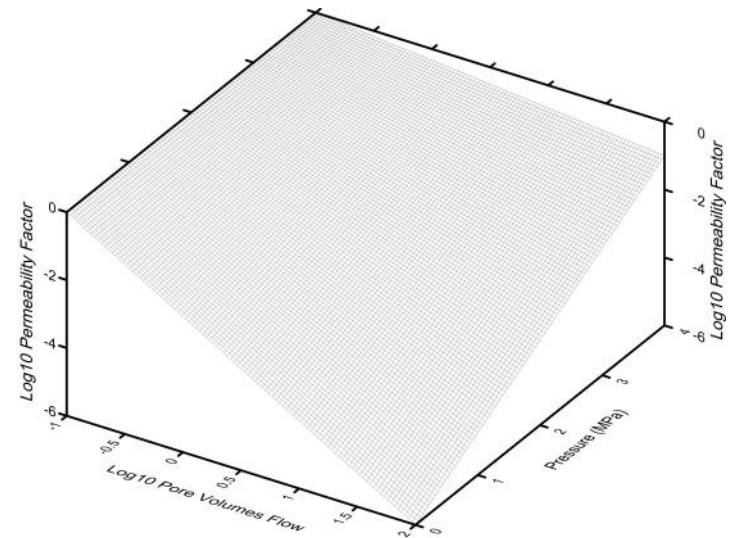
Evans et al. (2004)



Burnside (2010)

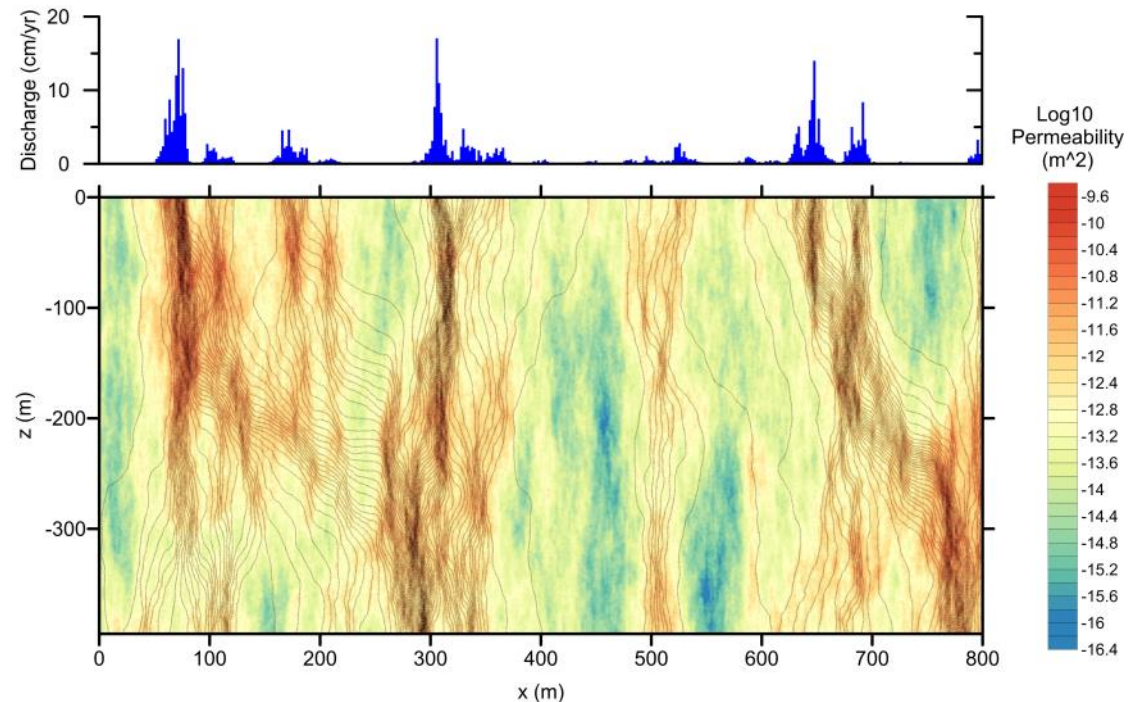
# Crystal Geyser Site: Grand Wash Fault Modeling

- Simplified, two-dimensional response surface defined for preliminary modeling example
- Permeability reduction due to calcite precipitation is a function of cumulative pore volume throughput of groundwater and as a function of fluid pressure
- Fluid pressure is taken as a gross proxy for chemical conditions in which higher calcite solubility is associated with higher fluid pressure (greater depth)



# Crystal Geyser Site: Grand Wash Fault Modeling

- Unconditional geostatistical simulation of initial permeability is semi-quantitatively consistent with geologic mapping of fault segments and alteration
- Initial simulated steady-state flow pattern is qualitatively similar to the spacing of springs and seeps along the Grand Wash fault with spacings of 100's of m between locations of groundwater discharge

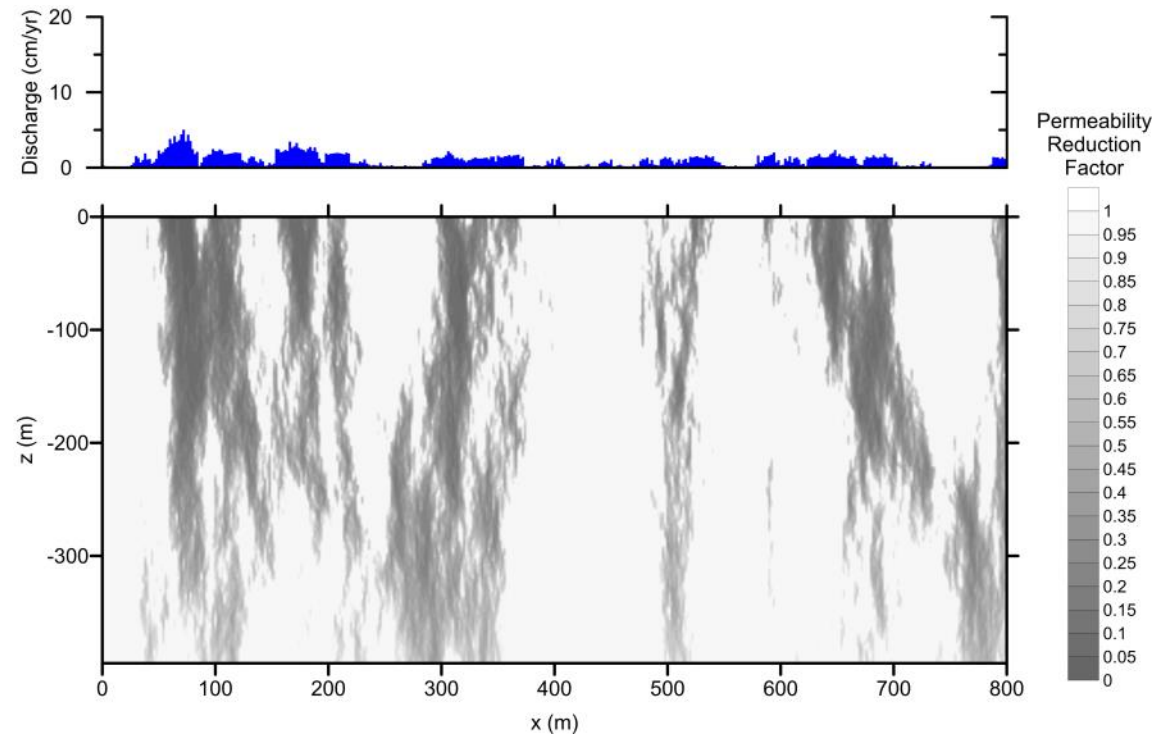


Time = 0

# Crystal Geyser Site: Grand Wash

## Fault Modeling

- Transient flow simulation includes explicit updating of the permeability field at each time step using the response surface shown previously
- Permeability is reduced by several orders of magnitude by calcite precipitation, primarily in the shallower high-flow channels
- Evolution of the flow field results in more dispersed groundwater discharge at the surface



Time = 1000 years

# Response Function based on Pore Scale Simulations

Influx conditions

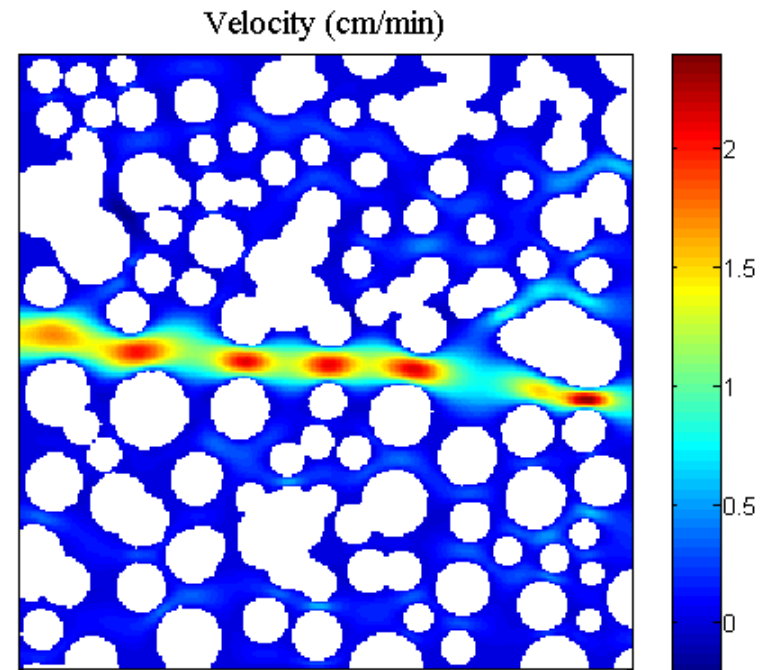
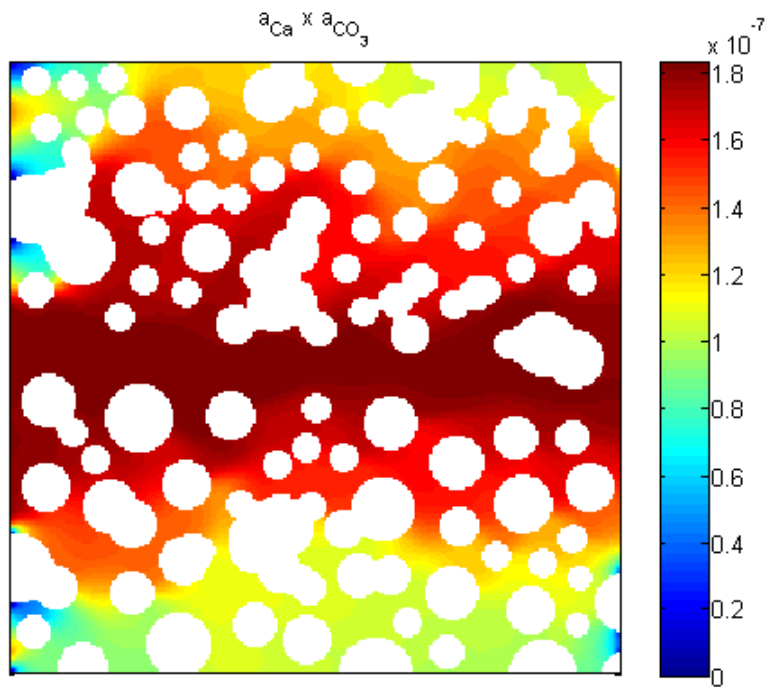
$[\text{Ca}^{2+}]_{\text{T}} = [\text{CO}_3^{2-}]_{\text{T}} = 20\text{mM}$

$\text{Pe} = 0.08, 0.8, 8$

$\text{Da} = 0.002, 0.02, 0.1$

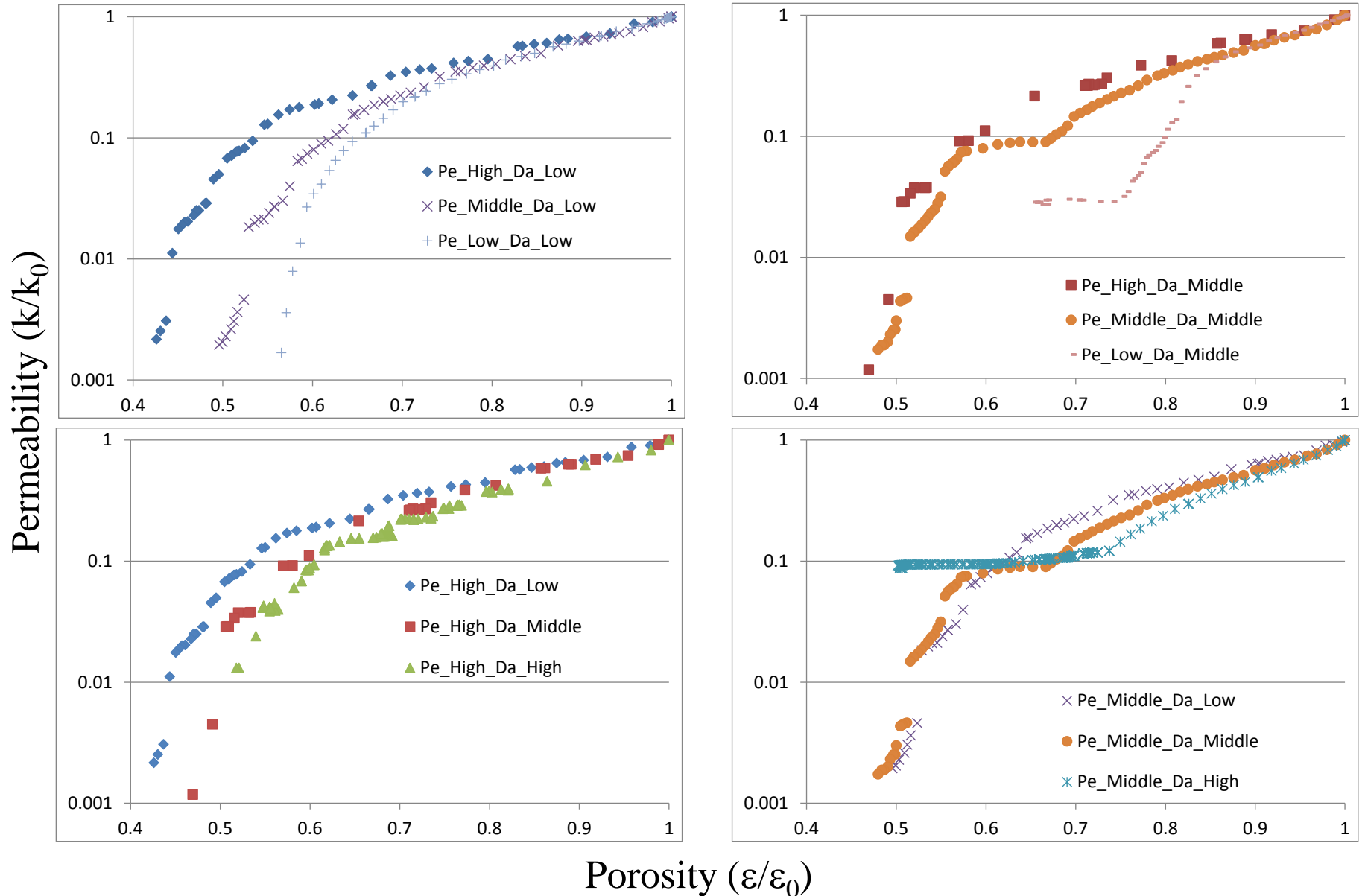
Speciation:  $\text{Ca}^{2+}$ ,  $\text{H}^+$ ,  $\text{CO}_3^{2-}$ ,  
 $\text{HCO}_3^-$ ,  $\text{H}_2\text{CO}_3$

No speciation:  $\text{Ca}^{2+}$ ,  $\text{CO}_3^{2-}$

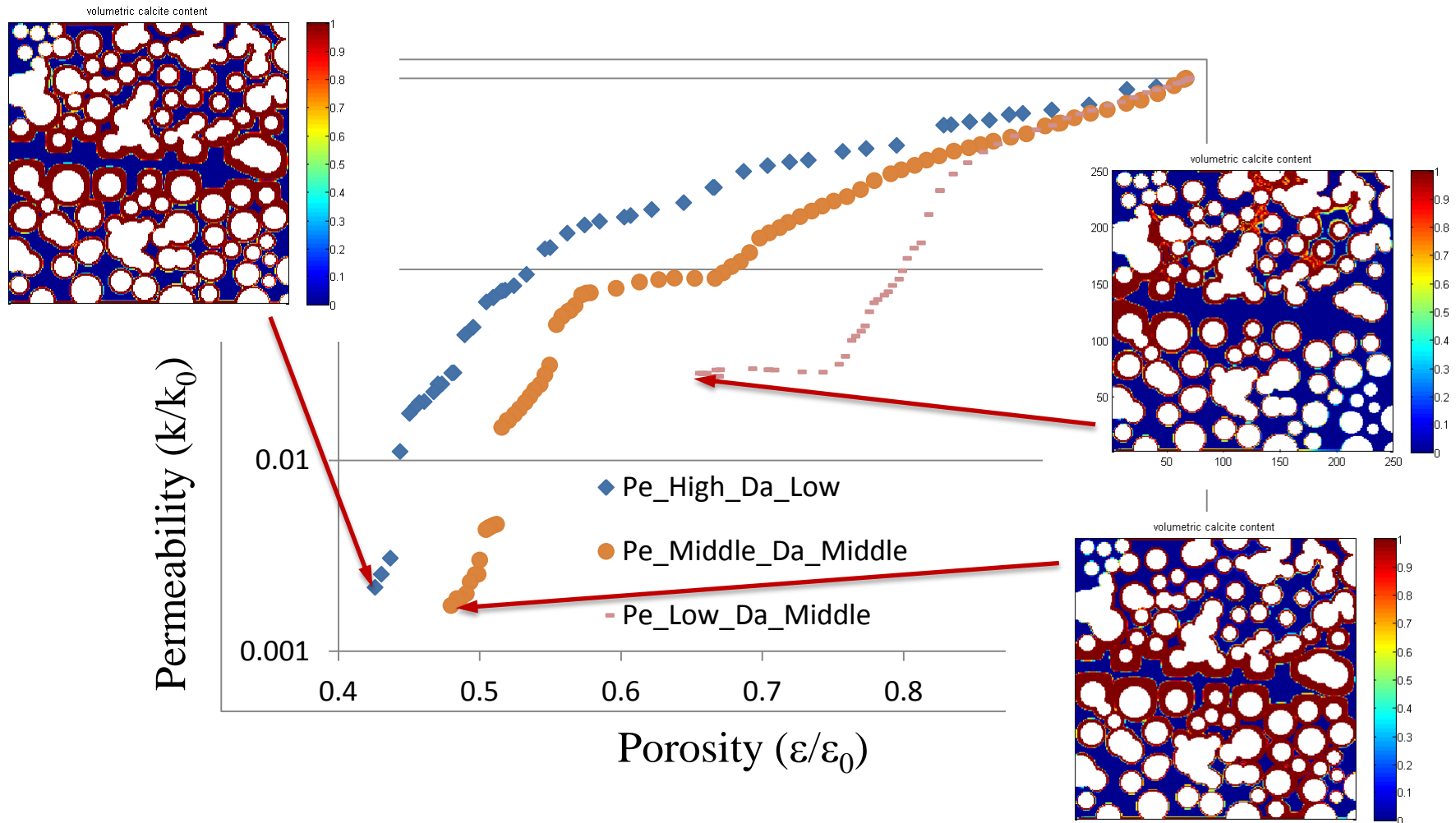


$K_{\text{sp, calcite}} = 3.3 \times 10^{-9} \text{ M}^2$

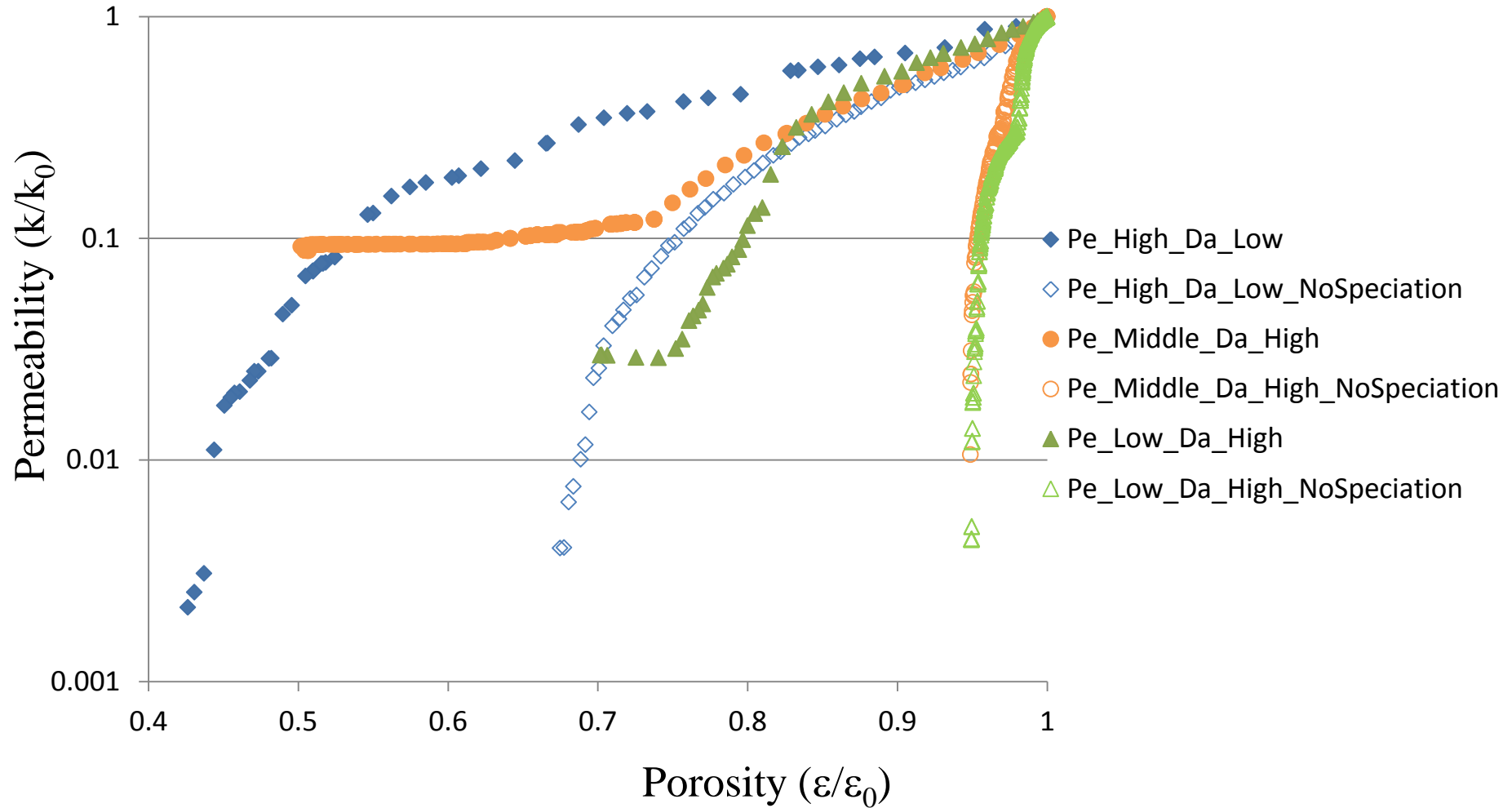
# Response Function based on Pore Scale Simulations



# Permeability-Porosity Relationships

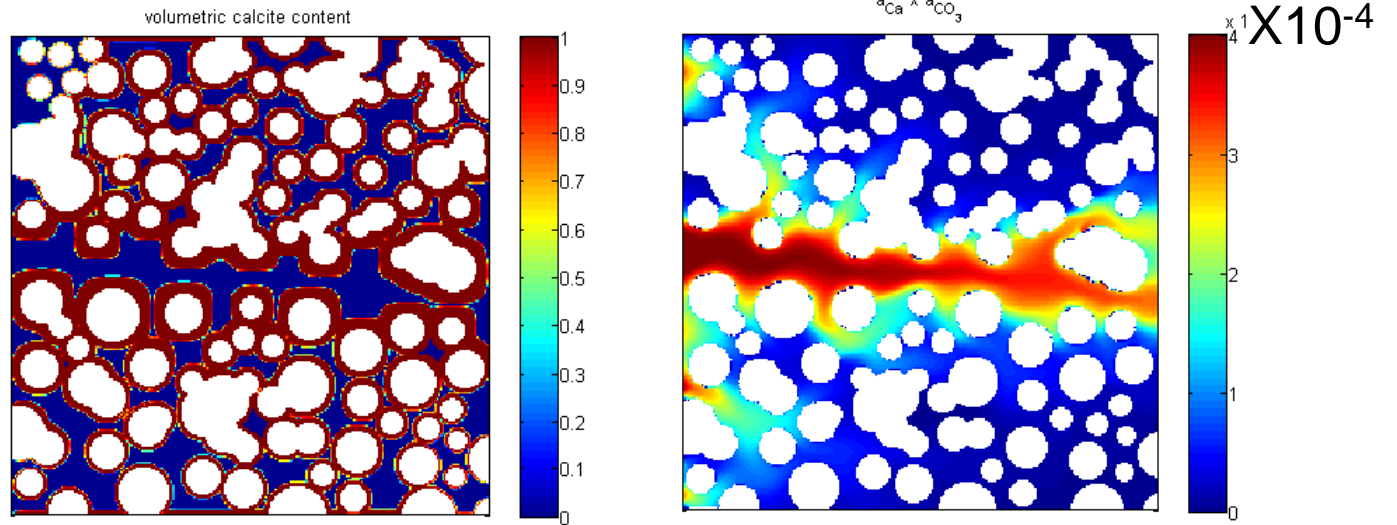


# Chemical Speciation

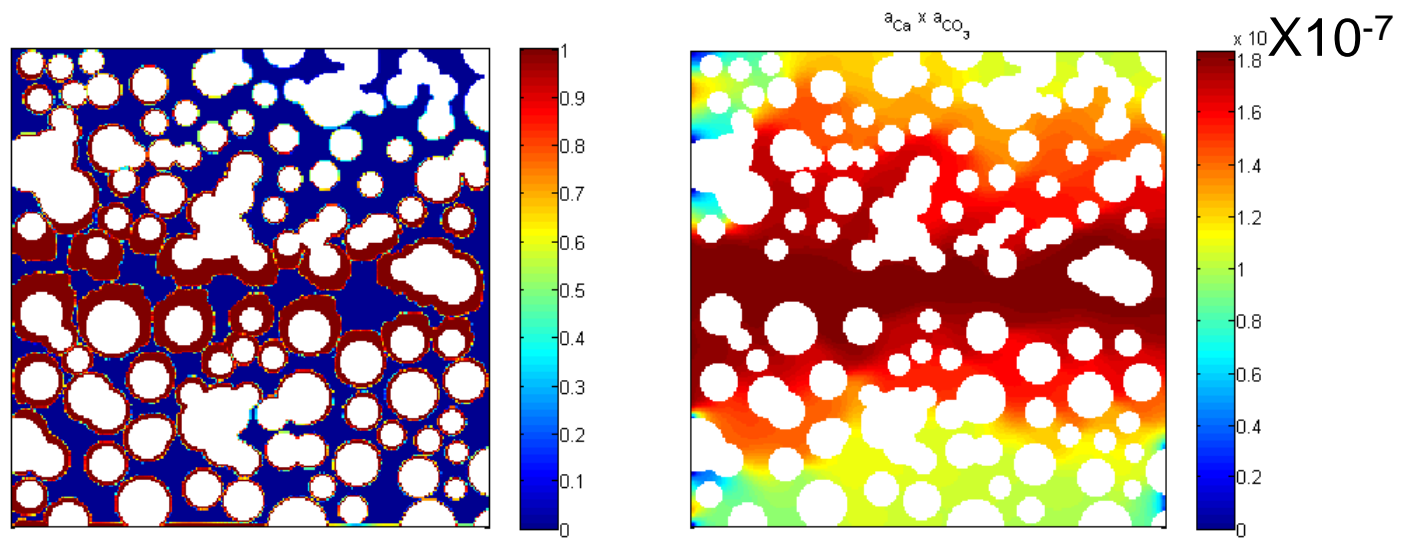


# High Pe & Low Da

Speciation

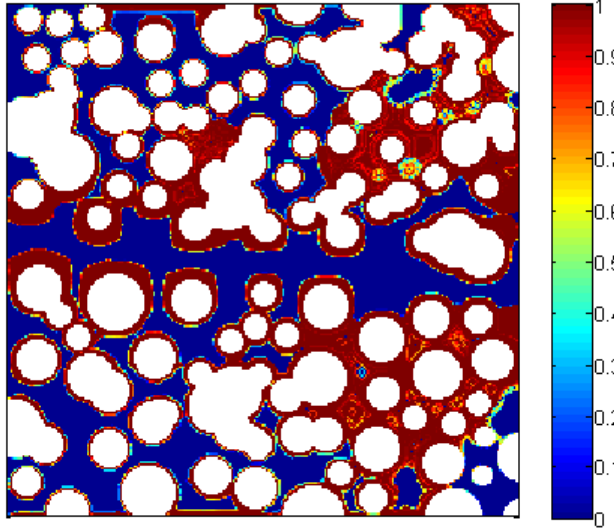


No speciation



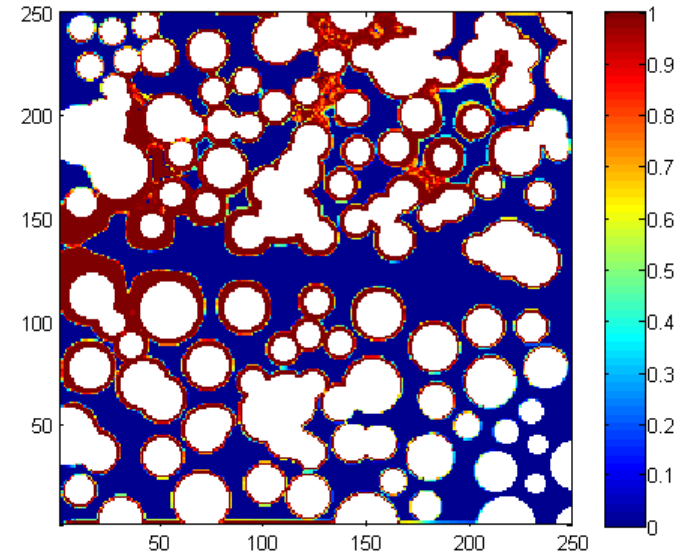
# Comparison

$Pe=0.8$ ;  $Da=0.1$

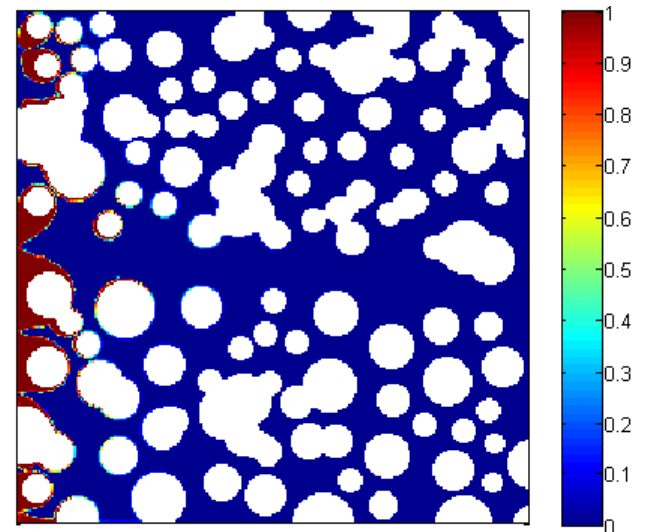
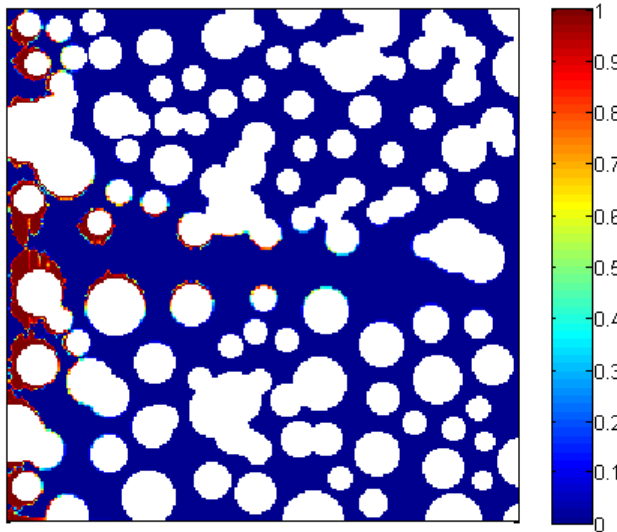


Speciation

$Pe=0.08$ ;  $Da=0.02$



No speciation



# Summary and Implications

- Vigorously tested pore-scale model can be used to develop a response function (or dimension reduced model) for permeability and porosity ( $k$ - $\epsilon$ ) relationships
- $k$ - $\epsilon$  relationships will be affected by solution chemistry, chemical reaction, pore structure configurations in addition to Pe and Da numbers
- Adaptive strategy to couple continuum and pore-scale using a response function approach will be tested

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