

# Liquid-vapor coexistence and critical point of $\text{Mg}_2\text{SiO}_4$ from *ab-initio* simulations

Joshua P. Townsend<sup>1</sup>, Gil Shohet<sup>2</sup>, Kyle R. Cochrane<sup>1</sup>

<sup>1</sup>High Energy Density Physics Theory, Sandia National Laboratories, Albuquerque, NM, 87185, USA

<sup>2</sup>Department of Aeronautics & Astronautics, Stanford University, Stanford, CA, 94305, USA

## Key Points:

- $\text{Mg}_2\text{SiO}_4$  vaporizes incongruently and produces Mg-poor and  $\text{O}_2$ ,  $\text{SiO}$  enriched vapor
- Critical point occurs at lower density and temperature than commonly used multi-phase EOS
- Giant impact simulations underestimate amount of supercritical material post-impact

---

Corresponding author: Joshua P. Townsend, [jptowns@sandia.gov](mailto:jptowns@sandia.gov)

12 **Abstract**

13 Hypervelocity impact-driven vaporization is characteristic of late-stage planet formation. Yet the  
 14 behavior and properties of liquid-vapor mixtures of planetary materials of interest are typically un-  
 15 known. Multiphase equations of state used in hydrodynamic simulations of planet impacts therefore  
 16 lack reliable data for this important phenomenon. Here, we present the first constraints on the liquid-  
 17 vapor critical point and coexistence phase boundary of  $Mg_2SiO_4$  computed from *ab-initio* molecular  
 18 dynamics simulations. We found that the vapor is depleted in magnesium and enriched in silica and  
 19 oxygen, while the coexisting liquid is enriched in magnesium and depleted in oxygen, from which  
 20 we infer vaporization is incongruent. The critical point was estimated from an equation of state fit  
 21 to the data. The results are in line with recent calculations of  $MgSiO_3$  and together confirm that  
 22 extant multiphase EOS models used in planetary accretion modeling significantly underestimate the  
 23 amount of supercritical material post-impact.

24 **1 Introduction**

25 Planets grow by accretion in which successive collisions of smaller bodies produce large ones.  
 26 The characteristic impact velocity for an Earth-mass planet is on the order of 10 km/s, and such  
 27 impacts generate significant amounts of hot liquid and vapor [Cuk and Stewart, 2012; Canup *et al.*,  
 28 2013; Nakajima and Stevenson, 2014; Lock and Stewart, 2017]. Beyond the Solar System, large  
 29 planet impact events have been inferred from astronomical observations of planets orbiting distant  
 30 stars [Meng *et al.*, 2014]. Therefore, the low-density and high-temperature properties of common  
 31 planetary materials like silicates and metals are of considerable interest to those studying planet  
 32 formation across a broad range of initial conditions.

33 Magnesium silicate  $Mg_2SiO_4$  is ubiquitous in the upper mantle of the Earth and other terrestrial  
 34 planets as well as pre-planetary and planetary nebulae elsewhere in the galaxy [Justtanont *et al.*,  
 35 1992; de Vries *et al.*, 2015]. In hydrodynamic simulations of giant-impact events a multiphase  
 36 equation of state (EOS) provides closure for mass, momentum, and energy conservation relations,  
 37 but both experimental data and theoretical models for low density high temperature  $Mg_2SiO_4$  are  
 38 lacking. Indeed, to-date there exist no experimental thermodynamic data of the critical behavior  
 39 for any magnesium silicate system. In the absence of any constraints from theory or experiment,  
 40 multiphase EOS models typically resort to overly simplistic coexistence models and therefore have  
 41 limited accuracy and may not produce quantitative predictions.

42 In this Letter, we report the results of a series of *ab-initio* density functional theory (DFT) based  
 43 molecular dynamics (DFTMD) calculations of  $\text{Mg}_2\text{SiO}_4$  that span liquid and vapor states (section  
 44 2.1 & 3.1). We show that vaporization occurs incongruently and that the chemical composition  
 45 of the coexisting liquid and vapor is sensitive to the bulk density and temperature (section 3.1).  
 46 We determine the critical point from an EOS model fit to the pressure-density-temperature states  
 47 compiled from DFTMD calculations (section 2.2 & 3.2). The rest of the coexistence curve was  
 48 determined from a fit to an analytic model for the phase boundary. We examine the sensitivity of  
 49 the estimated critical point to finite size effects and show the results from 224 atom calculations give  
 50 reasonable estimates of the critical point in the thermodynamic limit. We find that the critical point is  
 51 substantially different than those used in some previous hydrodynamic simulations of giant impacts,  
 52 and conclude that those simulations both underestimate the amount of supercritical material, and  
 53 over-estimate the thermodynamic span of liquid-vapor coexistence post-impact (section 4).

## 54 2 Methods

### 55 2.1 Molecular dynamics calculations

56 The gold-standard for computations of critical points is the Gibbs-ensemble Monte Carlo  
 57 (GEMC) method, in which the equilibrium thermodynamic properties of coexisting vapor and liquid  
 58 are simulated [Panagiotopoulos, 1987]. However, the limitations of this method when applied to  
 59 multicomponent systems motivates an exploration of alternatives for investigating critical behavior  
 60 in complex materials. This is especially important for systems in which vaporization may occur  
 61 incongruently, as expected for multicomponent systems like silicates and other planetary materials.  
 62 An Alternative to GEMC is molecular dynamics, in which the dynamics of the nuclei are simulated.  
 63 Molecular dynamics methods have been used to investigate liquid, vapor, coexistence and critical  
 64 behavior of both single and multicomponent materials with success [Guissani and Guillot, 1993;  
 65 Kresse and Hafner, 1997; de Koker *et al.*, 2009; Desjarlais, 2009; Faussurier *et al.*, 2009; Ghosh  
 66 and Karki; Xiao and Stixrude, 2018].

67 The goal of this study was to determine the liquid-vapor critical point and coexistence phase  
 68 boundaries of  $\text{Mg}_2\text{SiO}_4$  from *ab-initio* DFTMD calculations. To accomplish those goals we carried  
 69 out a series of DFTMD calculations on supercells spanning 5500-7500 K and 0.22-1.50 g/cm<sup>3</sup>, or  
 70 about 16-2 times expanded with respect to forsterite at ambient conditions. All DFTMD calculations  
 71 were performed using the VASP code (version 5.4.1), an implementation of Kohn-Sham density  
 72 functional theory (DFT) using a plane-wave basis, projector-augmented-wave pseudopotentials, and

73 periodic boundary conditions [Kresse and Hafner, 1993, 1994; Kresse and Furthmüller, 1996].  
 74 The irreducible Brillouin zone for those supercells was sampled at the Baldereschi mean value  
 75 ( $k=(0.25, 0.25, 0.25)$ ), and the electronic occupations were populated using a Fermi-Dirac distribution  
 76 according to the Mermin finite temperature formulation of DFT [Mermin, 1965]. The exchange  
 77 and correlation contributions to the energy were computed using the PBE generalized gradient  
 78 approximation [Perdew *et al.*, 1996]. The VASP PAW pseudopotentials were used with a plane wave  
 79 energy cutoff of 800 eV, which provided converged values of the stress tensor at the lowest density  
 80 we studied. Because the temperatures and densities considered in our study were not too high, we  
 81 used a magnesium pseudopotential with a  $3s^2$  valence configuration, while that of silicon and oxygen  
 82 were  $3s^23p^2$ , and  $2s^22p^4$ , respectively. All calculations were performed on cubic simulation cells  
 83 within the *NVT* ensemble using a velocity-scaling thermostat and a 1 femtosecond timestep for the  
 84 nuclear motion.

85 A system at a critical point is characterized by density fluctuations at all length scales. This is  
 86 problematic from the standpoint of atomistic simulations, which are necessarily finite. As a conse-  
 87 quence, finite size effects were a principal concern. We therefore performed DFTMD calculations in  
 88 supercells containing 28, 56, 112, and 224 atoms in order to quantify the finite size errors associated  
 89 with the critical point density, temperature, and pressure. Equilibration time of the system near the  
 90 critical point was another concern. Typically, we found that around 2-5 ps of simulated time was  
 91 sufficient to equilibrate the system, and that a further 4-5 ps of simulated time provided accurate  
 92 pressures. Equilibration was checked in a few cases by allowing some simulations to run for up to  
 93 100 ps in the 28 and 56 atom supercells.

## 94 2.2 EOS model and critical point estimation

95 The simplest equation of state (EOS) with a critical point is cubic in the density:

$$P(\rho, T) = (a_0 + a_1 T)\rho + (b_0 + b_1 T)\rho^2 + (c_0 + c_1 T)\rho^3 \quad (1)$$

96 where  $P$ ,  $\rho$ ,  $T$  are pressure, density, and temperature, respectively. Eqn. 1 is a generalized virial  
 97 equation of state and was selected for its flexibility and generality in describing both liquids and  
 98 gases. Cubic equations of state are commonly used to describe critical phenomena of a wide class  
 99 of materials [Wong and Sandler, 1992; Guissani and Guillot, 1993; Valderrama, 2003; Faussurier  
 100 *et al.*, 2009]. Regardless of the underlying EOS model, the critical point satisfies:

$$\left(\frac{\partial P}{\partial \rho}\right)_T = \left(\frac{\partial^2 P}{\partial \rho^2}\right)_T = 0 \quad (2)$$

which corresponds to the statement that the first order transition from liquid to vapor becomes second order at the critical point [Hansen and McDonald, 1990]. These criteria together with the EOS model yield an analytic solution for the critical point:

$$T_C = \frac{-B \pm \sqrt{B^2 - 4AC}}{2A} \quad (3)$$

$$\rho_C = -\frac{b(T_C)}{3c(T_C)} \quad (4)$$

$$P_C = P(\rho_C, T_C) \quad (5)$$

101 with  $A = a_1 c_1 - \frac{1}{3} b_1^2$ ,  $B = a_0 c_1 + a_1 c_0 - \frac{2}{3} b_0 b_1$ , and  $C = a_0 c_0 - \frac{1}{3} b_0^2$ .

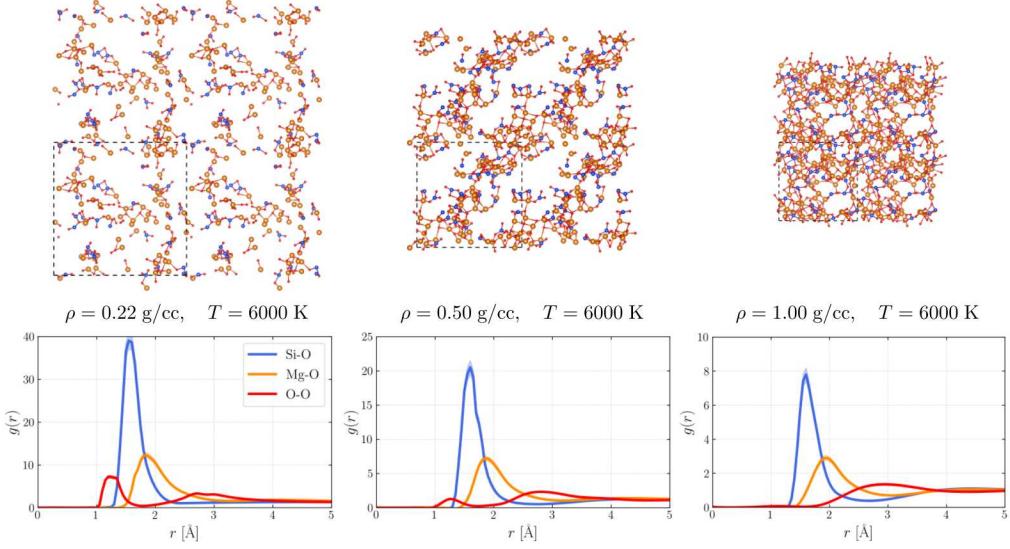
102 Our strategy is then straightforward: Run DFTMD simulations on a grid of  $\rho, T$  points for  
 103 systems of different sizes, then fit the subsequent  $P(\rho, T)$  data to the EOS model from which the  
 104 critical point is obtained for each system. The DFTMD data with best fit EOS models and critical  
 105 points for all systems studied are provided in the supplementary materials in fig. S1 and table S1.  
 106 The appeal of this approach is its simplicity. Our estimate of the critical point does not depend on  
 107 detailed knowledge of the internal distribution of liquid and vapor in the simulation cell, rather we  
 108 require only  $P(\rho, T)$ , which is an easily quantifiable bulk thermodynamic property which converges  
 109 rapidly with respect to the simulation time.

110 **3 Results**

111 **3.1 Chemical and structural properties in the vicinity of the critical point**

112 Direct inspection of the DFTMD trajectories indicates that the calculations spanned pure liquid,  
 113 coexisting liquid and vapor, and vapor states, as shown in fig. 1. Due to the chemical complexity of  
 114  $\text{Mg}_2\text{SiO}_4$ , we did not expect to be able to unambiguously label, in general, any particular collection  
 115 of atoms as belonging to either liquid or vapor. However, we found that the most abundant species  
 116 were easily identifiable small molecules such as O,  $\text{O}_2$ ,  $\text{SiO}$ ,  $\text{SiO}_2$ ,  $\text{MgO}$ , and Mg, which clearly  
 117 composed the vapor phase. The speciation of these small molecules for several isotherms is shown  
 118 in fig. 2 and the corresponding bond lifetimes are shown in fig. S3. At intermediate densities the  
 119 system tended to segregate into long-lived large globs of liquid and a variety of vapor species, and at  
 120 the highest density we studied,  $1.50 \text{ g/cm}^3$ , we find that the system is nearly completely liquid with  
 121 a small amount of transient molecular and atomic oxygen.

122 Broadly, where liquid and vapor coexisted, we found that the liquid was enriched in magne-  
 123 sium and depleted in silica and oxygen as compared to the vapor, which was enriched in oxygen  
 124 and silica and depleted in magnesium. For example, at  $6250 \text{ K}$  and  $0.50 \text{ g/cm}^3$  the composition



139 **Figure 1.** Selected snapshots of DFTMD trajectories (top) for the 224-atom system for various densities on  
140 6000 K isotherm, and corresponding spherically averaged pair distribution functions,  $g(r)$  (bottom). As the  
141 system is expanded coexistence of liquid and vapor are clearly observed. Magnesium, silicon, and oxygen are  
142 represented by orange, blue, and red spheres, respectively. The snapshots have been tiled to more clearly show  
143 the large-scale structure of the system (actual simulation regions marked with dashed lines).

125 of the liquid is  $\text{Mg}_{61.8 \pm 0.2}\text{Si}_{29.6 \pm 0.2}\text{O}_{116.0 \pm 0.5}$ , or normalized to silicon:  $\text{Mg}_{2.1 \pm 0.1}\text{Si}_{1.00}\text{O}_{3.9 \pm 0.2}$ .  
126 Similarly, for the vapor we found a composition of  $\text{Mg}_{2.2 \pm 0.2}\text{Si}_{2.4 \pm 0.2}\text{O}_{12.0 \pm 0.5}$ , or normalized to  
127 silicon:  $\text{Mg}_{0.9 \pm 0.1}\text{Si}_{1.0}\text{O}_{5.1 \pm 0.2}$ . Evidently, the liquid retained a composition of nearly stoichiometric  
128  $\text{Mg}_2\text{SiO}_4$ , while that of the vapor was somewhat different, in line with a recent study of  $\text{MgSiO}_3$   
129 at similar conditions [Xiao and Stixrude, 2018]. We also found that the onset of vaporization at  
130 all temperatures was characterized by the appearance of molecular and atomic oxygen, the latter of  
131 which, as expected, was more abundant at higher temperatures. The results suggest that vaporization  
132 of  $\text{Mg}_2\text{SiO}_4$  occurs incongruently (*i.e.* that the compositions of the coexisting liquid and vapor are  
133 different), similar to recent investigations of the critical behavior of other silicates [Kraus *et al.*,  
134 2012; Xiao and Stixrude, 2018; Connolly, 2016]. Although there are no experimental data on the  
135 critical behavior of  $\text{Mg}_2\text{SiO}_4$ , comparison of our results with experimental data on vaporization  
136 at lower pressures suggest that vaporization is complex; at lower pressure and temperature boiling  
137 of  $\text{Mg}_2\text{SiO}_4$  was previously reported to occur congruently [Nagahara *et al.*, 1994]. More recent  
138 experiments on iron-bearing olivine suggest incongruent vaporization [Costa *et al.*, 2017].

### 144 3.2 Critical Point Estimation & Uncertainty Quantification

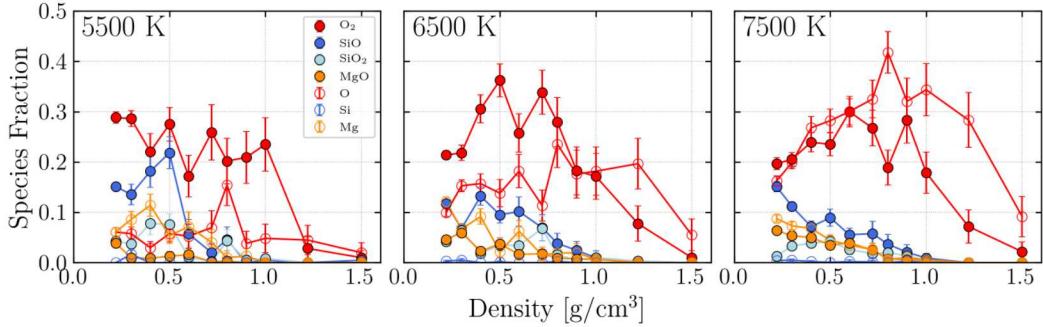
145 The critical point was computed analytically from a fit of the EOS given in eqn. 1 to the  
 146  $P(\rho, T)$  points from the DFTMD calculations. The 224 atom results are most representative of  
 147 the properties of the system in the thermodynamic limit, and so we report in the main text those  
 148 results unless otherwise indicated. The best-fit critical point values obtained are:  $\rho_c = 0.52 \pm 0.03$   
 149  $\text{g/cm}^3$ ,  $T_c = 6240 \pm 200 \text{ K}$ , and  $P_c = 1.3 \pm 0.2 \text{ kbar}$ . The data and corresponding best-fit EOS  
 150 are shown in fig. 3. In the coexistence region isotherms below the critical point clearly display  
 151 van der Waals loops. In the thermodynamic limit such curvature is unphysical because it implies a  
 152 negative compressibility and an imaginary sound speed. However, from the standpoint of atomistic  
 153 simulation, every point on the isotherm is stable and the curvature is in part due to finite size-effects  
 154 [Binder *et al.*, 2012]. This effect is readily observed from the results of the simulations by comparing  
 155 the curvature of the isotherms obtained from the different sized simulation cells: isotherms obtained  
 156 from simulations on the smaller cells have more pronounced curvature in the coexistence region  
 157 than those obtained from larger cells (fig. S1). Essentially, the result is that for the data presented  
 158 here, the local maxima and minima in the isotherms in the coexistence region do not correspond to  
 159 spinodal decomposition.

With one point on the liquid-vapor phase boundary in hand, we estimated the limits of co-existence below the critical point via the familiar Maxwell equal-area construction [Callen, 2006]. Famously, renormalization group theory applied to critical behavior in the 3d Ising model gives an analytic form for the phase boundaries for the vapor and liquid sides of the vapor dome:

$$\rho_v = \rho_c - \frac{1}{2} \left( C_1 x^\beta + C_2 x^{\beta+\Delta} \right) + C_3 x \quad (6)$$

$$\rho_l = \rho_c + \frac{1}{2} \left( C_1 x^\beta + C_2 x^{\beta+\Delta} \right) + C_3 x \quad (7)$$

160 with  $x \equiv \left( 1 - \frac{T}{T_c} \right)$ ,  $\beta = 0.325$ , and  $\Delta = \frac{1}{2}$  [Pathria and Beale, 2011; Hansen and McDonald, 1990].  
 161 Recently this approach has been employed by Xiao and Stixrude [2018] to describe coexistence in  
 162  $\text{MgSiO}_3$ . The phase boundary estimates from the Maxwell construction and subsequent analytic fit  
 163 for the 112 and 224 atom cells are shown in fig. 3.2. As expected, the vapor-side of the coexistence  
 164 region is poorly resolved, even in the largest simulation cell we studied. This is due to the finite size  
 165 of the simulations and because the isotherms are extrapolated below  $0.22 \text{ g/cm}^3$ . However, the liquid  
 166 side of the dome shows smaller finite size effects resulting in improved agreement between systems.  
 167 For that reason, only the points closest to the critical point on the vapor-side of the critical point were  
 168 included in the analytic fit, from which we obtained  $C_1 = 1.23 \pm 0.08$ ,  $C_2 = 2.28 \pm 0.35$ ,  $C_3 = 1.24 \pm$   
 169 0.14.

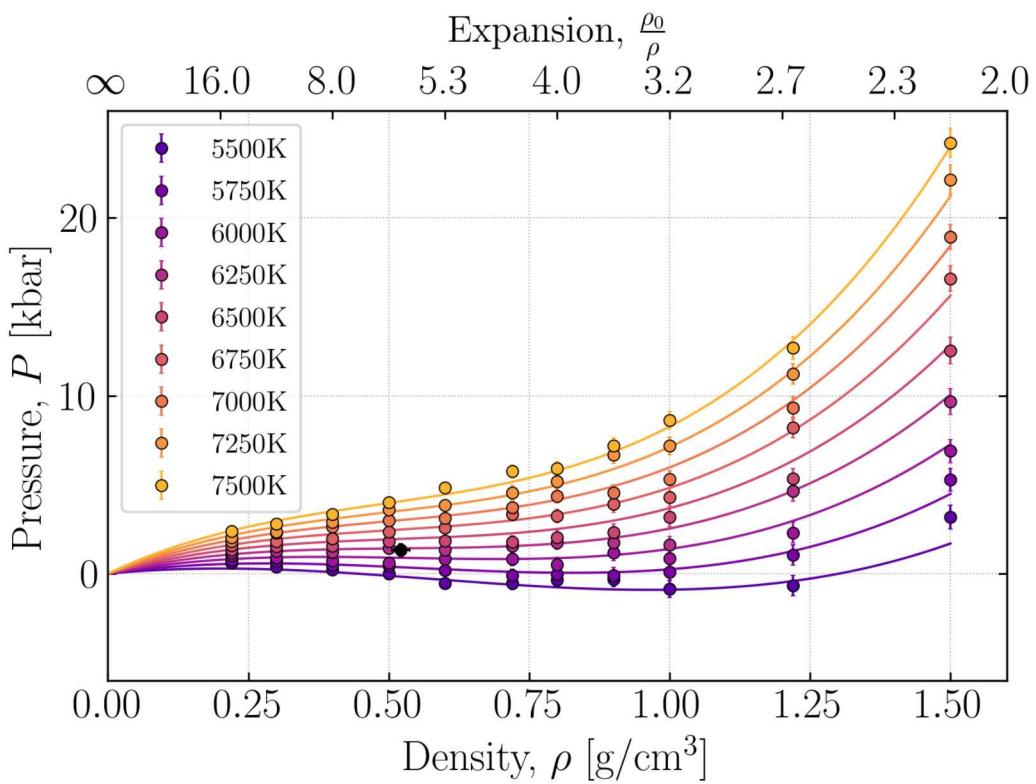


178 **Figure 2.** Abundance of selected light species in the vapor along 5500 K, 6500 K, and 7500 K isotherms.  
 179 The species fraction is the number of molecules of a particular type divided by the total number of molecules  
 180 in the cell.

170 Finally, it was important to compare the estimated critical points for systems of different sizes in  
 171 order to investigate the finite size effects associated with our estimation of the critical point. We find  
 172 that the critical points for the 112-atom and 224-atom cells are within their mutual uncertainties, as  
 173 shown in fig. 4. Combining the results from the 28-, 56-, 112-, and 224-atom cells and extrapolating  
 174 the critical point parameters yields an estimate of the critical point in the thermodynamic limit  
 175 (shown in fig. S2):  $\rho_c^{extrap} = 0.51 \pm 0.01 \text{ g/cm}^3$ ,  $T_C^{extrap} = 6110 \pm 26 \text{ K}$ , and  $P_c^{extrap} = 1.14 \pm 0.09$   
 176 kbar, which is within the uncertainties of the results from the 224 atom cell. We conclude that the  
 177 finite size effects in the critical point estimate in the 224 atom cell are small.

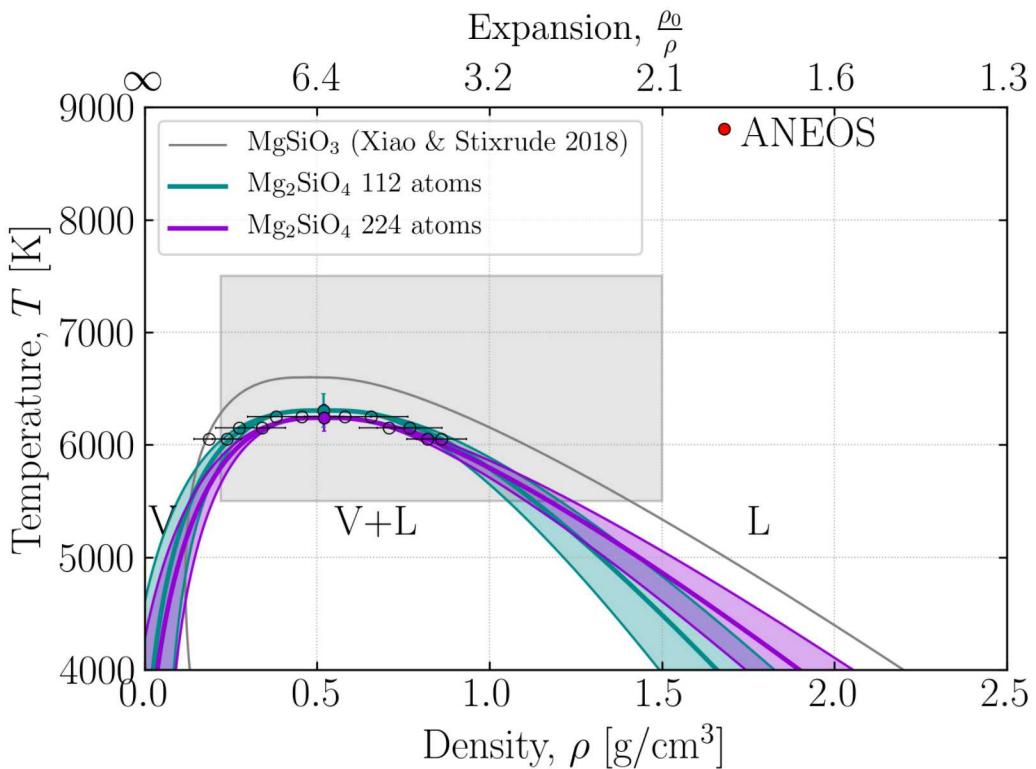
#### 187 4 Discussion & Conclusions

188 The major result of the present study establishes a first estimate of the critical behavior of  
 189  $\text{Mg}_2\text{SiO}_4$ , the most abundant phase in the Earth's outermost layers, as well as the Moon. We find  
 190 that the critical point and liquid-vapor phase boundary estimated from DFT-MD calculations differs  
 191 substantially from those in commonly used multiphase EOS models of giant impact and accretion  
 192 simulations (*cf.* [Canup \[2012\]](#); [Cuk and Stewart \[2012\]](#); [Nakajima and Stevenson \[2014\]](#); [Lock and](#)  
 193 [Stewart \[2017\]](#)). In the ANEOS model for forsterite, for example the critical point is at  $1.68 \text{ g/cm}^3$ ,  
 194  $8800 \text{ K}$ , and  $10 \text{ kbar}$ , while our results suggest a critical point at  $0.52 \text{ g/cm}^3$ ,  $6230 \text{ K}$ , and  $1.22 \text{ kbar}$   
 195 [[Stewart et al., 2020](#)]. This is significant as it suggests that nearly all accretionary collisions produce  
 196 some amount of vaporization [[Davies et al., 2020](#)]. In larger collisions characteristic of late-stage  
 197 accretion of Earth-sized planets, impact velocities often exceed  $10 \text{ km/s}$  [[Greenberg et al., 1978](#)]. The  
 198 heat and entropy liberated from such an event typically results in prodigious amounts of hot liquid,



181

**Figure 3.** EOS fit (lines) to DFTMD results (points) of 224 atom system with critical point (black circle).



182 **Figure 4.** (L)iquid-(V)apor phase boundaries obtained from Maxwell constructions from 224 and 112-atom  
 183 data sets (right). Filled circles mark critical points obtained from EOS. Unfilled circles mark phase boundary  
 184 estimates from the Maxwell constructions. Lines and shaded regions show phase boundary and  $1\sigma$  confidence  
 185 interval. The grey box shows the simulation domain. The red point marks the location of the critical point in  
 186 the tabulated EOS models used in some hydrodynamic simulations of giant impacts [Stewart *et al.*, 2019].

199 vapor, and supercritical fluid immediately post-impact [Lock and Stewart, 2017]. How the system  
200 evolves post-impact would obviously influence the compositions of any eventual planet or moon.  
201 Therefore, hydrodynamic simulations that rely on incorrect EOS models will underestimate the  
202 quantity of supercritical material, and overestimate the quantity of distinct liquid and vapor phases,  
203 which might subsequently influence the chemical compositions of the resulting bodies [Stewart *et al.*,  
204 2020].

205 There are several limitations in the present study which suggest possible avenues of future  
206 research. As with all DFT calculations, the choice of exchange correlation functional introduces a  
207 systematic bias in the results. Although we expect this bias to be relatively small for the problem  
208 studied here, the choice of functional on critical point parameters is an interesting topic for future  
209 study. It is notable that at intermediate densities the atomic fraction of the cell that is composed  
210 of vapor is only about 16 out of 224 atoms, about 7%, so larger simulations may better resolve  
211 the speciation in the vapor. Similarly, the vapor side of the Maxwell construction would likely be  
212 better resolved in much larger systems which may soon be possible through the use of sophisticated  
213 classical interatomic potentials such as the recently developed SNAP [Wood and Thompson, 2018].

214 However, the main limitation of both this and previous work on silicates is the assumption of  
215 single-component-like critical behavior. Although the EOS isotherms of the present study show all  
216 the trademark features of single component critical behavior,  $Mg_2SiO_4$  is a multicomponent system,  
217 and therefore may not have a single critical point. Despite this assumption, an in-depth study of  
218 the critical behavior of  $SiO_2$  by Connolly [2016] suggests that the width of the liquid-vapor phase  
219 boundary in  $SiO_2$  is relatively narrow and thus is likely smaller than the uncertainties on the densities  
220 of the coexisting liquid and vapor from our calculations, as shown in fig. 4. In accordance with other  
221 silica-bearing systems studied thus far, the assumption of single component-like critical behavior  
222 seems reasonable given the current state of knowledge of these materials, and the capabilities of  
223 current hydrodynamic simulation techniques [Kraus *et al.*, 2012; Connolly, 2016; Xiao and Stixrude,  
224 2018; Stewart *et al.*, 2020].

225 These limitations aside, the advantage of the present approach is that the estimation of the  
226 critical point does not depend on detailed knowledge of the atomic-scale structures in the simulation,  
227 but instead on the bulk thermodynamic properties of the mixed system. The fact that the present and  
228 previous studies of Mg-bearing silicate systems find remarkably similar critical points despite using  
229 entirely independent methods lends confidence to these approaches despite their assumptions.

230 In summary, the study of the critical behavior of planetary materials offers important constraints  
231 on multiphase EOS models needed for accurate studies of planet formation. For the  $\text{Mg}_2\text{SiO}_4$  system,  
232 recent experimental and theoretical studies have led to significant improvements in the EOS in high  
233 density, high pressure, high temperature regimes [Root *et al.*, 2018; Stewart *et al.*, 2019]. The present  
234 work extends these improvements to low pressure, low density, and moderate temperature.

235 **Acknowledgments**

236 The authors thank Michael P. Desjarlais, Luke Shulenburger, Raymond C. Clay III, and Sarah T.  
237 Stewart for many fruitful discussions which greatly improved the quality of this manuscript. G.S.  
238 was supported by the DOE SSGF program grant number DE-NA0003864. J.P.T. and K.R.C. ac-  
239 knowledge support from the Z Fundamental Science Program. Sandia National Laboratories is  
240 a multimission laboratory managed and operated by National Technology and Engineering Solu-  
241 tions of Sandia, LLC., a wholly owned subsidiary of Honeywell International, Inc., for the U.S.  
242 Department of Energy's National Nuclear Security Administration under contract DE-NA0003525.  
243 This paper describes objective technical results and analysis. Any subjective views or opinions that  
244 might be expressed in the paper do not necessarily represent the views of the U.S. Department of  
245 Energy or the United States Government. The data used in this manuscript can be obtained from  
246 <https://doi.org/10.5281/zenodo.3970968>.

247 **References**

- 248 Binder, K., B. J. Block, P. Virnau, and A. Tröster (2012), Beyond the van der waals loop: What can  
 249 be learned from simulating lennard-jones fluids inside the region of phase coexistence, *American*  
 250 *Journal of Physics*, 80, 1099–1109, doi:10.1119/1.4754020.
- 251 Callen, H. B. (2006), *Thermodynamics & an Intro. to Thermostatistics*, John Wiley & Sons.
- 252 Canup, R., A. Barr, and D. Crawford (2013), Lunar-forming impacts: High-resolution sph and  
 253 amr-cth simulations, *Icarus*, 222, 200 – 219, doi:<https://doi.org/10.1016/j.icarus.2012.10.011>.
- 254 Canup, R. M. (2012), Forming a moon with an earth-like composition via a giant impact, *Science*,  
 255 338, 1052.
- 256 Connolly, J. A. (2016), Liquid-vapor phase relations in the si-o system: A calorically constrained  
 257 van der waals-type model, *Journal of Geophysical Research: Planets*, 121, 1641–1666, doi:  
 258 10.1002/2016JE005059.
- 259 Costa, G. C., N. S. Jacobson, and B. F. Jr. (2017), Vaporization and thermodynamics of forsterite-rich  
 260 olivine and some implications for silicate atmospheres of hot rocky exoplanets, *Icarus*, 289, 42 –  
 261 55, doi:<https://doi.org/10.1016/j.icarus.2017.02.006>.
- 262 Cuk, M., and S. T. Stewart (2012), Making the moon from a fast-spinning earth: A giant impact  
 263 followed by resonant despinning, *Science*, 338, 1047.
- 264 Davies, E. J., P. J. Carter, S. Root, R. G. Kraus, D. K. Spaulding, S. T. Stewart, and S. B. Jacobsen  
 265 (2020), Silicate melting and vaporization during rocky planet formation, *Journal of Geophysical*  
 266 *Research: Planets*, 125(2), e2019JE006227, doi:10.1029/2019JE006227.
- 267 de Koker, N. P., L. Stixrude, and B. B. Karki (2009), Thermodynamics, structure, dynamics, and  
 268 freezing of mg<sub>2</sub>si<sub>4</sub> liquid at high pressure, *Geochimica et Cosmochimica Acta*, 72, 1427–1441,  
 269 doi:10.1016/j.gca.2007.12.019.
- 270 de Vries, B. L., K. Maaskant, M. Min, R. Lombaert, L. Waters, and J. Blommaert (2015), Micron-  
 271 sized forsterite grains in the pre-planetary nebula of iras 17150- 3224-searching for clues to the mys-  
 272 terious evolution of massive agb stars, *Astronomy & Astrophysics*, 576, A98, doi:10.1051/0004-  
 273 6361/201424843.
- 274 Desjarlais, M. P. (2009), Quantum molecular dynamics simulations for generating equation of state  
 275 data, *AIP Conference Proceedings*, 1161, 32–38, doi:10.1063/1.3241205.
- 276 Faussurier, G., C. Blancard, and P. L. Silvestrelli (2009), Evaluation of aluminum critical point using  
 277 an ab initio variational approach, *Phys. Rev. B*, 79, 134,202, doi:10.1103/PhysRevB.79.134202.
- 278 Ghosh, D. B., and B. B. Karki ().

- 279 Greenberg, R., J. F. Wacker, W. K. Hartmann, and C. R. Chapman (1978), Planetesimals to planets:  
280 Numerical simulation of collisional evolution, *Icarus*, 35, 1 – 26.
- 281 Guissani, Y., and B. Guillot (1993), A computer simulation study of the liquid–vapor coexistence  
282 curve of water, *The Journal of Chemical Physics*, 98(10), 8221–8235, doi:10.1063/1.464527.
- 283 Hansen, J.-P., and I. R. McDonald (1990), *Theory of simple liquids*, Elsevier.
- 284 Justtanont, K., M. J. Barlow, C. J. Skinner, and A. G. G. M. Tielens (1992), The Nature of Dust  
285 around the Post–Asymptotic Giant Branch Objects HD 161796 and HD 179821, *The Astrophysical  
286 Journal*, 392, L75, doi:10.1086/186429.
- 287 Kraus, R. G., S. T. Stewart, D. C. Swift, C. A. Bolme, R. F. Smith, S. Hamel, B. D. Hammel, D. K.  
288 Spaulding, D. G. Hicks, J. H. Eggert, and G. W. Collins (2012), Shock vaporization of silica  
289 and the thermodynamics of planetary impact events, *J. Geophys. Res. - Planet*, 117, E09,009,  
290 doi:10.1029/2012JE004082.
- 291 Kresse, G., and J. Furthmüller (1996), Efficient iterative schemes for ab initio total-energy calcula-  
292 tions using a plane-wave basis set, *Phys. Rev. B*, 54, 11,169.
- 293 Kresse, G., and J. Hafner (1993), Ab initio molecular dynamics for liquid metals, *Phys. Rev. B*, 47,  
294 R558.
- 295 Kresse, G., and J. Hafner (1994), Ab initio molecular-dynamics simulation of the liquid-metal-  
296 amorphous-semiconductor transition in germanium, *Phys. Rev. B*, 49, 14,251.
- 297 Kresse, G., and J. Hafner (1997), Ab initio simulation of the metal/nonmetal transition in expanded  
298 fluid mercury, *Phys. Rev. B*, 55, 7539–7548, doi:10.1103/PhysRevB.55.7539.
- 299 Lock, S. J., and S. T. Stewart (2017), The structure of terrestrial bodies: Impact heating, coro-  
300 nation limits, and synestias, *Journal of Geophysical Research: Planets*, 122, 950–982, doi:  
301 10.1002/2016JE005239.
- 302 Meng, H. Y. A., K. Y. L. Su, G. H. Rieke, D. J. Stevenson, P. Plavchan, W. Rujopakarn, C. M. Lisse,  
303 S. Poshyachinda, and D. E. Reichart (2014), Large impacts around a solar-analog star in the era of  
304 terrestrial planet formation, *Science*, 345(6200), 1032–1035, doi:10.1126/science.1255153.
- 305 Mermin, N. D. (1965), Properties of the inhomogeneous electron gas, *Phys. Rev.*, 137, A1441.
- 306 Nagahara, H., I. Kushiro, and B. O. Mysen (1994), Evaporation of olivine: Low pressure phase rela-  
307 tions of the olivine system and its implication for the origin of chondritic components in the solar  
308 nebula, *Geochimica et Cosmochimica Acta*, 58, 1951 – 1963, doi:[https://doi.org/10.1016/0016-7037\(94\)90426-X](https://doi.org/10.1016/0016-7037(94)90426-X).
- 310 Nakajima, M., and D. J. Stevenson (2014), Investigation of the initial state of the moon-  
311 forming disk: Bridging sph simulations and hydrostatic models, *Icarus*, 233, 259 – 267, doi:

- 312 <https://doi.org/10.1016/j.icarus.2014.01.008>.
- 313 Panagiotopoulos, A. Z. (1987), Direct determination of phase coexistence properties of flu-  
314 ids by monte carlo simulation in a new ensemble, *Molecular Physics*, *61*, 813–826, doi:  
315 10.1080/00268978700101491.
- 316 Pathria, R., and P. Beale (2011), *Statistical Mechanics 3rd ed.*, Academic Press, Boston.
- 317 Perdew, J., K. Burke, and M. Ernzerhof (1996), Generalized gradient approximation made simple,  
318 *Phys. Rev. Lett.*, *77*, 3865–3868, doi:10.1103/PhysRevLett.77.3865.
- 319 Root, S., J. P. Townsend, E. Davies, R. W. Lemke, D. E. Bliss, D. E. Fratanduono, R. G. Kraus,  
320 M. Millot, D. K. Spaulding, L. Shulenburger, S. T. Stewart, and S. B. Jacobsen (2018), The  
321 principal hugoniot of forsterite to 950 gpa, *Geophysical Research Letters*, *45*, 3865–3872, doi:  
322 10.1029/2017GL076931.
- 323 Stewart, S., E. Davies, M. Duncan, S. Lock, S. Root, J. Townsend, R. Kraus, R. Caracas, and  
324 S. Jacobsen (2020), The shock physics of giant impacts: Key requirements for the equations of  
325 state, *21st Biennial APS Conference on Shock Compression of Condensed Matter (SCCM19)*.
- 326 Stewart, S. T., E. J. Davies, M. S. Duncan, S. J. Lock, S. Root, J. P. Townsend, R. G. Kraus,  
327 R. Caracas, and S. B. Jacobsen (2019), The shock physics of giant impacts: Key requirements for  
328 the equations of state.
- 329 Valderrama, J. O. (2003), The state of the cubic equations of state, *Industrial & Engineering  
330 Chemistry Research*, *42*(8), 1603–1618, doi:10.1021/ie020447b.
- 331 Wong, D. S. H., and S. I. Sandler (1992), A theoretically correct mixing rule for cubic equations of  
332 state, *AIChE Journal*, *38*(5), 671–680, doi:10.1002/aic.690380505.
- 333 Wood, M. A., and A. P. Thompson (2018), Extending the accuracy of the snap interatomic potential  
334 form, *The Journal of Chemical Physics*, *148*, 241,721, doi:10.1063/1.5017641.
- 335 Xiao, B., and L. Stixrude (2018), Critical vaporization of mgsio3, *Proceedings of the National  
336 Academy of Sciences*, *115*, 5371–5376, doi:10.1073/pnas.1719134115.