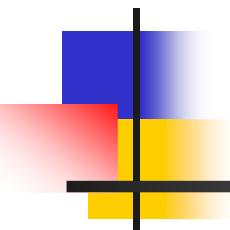


Pore-scale Simulation of Mixing-induced Calcium Carbonate Precipitation and Dissolution in a Microfluidic Pore Network



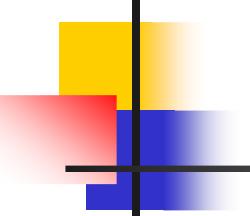
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PNNL, Aug. 9, 2011

Collaborators: Profs. Albert Valocchi and Charlie Werth (UIUC)
Dr. Tom Dewers (SNL)

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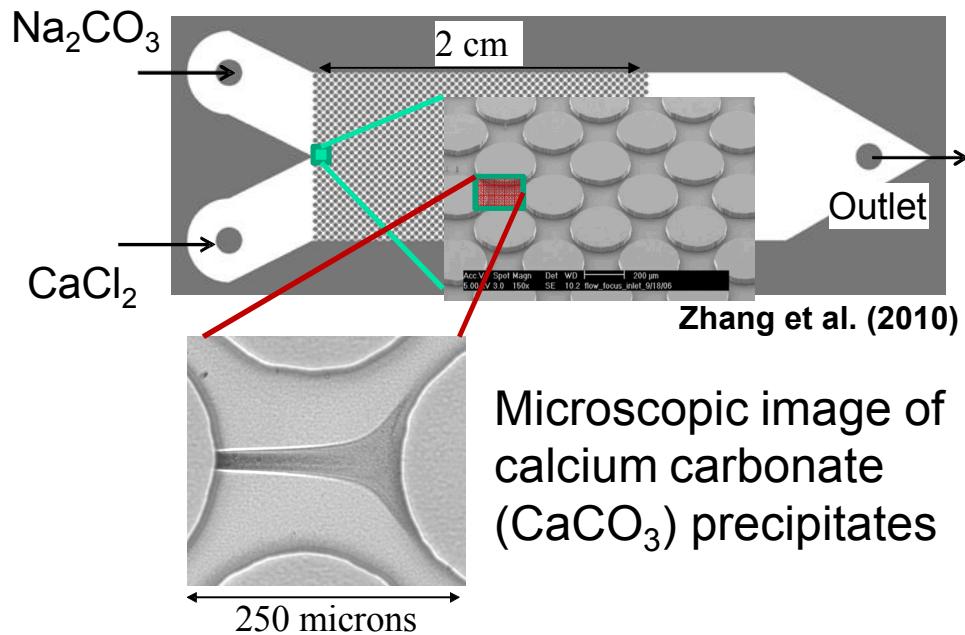


Motivations

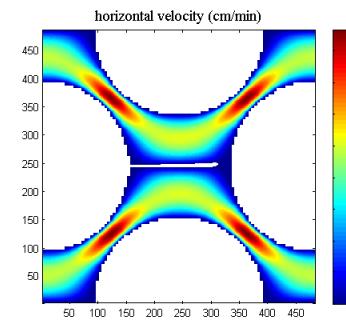
- Formation of Calcium Carbonate (CaCO_3) plays an important role in many natural and engineered systems
 - Geological CO_2 storage
 - Growth of calcite concretions in fractured and porous media
 - Biomineralization of products from microorganisms
 - Formation of scale in boilers and cooling towers
- Pore scale mixing and reaction can affect CO_2 injection efficiency and storage capacity
 - Reactions in the field (e.g., flowing conditions) can be limited by rates of transverse mixing
 - Recent development of in-situ measurement techniques for (sub) pore-scale reactive transport experiments provides a unique opportunity to test and validate pore-scale modeling approaches

Methods

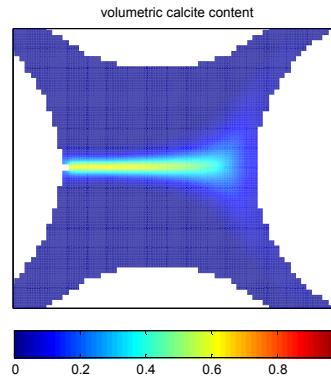
Experimental setup



Pore scale modeling



Water flow velocity in pore space
(resolution: 1 micron)



Simulation result of calcium carbonate (CaCO_3) precipitates

- Two solutions are mixing along the centerline and CaCO_3 precipitates
- Microscopic images are taken over time

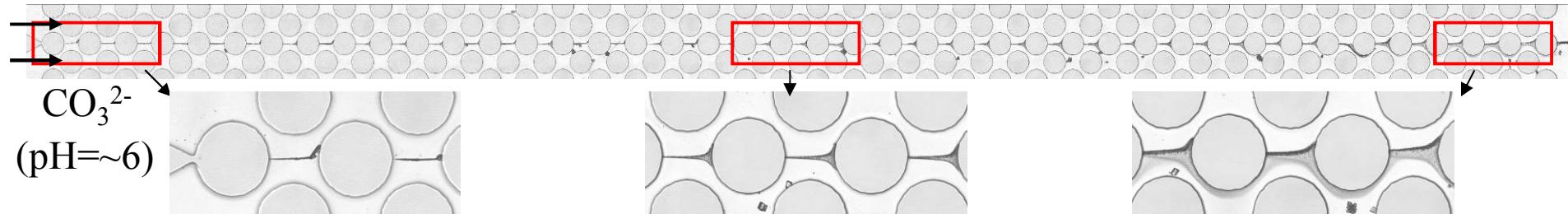
- Lattice Boltzmann Method for water flow
- Direct numerical simulation of CaCO_3 precipitation and dissolution

Precipitate morphology and growth rate

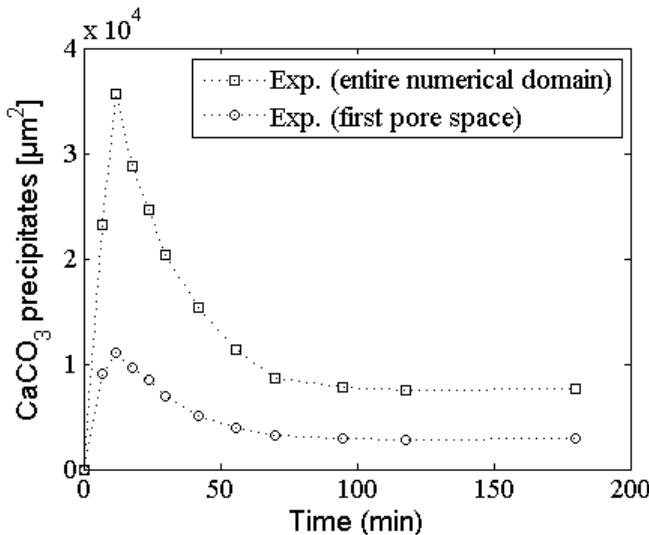
Ca²⁺ (pH=~11)

[Ca²⁺]_T=[CO₃²⁻]_T=25 mM at 2 hrs

Zhang et al., ES&T (2010)

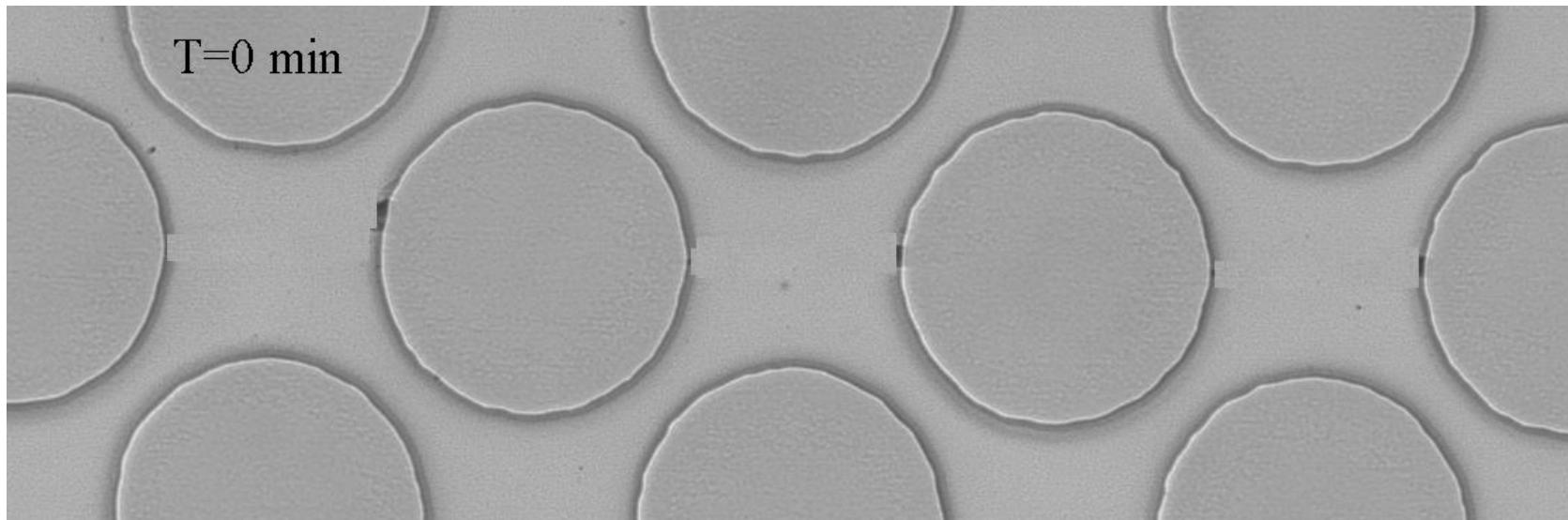


- Precipitation ~ along the centerline within one pore space transverse to the primary flow direction, with some large crystals off the centerline
- Width of the precipitate line ~ increase with distance from the inlet



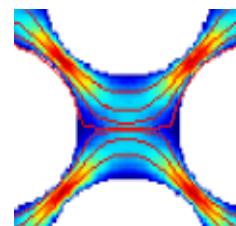
- Precipitate area increases rapidly due to fast precipitation
- and then decreases due to fast dissolution within 13 min
- until a relatively constant but lower plateau is reached over 3-4 hours

Experimental Results



CaCO₃ Precipitation and Dissolution

**Lattice Boltzmann Method:
Velocity field (u) at pore scale**



Finite Volume Method: Reactive transport at pore scale

$\nabla \cdot (\mathbf{u} \Psi_j - D \nabla \Psi_j) = 0$ in liquid phase (Ψ_j : total concentration of primary species j)

$$D \frac{\partial \Psi_j}{\partial \mathbf{n}} = -I_m \quad \text{on reactive surface}$$

$\Psi_j = C_j + \sum_{i=1}^{n_{eq}} \nu_{ji} C_i$ Chemical equilibrium in bulk fluid (e.g., H⁺, HCO₃⁻, ...)
Extend Debye-Hückel Equation for activity coefficients

$$I_m = -k_{cc} (1 - \Omega) = - (k_1 a_{H^+} + k_2 a_{H_2CO_3} + k_3) \left(1 - \frac{Q_{cc}}{K_{sp}} \right)$$

Heterogeneous reaction at mineral surfaces
 Ω = supersaturation index; Q_{cc} = ion activity product; K_{sp} = solubility product

Calcite Precipitation and Dissolution

$$\frac{\partial V_m}{\partial t} = \overline{V_m} a_m K_C \left(\frac{a_{Ca^{2+}} a_{CO_3^{2-}}}{K_{sp}} - 1 \right)$$

Calcite volume fraction (V_m) is updated explicitly over time
 $\overline{V_m}$ is the molar volume of calcite

Key features of modeling

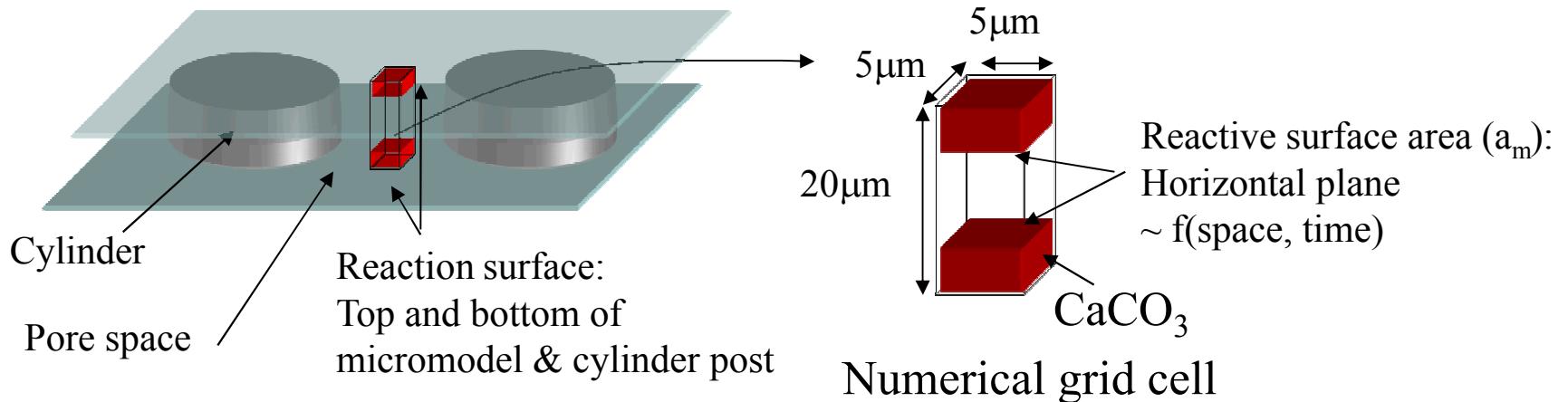
1. Flow field (solved by Lattice Boltzmann method)

- CaCO_3 volume fraction (θ) of a grid cell ($5\mu\text{m} \times 5\mu\text{m} \times 20\mu\text{m}$) is greater than a threshold value (e.g., 0.6), then no flow is allowed through the grid cell
- Diffusion is still allowed until the grid cell is fully occupied by calcite

2. Effective diffusion coefficient = $D_m * \text{tortuosity } (\tau)$

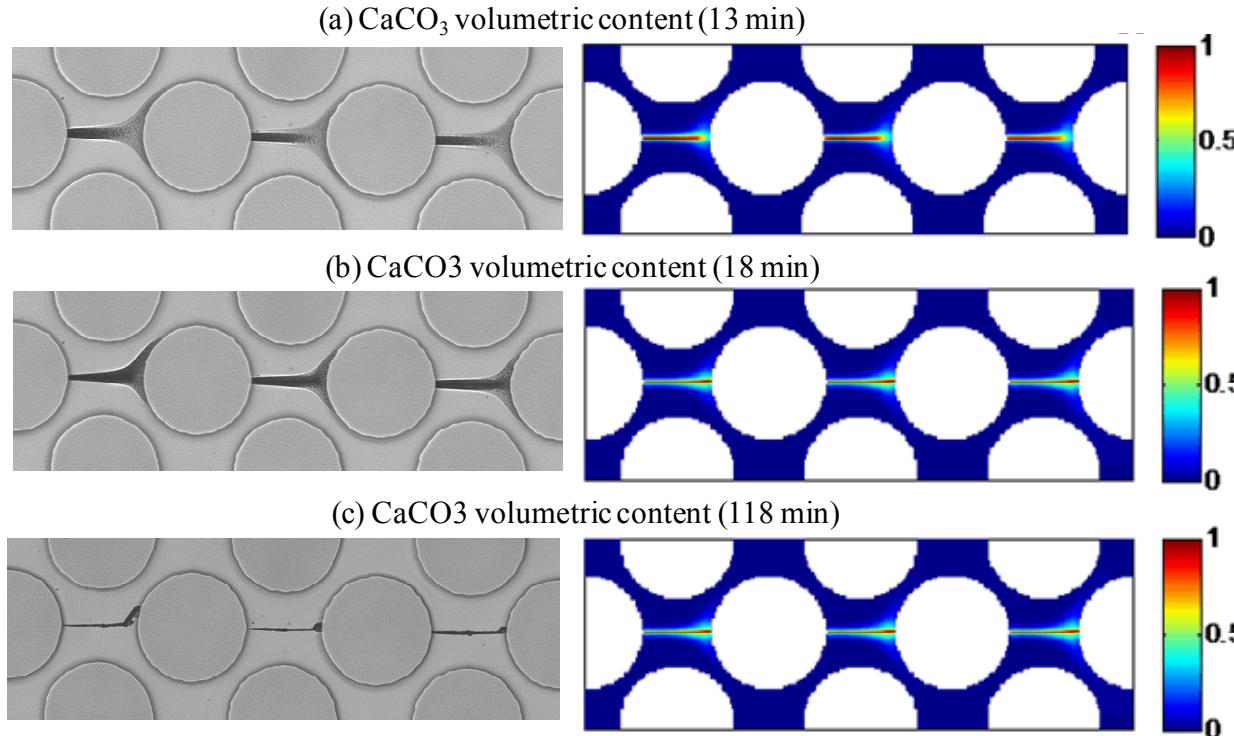
- $\tau(\theta) = (1-\theta)^n$ where $n \sim 0$ to 3

3. Quasi 3D grid cell for reactive surface



4. Effects of n , k_{cc} , and dissolution factor on precipitation and dissolution rate

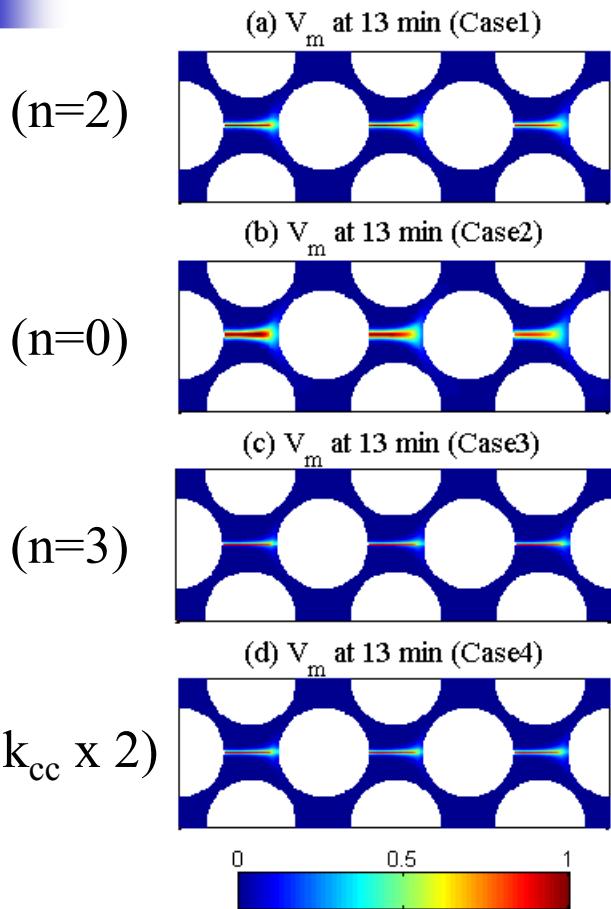
Model results: Reference case (25 mM)



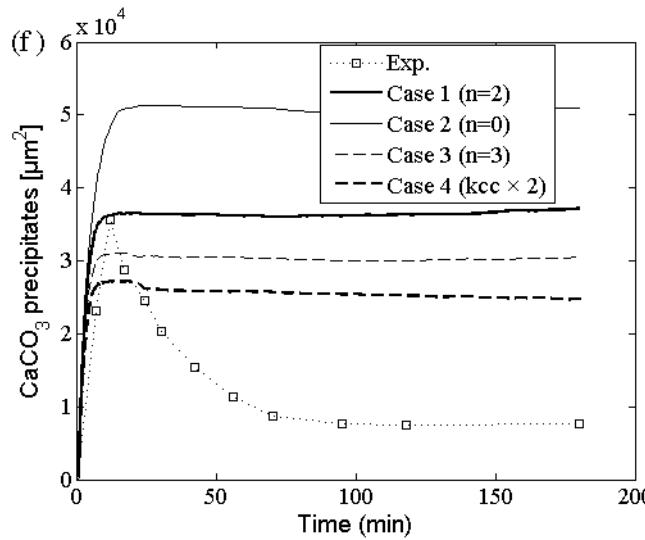
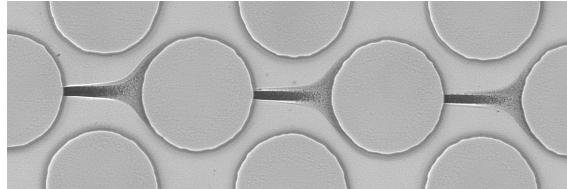
Reference case with literature values ($n=2$, k_{CC} from Chou et al. (1989)):

- Initial precipitation is modeled well
- Simulation was not able to capture the dissolution process after ~ 13 min
- Experimental results show CaCO₃ dissolution may be pH-dependent

Sensitivity: D_{eff} & k_{cc}



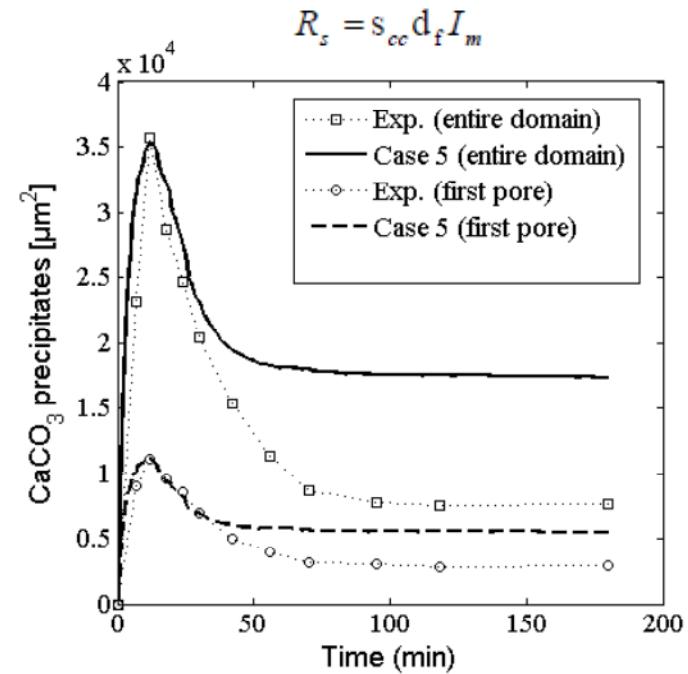
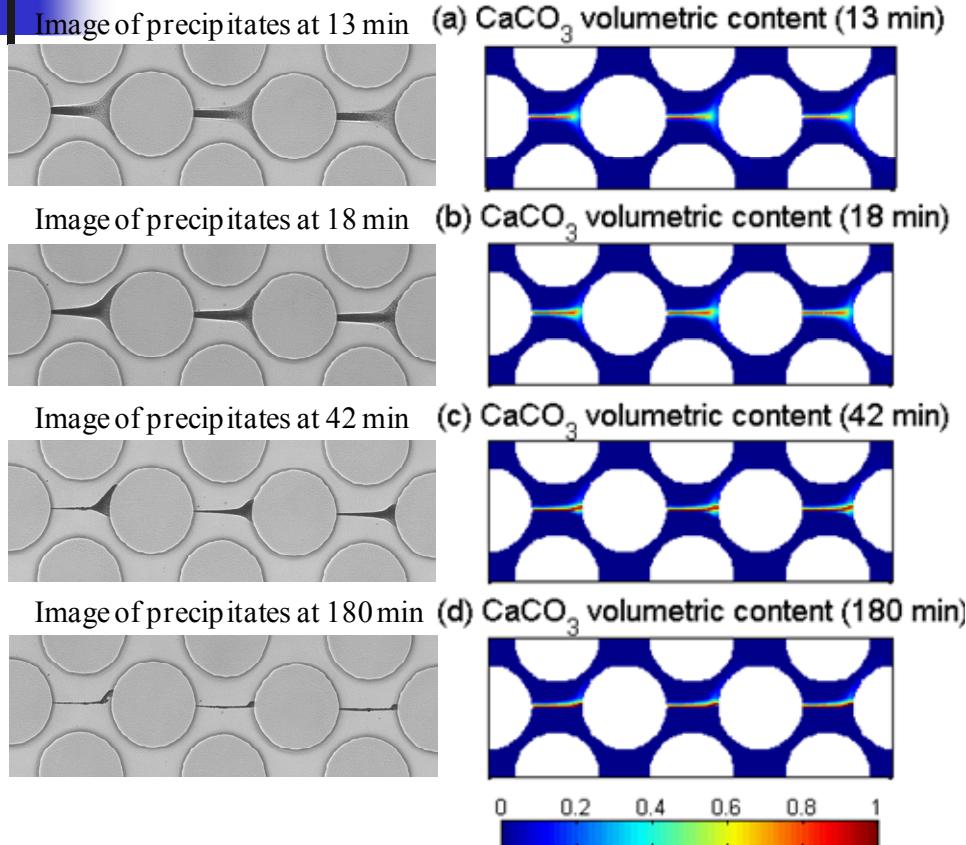
(e) Image of precipitates at 13 min



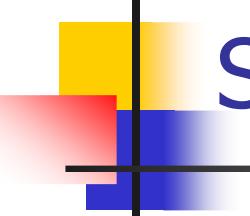
$$D_{\text{eff}} = D_m * (1-\theta)^n$$

- As the n value decreases (increases), D_{eff} decreases less (more) as V_m increases, resulting in more (less) precipitation compared to the experimental results
- As k_{cc} increases, precipitation occurs faster, particularly along the centerline, resulting in a reduction of diffusion (i.e., mixing) -> decreasing the maximum precipitate area

Dissolution factor (factor=300)

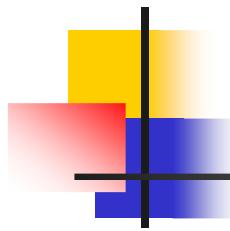


- Thickness and area of precipitate along the centerline after 13 min decreases more quickly, and the precipitate area matches the experimental data until 30 minutes
- Model predicts dissolution below the centerline well, but not above the centerline
- The need to include a higher dissolution factor may be attributed to an increase in the precipitate surface area and the recrystallization process during dissolution

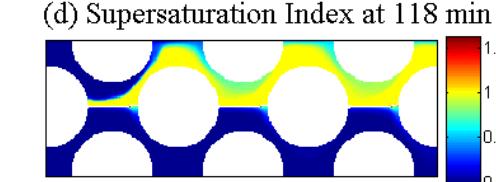
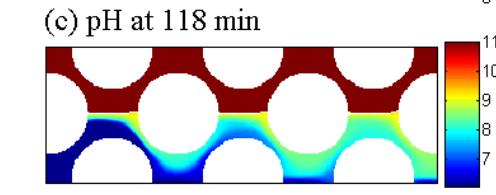
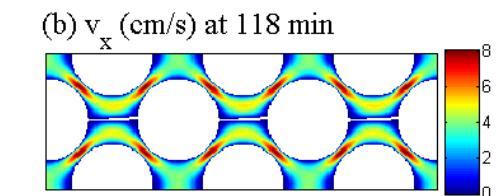
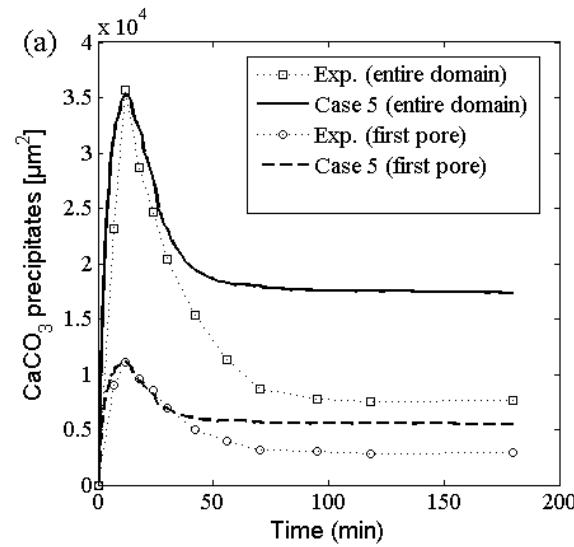
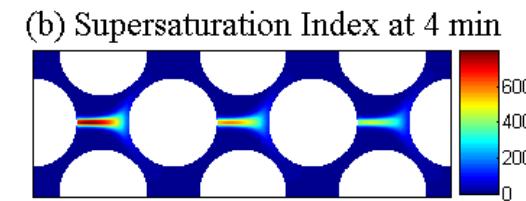
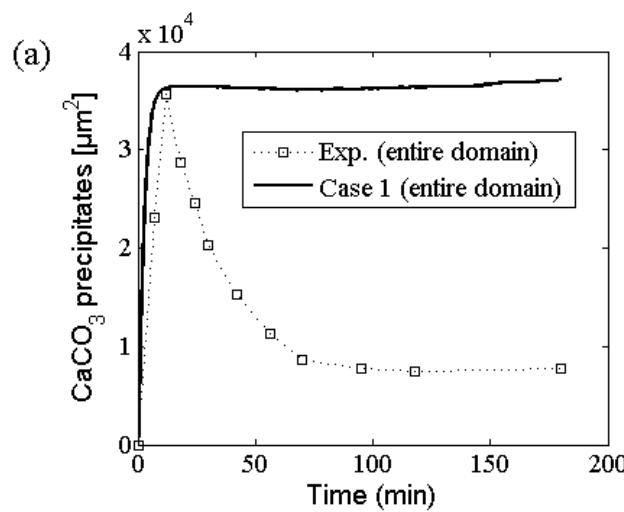


Summary and Challenges

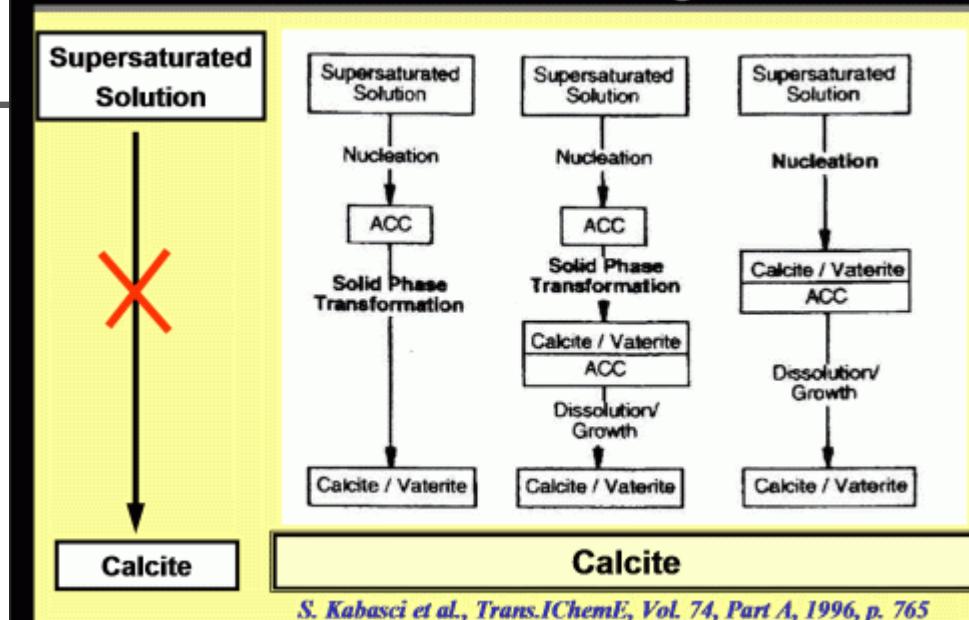
- Pore Scale LB-FVM qualitatively captures governing physics in transverse-mixing induced CaCO_3 formation
 - CaCO_3 formation and precipitate patterns
 - Pore blocking due to precipitation
- The effects of geochemical reactions and flow field change are coupled properly
- There is a need to account for the enhanced dissolution, possibly linking to reactive surface area at sub-micro scale and recrystallization processes
- Pore-scale modeling and experimental results will be used to test the validity of various upscaling (pore to continuum) and multi-scale (hybrid) methods, and to develop a new method of obtaining effective dispersion coefficient values and reactive surface area



Questions?



Ostwald's Rule of Stages



Modifications of Calcium Carbonate

CALCITE	$\rho = 2.8 \text{ g/cm}^3$	crystalline	thermod. stable
ARAGONITE	$\rho = 3.0 \text{ g/cm}^3$	crystalline	metastable
VATERITE	$\rho = 2.7 \text{ g/cm}^3$	crystalline	metastable
MONOHYDRATE	$\rho = \text{ n. a.}$	crystalline	metastable
HEXAHYDRATE	$\rho = 1.8 \text{ g/cm}^3$	crystalline	metastable
COLLOIDAL	$\rho = \text{ n. a.}$	amorphous	metastable