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Title: Probing processes at the base of the Antarctic Ice Sheet

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# Probing processes at the base of the Antarctic Ice Sheet

Carolyn Branecky Begeman, Los Alamos National Laboratory

**Los Alamos**  
NATIONAL LABORATORY  
EST. 1943

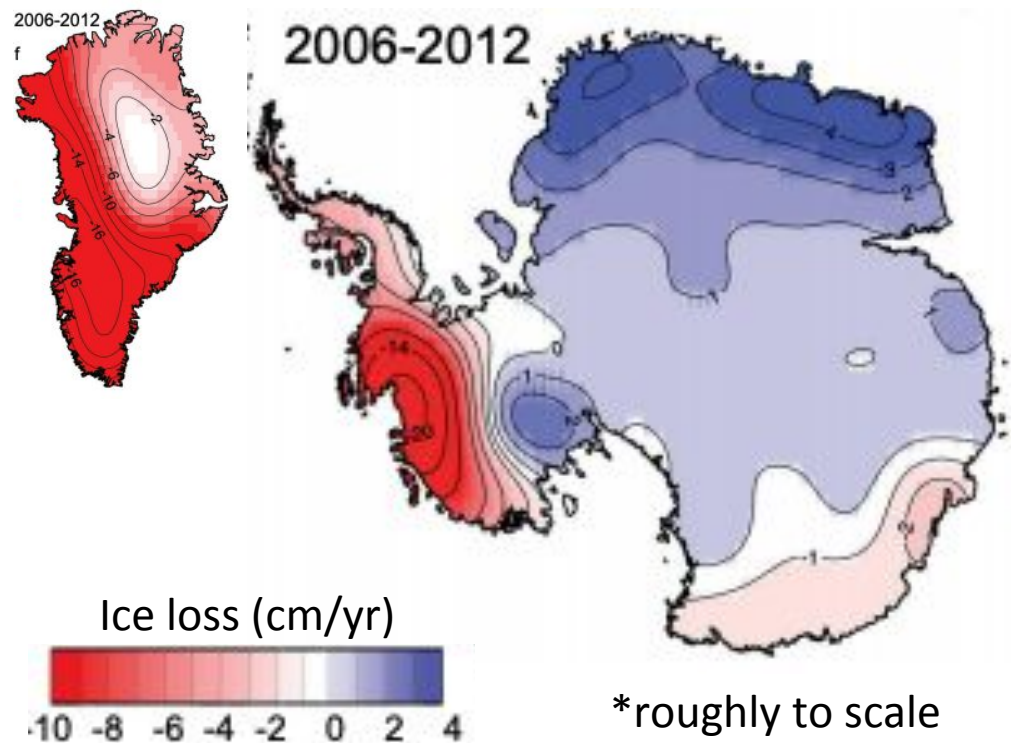
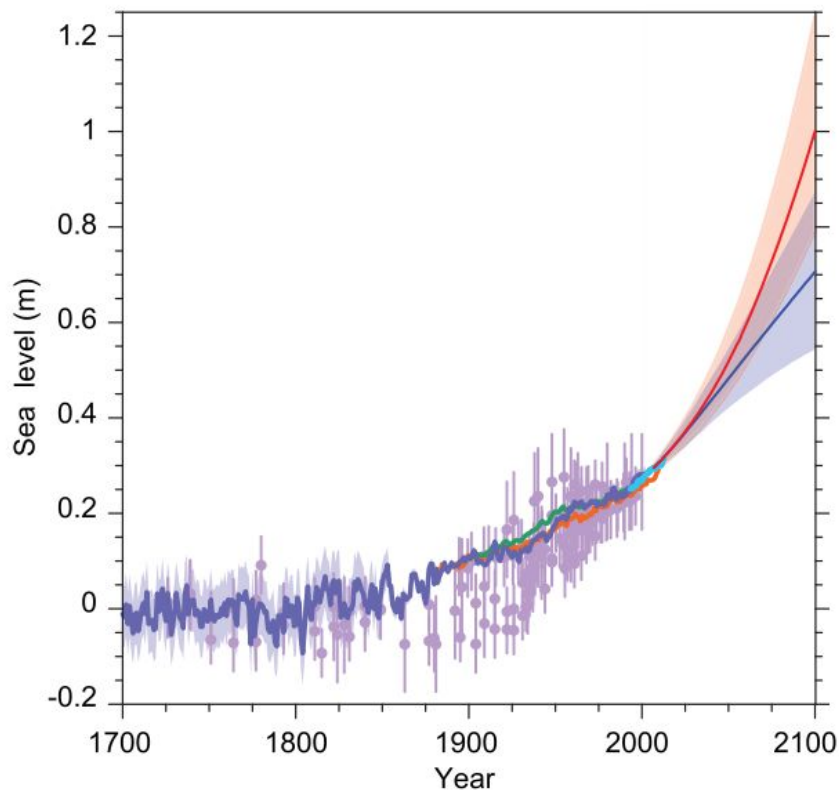


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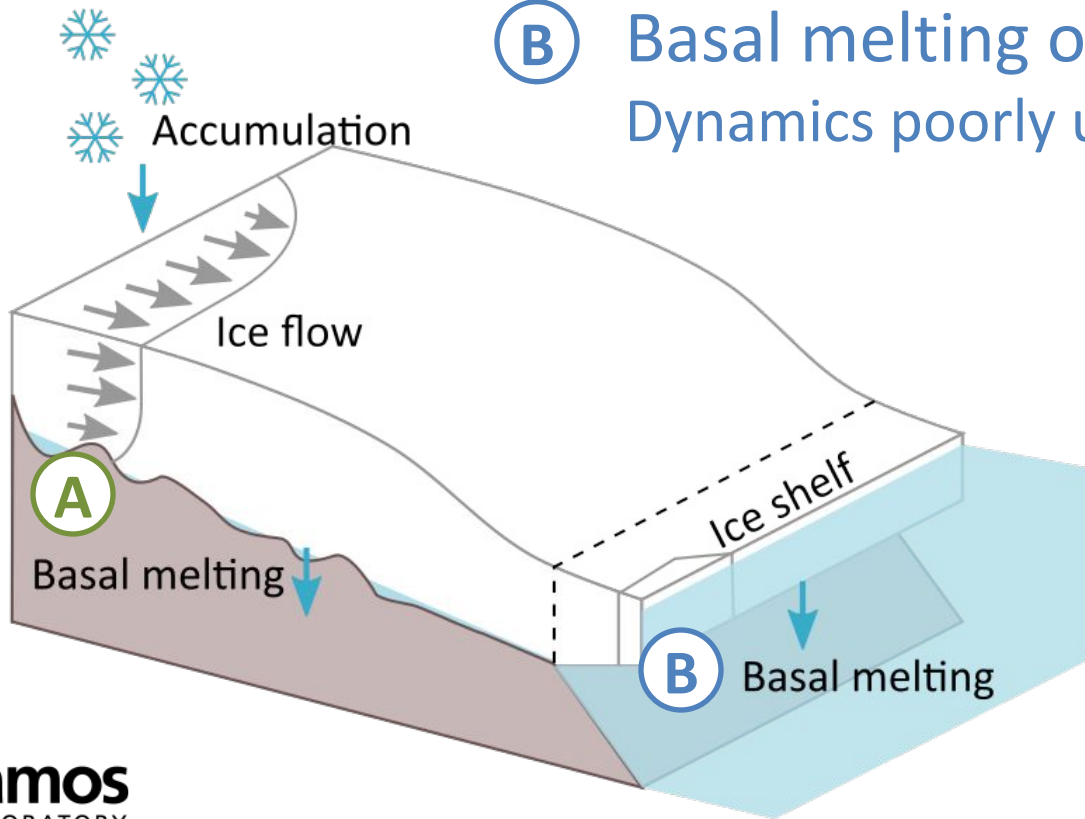
# How much will sea levels rise in the 21<sup>st</sup> century? due to ice sheet mass loss



IPCC AR5, reprinted from Velicogna, 2009

# Predicting ice mass loss with an ice sheet model

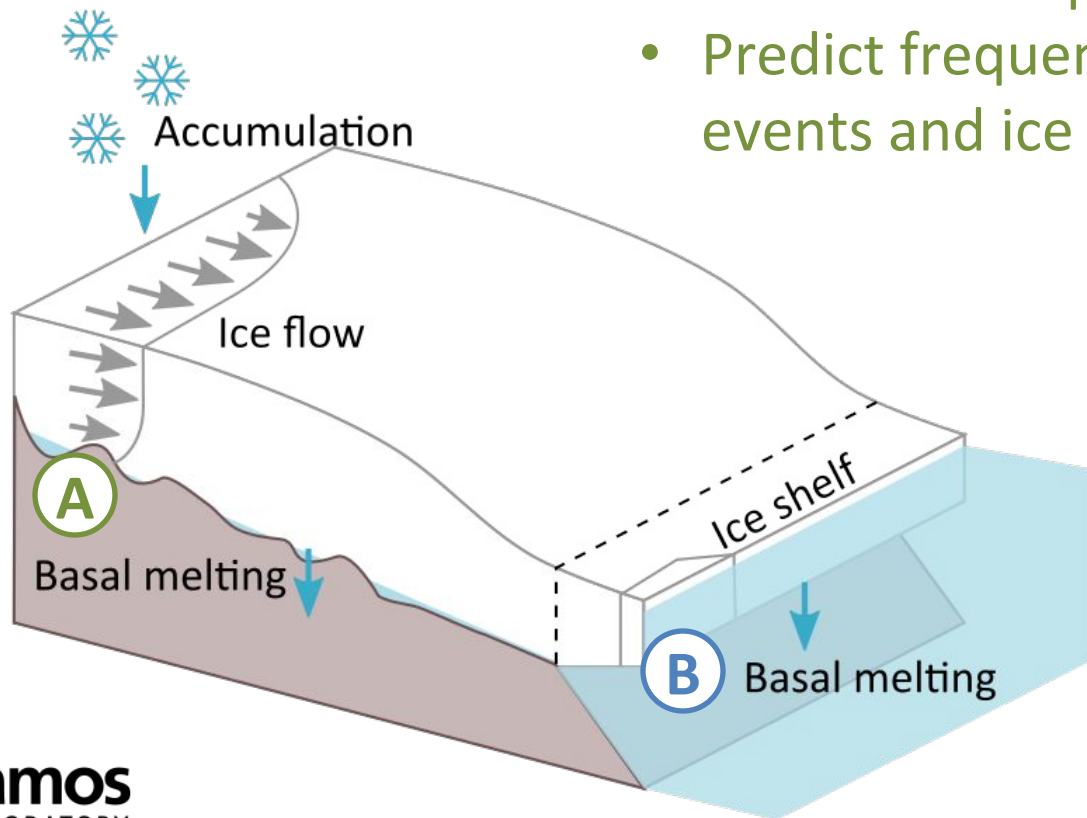
- ① Basal melting of ice on land:  
Dynamics understood but geologic information lacking
- ② Basal melting of ice over ocean:  
Dynamics poorly understood



# My work toward improving predictions of ice mass loss

## ① Basal melting of ice on land (PhD)

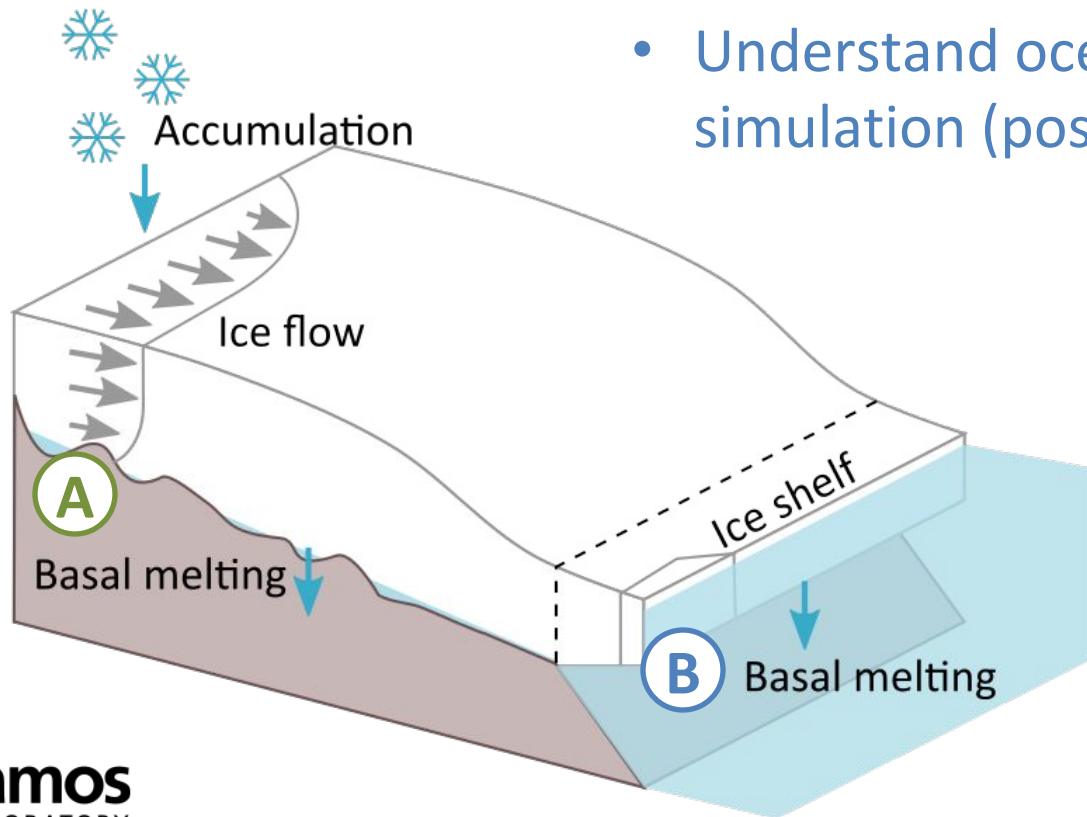
- **Constrain with observations of geothermal flux**
- Characterize spatial variability of geologic processes
- Predict frequency relationships between magmatic events and ice loss



# My work toward improving predictions of ice mass loss

## ② Basal melting of ice over ocean

- Constrain with observations of melt rate and local ocean conditions (PhD)
- Understand ocean dynamics through large-eddy simulation (postdoc)

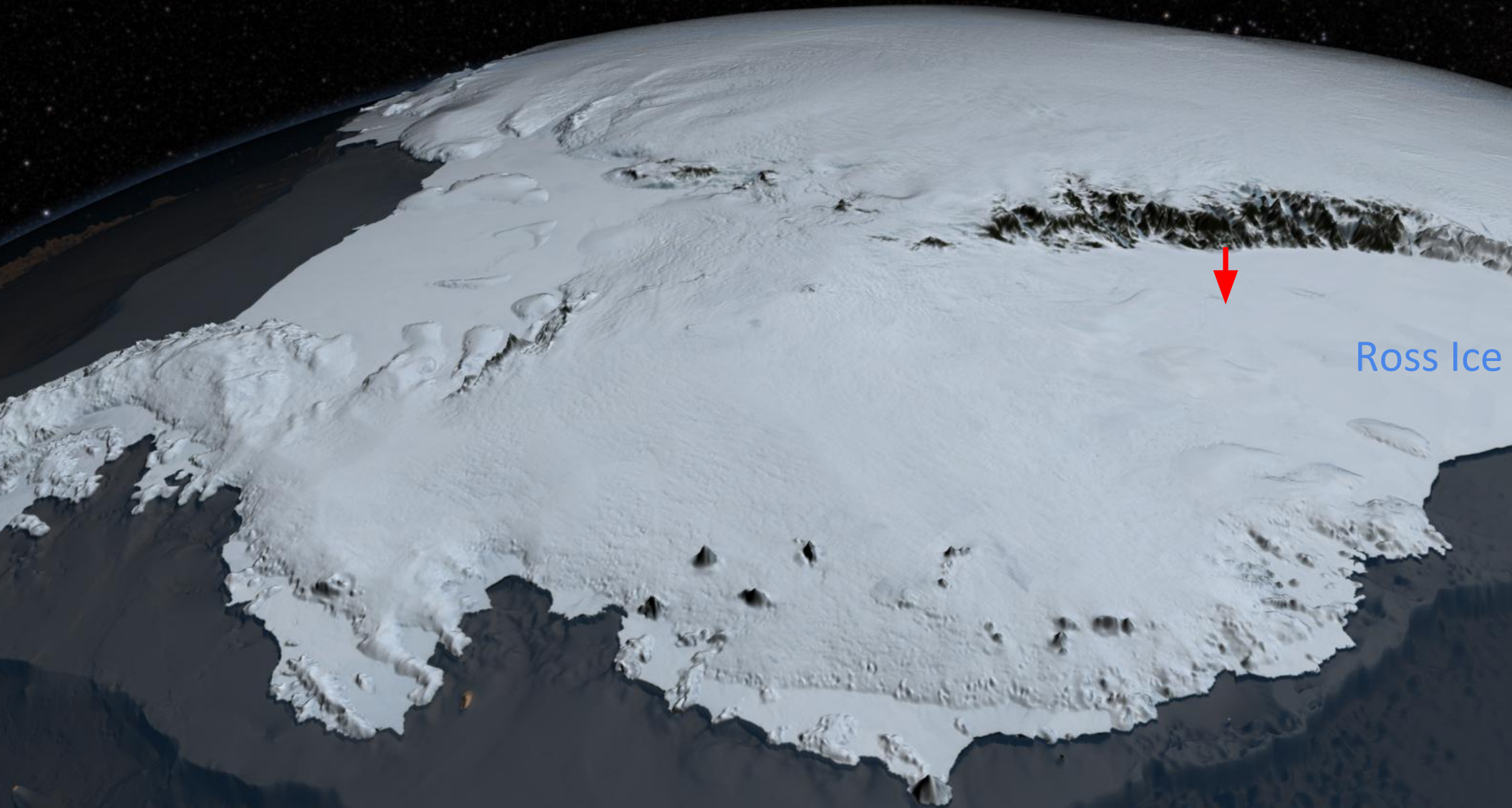


Fully-coupled interactions between ice sheet and ocean model



Quantify uncertainty in ice sheet mass balance from different sources

C. Begeman SSI Seminar Nov. 15 2019

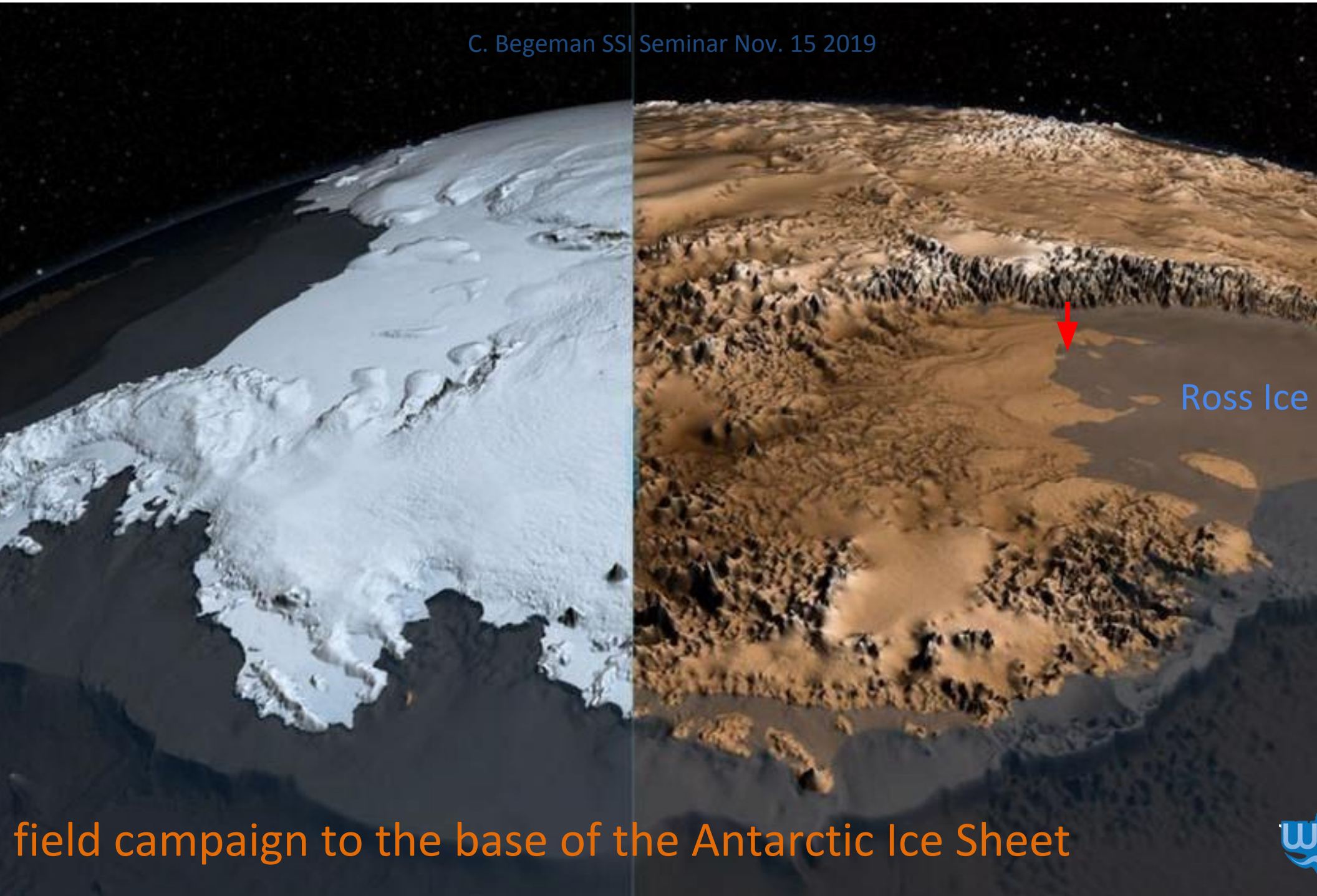


Ross Ice

field campaign to the base of the Antarctic Ice Sheet



C. Begeman SSI Seminar Nov. 15 2019



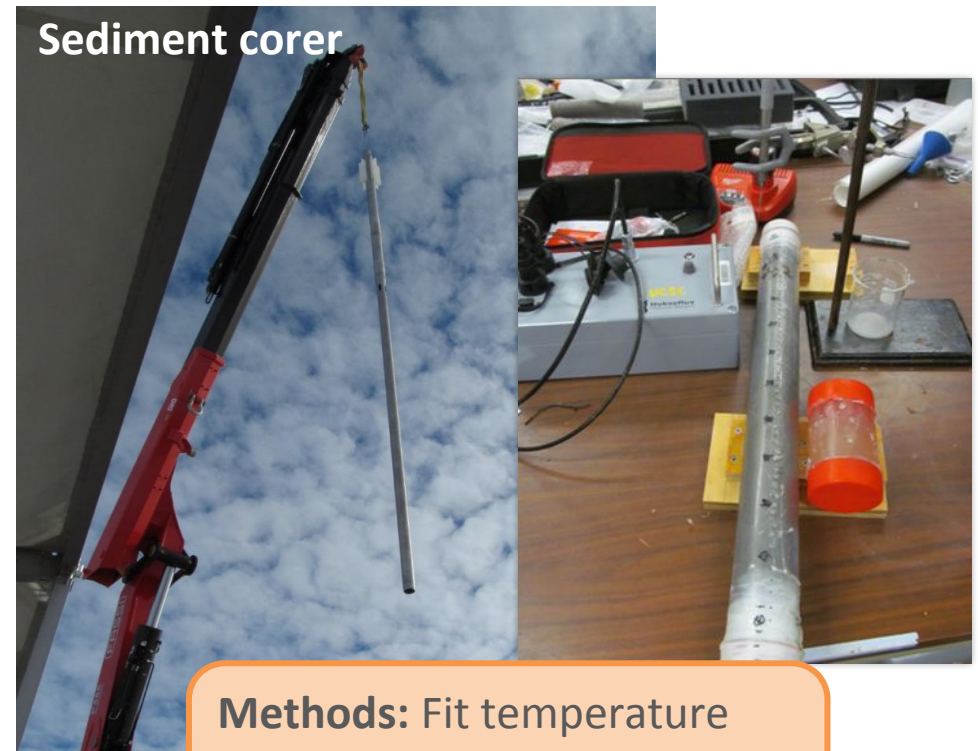
field campaign to the base of the Antarctic Ice Sheet





# Measuring geothermal heat flux

**G** = Temperature gradient x Thermal conductivity =  $88 \pm 7 \text{ mW m}^{-2}$   
in sediments of sediments  
n=2 n=40  
average continental G = 6



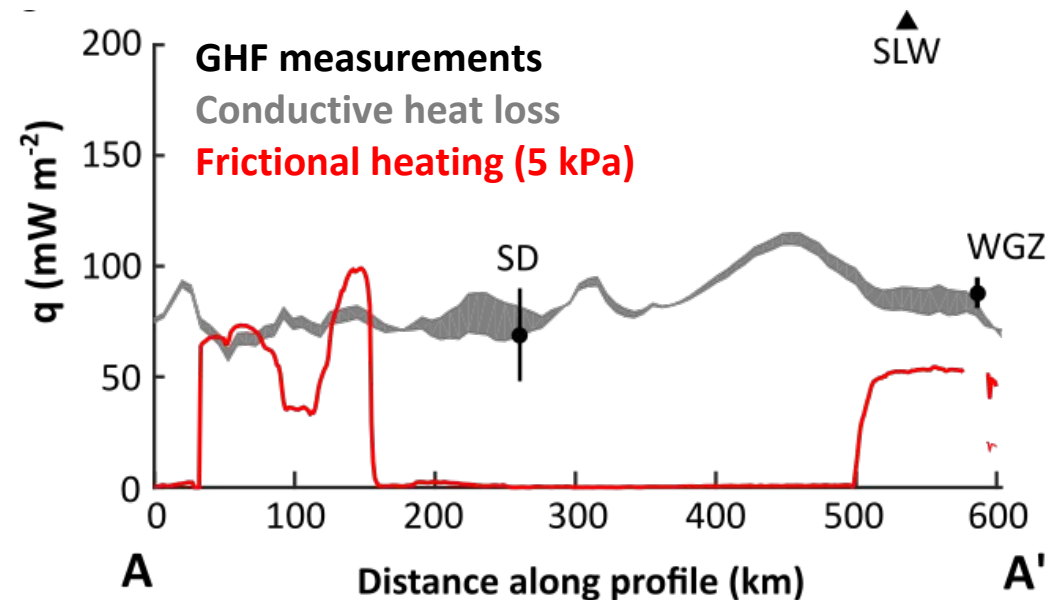
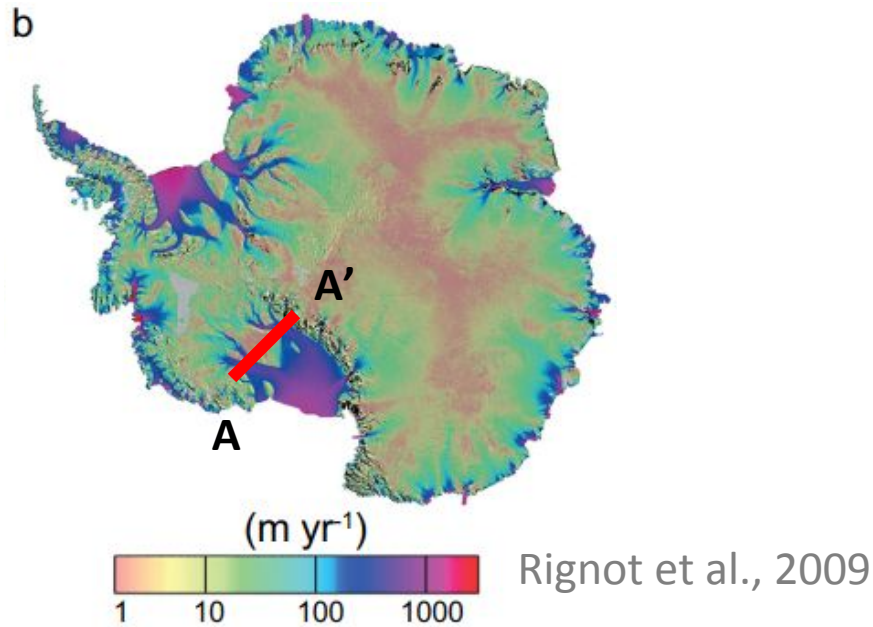
**Methods:** Fit temperature evolution after thermal pulse to an analytical model

# Is geothermal heat important to the basal heat budget?

$$\rho_i L m = k_s \frac{dT_s}{dz} + \tau_s u_i - k_i \frac{dT_i}{dz}$$

Basal melting      geothermal heat flux      frictional heat flux      conductive heat loss

Ice velocity

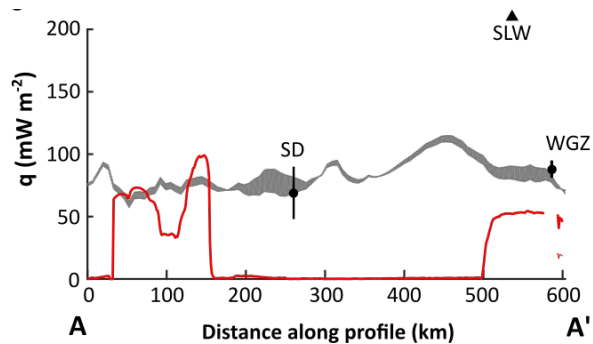
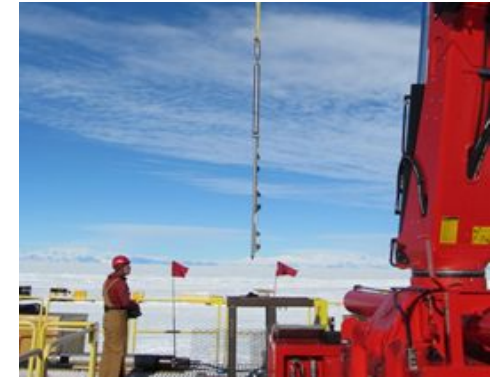


## Methods:

Analytical model for thermal advection and conduction in ice and Earth's crust

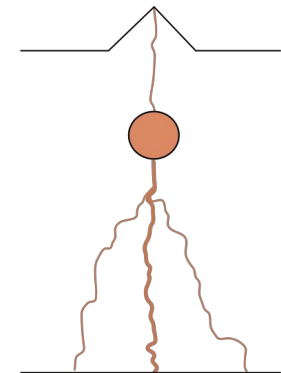
## ① Progress on geothermally-driven ice melting

I measured geothermal heat flux at a new site, which will help constrain future continent-wide maps

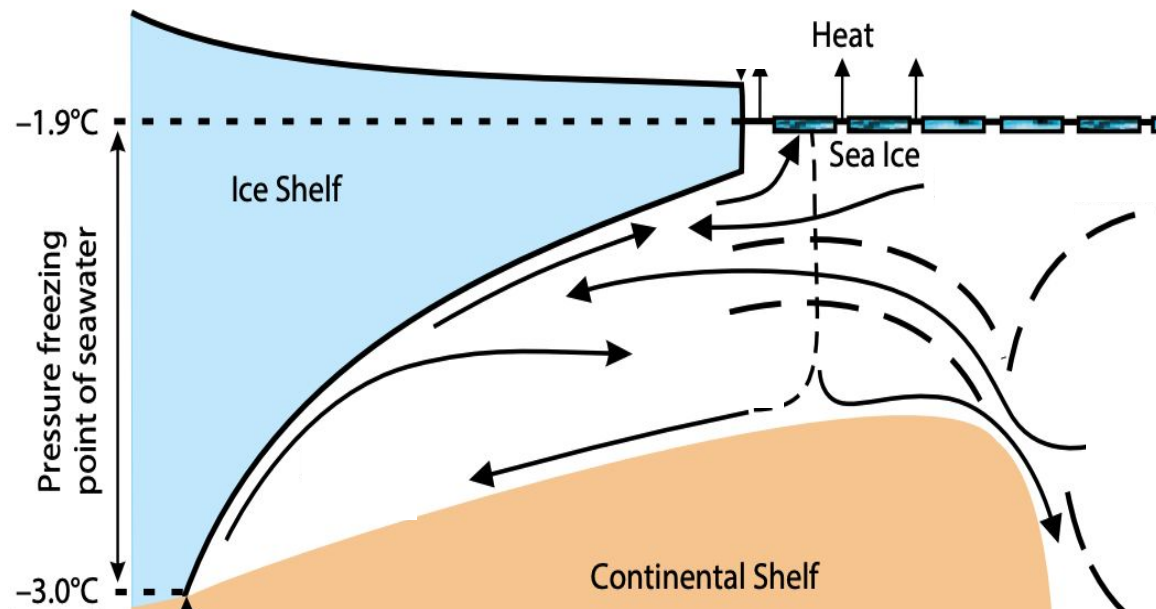
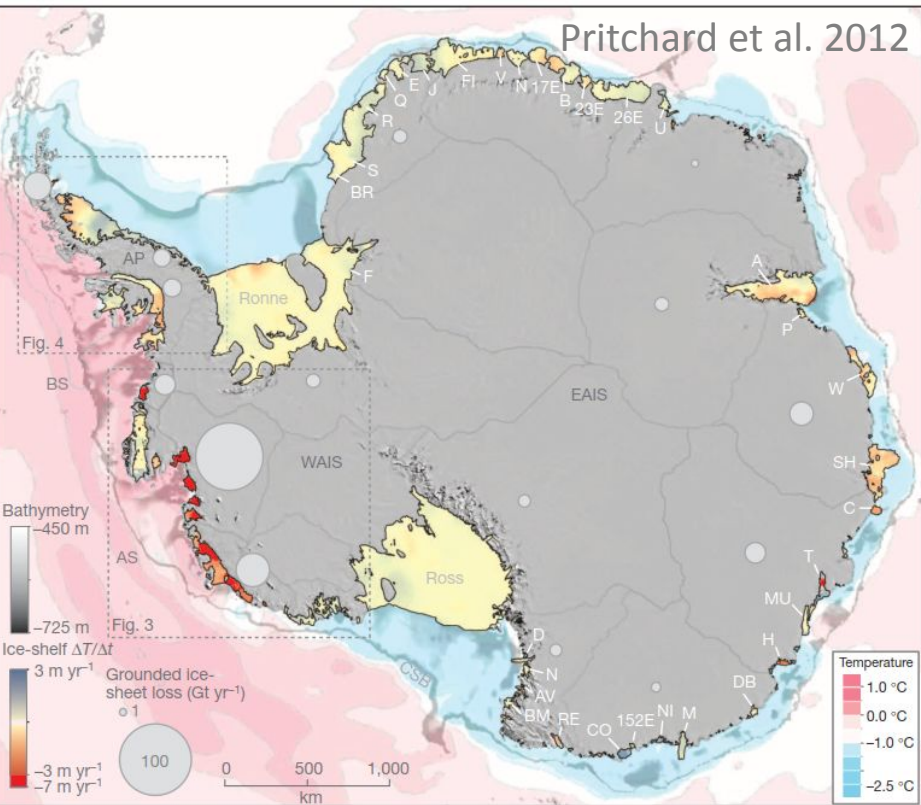


I found that geothermal heat flux is glaciologically important at this site

Magmatism can enhance ice loss in West Antarctica

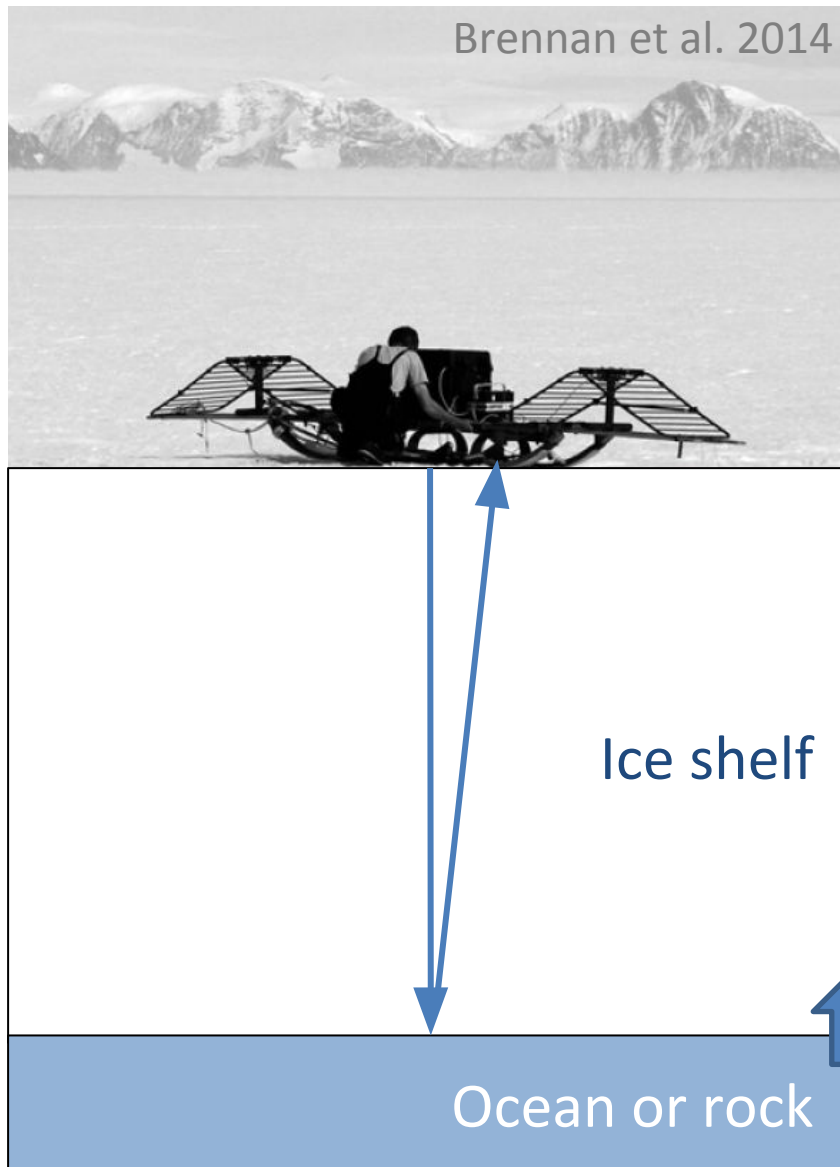


# B) What is the relationship between melting and ocean conditions



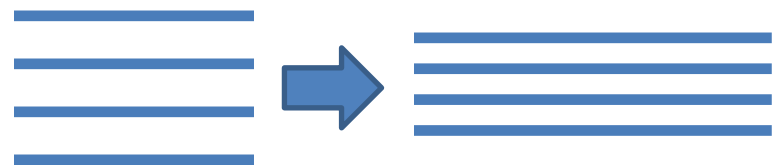
Dinniman et al. (2016)

# The cutting edge in ice melt measurements



Phase-sensitive FMCW radar  
305 MHz center frequency,  
160 MHz bandwidth

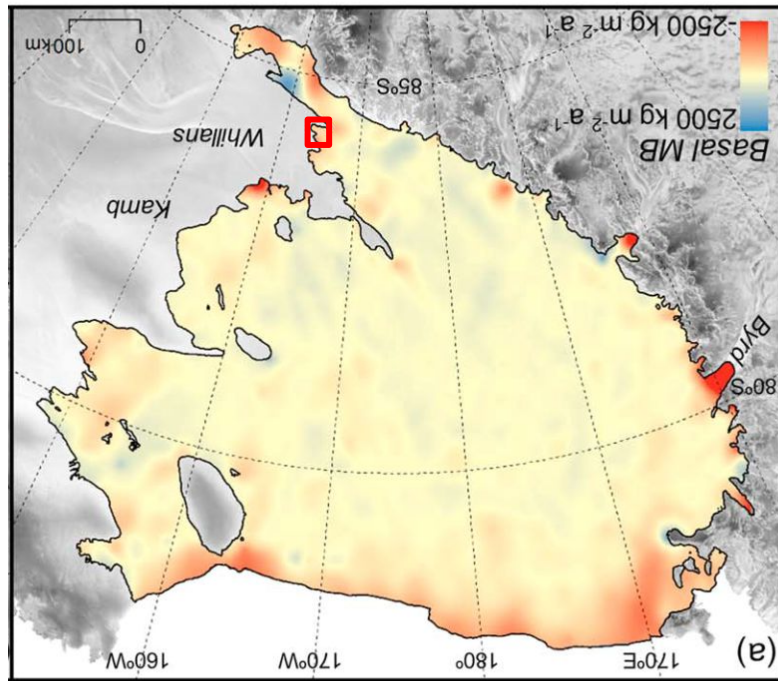
Can separate  
ice-shelf thinning by stretching



from  
ice-shelf thinning by melting

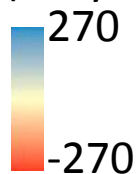
# Spatial and temporal melt variability remains largely unexplained

Ice-shelf average:  $25 \text{ cm yr}^{-1}$

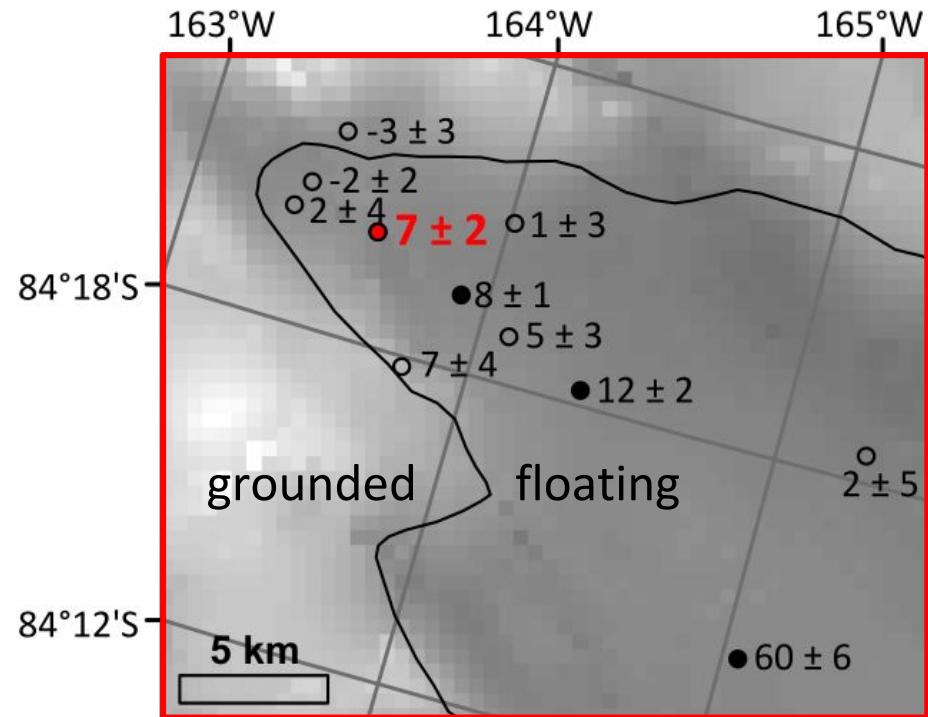


InSAR-derived melt rates  
Moholdt et al. 2014

Melt rate  
( $\text{cm yr}^{-1}$ )



Local:  $< 10 \text{ cm yr}^{-1}$

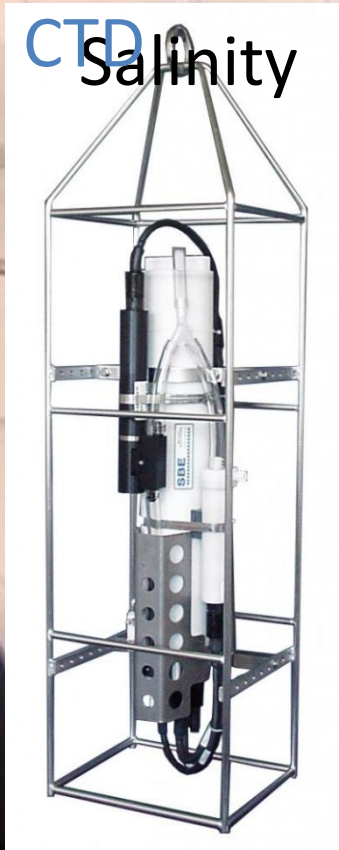


Phase-sensitive radar melt rates  
Begeman et al. 2018 GRL

# Oceanographic observations

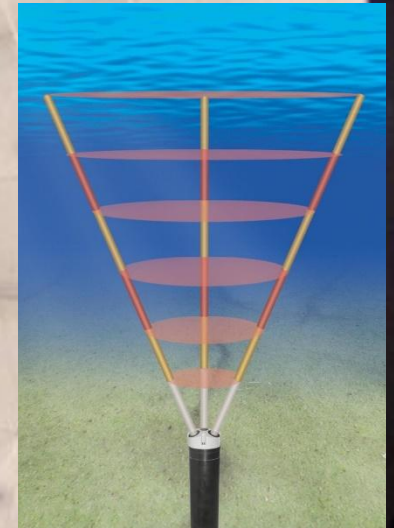
Temperature

CTD  
Salinity



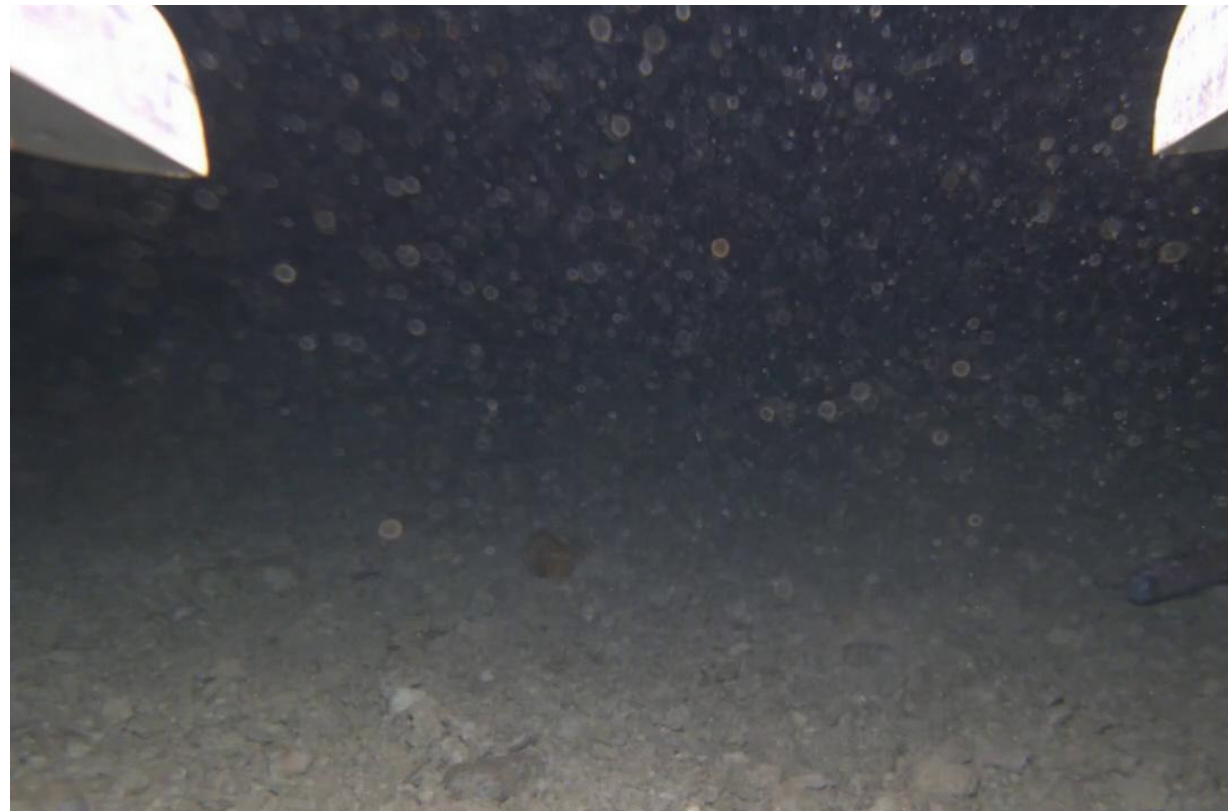
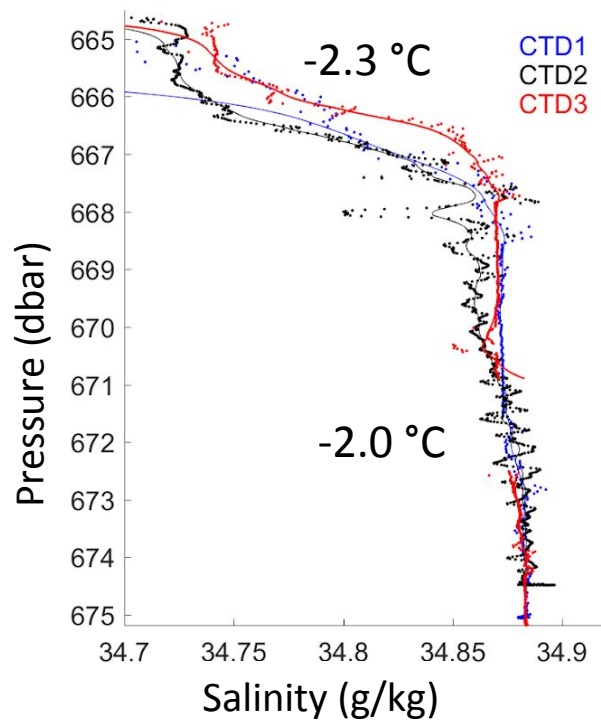
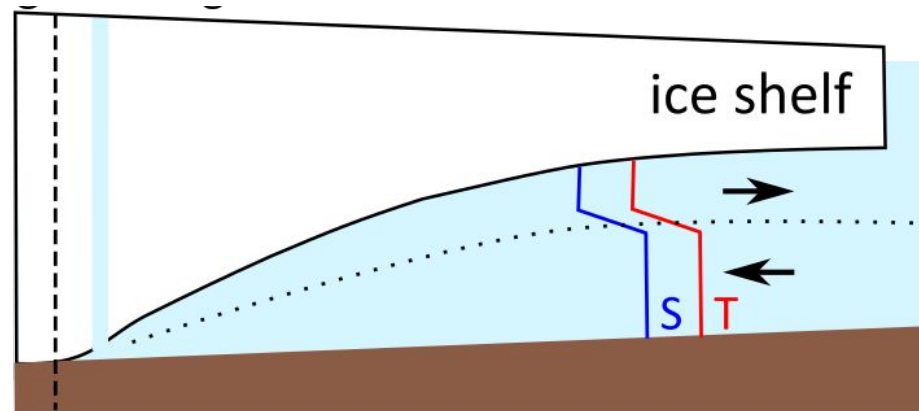
Jill Mikucki

Current  
velocity



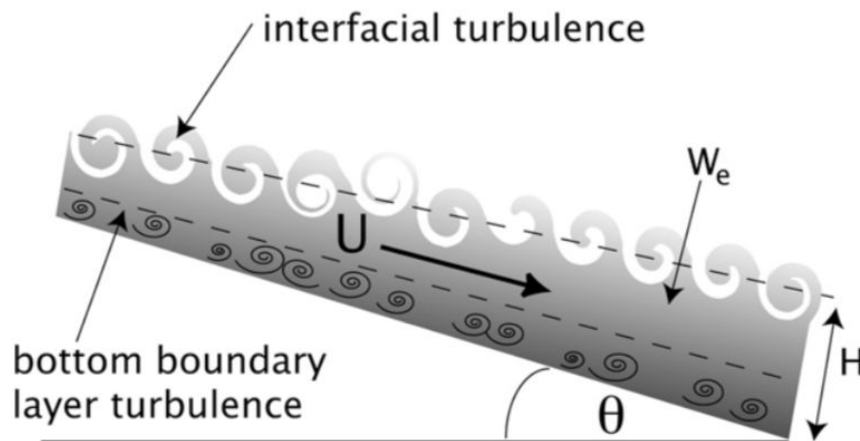
Tim Hodson

Current melt  
parameterization fails  
to accurately predict  
melt rates given  
ocean conditions

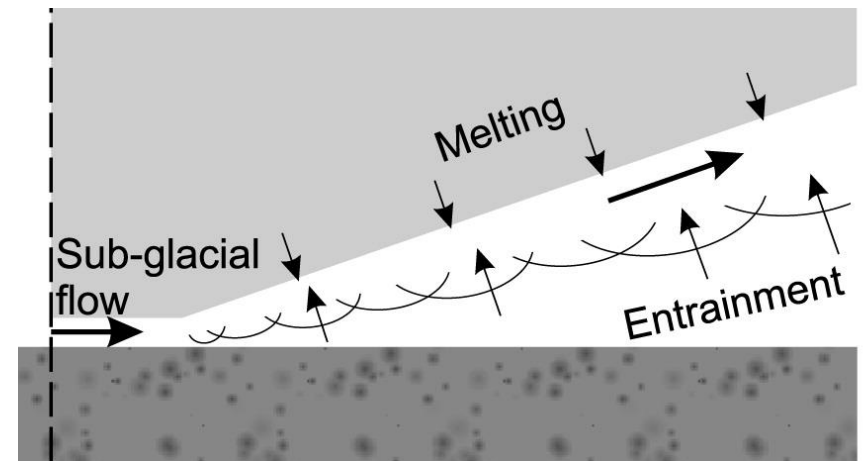


# Limited understanding of vertical heat and salt transport across the ice-shelf ocean boundary layer

“plume theory” postulates that the positively buoyant ice-shelf ocean boundary layer is similar to the negatively buoyant seafloor boundary layer



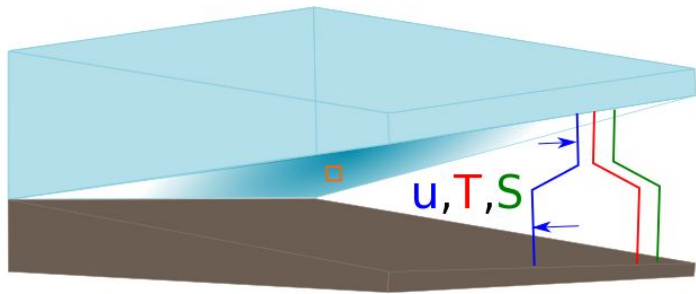
Wells et al. (2010) JPO



Jenkins (2011) JPO

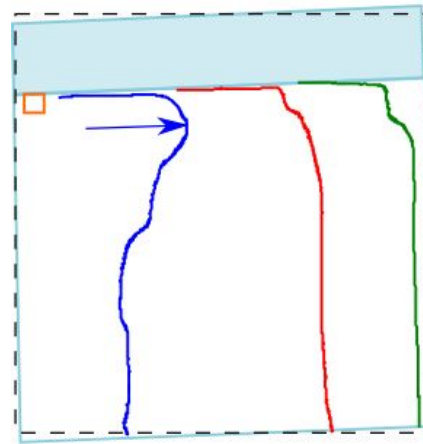
# A hierarchy of ocean models

General circulation model



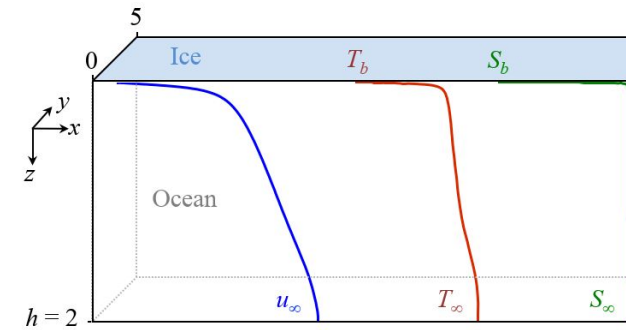
Grid  $\Delta z$ :  $>10$  m  
Eddy viscosity: fully parameterized

Large-eddy simulation



0.25 m  
sub-grid parameterization

Direct numerical simulation



0.05 m  
no parameterization

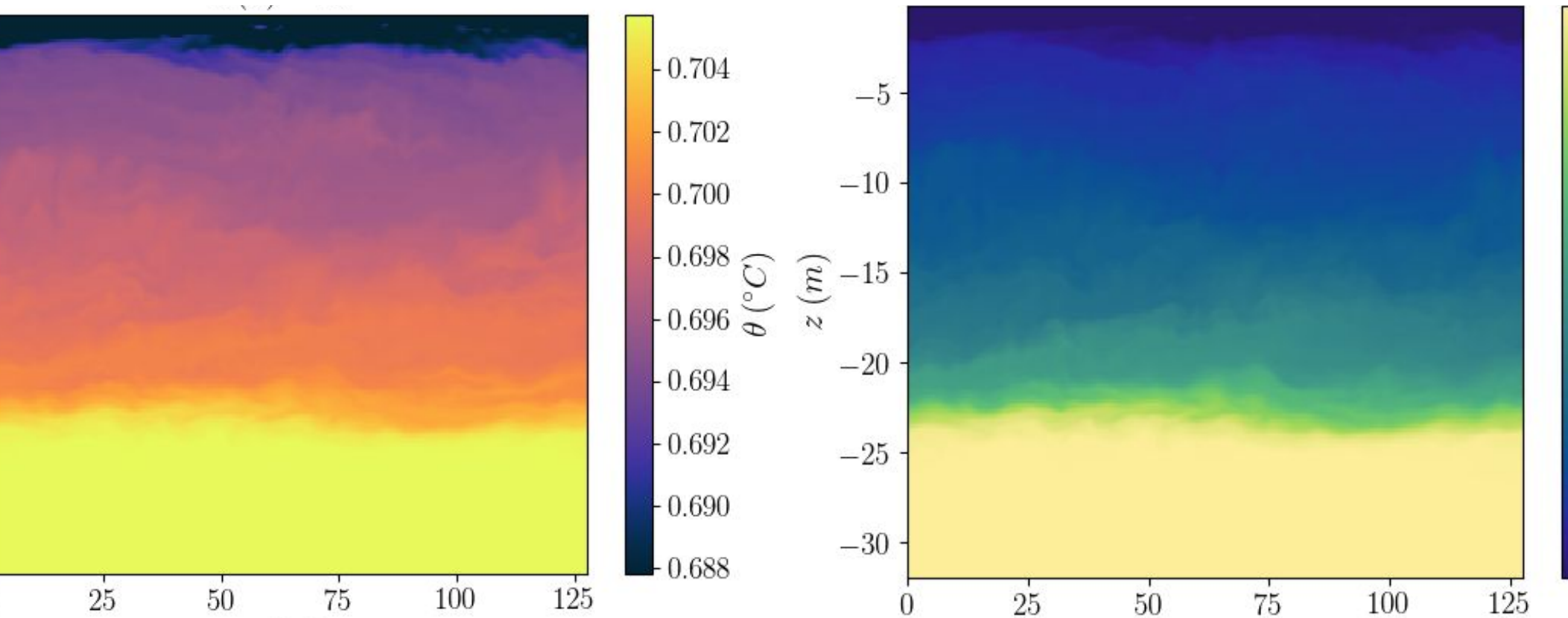
# Modeling the ice-shelf ocean boundary layer

## Parallelized Large-eddy Model, Leibniz Universitat Hanover

- Large-eddy models resolve the energy-containing scales of the flow and parameterize scales that are not resolved
- **Sloping surface at the top of the domain**
  - Gravity and coriolis are rotated
  - Drag formulation that accounts for stratification
- **Dynamically-melting ice**
  - Heat and salt fluxes based on sea ice observations and DNS
  - Nonlinear freezing point
- **Tidal forcing**
  - time-varying large-scale horizontal pressure gradients

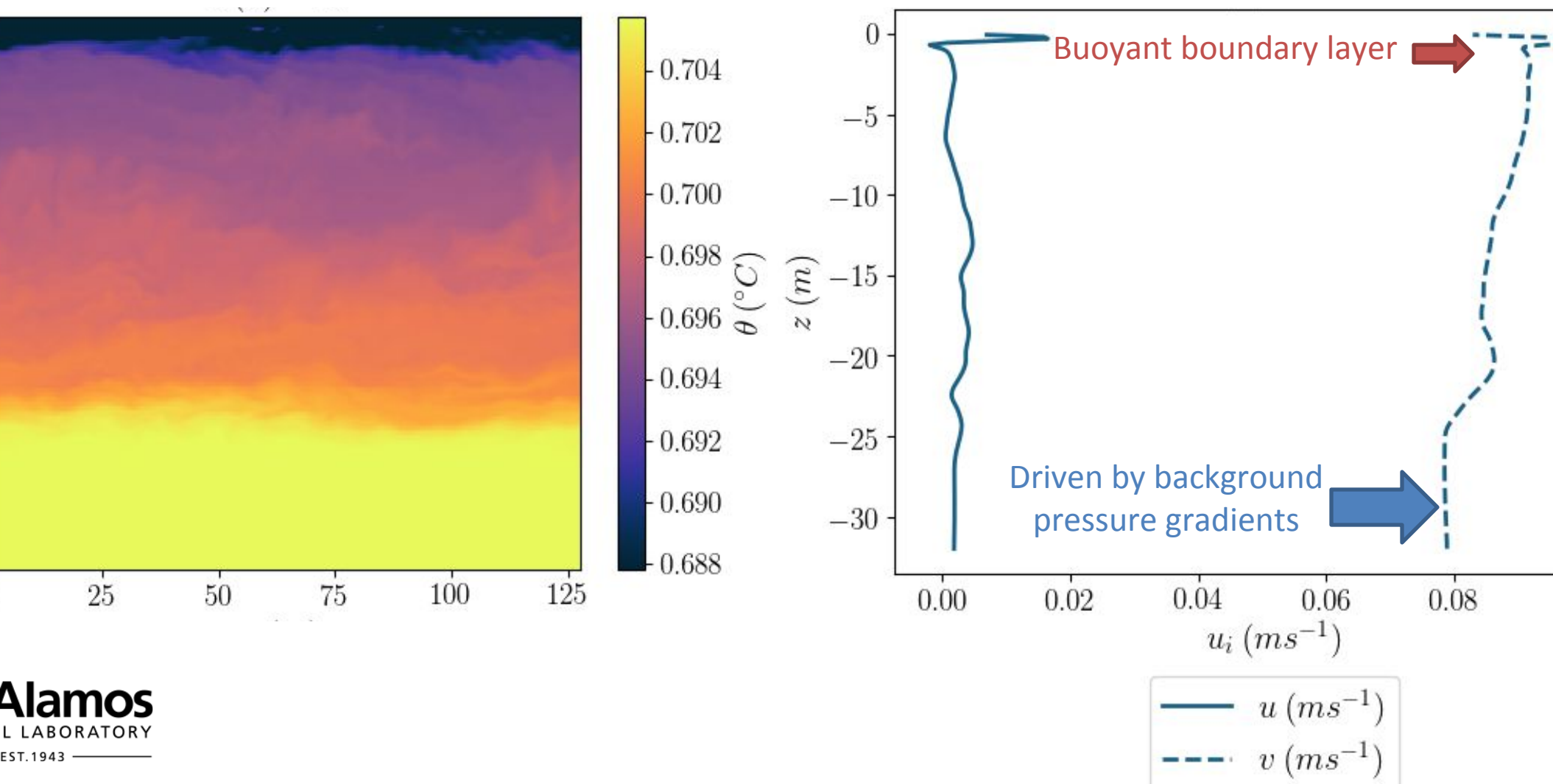


# Ice-shelf ocean boundary layer development in PALM



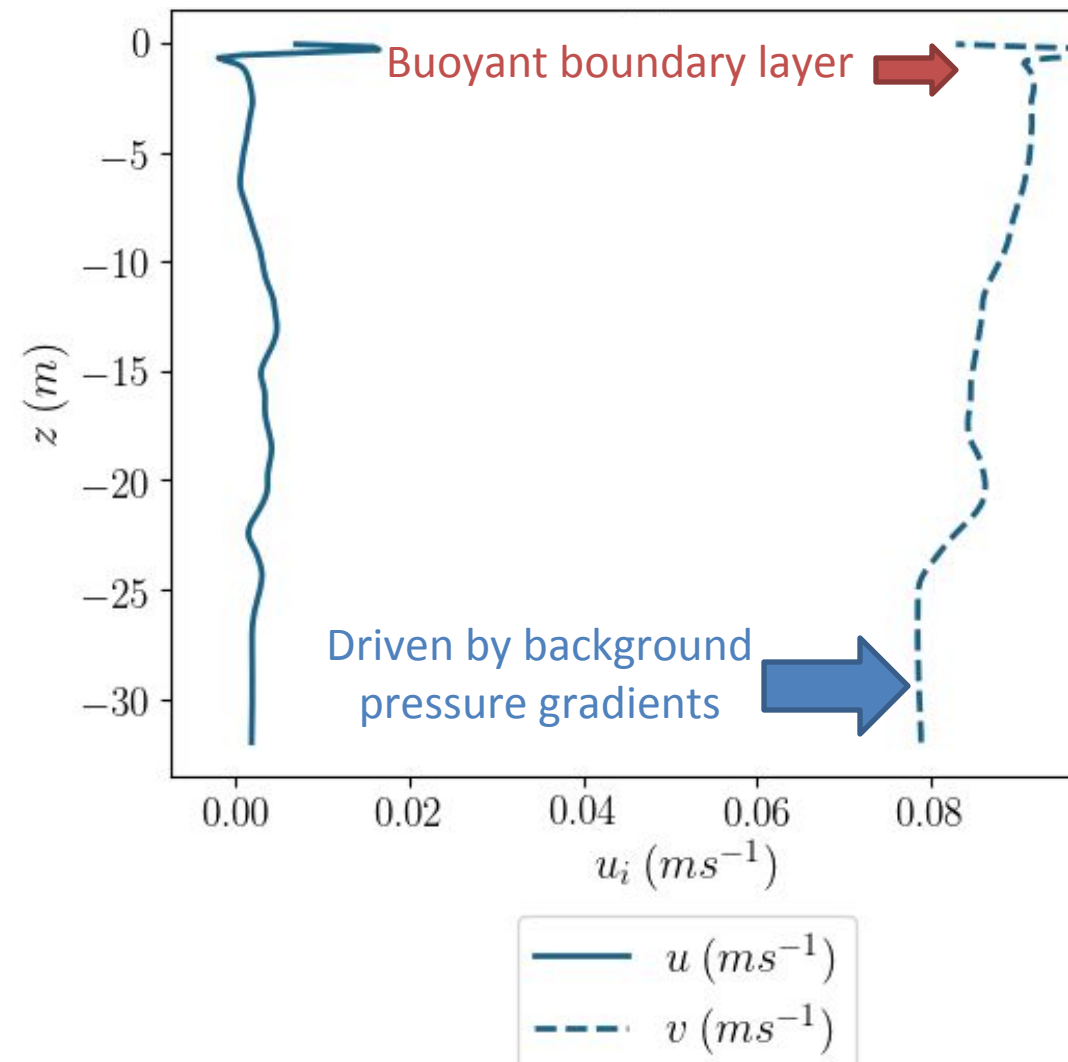
# Ice-shelf ocean boundary layer development in PALM

1.5 m yr<sup>-1</sup> melt rate



# Ice-shelf ocean boundary layer development in PALM

Momentum transport across the boundary layer is heavily influenced by stable stratification



# Ice-shelf ocean boundary layer development in PALM

Momentum transport across the boundary layer is heavily influenced by stable stratification

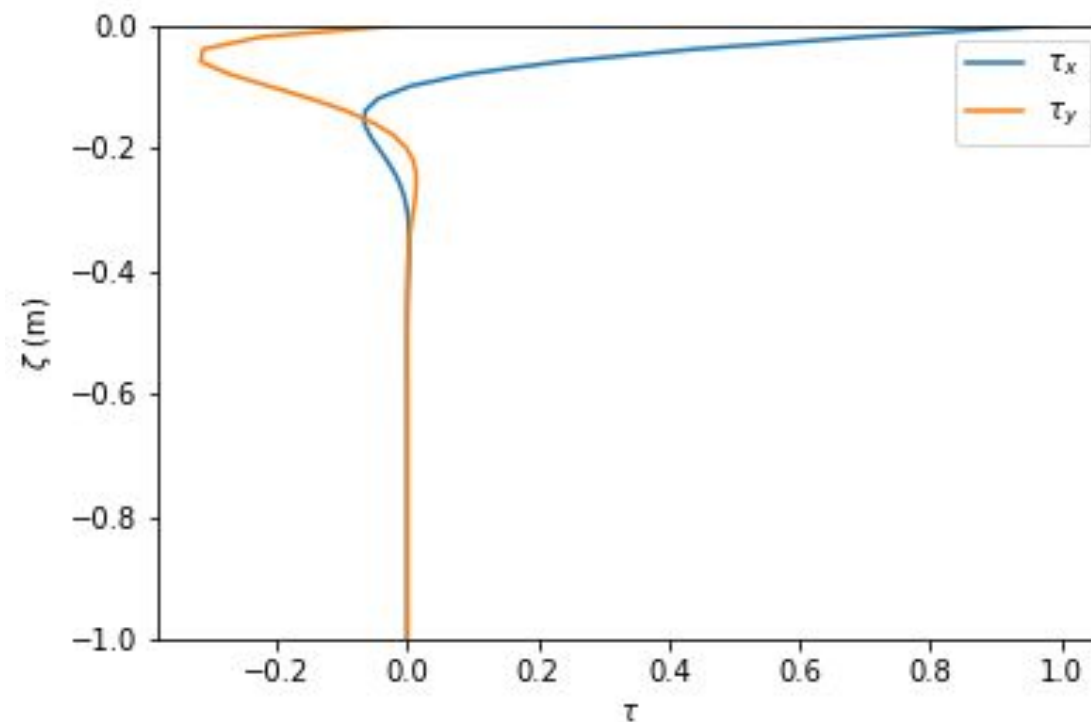
Monin-Obukhov Similarity Theory:

Scaled depth

$$\zeta = z/L_O$$

Monin-Obukhov length

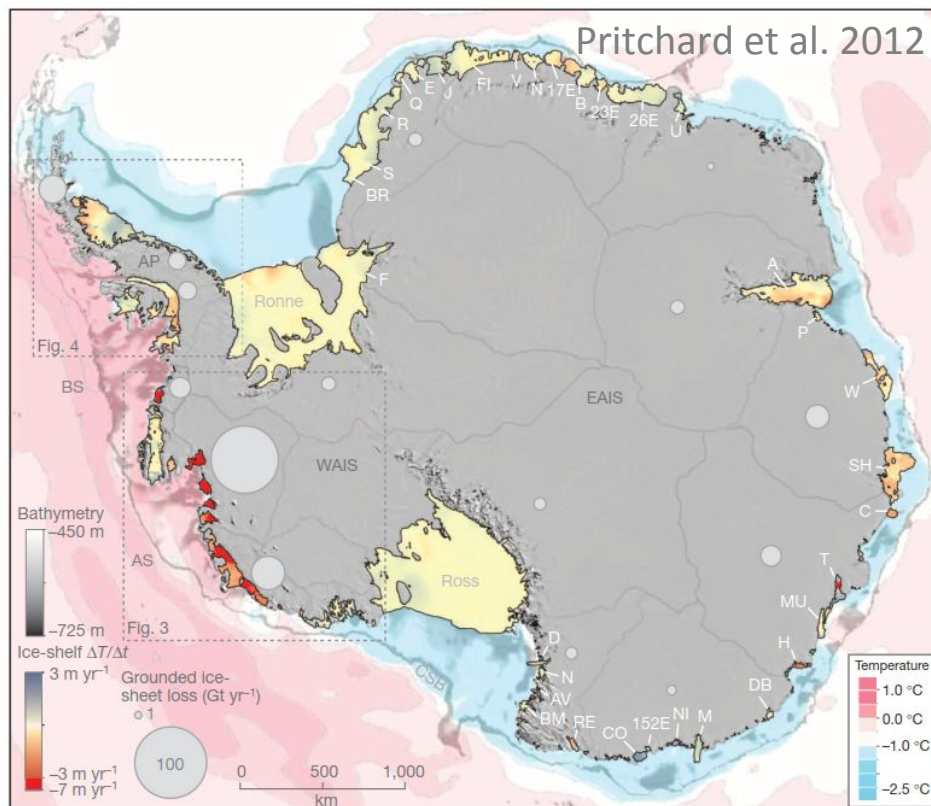
$$L_O = \frac{\rho u_*^3}{\kappa F_B}$$



Stress or vertical momentum flux

# What is the relationship between melt rates and ocean condition

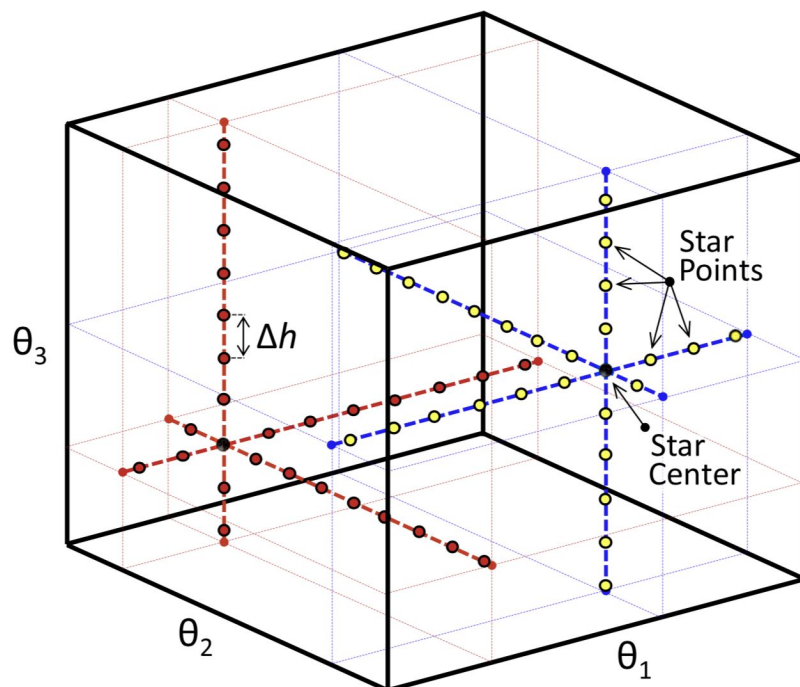
**Melt rate** =  $f$  (temperature, salinity, pressure, velocity, slope, basal roughness)



k degrees of freedom  
k is at least 7, not counting tidal frequencies

# What is the relationship between melt rates and ocean condition

**Melt rate** =  $f$  (temperature, salinity, pressure, velocity, slope, basal roughness)

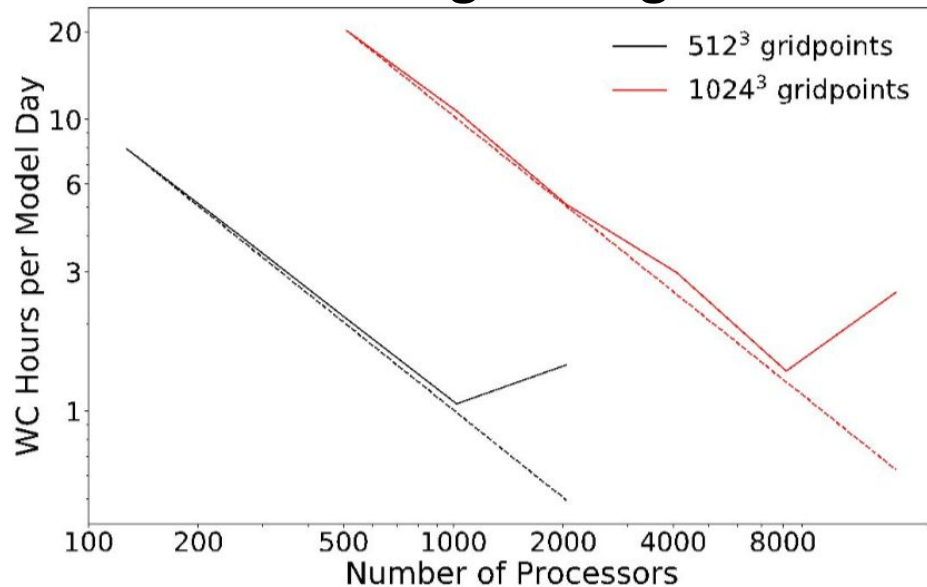


$k$  degrees of freedom  
 $k$  is at least 7, not counting tidal frequencies

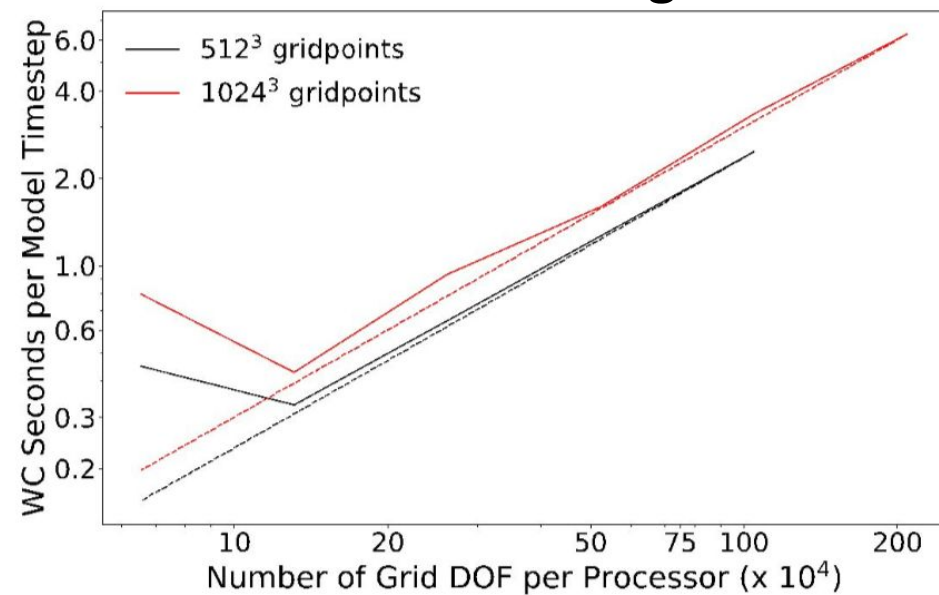
Sensitivity analysis:  
 $n$  samples of  $k$ -dimensional space  
Variance-based:  $n(10k+1)$   
Derivative-based:  $n(k+1)$

# The computational cost of large-eddy simulations

Strong scaling



Weak scaling



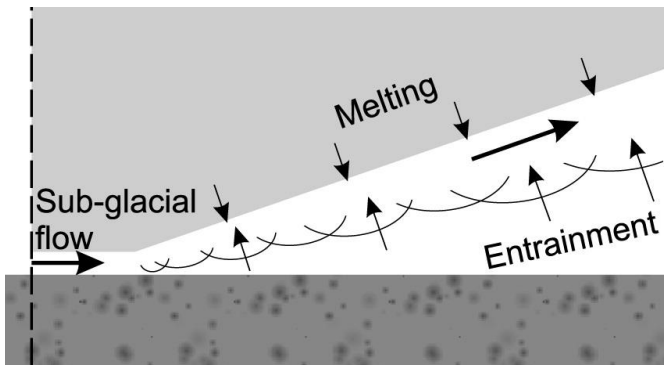
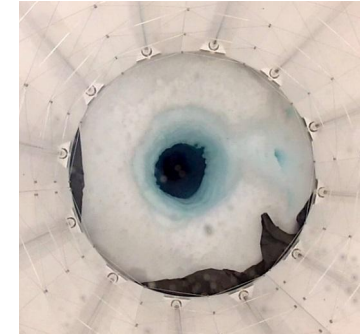
~30K core-hours per simulation

Allocated 6.6M core-hours on LANL machines

~100 simulations to constrain new melt parameterization

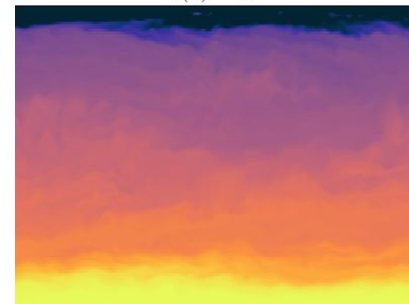
## ② Progress on the dynamics of ice-shelf melting

Observations of both ice-shelf melting and ocean conditions reveal that our existing melt parameterization is inaccurate



I developed a large-eddy simulation code to capture the unique configuration and forcing for an ice-shelf ocean boundary layer

I am running a suite of simulations to inform a new melt parameterization for global ocean models





# Thank you

Personal website **[carolynbegeman.weebly.com](http://carolynbegeman.weebly.com)**

DOE Earth System Model website **[e3sm.org](http://e3sm.org)**

Begeman, C., et al. 2018. Ocean stratification reduces melt rates at the grounding zone of the Ross Ice Shelf. *Journal of Geophysical Research, Oceans*.

Begeman, C., Tulaczyk, S., Fisher, A. 2017. Elevated and spatially geothermal flux below the West Antarctic Ice Sheet. *Geophysical Research Letters*.

Begeman, C. et al. 2019. Ice-shelf ocean boundary layer dynamics from large-eddy simulations. AGU Fall Meeting.

