

# Strain Localization and Elastic-Plastic Coupling During Deformation of Porous Sandstone

Thomas A. Dewers<sup>1\*</sup>, Kathleen A. Issen<sup>2</sup>, David J. Holcomb<sup>1,3</sup>, William A. Olsson<sup>1,3</sup>, and Mathew D. Ingraham<sup>1</sup>

<sup>1</sup>Sandia National Laboratories, Geomechanics Department, Albuquerque, NM

<sup>2</sup>Mechanical and Aeronautical Engineering, Clarkson University, Potsdam, NY

<sup>3</sup>Retired

November 11, 2013

\*Corresponding Author:

Thomas Dewers

[tdewers@sandia.gov](mailto:tdewers@sandia.gov)

Sandia National Laboratories

P.O. Box 5800 MS 0751

Albuquerque, NM 87185

505 845 0631 (phone)

505 844 7354 (fax)

## Abstract

Results of axisymmetric compression tests on weak, porous Castlegate Sandstone (Cretaceous, Utah, USA), performed at various mean stresses and covering a range of dilational and compactional behaviors, are examined for localization behavior. Assuming isotropy, bulk and shear moduli evolve as increasing functions of mean stress and Mises equivalent shear stress respectively, and as decreasing functions of work-conjugate plastic strains. Acoustic emissions events located during testing show onset of localization and permit calculation of observed shear and low-angle compaction localization zones, or bands, as localization commences. Total strain measured experimentally partitions into: A) elastic strain with constant moduli, B) elastic strain due to stress dependence of moduli, C) elastic strain due to moduli degradation with increasing plastic strain, and D) plastic strain. The third term is the elastic-plastic coupling strain, and though often ignored, contributes significantly to pre-failure total strain for brittle and transitional tests. Constitutive parameters and localization predictions derived from experiments are compared to theoretical predictions. In the brittle regime, predictions of band angles (angle

35 between band normal and maximum compression) demonstrate good agreement with observed  
36 shear band angles. Compaction localization was observed in the transitional regime in between  
37 shear localization and spatially pervasive compaction, over a small range of mean stresses. In  
38 contrast with predictions, detailed acoustic emissions analyses in this regime show low angle,  
39 compaction-dominated localization.

40

## 41 **1. Introduction**

42 Sandstones exhibit a range of deformational responses under stress conditions typical of  
43 the earth's subsurface. Understanding these responses, and quantifying their behavior  
44 mathematically (i.e., constitutively), is necessary to predict sandstone response for a variety of  
45 sub-surface geo-engineering pursuits such as oil and gas production and exploration, and  
46 geologic carbon storage [*de Waal and Smits, 1985; Rutqvist et al., 2009, Wawersik et al., 2000*].  
47 An important aspect of rate-independent sandstone deformation is the transition between  
48 localized deformation, in the form of shear and compaction zones or bands (i.e., concentrated  
49 tabular zones of strain that lack a discrete surface of discontinuity), and spatially pervasive pore  
50 collapse. The formation and development of localization can have important consequences for  
51 reservoir quality. Predictability of localized deformation from advanced constitutive modeling is  
52 thus a relevant consideration for waste injection or resource extraction [*Sulem et al., 1999; Sulem  
53 and Ouffroukh, 2006; Gajo and Bigoni, 2008*].

54 Localization in sandstones has been studied since the late 1970's [*Aydin, 1978; Aydin and  
55 Johnson, 1983; Antonellini and Aydin, 1995*], and recently identified variants in sandstone  
56 localization behavior include shear bands and compaction bands [*Olsson, 2001; Holcomb and  
57 Olsson, 2003; Wong et al., 2001; Klein et al., 2001, Issen and Rudnicki, 2000; 2001; Olsson et*

58 *al.*, 2002; *Detournay et al.*, 2003; *Bésuelle et al.*, 2003; *Haimson*, 2001; 2003; *Borja and Aydin*,  
59 2004; *Borja*, 2004; *Borja*, 2006; *Haimson and Lee*, 2004; *Baud et al.*, 2004; *Tembe et al.*, 2006;  
60 *Stanchits et al.*, 2007; *Rudnicki*, 2007; *Eichhubl et al.*, 2010; *Ingraham et al.*, 2013a; 2013b] and  
61 extension (dilation) bands [*Bésuelle et al.*, 2000; *Bésuelle*, 2001; *Du Bernard et al.*, 2002]. *Aydin*  
62 *et al.* [2006] note that the earliest reports of compaction bands in the field were by *Hill* [1989],  
63 and possibly the first experimental evidence of compactional deformation bands produced in the  
64 laboratory were by *Muhuri* [1994]. Recent detailed descriptions by *Eichhubl et al.* [2010] discern  
65 between pure compaction bands and shear-enhanced compaction bands.

66 While it is desirable to link constitutive response of laboratory samples to field examples,  
67 the disparate time scales and complex boundary conditions existing in field settings versus  
68 experiments make such comparisons difficult [*Holcomb et al.*, 2007]. As such, direct comparison  
69 of geological examples and laboratory experiments is fraught with uncertainty. It is relevant  
70 however to understand laboratory constitutive response in order to gain some insight on field  
71 observation. In axisymmetric compression experiments ( $\sigma_1 > \sigma_2 = \sigma_3$ , where the principal  
72 stresses are  $\sigma_1 \geq \sigma_2 \geq \sigma_3$ , positive in compression), compaction localization has been observed to  
73 form in high porosity sandstones (20-28%) [*Olsson*, 1999; *DiGiovanni et al.*, 2000; *Olsson and*  
74 *Holcomb*, 2000; *Wong et al.*, 2001; *Klein et al.*, 2001; *Holcomb and Olsson*, 2003; *Baud et al.*,  
75 2004; *Vajdova et al.*, 2004; *Stanchits et al.*, 2007]. Work by *Olsson* [1999], *Holcomb and*  
76 *Olsson* [2003] and *Vajdova et al.* [2004] demonstrate that localized compactive behavior results  
77 in order of magnitude changes in permeability.

78 This research examines mechanical response and localization behavior of a porous  
79 sandstone reservoir analogue rock, the Upper Cretaceous Castlegate Formation sandstone of east  
80 central Utah, USA [*McLaurin and Steel*, 2007], a dominantly fluvial sheet sand facies of the

81 Mesaverde Group. A suite of axisymmetric compression tests are conducted over a wide range  
82 of mean stresses under triaxial and constant-mean stress loading conditions. While compaction  
83 localization is of particular interest, Castlegate samples are also observed to fail via shear  
84 localization or uniform (non-localized) pore collapse, depending on mean stress. The  
85 localization theory of *Rudnicki and Rice* [1975] is applied toward predicting observed failure  
86 modes, and focus is placed on elastic-plastic coupling (i.e., modulus degradation with plastic  
87 strain) as an important consideration for constitutive modeling of porous sandstone deformation.  
88

## 89 **2. Theoretical Background**

### 90 *2.1 Constitutive Response and Elastic-Plastic Coupling*

91 To evaluate localization predictions from *Rudnicki and Rice* [1975], the total strain  
92 determined in experiments must be separated into elastic and inelastic components. This section  
93 describes the model used to conduct this strain separation. We start with common assumptions  
94 in elastic-plastic constitutive models, namely the separation of total strain into elastic and plastic  
95 parts (denoted by ‘*e*’ and ‘*p*’ superscripts)

$$\varepsilon_{ij}^t = \varepsilon_{ij}^e + \varepsilon_{ij}^p$$

96 (1)

97 where, upon unloading to zero stress, elastic strain vanishes and total strain equals plastic strain.

98 We additionally assume an elastic strain that follows from

$$\sigma_{ij} = C_{ijkl} \varepsilon_{kl}^e \quad (2)$$

100 With assumptions of isotropy and with usual definitions of invariants, (2) becomes

101 
$$\varepsilon^e = \frac{\sigma}{K}, \quad \gamma^e = \frac{\tau}{G} \quad (3)$$

102 For the axisymmetric compression (ASC) tests used here, mean stress  $\sigma = -I_1/3 = -\sigma_{kk}/3$  is  
 103 determined by  $\sigma = (\sigma_A + 2\sigma_L)/3$  and Mises shear stress,  $\tau = \sqrt{J_2} = \sqrt{\frac{1}{2}s_{ij}s_{ij}}$ , (where  $I_1$  and  $J_2$  are  
 104 the first and second invariants of the stress tensor, and the deviatoric stress is  $s_{ij} = \sigma_{ij} - \frac{1}{3}\sigma_{kk}\delta_{ij}$ )  
 105 is given by  $\tau = (\sigma_A - \sigma_L)/\sqrt{3}$ ,  $\sigma_A = \sigma_1$  is the axial stress and  $\sigma_L = \sigma_3 = \sigma_2$  is the confining  
 106 pressure. The elastic volumetric strain is found from  $\varepsilon^e = \varepsilon_A^e + 2\varepsilon_L^e$ , and the elastic shear strain  
 107 from  $\gamma^e = 2(\varepsilon_A^e - \varepsilon_L^e)/\sqrt{3}$ , during elastic portions of deformation, where  $\varepsilon_A$  is the axial strain and  
 108  $\varepsilon_L$  is the lateral strain. The bulk modulus and shear modulus are  $K$  and  $G$ , respectively.

109

110 Guided by experimental behavior shown later, we examine stress and plastic strain dependence  
 111 of elastic moduli as the following

112 
$$K = K(\sigma, \varepsilon^p), \quad G = G(\tau, \gamma^p) \quad (4)$$

113 Placing these into the former relations and expressing in incremental form, these become

114 
$$d\varepsilon^t = d\left(\frac{\sigma}{K(\sigma, \varepsilon^p)}\right) + d\varepsilon^p$$

115 and (5)

116 
$$d\gamma^t = d\left(\frac{\tau}{G(\tau, \gamma^p)}\right) + d\gamma^p$$

117 Expanding the total derivative yields

118 
$$d\varepsilon^t = \frac{d\sigma}{K} - \frac{\sigma}{K^2} \left( \frac{\partial K}{\partial \sigma} d\sigma + \frac{\partial K}{\partial \varepsilon^p} d\varepsilon^p \right) + d\varepsilon^p$$

119 and (6)

120 
$$d\gamma^t = \frac{d\tau}{G} - \frac{\tau}{G^2} \left( \frac{\partial G}{\partial \tau} d\tau + \frac{\partial G}{\partial \gamma^p} d\gamma^p \right) + d\gamma^p$$

121 From this, there are four strain increments that we label

122 
$$d\varepsilon^A = \frac{d\sigma}{K}; d\varepsilon^B = -\frac{\sigma}{K} \frac{\partial K}{\partial \sigma} d\sigma; d\varepsilon^C = -\frac{\sigma}{K} \frac{\partial K}{\partial \varepsilon^p} d\varepsilon^p; d\varepsilon^D = d\varepsilon^p$$

123 (7)

124 
$$d\gamma^A = \frac{d\tau}{G}; d\gamma^B = -\frac{\tau}{G} \frac{\partial G}{\partial \tau} d\tau; d\gamma^C = -\frac{\tau}{G} \frac{\partial G}{\partial \gamma^p} d\gamma^p; d\gamma^D = d\gamma^p$$

125 where the superscript ‘A’ refers to strain due to change in stress at a constant modulus, ‘B’ is due  
 126 to the stress dependence of the modulus, ‘C’ is the coupled strain due to plastic strain  
 127 dependence of the modulus, and ‘D’ is the plastic strain. Strains A, B and C are mechanically  
 128 recovered upon unloading to zero stress, and only strain D (plastic) is not recovered.

129 Following *Rudnicki* [1984], we distinguish between an “increment of inelastic strain”  
 130 which arises from only irrecoverable (plastic) strain,

131 
$$d\varepsilon^P = d\varepsilon^D \text{ and } d\gamma^P = d\gamma^D, \quad (8)$$

132 versus an “inelastic increment of strain” that includes both irrecoverable plastic strain and strain  
 133 arising from degradation of elastic moduli (and as such combines C and D strains). The inelastic  
 134 increments of strain are required to evaluate localization predictions [*Rudnicki and Rice*, 1975;  
 135 *Rudnicki* 1984]. Consistent with *Rudnicki* [1984] after *Hill and Rice* [1973], we use the  
 136 following notation for inelastic increments of strain:

137 
$$d^P \varepsilon = d\varepsilon^C + d\varepsilon^D \text{ and } d^P \gamma = d\gamma^C + d\gamma^D \quad (9)$$

138 If the elastic moduli do not evolve with stress, then the B strain increment vanishes and (7) is  
 139 equivalent to equation (2) of *Collins* [2002] and to equation (1) (with different sign convention)  
 140 of *Rudnicki* [1984].

141

## 142 **2.2 Localization Theory**

143 The strain localization framework of *Rudnicki and Rice* [1975] predicts conditions for the  
144 onset of localization in the form of a planar band of deformation, consisting of shear parallel to  
145 the band and/or compaction or dilation normal to the band. Localization is attributed to a  
146 constitutive instability in an initially uniform, inelastic deforming, pressure-dependent material;  
147 formation of a planar band of localized strain is an alternative solution to continued  
148 homogeneous deformation. The localization condition was found to be  $\det |n_i L_{ijkl} n_l| = 0$ , where  $n_i$   
149 are the components of the band normal and  $L_{ijkl}$  is the modulus tensor in the constitutive relation  
150  $d\sigma_{ij} = L_{ijkl} d\varepsilon_{kl}$ . Using the localization condition, and assuming that the hardening modulus  
151 decreases with continued loading, *Rudnicki and Rice* [1975] determined the predicted band  
152 orientation and the expression for the critical hardening modulus. Here, we present the relevant  
153 expressions, specialized for axisymmetric compression testing. Shear bands are predicted when

$$154 \quad -\sqrt{3} < \beta + \mu < +\sqrt{3} \left( \frac{2-\nu}{1+\nu} \right), \quad (12)$$

155 where  $\beta$  is the dilation coefficient (local slope of the plastic potential, Figure 1), which is positive  
156 for dilatant response, and negative for compactant response. Values for the dilation coefficient  
157 are determined from inelastic increments of strain,  $\beta = -d^p \varepsilon / d^p \gamma$ , at the onset of localization.  
158 The friction factor is  $\mu$ , which is the local slope of the yield surface (see Figure 1). At low to  
159 intermediate mean stresses,  ${}^p \gamma$  is used to track yield surface evolution; therefore, the initial and  
160 subsequent yield surfaces are contours of constant  ${}^p \gamma$ . Values for  $\mu$  are determined by  
161 calculating the local slope of the yield surface at the onset of localization. Poisson's ratio for  
162 elastic unloading is  $\nu$ .

163

164 For shear bands, the predicted band angle is given by [Rudnicki and Olsson, 1998]:

165

$$166 \quad \theta = \frac{\pi}{4} + \frac{1}{2} \arcsin \left[ \frac{2}{3\sqrt{3}} (1+\nu)(\beta + \mu) - \frac{1}{3}(1-2\nu) \right], \quad (13)$$

167

168 where the band angle is defined as the angle between the band normal and the direction of

169 maximum compression. The onset of localization is predicted to occur when the hardening

170 modulus,  $h$ , (slope of the shear stress – plastic shear strain curve at constant mean stress)

171 decreases to a critical value,  $h_{cr}$ . The expression for the critical hardening modulus for shear

172 bands is [Rudnicki and Rice, 1975; Issen and Challa, 2008]

173

$$174 \quad h_{cr}^s = \frac{G(1+\nu)}{9(1-\nu)} \left\{ (\beta - \mu)^2 - \frac{(1-\nu)}{2} \left[ \sqrt{3} + (\beta + \mu) \right]^2 \right\}. \quad (14)$$

175

176 If  $(\beta + \mu) \leq -\sqrt{3}$ , then compaction bands,  $\theta = 0^\circ$ , are predicted, and the expression for the

177 critical hardening modulus is [Issen and Rudnicki, 2000]

178

$$179 \quad h_{cr}^c = -\frac{4G(1+\nu)}{9(1-\nu)} \left( \beta + \frac{\sqrt{3}}{2} \right) \left( \mu + \frac{\sqrt{3}}{2} \right). \quad (15)$$

180

181 If  $(\beta + \mu) \geq \sqrt{3}(2-\nu)/(1+\nu)$ , then dilation bands,  $\theta = 90^\circ$  are predicted. However, since

182 dilation bands were not observed in the Castlegate experiments described below, the expression

183 for the hardening modulus is not provided, but can be found in other works [e.g., *Issen and*  
184 *Rudnicki, 2000; Issen and Challa, 2008*].

185 The above predictions are appropriate for low and intermediate mean stress tests, where  
186 significant plastic shear strain occurs. However, for high porosity sandstone, a yield surface cap  
187 exists [*Wong et al., 1992; Wong et al., 1997; Olsson, 1999; Baud et al., 2006; Tembe et al., 1997,*  
188 *Ingraham et al., 2013b*] and under hydrostatic pressure, compaction is accompanied by  
189 significant plastic volume strain [*Zhang et al., 1990*] and little plastic shear strain. Therefore,  
190 extending the yield surface as a contour of constant  $^P\gamma$  to intersect the mean stress axis is  
191 problematic. A simple way to circumvent this problem is to formulate a separate constitutive  
192 framework for the cap, which depends on inelastic volume strain, instead of inelastic shear strain  
193 [e.g., *DiMaggio and Sandler, 1971; Issen, 2002; Baud et al., 2006*]. To simplify our data  
194 analysis, we choose the cap model derived by *Issen [2002]*, as the resulting localization  
195 conditions parallel those of *Rudnicki and Rice [1975]*. For the cap, the initial and subsequent  
196 yield surfaces are contours of constant  $^P\varepsilon$ ; values for  $\mu$  are still determined by calculating the  
197 local slope of the yield surface at the onset of localization. Determination of values for  $\beta$   
198 remains unchanged. However, the relevant hardening modulus is now the “bulk” hardening  
199 modulus,  $k$ , (slope of the mean stress – plastic volume strain curve at constant shear stress).  
200 Using this cap mode, when  $\beta$  and  $\mu$  are both negative (as expected for compacting material),  
201 equations (12) and (13) are still applicable. The critical bulk hardening modulus is  
202  $k_{cr} = h_{cr} / (\beta\mu)$ , using the appropriate  $h_{cr}$  from equation (14) or (15) [*Issen, 2002*].

203

### 204 3. Experimental Methods

#### 205 3.1 Axisymmetric Compression Experiments

206  
207 There have been numerous studies examining deformation of Castlegate sandstone  
208 [Vinegar *et al.* 1991; Olsson, 1999; Olsson and Holcomb, 2000; Olsson, 2001; Olsson *et al.*,  
209 2002; Holcomb and Olsson, 2003; Coop and Willson, 2003; Holcomb *et al.*, 2004; Ingraham *et*  
210 *al.*, 2013a; 2013b], as it may serve as an outcrop analogue for porous reservoir sandstones from  
211 similar depositional environments. Here, an experimental data set of Castlegate behavior is  
212 compiled to compare localization occurrences to theoretical predictions. Coring, sample  
213 preparation and jacketing procedures were performed as described by Olsson [1999]. Right-  
214 cylindrical samples were prepared by coring parallel to the bedding plane and grinding to a  
215 nominal 5.08 cm diameter by 12.7 cm length. All samples were cored from the same block of  
216 Castlegate in an attempt to minimize sample-to-sample variability. As discussed by Holcomb  
217 and Olsson [2003], coring parallel to the bedding plane eliminates the possibilities that  
218 compaction bands are a product of variation in material properties across the bedding planes or  
219 simply collapse parallel to bedding.

220 The sandstone cylinders were jacketed with 100 micron copper foil, with a copper cap at  
221 each end. To minimize end frictional effects, a stearic acid-petroleum jelly lubricant was used  
222 between the copper cap and the sandstone, and a teflon sheet coated with molybdenum  
223 disulphide was placed between the outer copper caps and upper and lower pistons [Holcomb and  
224 Olsson, 2003].

225 Experiments were conducted using a servo-controlled test frame under axial strain control  
226 at  $10^{-5}$  to  $10^{-6}$  s<sup>-1</sup> nominal strain rates, using total axial strain as the control variable, to minimize  
227 effects of machine stiffness and to better control strain rates during localization. Axial strain ( $\epsilon_A$ )

228 measurements were made via linear variable differential transformer (LVDTs) mounted end to  
229 end on the sample, and lateral strains ( $\epsilon_L$ ) were measured using a ring gage mounted near the  
230 middle of the sample. Axial load was measured via an external load cell. During each loading  
231 path, mechanical data acquired include axial load, axial displacement, confining pressure, pore  
232 pressure (where applicable), radial displacement, and acoustic emissions. Stress states for  
233 samples deformed to large strains and higher mean stresses were corrected for changes in cross  
234 sectional area; for all samples that underwent localization, this correction amounts to less than 1  
235 MPa and so was not applied in these cases.

236 Samples were subjected to a variety of loading paths: uniaxial, hydrostatic, hydrostatic  
237 followed by triaxial compression, and hydrostatic followed by constant mean stress loading.  
238 Seventeen experiments, consisting of these different loading paths, are summarized in Table 1.  
239 Results of seventeen such experiments are plotted in Figure 2, which shows axial stress – axial  
240 strain and – lateral strain behavior (in each of Figure 2A and B, the curves to the right plot axial  
241 strains, while the curves to the left plot lateral strains). These experiments are also portrayed in  $\tau$   
242 –  $\sigma$  stress space (Figures 2C-D). Figure 2A and C are for tests which exhibited some form of  
243 localization, while Figures 2B and D show results of tests which did not experience any  
244 localization behavior.

245 In most of the constant mean stress tests and in the uniaxial tests, unloading-reloading  
246 cycles (visible as stress decreases in the curves in Figures 2A-B) were performed to track the  
247 evolution of elastic properties of the sandstone. This permitted quantifying the important role of  
248 plastic deformation in modifying elastic parameters (elastic-plastic coupling as discussed above).  
249 During typical unloading cycles, stresses were decreased by about 10% of expected peak stress  
250 (for unloading during constant mean stress tests,  $\sigma_L$  was increased as  $\sigma_A$  was decreased, so as to

251 change only  $\tau$ ), before reloading and continuing the loading portion of the test. During unloading  
252 loops, stress and strain exhibit hysteresis as a result of the use of an external load cell. Friction  
253 between piston and piston seal (where the piston enters the pressure vessel) delays transmission  
254 of load from the loading system to the sample, both on unloading and reloading. To surmount  
255 this, only portions of unloading loops, wherein friction has been overcome, were used to infer  
256 elastic properties. Only pre-peak unloading-reloading cycles were used in determining elastic  
257 moduli evolution, as after peak stress is achieved (corresponding to onset of shear or compaction  
258 localization), sandstone specimens can no longer be reasonably approximated as homogeneous.  
259

### 260 ***3.2 Acoustic Emissions***

261 For the majority of the experiments, acoustic emissions (AE) were recorded using arrays  
262 of three to fourteen transducers soldered to the sample copper jacket, in the manner described by  
263 *Olsson and Holcomb* [2000]. AE signals were amplified by 60 db in the frequency band from  
264 100 kHz to 1.2 MHz. A full wave form recording system capable of acquiring about 5 events per  
265 second recorded events detected by five or more transducers. Signals were digitized at 8 bit  
266 resolution and 32 MHz for later use in event location. It is estimated that the recorded events  
267 represent approximately 10% of the total event set. Recorded events were processed by means of  
268 “picker” software to determine arrival times for the signal from an event at each transducer, and  
269 the simplex algorithm of *Olsson and Holcomb* [2000] was used in determining locations from the  
270 first arrivals. A separate event counter system triggered by one event detector typically records  
271 several hundred thousand events during an experiment. Event locations were used to determine  
272 orientation of localization features.

273

## 274 **4. Experimental Results**

### 275 **4.1 Stress-Strain Response**

276 Examples of Castlegate behavior under different loading conditions are shown in Figure 3.  
277 Black lines plot axial stress – axial strain curves, while grey lines plot AE rate. Uniaxial loading  
278 behavior is shown in Figure 3A (experiment 41; no AE was recorded for this test), resulting in a  
279 stress peak and formation of high-angle shear localization. Constant mean stress loading ( $\sigma$   
280 equal to 30 MPa) is shown in Figure 3B (experiment 22), resulting in shear localization and a  
281 less distinct stress peak. Constant mean stress loading ( $\sigma$  equal to 70 MPa), with dominantly  
282 compactional localization is shown in Figure 3C (experiment 15). Constant mean stress loading  
283 ( $\sigma$  equal to 90 MPa, experiment 21), resulting in no localization, is shown in Figure 3D. Note  
284 that in Figures 3B-D, the initial loading curves contain a segment of hydrostatic loading,  
285 followed by an increase in differential stress at constant mean stress. The post-peak plateau in  
286 Figure 3C is characteristic of compaction localization and propagation [*Olsson and Holcomb*,  
287 2000] and it is interesting that no stress drops are discernable along the stress plateau, as occur in  
288 other examples of experimental compaction localization [e.g. *Baud et al.*, 2006].

289 In Figures 3B and C, it is evident in all experiments that some acoustic activity, suggestive  
290 of plastic deformation, initiated from the start of loading. Peak AE activity, for samples which  
291 underwent shear or compaction localization, occurs near peak stress (where localization  
292 commences as per *Olsson and Holcomb*, [2000]) and then tapers off. In experiment 21 (Figure  
293 3D), in which no localization occurred, AE activity peaks near peak stress and dies off with  
294 subsequent loading, while in experiment 15 (compaction localization, Figure 3C), continued AE  
295 activity attends the development and propagation of zones of compaction localization. These  
296 results are consistent with the results of Ingraham et al. [2013].

297

## 298 *4.2 Acoustic Emissions Analyses*

299 Historically, determinations of localization were made post-test by measuring the slope of  
300 the macroscopic fracture, often made evident by the deformation of the jacket. For shear  
301 localization, this is reasonably valid because once shearing begins on a plane, it is difficult for  
302 the angle of localization to change due to the kinematic constraints of the now-discrete blocks of  
303 material sliding past each other. It is possible for other zones of localization to form, however,  
304 and then determination of initial bands of localization is impossible post-test. Because the  
305 localization analysis only applies at the point where the localization condition is first satisfied, it  
306 is important to know the band angle prior to significant deformation and possible formation of  
307 secondary features. The problem is compounded for low angle bands which often do not visibly  
308 deform the jacket or sample, requiring microscopic examination to determine localization band  
309 characteristics.

310 Acoustic emissions are elastic waves produced by grain-scale failures and motions that are  
311 assumed to be discrete microscopic sources of continuum macroscopic phenomena described by  
312 localization theories. If deformation is localized, then it follows that the locations where AE  
313 occurs should similarly be localized and can be used to track and characterize spatial character of  
314 the localization process. This technique has the advantage of being able to observe the interior of  
315 the sample during deformation without interfering with the experiment; it can detect features that  
316 are obscured or destroyed by the progressive deformation of the sample after localization occurs.

317 Deforming one specimen can easily result in 100,000 recorded AE, of which perhaps 50%  
318 can be located to an accuracy of 2-3 mm. By plotting the locations as a series of time slices, like  
319 a movie, the spatial and temporal evolution of the causative deformations can be observed. This

320 was the approach used by *Olsson and Holcomb* [2000] to confirm that low angle zones of  
321 localized deformation, termed compactional bands by them, were forming and propagating  
322 through Castlegate sandstone. In that work, clouds of AE were described as relatively flat-lying,  
323 i.e., having a band angle of zero degrees. In the current work, a more thorough analysis was  
324 done of event cloud orientation to facilitate comparison with calculated band angles discussed  
325 later.

326         Figures 4, 5, and 6 are examples of AE results for tests, arranged in order of increasing  
327 mean stress, which produced shear localization, compaction localization, and no localization,  
328 respectively. In each figure, the lower panel plots axial stress as a function of axial strain, along  
329 with the band angle computed for the cloud of AE events that occurred during sequential time  
330 slices of the experiment. At the top of each figure are three examples, captured from the full  
331 results, showing the AE locations in the cylindrical specimen during three time slices, with a  
332 vector (in grey) showing the normal to the best fit plane. Black dots correspond to events in the  
333 upper half of the specimen that were used to compute the best-fit plane. Gray dots are events that  
334 occurred in the lower half of the specimen during the same time slice. Because localization  
335 frequently results in bands that are simultaneously forming in the top and bottom halves of the  
336 specimen, it was necessary to restrict the analysis to half of the sample at a time. Arrows  
337 pointing to the stress – strain curve indicate the portion of the test of which the examples are  
338 representative: pre-peak, near-peak and post-peak. In all figures, the plot has been rotated about  
339 the cylinder axis until the view direction is along the strike of the fitted plane, emphasizing the  
340 slope.

341         Results from a test at low mean stress, seen in Figure 4, show a diffuse band of events  
342 formed in the upper half of the sample near peak stress, and evolved post-peak into a band with a

343 well defined slope. The band angle evolved from near  $90^\circ$ , a numerical artifact as discussed  
344 below, to about  $50^\circ$  as the stress dropped post peak. Near the stress peak, it is clear that AE  
345 activity had localized, but there was no obvious shear plane until the sample shortens by about  
346 0.007 strain units post-peak. A possible explanation for this phenomenon is that when the  
347 localization condition is satisfied in the cylindrically symmetric stress state, the band angle with  
348 respect to the maximum principal stress is determined, but not the azimuthal angle. As a result,  
349 localization initiates on all possible planes with that band angle, resulting in a rotationally  
350 invariant cloud of events, which resolve into one dominant band as deformation continues.

351         A low angle compactional band formed in sample 15, which was tested at an intermediate  
352 mean stress (see Figure 5). Prior to peak stress, events were uniformly distributed with no  
353 meaningful band angle calculated. At peak, a well defined band formed near the upper end of the  
354 sample, accompanied by a nearly complete cessation of activity in the lower half of the sample.  
355 The band angle was computed to be approximately  $10^\circ$  and remained in the range of  $5^\circ$  to  $15^\circ$  as  
356 the band of AE activity propagated down the sample.

357         Figure 6 presents results for an experiment where no localization occurred. Pre-peak  
358 (left-most AE location plot), the events are uniformly distributed as would be expected for  
359 homogeneous deformation accompanied by plastic shear and volume strain associated with pore  
360 collapse. Near the peak (center plot) there is some indication of localization deformation as  
361 shown by the denser concentration of events in the lower half of the cylinder, but no real band or  
362 zone forms. As deformation continues, the event distribution becomes relatively uniform,  
363 consistent with the interpretation that localization did not occur in this sample. Note that no plane  
364 normal is plotted, even though the algorithms computed a good fit to a plane with a dip angle of  
365 approximately  $90^\circ$ , as shown in the lower portion of Figure 6. Visually it is clear that the best fit

366 plane would indeed run parallel to the axis of the cylinder as a result of the 2.5 to 1 aspect ratio  
367 forced on the event cloud by the confines of the cylindrical specimen. However, for our  
368 purposes, this result is meaningless and attributable to sample geometry.

369 Acoustic emissions analyses were conducted for all specimens for which AE events were  
370 collected. For those specimens where localization occurred, the localization band angles were  
371 determined as described above; these observed angles are shown in Table 2. In some specimens,  
372 a low angle compaction localization developed at each end of the specimen; therefore, two  
373 angles are reported.

374

## 375 **5. Strain Partitioning Into Elastic and Plastic Portions.**

### 376 *5.1 Evolution of Bulk Moduli*

377 To examine evolution of elastic moduli with stress and plastic strain, we first look at  
378 Castlegate hydrostatic consolidation (experiment 38), shown in Figure 7A. The mean stress –  
379 volume strain curve portrays a typical behavior for porous sands and sandstones [Zhang *et al.*,  
380 1990; Karner *et al.*, 2003; 2005], including an initial concave-up pressure-strain segment  
381 (nonlinear elastic), a first inflection followed by a slightly convex segment (initial and  
382 subsequent plastic yielding), a second inflection followed by a roll-over with large strain  
383 accumulation (pore collapse), and a final convex segment with increasing slope (compaction and  
384 pore crush-out). Four unloading loops and the final unloading segment in this experiment aid in  
385 interpreting elastic moduli evolution during the progressive loading. Evident by the difference in  
386 slopes of the initial loading and final loading segments is an increase in elastic compliance with  
387 accumulating plastic strain, that is, elastic-plastic coupling as discussed above.

388 For non-linear elasticity, it is important to distinguish between stress dependent secant  
389 moduli, such as given in (3), with tangent moduli, which are local slopes of stress-strain curves.  
390 We denote secant bulk and shear moduli as  $K$  and  $G$  respectively, and tangent bulk and shear  
391 moduli as  $K_{tan}$  and  $G_{tan}$ , respectively. Tangent bulk moduli,  $K_{tan}$ , are determined from local  
392 slopes of mean stress-volume strain curves in elastic loading/unloading regimes, which include  
393 pre-yield initial loading, during unload-reload cycles, and final unloading. For the initial, final,  
394 and unloading segments of the stress-strain curve for experiment 38 (Figure 7A), tangent moduli  
395 are plotted as symbols in Figure 7B. Values of  $K_{tan}$  increase initially due to stress dependence,  
396 and then decrease relative to the values of  $K_{tan}$  determined during the initial unloading loop,  
397 suggestive of plastic strain dependence due to yielding initiating at ~50 MPa. Combining the first  
398 unloading moduli with the pre-50 MPa moduli suggests a  $K(\sigma)$  relation wherein  $K_{tan}$  increases as  
399 an exponential function of  $\sigma$ , similar to that determined for Berea sandstone by *Kahn et al.*  
400 [1991]. The  $K_{tan}$  calculated from the second unloading loop displays a slope that is less than  
401 values from the first unloading loop, as are moduli determined from the final unloading. Note  
402 that the final unloading moduli also indicate an exponential  $K_{tan}(\sigma)$  relation, but with lesser  
403 slopes, suggesting that moduli at similar stress conditions decrease as a function of increasing  
404 plastic strain. Slopes and zero-stress intercepts appear to both decrease monotonically as a  
405 function of increasing plastic strain. These relationships are also found during hydrostatic  
406 loading/unloading portions of other experiments (not shown here).

407 In an attempt to capture these functional relationships in as simple and tractable a  
408 mathematical form as possible, we suggest the following for a stress- and strain-dependent bulk  
409 modulus,  $K = K(\sigma, \varepsilon^p)$ :

$$410 \quad K = K_0 \left( 1 + K_1 \sigma - K_2 e^{-K_3 K_0 \sigma} \right) \quad (16)$$

411

412 where  $K_1$ ,  $K_2$  and  $K_3$  are constants and  $K_0$  is a parameter that decreases as a function of plastic  
413 volume strain,  $\varepsilon^p$ . Here we have chosen the work-conjugate stress and strain variables (i.e., an  
414 increment of plastic work for an isotropic material given by  $dW = \sigma d\varepsilon^p + \tau d\gamma^p$ ) to express the  
415 stress and strain dependence of  $K$ . This stress dependence is suggested from work on  
416 compression- and shear-wave velocities in porous sandstones. For example *Kaselow and Shapiro*  
417 [2004] provide a pore-scale interpretation in which the exponential term in (16) derives from  
418 compliant microcracks and cracked-shaped pores, and the linear term derives from more rounded  
419 pores; these are termed compliant and stiff porosity, respectively. The plastic volume strain  
420 dependence we apply is based on empirical observation of Castlegate behavior.

421 The form of  $K_{tan}$  applicable to elastic portions of stress – strain curves, which include initial  
422 loading, unload-reload cycles, and final unloading (whereupon the increment of plastic strain  
423  $d\varepsilon^p$  vanishes) can be found by combining (16) with the first of equations (6) and assembling  
424 terms such that  $d\varepsilon^e = d\sigma/K_{tan}$ , where

$$425$$

$$426 \quad K_{tan} = \left\{ \frac{1}{K} - \frac{\sigma}{K^2} (K_0^2 K_2 K_3 e^{-K_3 K_0 \sigma}) \right\}^{-1} \quad (17)$$

427  
428 Values for the parameter set  $\{K_0, K_1, K_2, K_3\}$  can be determined by either curve fitting secant  
429 moduli to (3) or tangent moduli to (6). The procedure we adopt is to (1) fit the final unloading  
430 curve (Fig. 7A) to find  $\{K_0, K_1, K_2, K_3\}$ , and then fit unload-reload cycles with these values of  
431  $K_1, K_2$ , and  $K_3$  to find a unique  $K_0$  applicable to each cycle and initial loading; and (2) determine  
432 a smooth function for  $K_0 = f(\varepsilon^p)$  to interpolate between unload loops. As  $\varepsilon^p$  is initially unknown,  
433 a simple procedure for interpolation is to first find  $K_0 = f(\varepsilon^t)$ , that is, as a function of total strain.  
434 While there are certainly a number of functions that we could use, a simple form for  $K_0$  that we  
435 have found to apply to all experimental results presented here is

436 
$$K_0 = K_0' (1 - K_4 e^{-K_5 / \epsilon_p}) \quad (18)$$

437 where, for initial interpolating, the  $\epsilon^p$  is replaced by  $\epsilon^l$ . This expression is applied to each  
 438 experiment with the same values for  $K_1, K_2, K_3, K_4, K_5$  and a  $K_0$  that evolves monotonically with  
 439 plastic volume strain but is constant during unloading loops. Best-fit  $K_i$  parameters for initial  
 440 loading, unload-reload cycles, and final unloading for non-localized samples are determined for  
 441 each experimental mean stress-volume strain path (that is, we fit the secant modulus) by  
 442 applying the nonlinear Levenberg-Marquardt regression algorithm, using a commercially  
 443 available software package. To determine plastic volume strains applicable to a given unloading  
 444 loop, each unloading stress – strain curve is extrapolated to zero stress, and the plastic volume  
 445 strain is equated to the intercept of the volume strain axis.

446 To demonstrate the viability of our model for  $K$ , calculated tangent moduli are shown as the  
 447 solid black line in Figure 7B for experiment 38 using the parameterization scheme as  $\{K_0', K_1,$   
 448  $K_2, K_3, K_4, K_5\}$  equal to  $\{2814, 0.0057, 0.79, 1.54e-5, 0.98, 0.067\}$ . Compared to experimental  
 449 values (symbols in Figure 7B), this model for  $K_{tan}$  does well at the latter portions of the  
 450 experiment and for initial loading, but slightly overestimates moduli for the second unload loop,  
 451 and a slightly lessened slope for the first unload loop. For comparison, the calculated secant  
 452 moduli  $K$  for the same  $K_i$  values is shown by the grey line in Figure 7B; values of the secant  
 453 moduli are approximately half those of the tangent moduli for much of the stress path.

454

## 455 ***5.2 Partitioning of Total Volume Strain Into Elastic and Plastic Parts***

456 Having developed a model for the evolution of  $K$  during loading and plastic yielding, it is  
 457 a simple matter to calculate elastic volume strains by combining equations (3), (16) and (18),  
 458 and, by subtracting elastic from total strains, to calculate plastic volume strain. This is shown

459 graphically in Figure 7C. The solid black line is total (measured) volume strain, the dark grey  
460 line is calculated elastic strain ( $\varepsilon^A + \varepsilon^B + \varepsilon^C$ ), and, by difference, the dashed black line shows  
461 plastic volume strain ( $\varepsilon^D$ ). The light grey line, which shows no hysteresis, plots the elastic strain  
462 that would manifest if no modulus degradation occurred ( $\varepsilon^A + \varepsilon^B$ ). It is immediately evident that  
463 elastic-plastic coupling ( $\varepsilon^C$ , the difference between the dark grey and light grey curves), caused  
464 by the plastic volume strain dependence of the bulk modulus, accounts for approximately half of  
465 the total strain for much of the loading path. Similar results were found for some lithofacies of  
466 the Mt. Simon Sandstone [Dewers *et al.*, 2013]. For example, at peak mean stress of 400 MPa,  
467 the total measured strain ( $\varepsilon^T$ ) is about 0.18 strain units, plastic strain ( $\varepsilon^D$ ) is about 0.05 strain  
468 units, and the elastic volume strain ( $\varepsilon^A + \varepsilon^B + \varepsilon^C$ ) is 0.14 strain units. The elastic strain with no  
469 modulus degradation ( $\varepsilon^A + \varepsilon^B$ ) would amount to a little less than 0.05 strain units, which implies  
470 that the coupling elastic strain ( $\varepsilon^C$ ) and inelastic strain ( ${}^p\varepsilon = \varepsilon^C + \varepsilon^D$ ) would, at peak load,  
471 amount to approximately 0.09 and 0.14 strain units, respectively. The coupling strain then,  
472 accounts for a significant portion of total strain accumulated during the experiment, as well as  
473 much of the shape of the overall hydrostat.

474 Figure 7D plots equation (18) as a dashed line, for the plastic strain path shown in Fig. 7C;  
475 the solid symbols are plotted for each unload-reload cycle. This suggests that much of the  
476 modulus degradation, with largest drop in modulus value, occurs between the second and third  
477 unload-reload cycle. This is reasonable, corresponding to the largest accumulation of plastic  
478 strain and accordingly, the “knee” of the hydrostat. Values for  $\{K_4, K_5\}$  applicable to using  $\varepsilon^p$   
479 are  $\{0.68, .003\}$ .

480

### 481 ***5.3 Evolution of Shear Moduli and Partitioning of Total Shear Strain***

482 Secant and tangent shear moduli,  $G$  and  $G_{tan}$ , can found by similar methods as done for  
 483 bulk moduli. As an example we show evolution of shear moduli as functions of plastic shear  
 484 strain,  $\gamma^p$ , and shear stress,  $\tau$ , for experiment 42, a constant mean stress test at  $\sigma$  equal to 150  
 485 MPa. The stress path for 42 (Figure 8A) includes hydrostatic loading to 150 MPa, followed by a  
 486 constant mean stress loading to  $\tau$  equal to 70 MPa with eight unload loops. Sample 42  
 487 experienced no localization, which enable us to determine shear moduli from final unloading as  
 488 well.

489 Figure 8B shows tangent moduli from the unload loops and final unloading, and is  
 490 suggestive of a shear modulus that increases with  $\tau$  at constant plastic shear strain, and decreases  
 491 as a function of plastic shear strain,  $\gamma^p$ . (Note that we observe a clear mean stress dependence of  
 492 shear moduli, but as most of our results are for constant mean stress tests, we focus on the  
 493 observed shear stress dependence.) In analogy with our model for  $K$ , we choose to express the  
 494 stress and plastic strain dependence of  $G$  in terms of the work-conjugate variables  $\gamma^p$  and  $\tau$  as the  
 495 following:

$$496 \quad G = G_0' \left( 1 + G_4 e^{-G_5/\gamma^p} \right) (1 + G_1 \tau). \quad (19)$$

497  
 498  $G_0$  and  $G_1$  parameters are determinable via a non-linear regressions procedure as was done for  $K$ .  
 499 We initially fit the stress dependence from the final unloading to give  $G_1$ , as well as final  $G_0$ , and  
 500 then apply the same  $G_1$  to the unload loops. As was done in the case of bulk moduli, we first use  
 501 the total strain in (19), replacing  $\gamma^p$  with  $\gamma^t$ , as a function that allows us to interpolate between  
 502 unload loops. In Figure 8B, we compare calculated  $G_{tan}$  (symbols) to (19) using the  
 503 parameterization scheme of  $\{G_0', G_1, G_2, G_4, G_5\}$  equal to  $\{7611, 0.0071, 0.60, 0.0021\}$  with  
 504 overall good agreement.

505 By combining (3) with (19), we then calculate elastic shear strains ( $\gamma^A + \gamma^B + \gamma^C$ ) shown by  
506 the dark gray line in Figure 8C, and then by difference with total strain, the plastic shear strain  $\gamma^P$   
507 (or  $\gamma^D$ ) shown by the dashed black line in Figure 8C. The calculated elastic strain ( $\gamma^A + \gamma^B$ ) with  
508 no plastic strain dependence (i.e.,  $G_4, G_5$  in equation 19 equal to zero) is shown by the light grey  
509 line. In contrast with the hydrostat in Figure 7A, it is evident that plastic shear strain dominates  
510 measured values compared to the elastic strain, and the elastic strain associated with elastic  
511 plastic coupling ( $\gamma^C$ ; the light grey curve subtracted from the dark grey curve) is roughly twice  
512 that of elastic strain with no coupling ( $\gamma^A + \gamma^B$ ).

513 Evolution of the  $G_0$  parameter with plastic shear strain  $\gamma^P$  is shown in Figure 8D. As was  
514 seen Figure 7C for the hydrostatic case, the biggest drop in shear modulus is associated with the  
515 “knee” of the stress strain curve, corresponding to the accumulation of plastic shear strain  
516 between the second and fourth unload-reload cycle.

517

#### 518 ***5.4 Partitioning of Strain for Experiments Resulting in Localization***

519 Following the procedures in the last two sections, results of partitioning of total strain into  
520 elastic and plastic portions for three experiments, up to peak stress where localization  
521 commenced, are shown in Figures 9A-F. These are experiment 41 (Figures 9A-B), loaded under  
522 uniaxial conditions to a peak  $\tau$  of 8.7 MPa; experiment 22 (Figures 9C-D), loaded to 25 MPa  
523 mean stress hydrostatically, and then loaded to a peak  $\tau$  of ~28 MPa at constant mean stress; and  
524 experiment 15 (Figures 9E-F), loaded to a mean stress of 70 MPa hydrostatically and then under  
525 triaxial conditions to a peak  $\tau$  of ~50 MPa. Solid black lines in Figure 9 refer to total strain,  
526 dashed lines plot calculated plastic strains ( $\varepsilon^D$  and  $\gamma^D$ ), dark grey lines plot elastic strains  
527 including elastic-plastic coupling ( $\varepsilon^A + \varepsilon^B + \varepsilon^C$  and  $\gamma^A + \gamma^B + \gamma^C$ ), and light grey lines are elastic

528 strains ignoring the coupling strain terms ( $\varepsilon^A + \varepsilon^B$  and  $\gamma^A + \gamma^B$ ). In all three cases, changes in  
529 bulk moduli accompanying plastic volume strain were deemed so small as to be negligible, and  
530 were ignored in the calculation of plastic volume strain. However, elastic-plastic coupling in the  
531 determination of elastic shear strain was found to be extremely important ( $\gamma^C$  is roughly 25%,  
532 50% and 66% of the total strain for experiments 41, 22 and 15, respectively), suggesting that a  
533 significant amount of pre-peak shear deformation in porous sandstones is attributable to elastic-  
534 plastic coupling. This observation also applies to experiments conducted at higher mean stresses  
535 which experienced no localization, although we do not include these results here.

536 This strain partitioning procedure was applied to eleven experiments which experienced  
537 localization (these are 3, 8, 13, 14, 15, 16, 17, 20, 22, 40, and 41 in Table 1 and Figure 2A and C)  
538 and two experiments at higher mean stress which exhibited no localization (these are 21 and 42  
539 in Table 1 and Figure 2B and D). The partitioned  $A$ ,  $B$ ,  $C$ , and  $D$  volume and shear strains so  
540 calculated, and in particular the inelastic strains ( $C+D$ ), are then used in determining constitutive  
541 parameters  $\beta$  and  $\mu$ , needed to evaluate localization theory predictions in the next section. Note  
542 that the triaxial experiments 13, 17, and 3 had no unload-reload cycles from which to determine  
543 the evolution of  $G$ , so we applied the  $G(\tau, \gamma)$  parameterization from 15 to these tests. As these  
544 were triaxial tests and as such cover a range of mean stress, this likely introduces some degree of  
545 uncertainty in strain partitioning for these experiments as  $G$  is a function of  $\sigma$  as well as  $\tau$ .

## 546 **6. Comparison with Localization Theory**

### 547 **6.1 Calculating $\mu$ : Yield Surfaces as Contours of Constant $^p\gamma$ and $^p\varepsilon$**

548 As discussed in Section 2, predictions of localization as per *Rudnicki and Rice* [1975]  
549 depend on values of the constitutive parameters  $\beta$ ,  $\mu$ , and hardening modulus  $h$ . *Holcomb and*

550 *Rudnicki* [2001] determined yield surfaces as contours of plastic shear strain,  $\gamma^p$ ; analogously,  
 551 we use contours of inelastic shear strain,  ${}^p\gamma$ , which accounts for elastic-plastic coupling. Yield  
 552 cap surfaces are determined as contours of inelastic volume strain,  ${}^p\varepsilon$ , in analogy with *Issen's*  
 553 [2002] use of plastic volume strain  $\varepsilon^p$ , but accounting for elastic-plastic coupling. In this section  
 554 we calculate  $\mu$  as local slopes of shear and cap yield surfaces at peak stress conditions, where  
 555 localization occurs.

556 For the shear surface, it is first necessary to determine a failure surface as the loci of peak  
 557 loading stress conditions. These are shown in Figure 10A (black squares) for five experiments  
 558 that displayed localization. A non-linear mean stress-dependent failure surface suggested by  
 559 *Brannon et al.* [2009] fits the peak stress data to an excellent degree, as shown by the dashed line  
 560 in Figure 10A (corresponding to  $\{a_1, a_2, a_3, a_4\}$  equal to  $\{48, 0.03, 46, 0.096\}$ ):

$$561 \quad \tau = a_1 - a_3 e^{-a_2 \sigma} + a_4 \sigma \quad (20)$$

562 This implies a value of  $\mu$  equal to

$$563 \quad \mu = \frac{\partial \tau}{\partial \sigma} = a_2 a_3 e^{-a_2 \sigma} + a_4 \quad (21)$$

564 For shear yield surfaces that apply to localization (i.e., at peak stresses) we are looking for  
 565 contours of constant  ${}^p\gamma$  that asymptote at the peak stress for the experiment of interest, in order to  
 566 calculate slopes and  $\mu$  for that experiment. Thus we apply the same form of (20) but fit contours  
 567 of  ${}^p\gamma$  that apply to the value of  ${}^p\gamma$  at failure (localization). Results for four experiments (only  
 568 results for 41, 22, 20, and 15 are shown for clarity in plotting) are shown in Figure 10A.  
 569 Corresponding values of  $\mu$  found from the curve fits and (21) are given in Table 2.

570 Similarly, for cap surfaces, we seek an elliptical-like equation to fit as contours of constant  
 571 inelastic volume strain  ${}^p\varepsilon$ . We follow the form used by *Brannon et al.* [2009]:

572 
$$\tau = \tau_f \left\{ 1 - \frac{\sigma - \sigma_f}{\sigma_h - \sigma_f} \right\}^{\frac{1}{2}} \quad (22)$$

573 where  $\sigma_f$  and  $\tau_f$  are equal to mean stress and shear stress values at the intersection of the cap  
 574 surface with the failure surface and  $\sigma_h$  is the mean stress where the cap intersects the hydrostat.  
 575 To express the cap surface as a contour of constant  ${}^p\varepsilon$ , we apply (22) with  $\tau_f$  and  $\sigma_f$  as fit  
 576 parameters using values of  $\sigma_h$  determined experimentally. The results for 14 and 15, experiments  
 577 which failed at higher mean stresses via a low-angle compaction localization (i.e., as shown in  
 578 Figure 5B for 15), and 21, which experienced no localization (as shown in Figure 5C), are shown  
 579 in Figure 10B. The  $\mu$  value is calculated as the slope of (22) where the cap meets the failure  
 580 surface, found from (22) to be:

581 
$$\mu = \frac{\partial \tau}{\partial \sigma} = \frac{\tau_f}{2} \left\{ 1 - \frac{\sigma - \sigma_f}{\sigma_h - \sigma_f} \right\}^{-\frac{1}{2}} \left\{ \frac{1}{\sigma_h - \sigma_f} \right\} \quad (23)$$

582 Corresponding values of  $\mu$  for the cap yield surface are given in Table 2 for 14, 13, 15, 17,  
 583 and 21. Samples 13 and 17 experienced localization at nearly the same  ${}^p\varepsilon$  values as 15 (0.0050  
 584 and 0.0055 for 13 and 17, respectively, compared to 0.0044 for 15), so we apply the slope  
 585 determined from 15 to these experiments.

## 586 **6.2 Calculating the $\beta$ Constitutive Parameter**

587 The dilation coefficient is a function of inelastic increments of volume and shear strain as  
 588 per *Rudnicki and Rice* [1975] and *Rudnicki* [1984] and is given by  $\beta = -d{}^p\varepsilon/d{}^p\gamma$ . The  $\beta$  value is  
 589 found from local slope of inelastic strain  ${}^p\varepsilon$  versus  ${}^p\gamma$  plot at peak stress, approximately when  
 590 localization commences. Recall that inelastic strains as defined here include the C and D  
 591 components of strain. Plots of  ${}^p\varepsilon$  versus  ${}^p\gamma$  for experiments 40, 41, 22, 20, 14, 15, 17, 13, 3 and

592 21 (in order of increasing mean stress) are shown in Figure 11. Values of  $\beta$  at peak load are  
593 given in Table 2.

594 Experiments that experienced shear localization (40, 41, 22, and 20) show failure at  
595 progressively larger values of  ${}^p\gamma$ , with decreasing positive values of  $\beta$ , correlated to increasing  
596 mean stress. Experiments that experienced compaction localization have increasingly negative  
597 values of  $\beta$  that correlate with increasing values of  ${}^p\varepsilon$ . Sample 21, which did not experience  
598 localization, is a constant  $\sigma$  test conducted at higher values of  $\sigma$  than the other tests, and so  
599 experienced more  ${}^p\varepsilon$  along the initial hydrostatic loading portion. This is why that experiment  
600 shows a higher  ${}^p\varepsilon$  value at the zero- ${}^p\gamma$  intercept than the other tests. Sample 3 is a triaxial  
601 experiment that localized at the highest  $\sigma$  of any of the other experiments that exhibited  
602 compaction localization. Between the  $\sigma$  values for samples 3 and 21 (between 80 and 90 MPa)  
603 there is a transition in behavior from compaction localization to pervasive compaction. Between  
604 the  $\sigma$  values for samples 20 and 14 (between 30 and 50 MPa), there is a transition in behavior  
605 from shear band localization to compaction localization.

### 606 **6.3 Predicted Band Angles and Hardening Moduli $h_{cr}$**

607 Results of  $\beta$  and  $\mu$  calculated for the eleven localizing samples of Figure 10A and B and Figure  
608 11, plus the non-localizing experiment 21 for comparison, are given in Table 2, along with the  
609 shear modulus and calculated Poisson's ratio (found from  $\nu = (3K - 2G) / [2(3K + G)]$ ), both  
610 determined just pre-peak. These values of  $\beta$  and  $\mu$  are, to our knowledge, the first to be  
611 determined using the full increment of inelastic strains and allowing for elastic-plastic coupling.

612 Critical hardening moduli for the shear surface, equations (5) and (6), and for the cap  
613 surface (where  $k_{cr} = h_{cr} / (\beta\mu)$  after *Issen* [2002]), using equations (5) and (6)), are given in  
614 Table 2 for the localizing experiments, along with the predicted band angles from equation (4).

## 615 **7. Discussion**

616 Observed angles of shear banding and compaction localization are given in Table 2 and  
617 can be compared to predicted band angles from the Rudnicki and Rice (1975) formalism, also  
618 shown in Table 2. The two uniaxial tests (40 and 41) formed shear bands at angles of  $67^\circ - 70^\circ$   
619 (Table 2). The overall specimen response was dilatant ( $\beta \approx 0.92$  and  $1.01$  for 41 and 40  
620 respectively) and the predicted band angles for both tests are  $73 - 76^\circ$ . The two constant mean  
621 stress tests at low mean stress (specimens 20 and 22) formed shear bands at angles of  $30^\circ - 48^\circ$ .  
622 The response was still dilatant ( $\beta \approx 0.17$  and  $0.26$  for 20 and 22 respectively), but less so than the  
623 uniaxial tests, and the predicted band angles are  $49^\circ$  and  $50^\circ$ . Thus, at low mean stresses, the  
624 predicted shear band angle demonstrates reasonable agreement with observed band angles,  
625 although predicted band angles are consistently slightly larger than observed angles. However,  
626 the predicted critical hardening modulus is strongly negative, in contrast with the observation of  
627 localization at/near peak stress, where  $h \approx 0$ . Predicted band angles for compactant experiments  
628 14, 13, 15, 17, and 21, with negative values of  $\beta$  and values of  $\mu$  calculated using shear yield  
629 surfaces as contours of constant  ${}^p\gamma$ , consistently over-predict band angles ( $35-40^\circ$  compared to 4-  
630  $25^\circ$  as observed). Band angles predicted using the cap surfaces as contours of constant  ${}^p\varepsilon$ , shown  
631 in the lower portion of have overall better agreement ( $10-23^\circ$  calculated compared to 4- $25^\circ$   
632 observed). Unlike the shear band tests, for compaction localization, the predicted critical  
633 hardening moduli are all positive, but this still contrasts with the observation of localization

634 at/near peak stress, where  $h \approx 0$ . There is one transitional specimen (14), at 50 MPa mean stress,  
635 that exhibited overall compactant response ( $\beta = -0.23$ ) for a predicted band angle of  $40^\circ$ . This  
636 specimen formed a low angle band,  $8^\circ - 12^\circ$ , even though the overall response was less  
637 compactant than the experiments exhibiting compaction localization.

638 As expected and as observed by others, at higher mean stresses beyond a threshold value,  
639 no localization occurs. This threshold is approximately 90 MPa for the Castlegate specimens in  
640 this study. At this mean stress, the overall response was compactant, with  $\beta$  equal to  $-0.51$ , but  
641 this value is much less than experiments at slightly lower mean stresses (70-79 MPa) that did  
642 exhibit compaction localization. Therefore, there appears to be a relatively small range of  
643 intermediate mean stresses for which the material response becomes strongly compactant leading  
644 to compaction localization. Below this small range, shear bands form and above it no  
645 localization occurs.

## 646 **8. Summary and Conclusions**

647 A suite of axisymmetric compression tests ( $\sigma_1 > \sigma_2 = \sigma_3$ ) were presented over a wide  
648 range of mean stresses and deformational behaviors, for Castlegate sandstone, an analog for  
649 porous sandstone reservoirs for oil production or subsurface carbon storage. Results were used  
650 to develop a complete set of constitutive parameters, including stress and plastic strain  
651 dependence of the elastic moduli. Acoustic emissions (AE) events were utilized to identify  
652 strain localization. Constitutive parameters and AE results permitted evaluation of theoretical  
653 predictions for strain localization.

654 At low mean stress, shear localization occurred, while at high mean stress, homogeneous  
655 compaction without localization (pore collapse) was observed. At a small range of intermediate  
656 mean stresses, compaction localization develops. For localization modes, acoustic emissions

657 events confirmed that band formation generally begins at/near the peak in the axial stress – axial  
658 strain curve and enabled determination of localization band angle. For shear bands, the observed  
659 band angles (angle between the band normal and the direction of maximum compression) were  
660  $30^\circ - 70^\circ$ , and decreased with increasing mean stress. For intermediate stress tests, compaction  
661 localization occurred. Previously this localization mode was reported as compaction bands  
662 (theoretically, a pure compaction band angle is  $0^\circ$ ). However, the more detailed AE analyses  
663 performed in this work reveals that these features are actually low angle bands,  $4^\circ - 25^\circ$ ; the lack  
664 of shear offset suggests that they are dominantly compactant. These observed angles were used  
665 to evaluate theoretical predictions using the framework of *Rudnicki and Rice* [1975].

666 An important finding is the relative importance of elastic-plastic coupling via the plastic  
667 strain dependence of elastic moduli in Castlegate Sandstone elastic-plastic deformation. Bulk and  
668 shear moduli evolve as increasing functions of mean stress and Mises equivalent shear stress,  
669 respectively, and as decreasing functions of the work-conjugate plastic strains. Total strain  
670 increments were partitioned into four parts: A) elastic strain (constant moduli), B) strain due to  
671 stress dependence of elastic moduli, C) strain due to moduli degradation with increasing plastic  
672 strain (“coupling strain”), and D) plastic strain. The sum of the first three parts is the elastic  
673 strain increment; the plastic strain increment, part *D*, is total strain less elastic strain. The sum of  
674 the third and fourth parts (*C* and *D*) is referred to as the increment of inelastic strain, and was  
675 used in the calculation of constitutive parameters. The friction factor,  $\mu$ , is the slope of the yield  
676 surface, which is a contour of constant inelastic shear strain (for the cap,  $\mu$  is determined as a  
677 contour of constant inelastic volume strain). The dilation coefficient,  $\beta$ , is the negative of the  
678 slope of the inelastic volume strain – inelastic shear strain curve.

679 At low mean stresses, overall specimen response was dilatant ( $\beta > 0$ ) and the predicted  
680 range of band angles of  $40^\circ - 73^\circ$  are in good agreement with observed band angles of  $30^\circ - 70^\circ$ ,  
681 although predicted angles were consistently slightly higher than the observed. At intermediate  
682 mean stresses, the response was compactant ( $\beta \approx -0.5$ ) and observed band angles were  $4^\circ - 25^\circ$ .  
683 Using a shear yield surface calculated as a contour of inelastic shear strain, the yield surface  
684 slope is positive and predicted band angles were  $35^\circ - 40^\circ$ ; this procedure thus over-predicted  
685 band angles significantly. Alternatively, using a cap yield surface calculated as a contour of  
686 inelastic volume strain, the yield surface slope is negative and lower angle bands were predicted  
687 ( $10^\circ - 23^\circ$ ), in better agreement with observed values. At low mean stresses, the predicted critical  
688 hardening modulus is strongly negative (implying post-peak localization), while for intermediate  
689 mean stresses, it is positive (pre-peak localization); these both contrast with AE results, which  
690 show that localization occurs at/near peak stress where the hardening modulus is zero.

691 In conclusion, the *Rudnicki and Rice* [1975] theoretical treatments show a reasonable  
692 predictive ability with respect to Castlegate Sandstone deformation. Elastic-plastic coupling is an  
693 important consideration for constitutive behavior of weak porous sandstones, and can be a major  
694 component of elastic-plastic behavior in constitutive modeling for subsurface engineering  
695 endeavors.

696

## 697 **Acknowledgements**

698 The authors gratefully acknowledge funding from the U.S. Department of Energy, Office of  
699 Science, Basic Energy Sciences Program and the National Science Foundation Award EAR-  
700 0711346 to Clarkson University (for KI). We thank Tom Pfeifle and Joe Bishop for comments  
701 and their work on improving the manuscript, and Rebecca Brannon for championing our work on

702 elastic-plastic coupling. Sandia National Laboratories is a multi-program laboratory managed  
703 and operated by Sandia Corporation, a wholly owned subsidiary of Lockheed Martin  
704 Corporation, for the U.S. Department of Energy's National Nuclear Security Administration  
705 under contract DE-AC04-94AL85000.

706

## 707 **References**

708 Antonellini, M., and A. Aydin (1995), Effect of faulting on fluid flow in porous sandstones:  
709 geometry and spatial distribution, *American Assoc. Petrol. Geol.* 79, 642–671.

710 Aydin, A. (1978), Small faults formed as deformation bands in sandstone, *Pure and Applied*  
711 *Geophysics* 116, 913–930.

712 Aydin, A., and A. M. Johnson (1983), Analysis of faulting in porous sandstones, *J. Struct. Geol.*  
713 5, 19–31.

714 Aydin, A., R.I. Borja, and P. Eichhubl (2006), Geological and mathematical framework for  
715 failure modes in granular rock, *Journal of Structural Geology* 28, 83-98

716 Baud, P., E. Klein, and T.-f. Wong (2004), Compaction localization in porous sandstones: spatial  
717 evolution of damage and acoustic emission activity, *Journal of Structural Geology* 26, 603-  
718 624.

719 Baud, P., V. Vajdova, and T.-f. Wong (2006), Shear-enhanced compaction and strain  
720 localization: Inelastic deformation and constitutive modeling of four porous sandstones, *J.*  
721 *Geophys. Res.*, 111, B12401, doi:10.1029/2005JB004101.

722 Bernabe', Y., D. Fryer, and R. Shively (1994), Experimental observations of the elastic and  
723 inelastic behaviour of porous sandstones, *Geophys. J. Int.*, 117, 403-418.

724 Bésuelle, P. (2001), Compacting and dilating shear bands in porous rock: theoretical and  
725 experimental conditions, *J. Geophys. Res.*, 106, 13435–13442.

726 Bésuelle, P., J. Desrues, and S. Raynaud (2000), Experimental characterisation of the localisation  
727 phenomenon inside a Vosges sandstone in a triaxial cell, *Int. J. Rock Mech. Min. Sci.*, 37,  
728 1223-1237.

729 Bésuelle, P., P. Baud, and T.-f. Wong (2003), Failure mode and spatial distribution of damage in  
730 Rothbach Sandstone in the brittle-ductile transition, *Pure Applied Geophysics* 160, 851-  
731 868.

732 Borja, R. (2006), Conditions for instabilities in collapsible solids including volume implosion  
733 and compaction banding, *Acta Geotechnica*, 1:107-122, DOI 10.1007/s11440-006-0012-x.

734 Borja, R., and A. Aydin (2004), Computational modeling of deformation bands in granular  
735 media. I. Geological and mathematical framework, *Comput. Methods Appl. Mech. Engrg.*  
736 193, 2667-2698.

- 737 Borja, R. (2004), Computational modeling of deformation bands in granular media. II:  
738 Numerical simulations, *Comput. Methods Appl. Mech. Engrg.* 193, 2699-2718.
- 739 Brannon, R.A., A.F. Fossum, and O.E. Strack (2009) KAYENTA: Theory and User's Guide.  
740 Sandia National Laboratories, SAND2009-2282. Albuquerque, NM.
- 741 Collins, I., and G. Houlsby (1997), Application of thermomechanical principles to the modeling  
742 of geotechnical materials, *Proc. R. Soc. Lond. A*, 453, 1975-2001.
- 743 Collins, I. (2002), Associated and non-associated aspects of the constitutive laws for coupled  
744 elastic/plastic materials. *The Int. J. of Geomechanics* 2, 259-267.
- 745 Coop, M., and S. Willson (2003), Behavior of hydrocarbon reservoir sands and sandstones, *J.*  
746 *Geotechnical and Geoenvironmental Eng.* 129, 1010-1019.
- 747 De Waal, J.A. and R.M.M. Smits (1985), Prediction of reservoir compaction and surface  
748 subsidence: Field application of a new model, *SPE Ann. Tech. Conf. & Exhib.*, 14214, Las  
749 Vegas, NV.
- 750 Dewers, T.A., Newell, P., Broome, S., Heath, J., and S. Bauer (2013), Geomechanical behavior  
751 of Cambrian Mount Simon Sandstone reservoir lithofacies, *Int. J. Greenhouse Gas Control*,  
752 in press.
- 753 Detournay, C., P. Cundall, and J. Parra (2003), A study of compaction band formation with the  
754 double yield model, in *FLAC and Numerical Modeling in Geomechanics: Proceedings of*  
755 *the Third International FLAC Symposium*, edited by B. et Al., pp. 27-33, Swets &  
756 Zeitlinger.
- 757 DiGiovanni, A., J. Fredrich, D.J. Holcomb, and W. A. Olsson (2000), Micromechanics of  
758 compaction in an analogue reservoir sandstone, in *Proc. North American Rock Mech.*  
759 *Symp.*, edited by Girard, J., Liebman, M., Breeds, C. , and T. Doe, Balkema, pp. 1153-  
760 1158, July 13.
- 761 DiMaggio, F.L. and I.S. Sandler (1971), Material model for granular soils, *J. Eng. Mech. Div.*,  
762 *ASCE*, 97, 935-950.
- 763 Du Bernard, X., P. Eichhubl, and A. Aydin (2002), Dilation bands: A new form of localized  
764 failure in granular media, *Geophys. Res. Lett.*, 29(24), 2176, doi:10.1029/2002GL015966.
- 765 Eichhubl, P., J.N. Hooker, and S.E. Laubach (2010), Pure and shear-enhanced compaction bands  
766 in Aztec Sandstone, *Journal of Structural Geology* doi:10.1016/j.jsg.2010.02.004.
- 767 Gajo, A., and D. Bigoni (2008), A model for stress and plastic strain induced nonlinear,  
768 hyperelastic anisotropy in soils, *Int. J. Numerical Analytical Methods in Geomechanics* 32,  
769 833-861, doi:10.1002/nag.648
- 770 Gajo, A., D. Muir Wood, and D. Bigoni (2007), On certain critical material and testing  
771 characteristics affective shear band development in sand, *Geotechnique* 57, 449-461.
- 772 Haimson, B., and H. Lee (2004), Borehole breakouts and compaction bands in two high-porosity  
773 sandstones, *Int. J. Rock Mech. & Min. Sci.*, 41, 287-301.
- 774 Haimson, B. C. (2001), Fracture-like borehole breakouts in high-porosity sandstone: are they  
775 caused by compaction bands?, *Phys. Chem. Earth (A)*, 26, 15-20.

- 776 Haimson, B. C. (2003), Borehole breakouts in Berea sandstone reveal a new fracture mechanism,  
777 *Pure Appl. Geophys.*, 160, 813–831.
- 778 Hill, R.E. (1989), Analysis of Deformation Bands in the Aztec Sandstone, Valley of Fire State  
779 Park, Nevada. M.S. thesis, University of Nevada, Las Vegas.
- 780 Hill, R. and J.R. Rice (1973), Elastic potentials and the structure of inelastic constitutive laws,  
781 *SIAM J. Appl. Math.* 25, 448-461.
- 782 Holcomb, D. J., and W. A. Olsson (2003), Compaction localization and fluid flow, *J. Geophys.*  
783 *Res.*, 108, 2290, doi:10.1029/2001JB000813.
- 784 Holcomb, D., and J. Rudnicki (2001), Inelastic constitutive properties and shear localization in  
785 Tennessee marble, *Int. J. Numerical Analytical Methods Geomechanics*, 25, 109-129.
- 786 Holcomb, D., W. Olsson, and J. Rudnicki (2004), Compaction localization: effects on  
787 permeability and reservoir mechanics, paper presented at AAPG Hedberg Conference,  
788 Austin, TX.
- 789 Holcomb, D., J. Rudnicki, K. Issen, and K. Sternlof (2007), Compaction localization in the earth  
790 and the laboratory: state of the research and research directions, *Acta Geotechnica* 2, 1-15.
- 791 Ingraham, M.D., Issen, K.A., and Holcomb, D. (2013a), Use of acoustic emissions to investigate  
792 localization in high-porosity sandstone subjected to true triaxial stress, *Acta*  
793 *Geotechnica* 8, 645-663.
- 794 Ingraham, M.D., Issen, K.A., and Holcomb, D. (2013b), Response of Castlegate sandstone to  
795 true triaxial states of stress, *J. Geophys. Res.*, 118, 536-552, doi:10.1002/jgrb.50084.
- 796 Issen, K.A. (2002), The influence of constitutive models on localization conditions for porous  
797 rock, *Eng. Fracture Mech.*, 69, 1891-1906.
- 798 Issen, K.A. and Challa, V. (2008), Influence of the intermediate principal stress on the strain  
799 localization mode in porous sandstone, *Journal of Geophysical Research*,  
800 113(B02103): doi:10.1029/2005JB004008.
- 801 Issen, K.A., and J.W. Rudnicki (2000), Conditions for compaction bands in porous rock, *J.*  
802 *Geophys. Res.*, 105, 21529-21536.
- 803 Issen, K.A., and J.W. Rudnicki (2001), Theory of compaction bands in porous rock, *Phys. Chem.*  
804 *Earth (A)*, 26(1-2), 95-100.
- 805 Karner, S., J. Chester, F. Chester, A. Kronenberg, and A. Hajash (2005), Laboratory deformation  
806 of granular quartz sand: Implications for the burial of clastic rocks, *Am. Assoc. Petrol.*  
807 *Geologists Bull.* 89 (5), 603-625.
- 808 Karner, S., F. Chester, A. Kronenberg, and J. Chester (2003), Subcritical compaction and  
809 yielding of granular quartz sand, *Tectonophysics*, 377, 357-381.
- 810 Khan, A., Y. Xiang, and S. Huang (1991), Behavior of Berea Sandstone under confining  
811 pressure Part I: Yield and failure surfaces, and nonlinear elastic response, *Int. J.*  
812 *Plasticity*, 7, 607-624.
- 813 Klein, E., P. Baud, T. Reuschle, and T-f. Wong (2001), Mechanical behaviour and failure mode  
814 of Bentheim sandstone under triaxial compression, *Phys. Chem. Earth (A)*, 26, 21–25.

- 815 McLaurin, B.T. and R.J. Steel (2007), Architecture and origin of an amalgamated fluvial sheet  
816 sand, lower Castlegate Formation, Book Cliffs, Utah, *Sedimentary Geology*, 197, 291-311.
- 817 Mollema, P. N., and M. A. Antonellini (1996), Compaction bands: A structural analog for anti-  
818 mode I cracks in aeolian sandstone, *Tectonophysics*, 267, 209–228.
- 819 Muhuri, S.K. (1994), Experimental and Microscopic Study of the Transition From Brittle to  
820 Ductile Behavior in the Porous Berea Sandstone as a Function of Confining Pressure. M.S.  
821 thesis, University of Oklahoma, Norman, Oklahoma.
- 822 Olsson, W. A. (2001), Quasistatic propagation of compaction fronts in porous rock, *Mech. Mat.*,  
823 33, 659-668.
- 824 Olsson, W. A. (1999), Theoretical and experimental investigation of compaction bands in porous  
825 rock, *J. Geophys. Res.*, 104, 7219-7228.
- 826 Olsson, W. A., and D. J. Holcomb (2000), Compaction localization in porous rock, *Geophys.*  
827 *Res. Letters*, 27, 3537–3540.
- 828 Olsson, W., D.J. Holcomb, and J. Rudnicki (2002), Compaction localization in porous sandstone:  
829 Implications for Reservoir Mechanics, *Oil and Gas Science and Technology, Rev. IFP*  
830 57, 591-599.
- 831 Rudnicki J.W. (1984), A class of elastic-plastic constitutive laws for brittle rock, *Journal of*  
832 *Rheology*, 28(6), 759-778.
- 833 Rudnicki, J. (2004), Shear and compaction band formation on an elliptic yield cap, *J.*  
834 *Geophysical Research*, 109, B03402.
- 835 Rudnicki, J. (2007), Extension of compaction bands in porous sandstones. Euro-conference of  
836 Rock Physics and Geomechanics on Natural Hazards: Thermo-Hydro-Mechanical  
837 Processes in Rocks, Erice, Italy.
- 838 Rudnicki, J.W., and J.R. Rice (1975). Conditions for the localization of deformation in pressure  
839 sensitive dilatant materials. *J. Mech. Phys. Solids*, 23, 371-394.
- 840 Rudnicki, J. W., and W.A. Olsson (1998), Reexamination of Fault Angles Predicted by Shear  
841 Localization Theory, *Int. J. Rock Mech. Min. Sci.* 35, 512-513.
- 842 Rutqvist J., D. Vasco, and L. Myer (2009), Coupled reservoir-geomechanical analysis of CO2  
843 injection and ground deformations at In Salah, Algeria, *Int. J. Greenhouse Gas Control*,  
844 doi:10.1016/j.ijggc.2009.10.017.
- 845 Stanchits, S., J. Fortin, Y. Gueguen, and G. Dresen (2007), Initiation and propagation of  
846 compaction bands in dry and wet sandstone. Euro-conference of Rock Physics and  
847 Geomechanics on Natural Hazards: Thermo-Hydro-Mechanical Processes in Rocks, Erice,  
848 Italy.
- 849 Sternlof, K.R., J.R. Chapin, D.D. Pollard, and L.J. Durlofsky (2004), Permeability effects of  
850 deformation band arrays in sandstone: *AAPG Bulletin*, 88, 1315-1329.
- 851 Sulem, J., and H. Ouffroukh (2006), Hydromechanical behaviour of Fountainebleau Sandstone,  
852 *Rock Mechanics Rock Engineering*, 39 (3), 185-213.

- 853 Sulem, J., I. Vardoulakis, E. Papamichos, A. Oulahna, and J. Tronvoll (1999), Elasto-plastic  
854 modeling of Red Wildmoor sandstone, *Mechanics of Cohesive-Frictional Materials*, 4,  
855 215-245.
- 856 Tembe, S., V. Vajdova, T.-f. Wong, and W. Zhu (2006), Initiation and propagation of strain  
857 localization in circumferentially notched samples of two porous sandstones, *J. Geophys. Res.*,  
858 111, B02409, doi:10.1029/2005JB003611.
- 859 Tembe, S., Vajdova, V., Baud, P., Zhu, W., and Wong, T.-f. (1997), A new methodology to  
860 delineate the compactive yield cap for two porous sandstones under undrained condition.  
861 *Mechanics of Materials*, 39, 513-523.
- 862 Vajdova, V., P. Baud, P., and T.-f. Wong (2004), Permeability evolution during localized  
863 deformation in Bentheim sandstone, *J. Geophys. Res.*, 109, B10406,  
864 doi:10.1029/2003JB002942.
- 865 Vinegar, H., J. DeWaal, and S. Wellington (1991), CT studies of brittle failure in Castlegate  
866 Sandstone. *Int. J. Rock Mechanics Mining Sciences & Geomech. Abstracts* 28, 441-448.
- 867 Wawersik, W. R., J. W. Rudnicki, P. Dove, J. Harris, J. M. Logan, L. Pyrak-Nolte, F. M. Orr, P.  
868 J. Ortoleva, F. Richter, N. R. Warpinski, J. L. Wilson, and T.-f. Wong (2000), Terrestrial  
869 sequestration of CO<sub>2</sub>: An assessment of research needs, *Advances in Geophysics*, 43, 97-  
870 177.
- 871 Wong, T.-f., H. Szeto, and J. Zhang (1992), Effect of loading path and porosity on the failure  
872 mode of porous rocks. *Appl. Mech. Rev.*, 45(8), 281-293.
- 873 Wong, T.-f., C. David, and W. Zhu (1997), The transition from brittle faulting to cataclastic flow  
874 in porous sandstones: Mechanical deformation. *J. Geophysical Research*, 102, 3009-3025.
- 875 Wong, T.-f., Baud, P., and E. Klein (2001), Localized failure modes in a compactant porous rock,  
876 *Geophys. Res. Lett.*, 28, 2521–2524.
- 877 Zhang, J., T.-f. Wong, and D. Davis (1990), Micromechanics of pressure-induced grain crushing  
878 in porous rocks, *Journal of Geophysical Research*, 95, 341-352.
- 879

## 880 **Figure Captions**

881 **Figure 1.** Idealized yield surface in space of Mises equivalent shear stress  $\tau$  and mean stress  $\sigma$ .  
882 Also delineated is a portion of the plastic potential, and definitions of the constitutive parameters  
883  $\beta$  and  $\mu$ . The vertically oriented arrow normal to the plastic potential surface refers to the  
884 direction of inelastic increment of strain as defined in the text. The dashed arrow shows the  
885 loading path in stress space.

886 **Figure 2. A and B.** Axial and radial stress-strain pathways shown for Castlegate experiments  
887 exhibiting localization (**A**) and no localization (**B**). Curves on the right hand side of plots  
888 correspond to axial strains; those on the left correspond to radial strains. **C and D.** Loading  
889 pathways in  $\sigma - \tau$  space are shown for uniaxial, triaxial, and constant mean stress experiments  
890 for samples with localization (**C**) and no localization (**D**).

891 **Figure 3.** Castlegate axial stress-strain curves with acoustic emissions for **A.** uniaxial loading  
892 path resulting in shear localization (experiment 41); **B.** constant mean stress loading (experiment  
893 22) resulting in shear localization; **C.** constant mean stress loading (experiment 15) resulting in  
894 compaction localization; and **D.** constant mean stress loading path (experiment 21) resulting in  
895 spatially pervasive pore collapse.

896 **Figure 4.** Acoustic Emissions (AE) results for specimen 22 (shear localization). The lower panel  
897 shows axial stress (thin black line) and calculated band angle (thick gray line) vs. axial strain. In  
898 the upper panel, acoustic emissions events (black and gray dots) are shown for three portions of  
899 the test indicated by the arrows. Only the black dot events were used to calculate the band angle.

900 **Figure 5.** AE results for specimen 15 (compaction localization). The lower panel shows axial  
901 stress (thin black line) and calculated band angle (thick gray line) vs. axial strain. In the upper  
902 panel, acoustic emissions events (black and gray dots) are shown for three portions of the test  
903 indicated by the arrows. Only the black dot events were used to calculate the band angle.

904 **Figure 6.** AE results for specimen 21 (no localization). The lower panel shows axial stress (thin  
905 black line) and calculated band angle (thick gray line) vs. axial strain. In the upper panel,  
906 acoustic emissions events (black and gray dots) are shown for three portions of the test indicated  
907 by the arrows.

908 **Figure 7. A.** Mean stress-volume strain path for hydrostatic compression experiment 38,  
909 including four unloading loops and final unloading. **B.** Observed (symbols) and model prediction  
910 (solid line) of tangent bulk modulus evolution for experiment 38, including both mean stress and  
911 plastic volume strain dependence. The solid grey line is the calculated secant bulk modulus. **C.**  
912 Partitioning of total volume strain in **A** (solid black line) into elastic (solid dark grey line) and  
913 plastic (dashed black line) portions. The light grey curve shows the elastic response ignoring  
914 elastic-plastic coupling. Nearly half of the total strain arises from elastic-plastic coupling, i.e., the  
915 plastic volume strain dependence of the bulk modulus. **D.** Evolution of the  $K_0$  parameter for the  
916 bulk modulus model, used to quantify the modulus degradation due to plastic volume strain  
917 dependence.

918 **Figure 8. A.** Measured shear stress-strain plot for Castlegate sandstone (experiment 42)  
919 deformed at a constant mean stress of 140 MPa. **B.** Observed tangent shear moduli (symbols)  
920 and model prediction (solid line) of shear modulus evolution for experiment 42, including stress  
921 and plastic shear strain-dependence. **C.** Partitioning of total measured shear strain in **A** (solid  
922 black line) into elastic (solid dark grey line) and plastic (dashed black line) portions. The light  
923 grey curve shows the elastic response ignoring elastic-plastic coupling. In this case, about 20%  
924 of the total strain arises from elastic-plastic coupling, i.e., the plastic shear strain dependence of  
925 the shear modulus. **D.** Evolution of the  $G_0$  parameter for the shear modulus model, used to  
926 quantify the modulus degradation due to plastic shear strain dependence.

927 **Figure 9.** Partitioning of total measured strain (black curves) into elastic (dark grey line) and  
928 plastic (dashed black line) portions for three Castlegate experiments, shown as mean stress vs.  
929 volume strain, and shear stress vs. shear strain: **A** and **B.** Experiment 41; **C.** and **D.** Experiment  
930 22; **E.** & **F.**, Experiment 15. The light grey curve in all cases shows the calculated elastic strain if  
931 elastic-plastic coupling is ignored.

932 **Figure 10. A.**  $\tau$  -  $\sigma$  plots contoured for constant inelastic shear strain. This includes both total  
933 plastic strain and that due to elastic plastic coupling and is equal to the sum of the inelastic  
934 increment of shear strain. Plotted are data from experiments 41, 22, 20, 15, and 21. Inelastic  
935 shear strain contour values are equal to the value at peak for each individual experiment. **B.**  $\sigma$  -  $\tau$   
936 plots contoured for inelastic volume strain (sum of inelastic increments of volume strain, which  
937 includes plastic volume strain and that due to elastic plastic coupling). Includes data from  
938 experiments 14, 15, and 21. Also shown as dashed line in both **A** and **B** is the failure surface: the  
939 locus of peak yield stresses for experiments that localized are shown as solid squares in **A**.

940

941 **Figure 11. A.** Plot of inelastic volume strain  ${}^p\varepsilon$  versus inelastic shear strain  ${}^p\gamma$ , for all  
942 axisymmetric triaxial compression experiments performed with Castlegate sandstone. The  
943 constitutive parameter  $\beta$  is determined as the negative of the local slopes of these curves.  
944 Experiments that resulted in shear localization have positive values of  $\beta$ , while those that  
945 experienced compaction localization have negative  $\beta$ .

946

947

948

949 **Tables**

950 **Table 1.** Summary of experiments, including loading configuration, stress conditions, peak  
 951 stresses, and failure mode

Experiment	Loading type	$\sigma$	Peak $\sigma_3^1$	Peak $\sigma_1^1$	Maximum $\sigma^1$	Maximum $\tau^1$	Localization mode
38	hydrostatic	-	400	400	400	-	
41	uniaxial	0	0	14.5	4.9	8.4	shear band
40	uniaxial	0	0	14.4	4.8	8.4	shear band
22	constant $\sigma$	25	8.5	56	25	28	shear band
20	constant $\sigma$	30	11	68	30	32	shear band
14	constant $\sigma$	50	26	98	50	42	low-angle shear b.
15	constant $\sigma$	70	41	127	70	52	compactional band
21	constant $\sigma$	90	60	150	90	56	no banding
23	constant $\sigma$	110	83	160	110	58	no banding
42	constant $\sigma$	150	110	233	150	72	no banding
44	constant $\sigma$	185	143	266	185	73	no banding
43	constant $\sigma$	195	150	284	195	77	no banding
13	triaxial	45	45	123	71	45	compactional band
17	triaxial	45	45	121	70	43	compactional band
8	triaxial	45	45	132	74	50	compactional band
3	triaxial	50	50	137	79	51	compactional band
16	triaxial	45	45	125	71	46	compactional band

Notes:  
<sup>1</sup> Mpa

952  
 953  
 954  
 955  
 956  
 957

958  
959

**Table 2.** Summary of pre-peak constitutive parameters and calculated angles

<b>Shear Surface</b>							
Exper#	$\beta$	$\mu$	$\nu$	$G^1$	$h_{cr}/G$	$\theta$ , predicted <sup>2</sup>	$\theta$ , observed <sup>2</sup>
41	0.92	1.23	0.23	1267	-1.01	73	67-70
40	1.01	1.23	0.23	1635	-1.07	76	67-69
22	0.17	0.68	0.16	2410	-0.39	49	30-35
20	0.26	0.63	0.15	4885	-0.42	50	45-48
14	-0.23	0.35	0.17	5273	-0.17	40	8-12
13	-0.58	0.26	0.20	4424	-0.22	35	10-20
15	-0.41	0.26	0.18	4598	-0.32	37	4-8; 5-25
17	-0.59	0.26	0.21	4528	-0.21	35	5-8; 19-20
21	-0.51	0.20	0.18	5912	-0.24	35	no local.
<b>Cap Surface</b>							
Exper#	$\beta$	$\mu$	$\nu$	$G^1$	$k_{cr}/G$	$\theta$ , predicted <sup>2</sup>	$\theta$ , observed <sup>2</sup>
14	-0.23	-0.82	0.17	5273	0.131	23	8-12
13	-0.58	-1.02	0.20	4424	0.051	10	10-20
15	-0.41	-1.02	0.18	4598	0.108	15	4-8; 5-25
17	-0.59	-1.02	0.21	4528	0.049	10	5-8; 19-20
21	-0.51	-0.88	0.18	5912	0.007	16	no local.
Notes:							
<sup>1</sup> MPa							
<sup>2</sup> degrees							

960

Figure 1

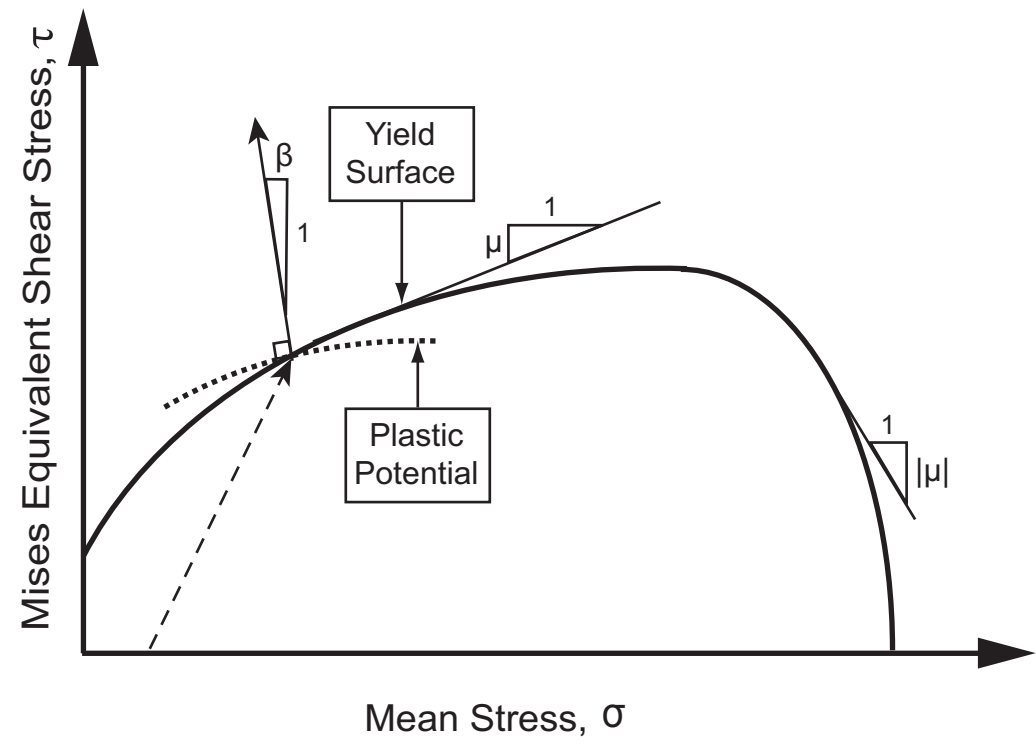


Figure 2

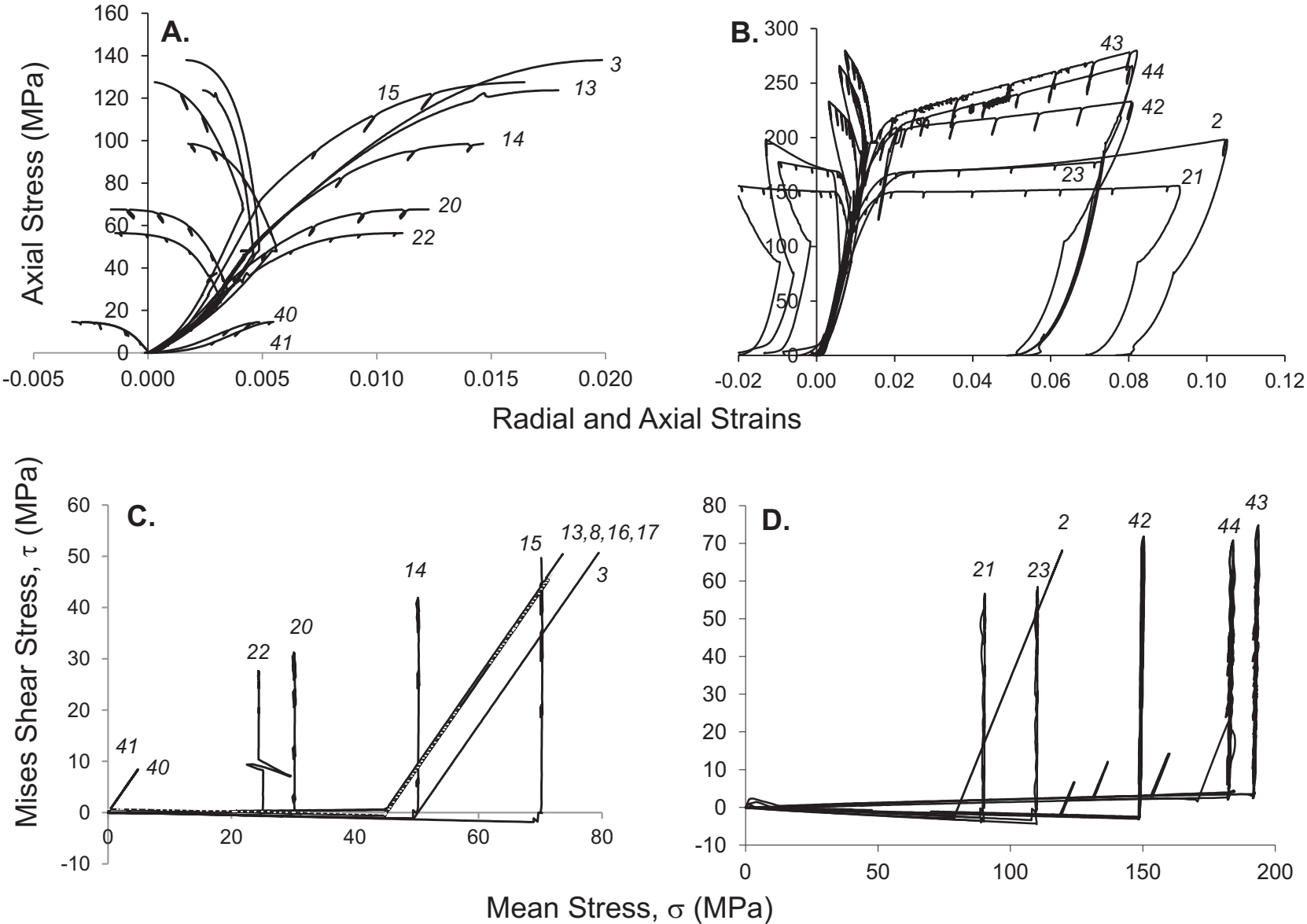


Figure 3

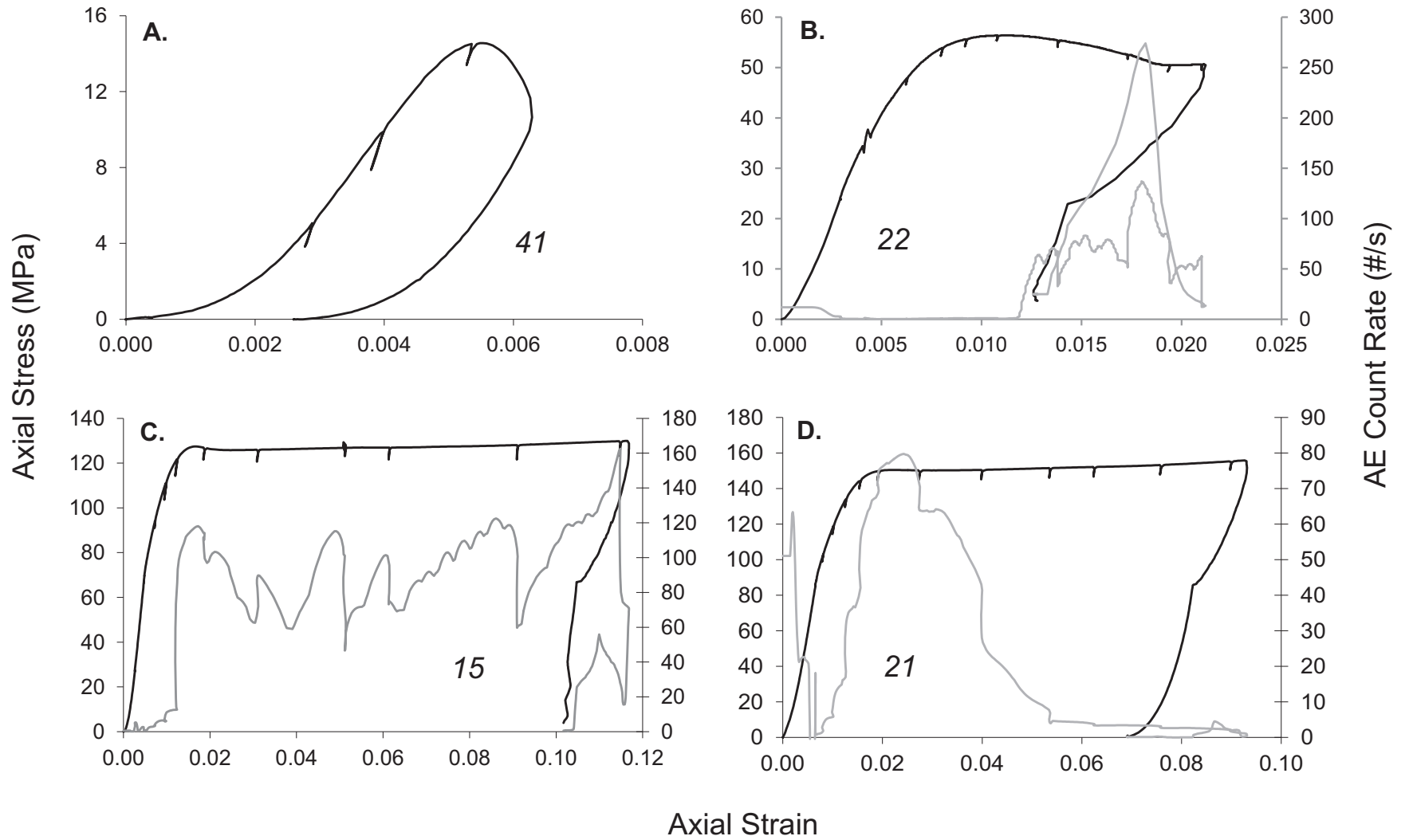


Figure 4

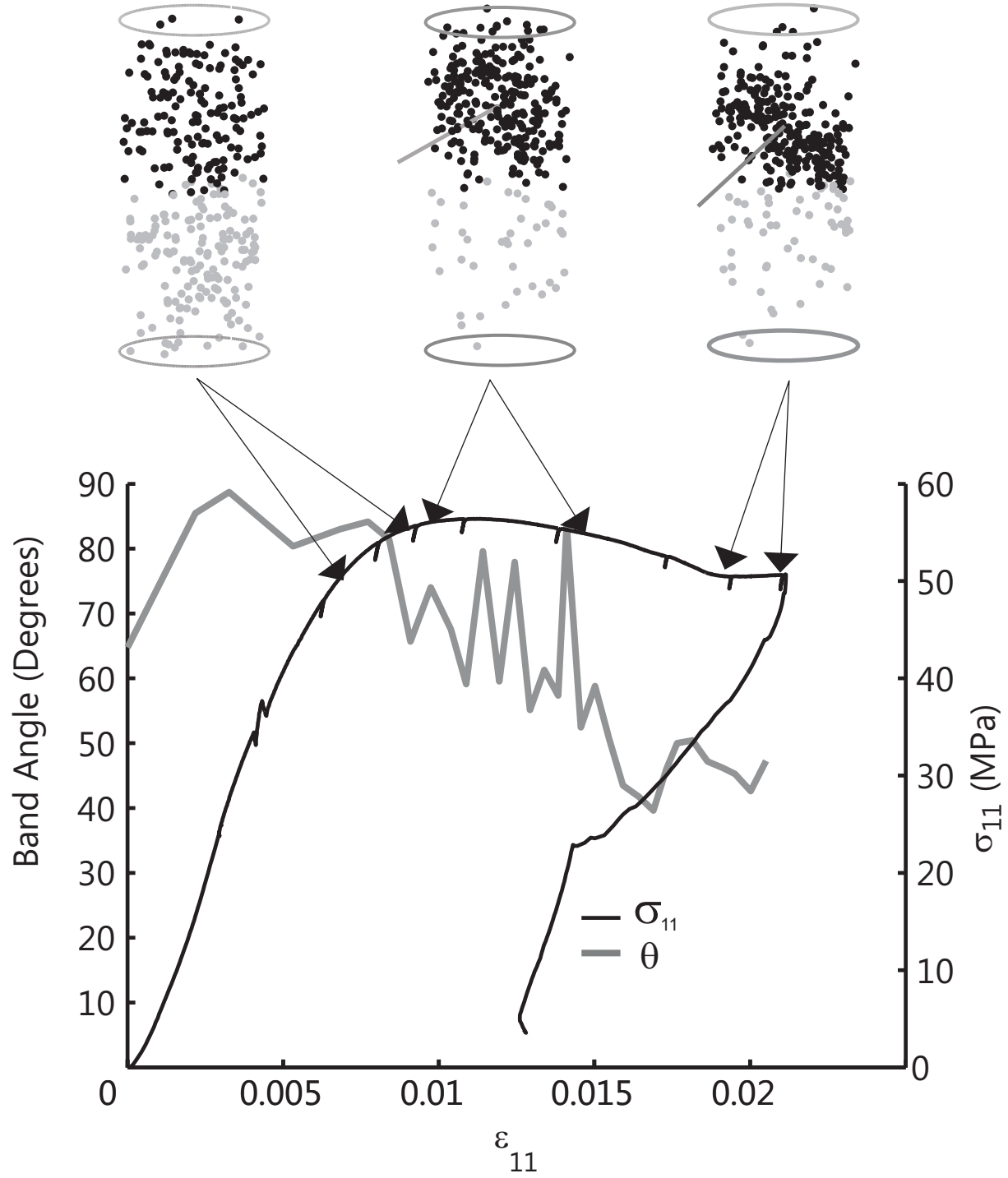


Figure 5

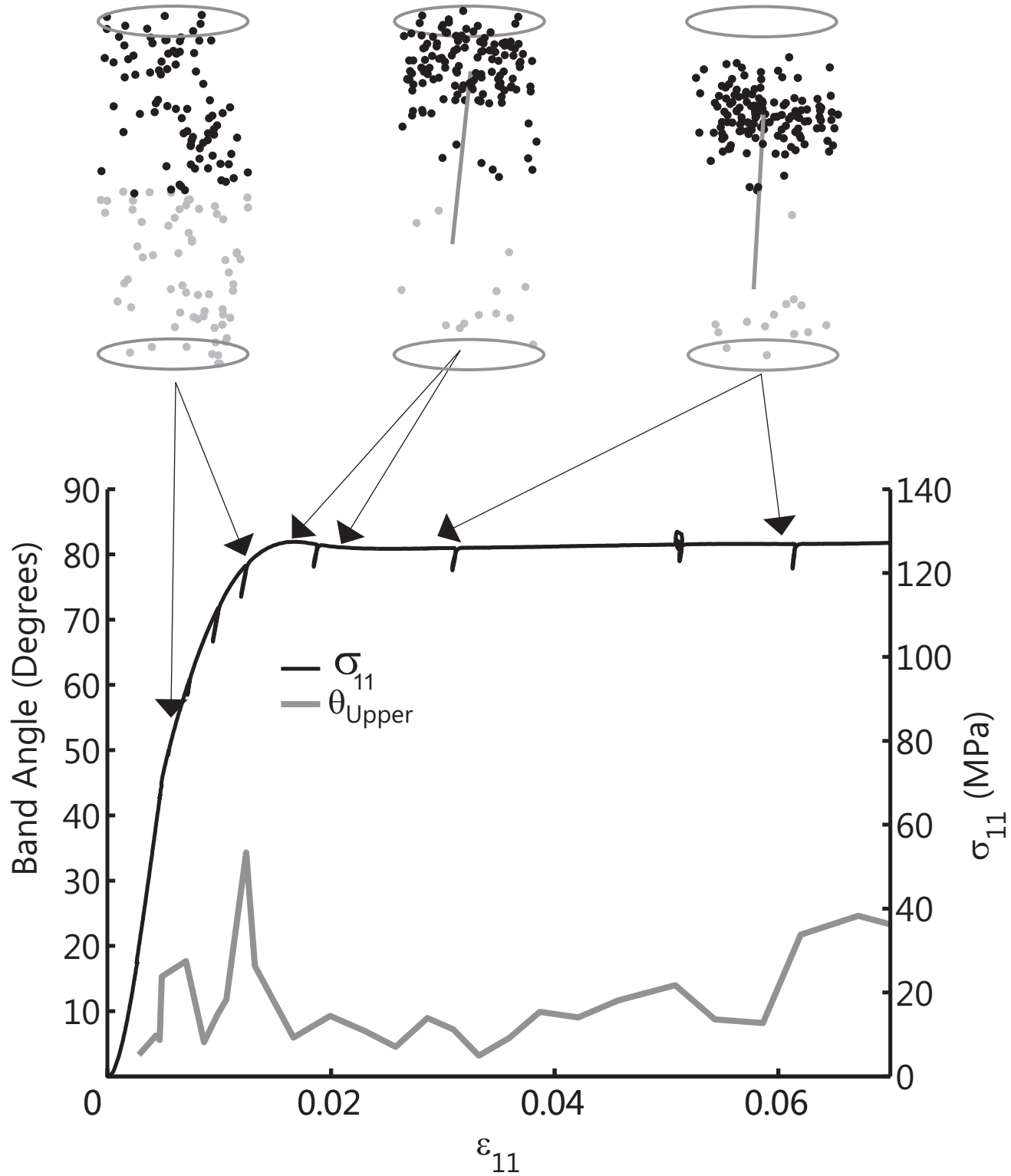


Figure 6

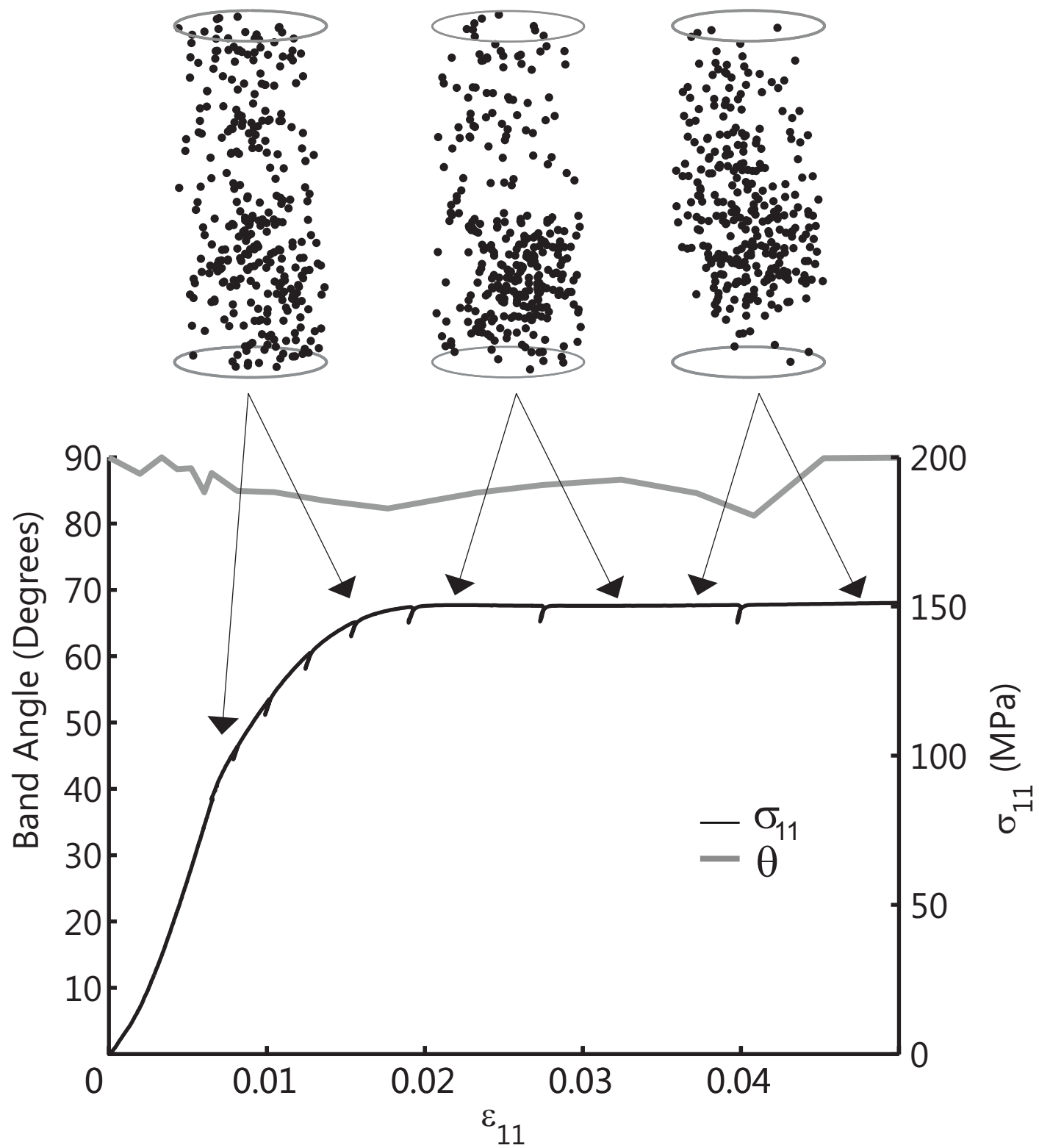


Figure 7

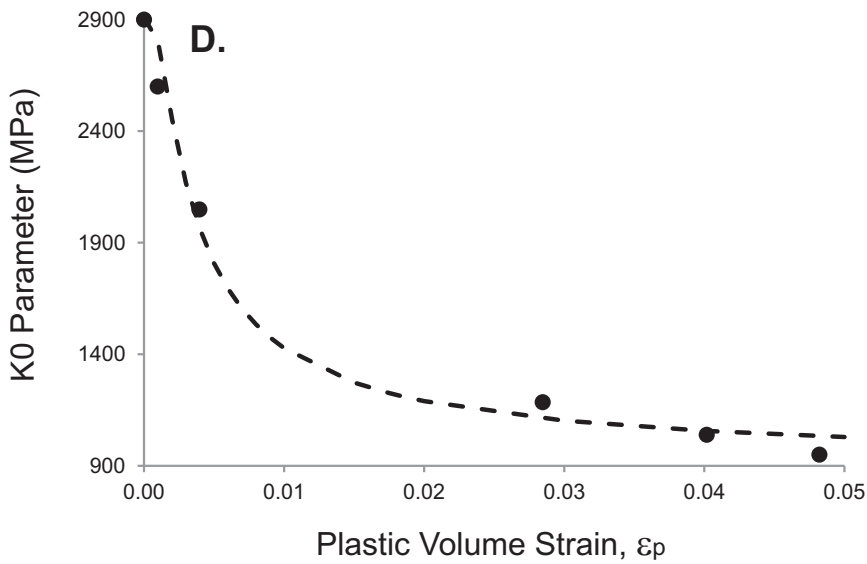
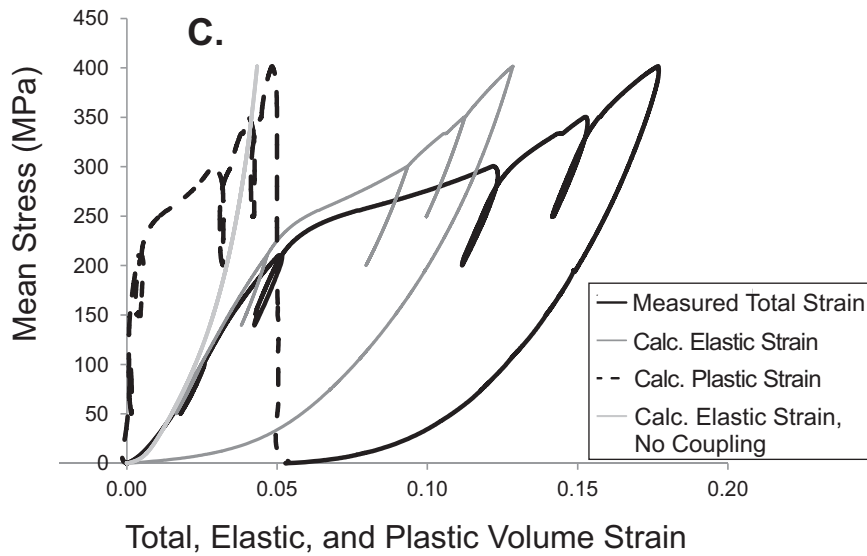
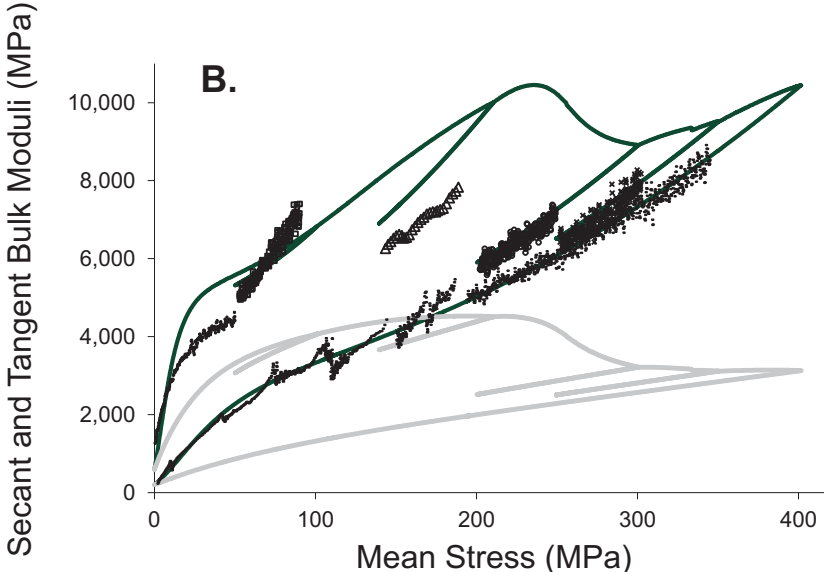
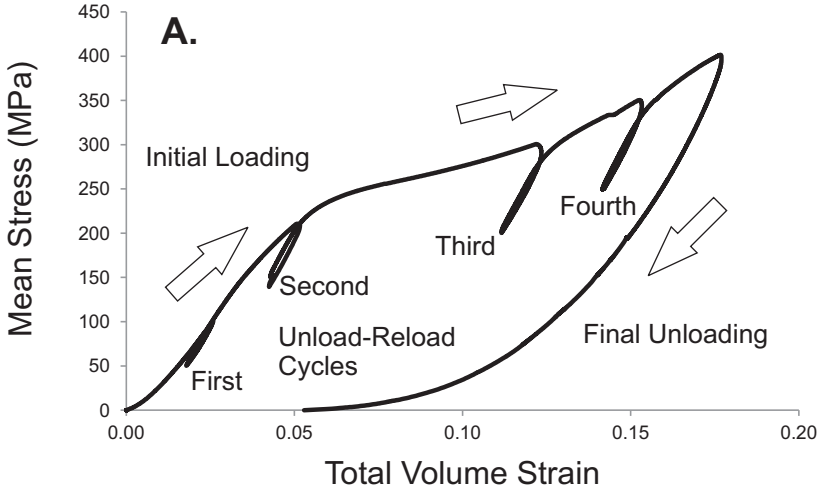


Figure 8

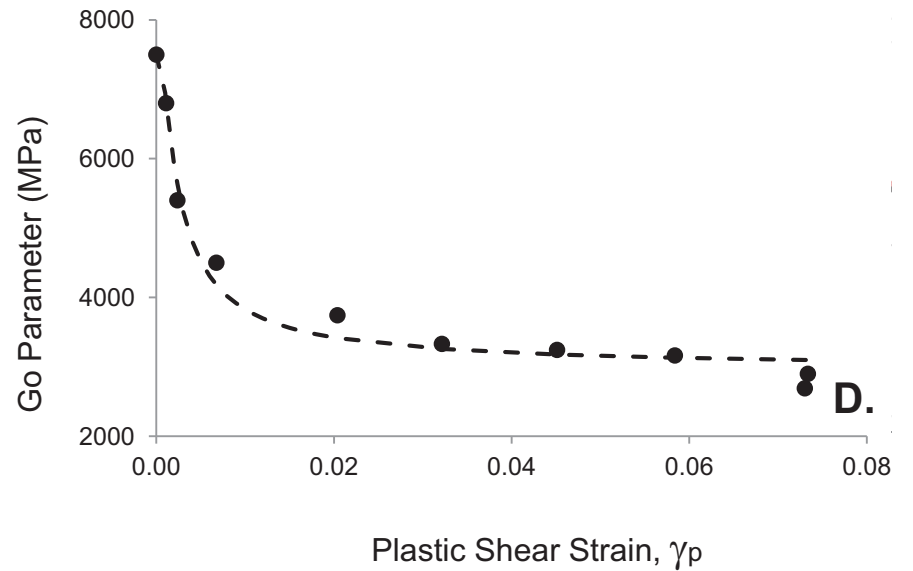
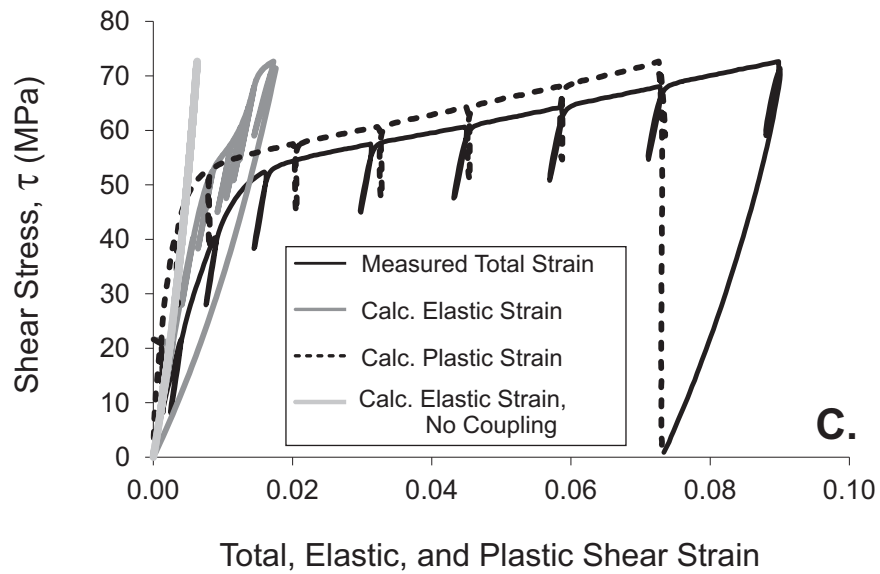
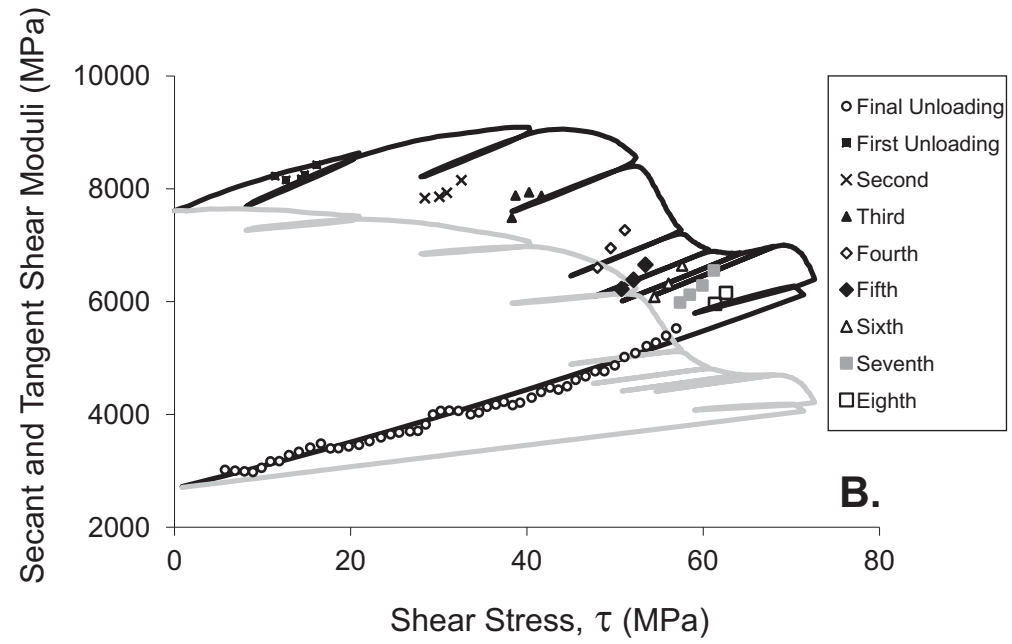
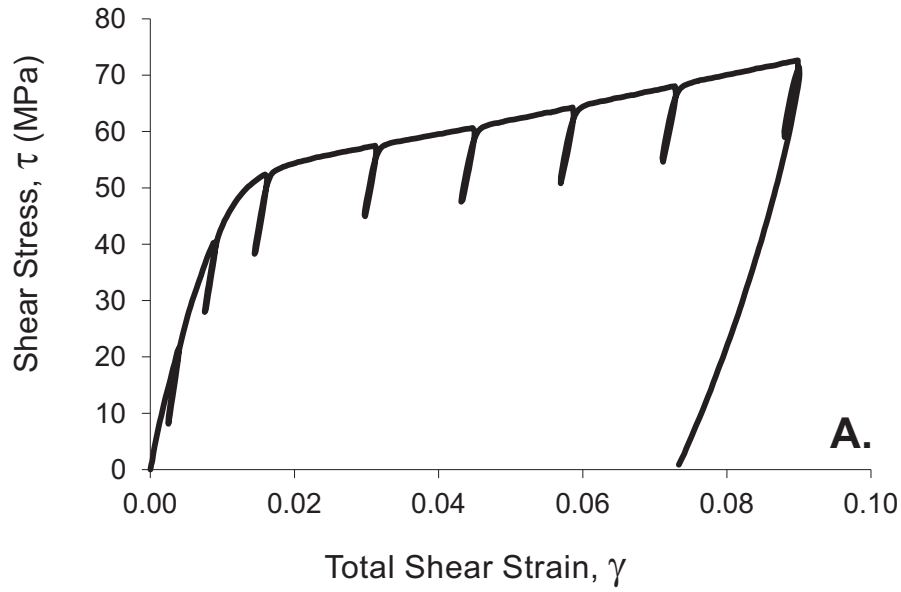


Figure 9

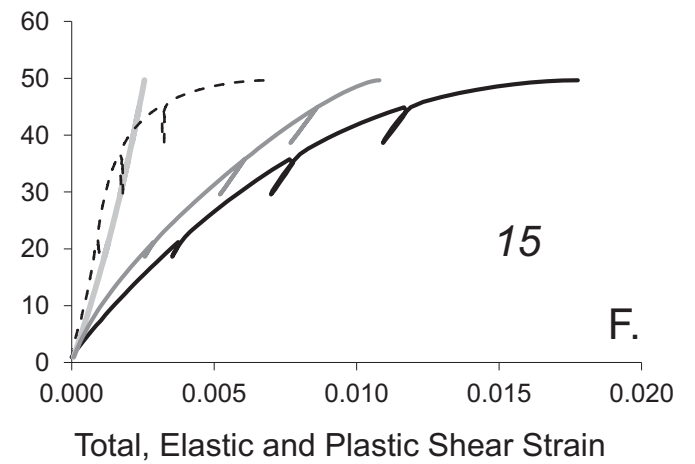
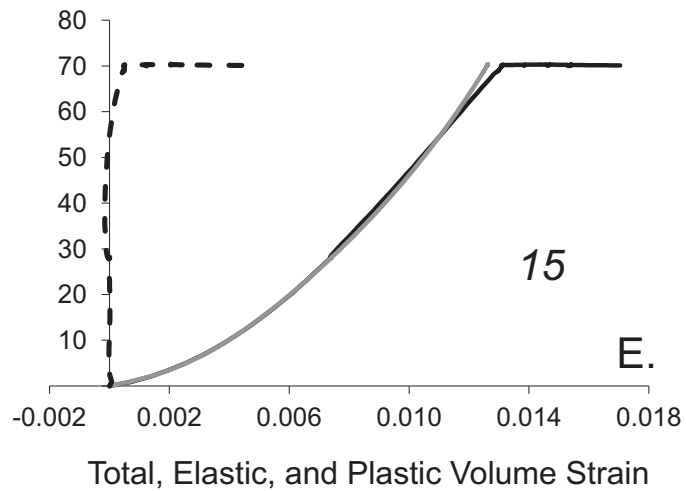
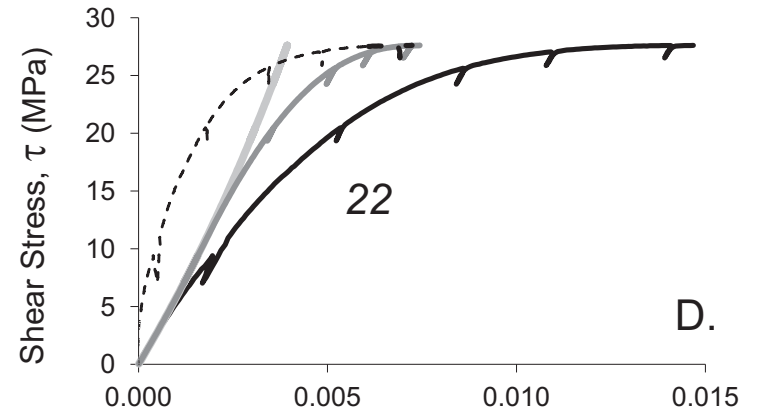
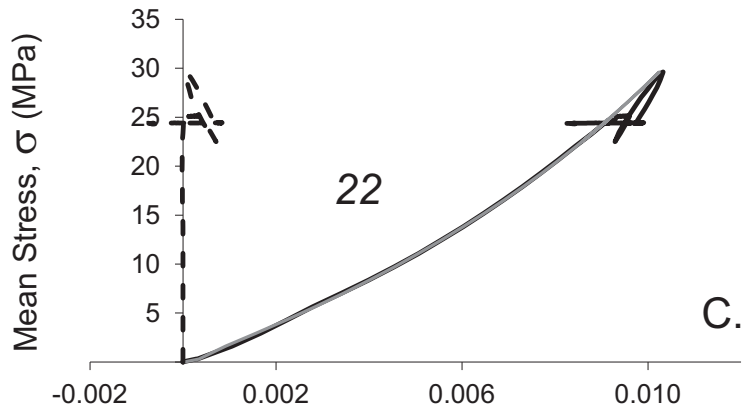
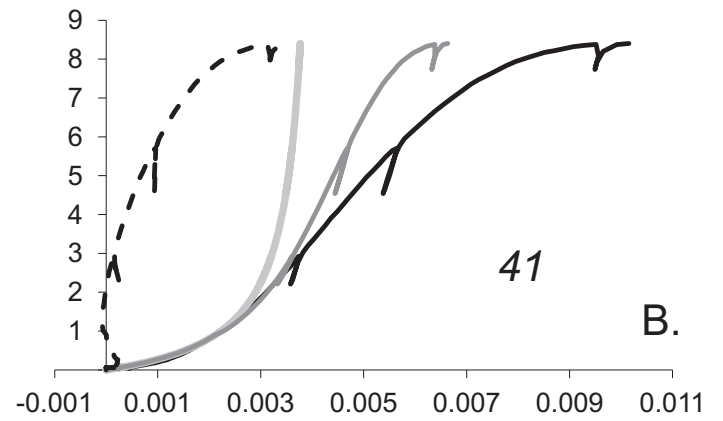
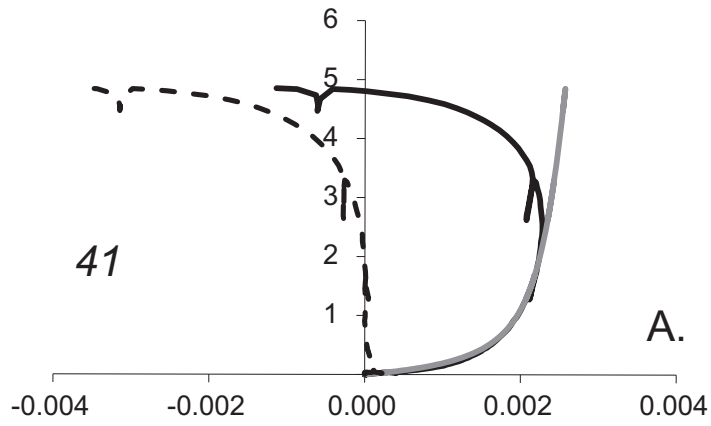


Figure 10

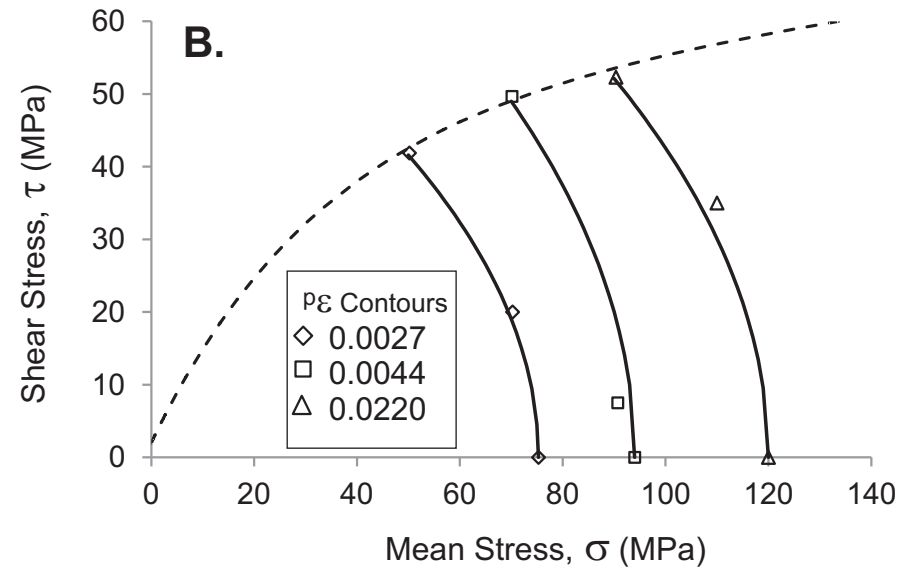
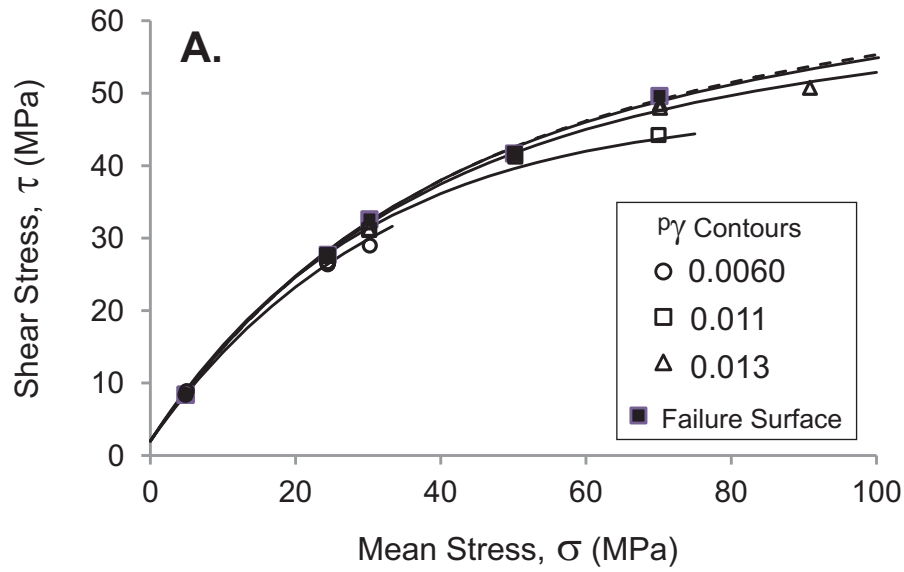


Figure 11

