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Title: Impacts of Marine Ecodynamics on the Dimethyl Sulfide(DMS) Production

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Impacts of Marine Ecodynamics on the Dimethyl Sulfide(DMS) Production

Shanlin Wang

Los Alamos National Laboratory

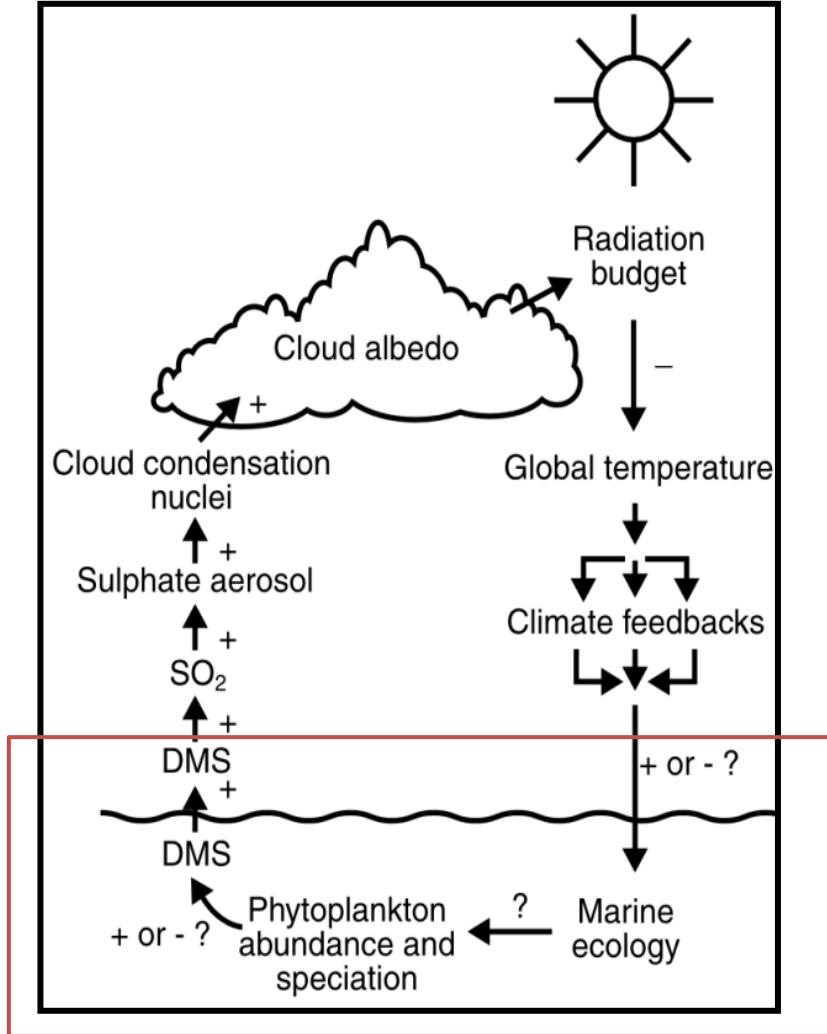
Thanks to Philip Cameron-Smith (LLNL),
Scott Elliott, Mathew Maltrud (LANL)

Mar 4, 2016

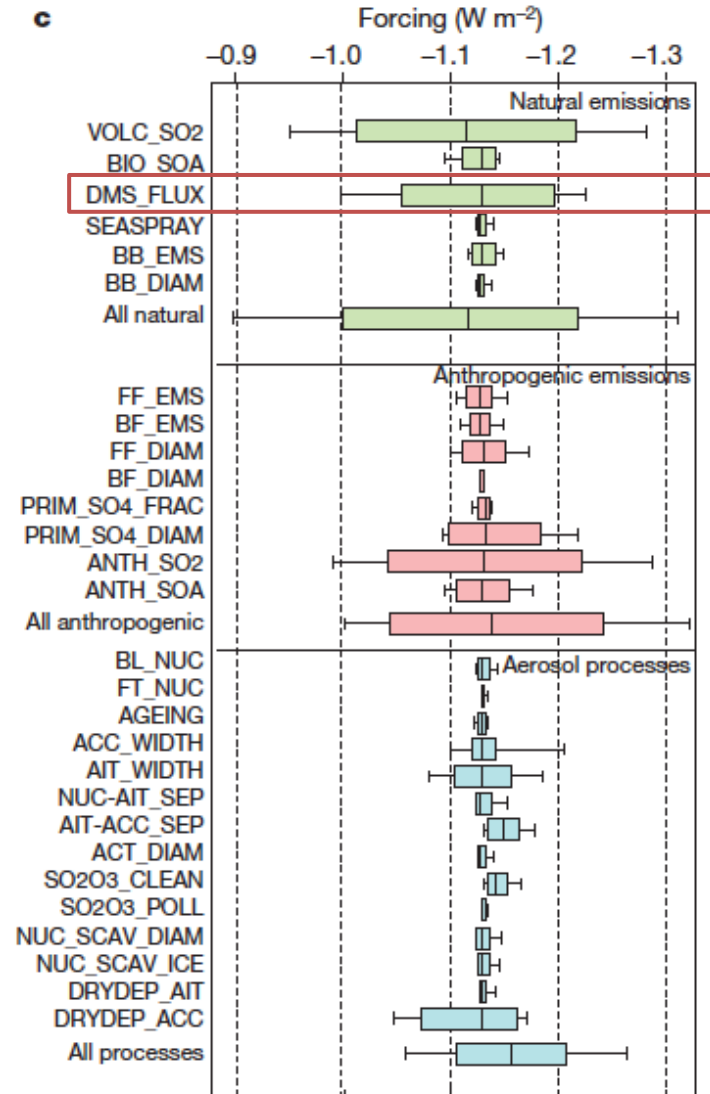


CCN, cloud-condensation nuclei; DOM, dissolved organic material; DMSO, dimethyl sulphoxide; MeSH, methanethiol; MPA, mercaptopropionate; MMPA, methylmercaptopropionate; MSA, methanesulphonic acid

Importance of DMS



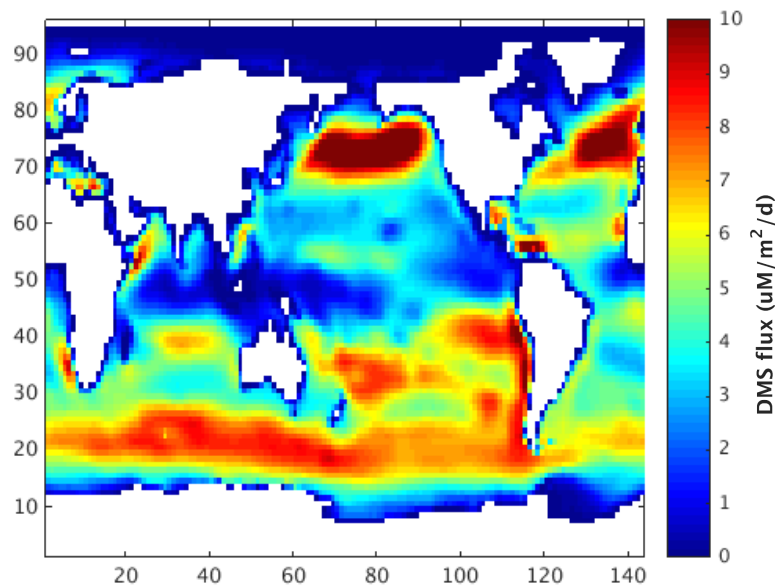
The CLAW Hypothesis
(Charlson et al., 1987)



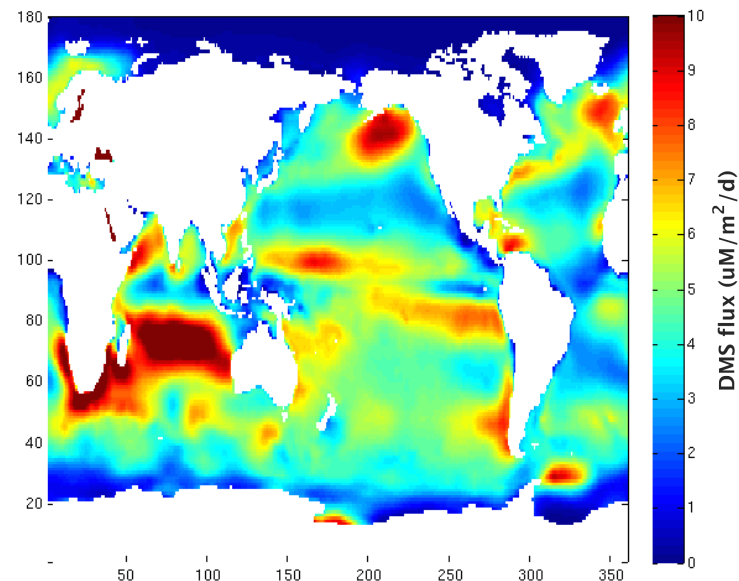
Magnitude and sources of uncertainty in aerosol first indirect forcing (Carslaw et al., 2013)

- These all point out the importance of improving DMS simulations.

DMS emissions



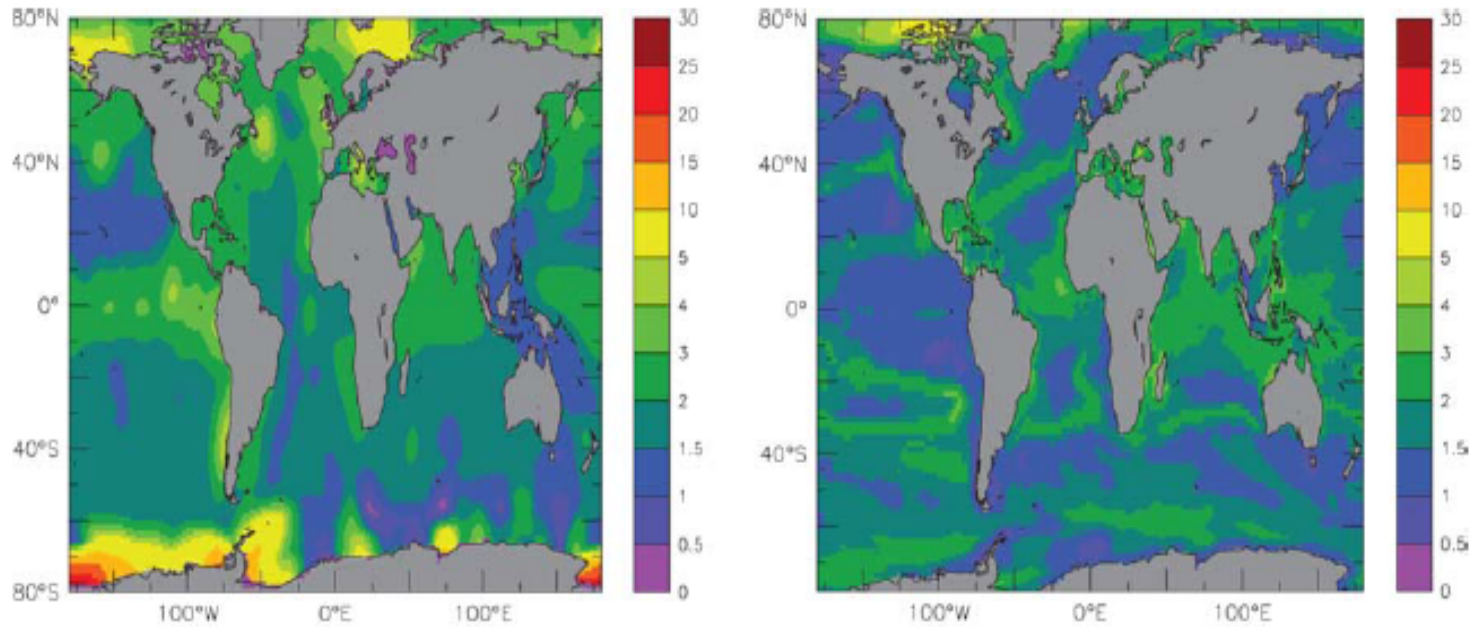
CAM DMS annual emission



Offline data-based emission

- DMS annual emission is 37.1 Tg/yr in data forced simulation
- Previous data-based estimates suggest annual DMS fluxes of 17.6 – 34.4 Tg/yr (Lana et al., 2011) and 15.0 – 29.4 Tg/yr (Kettle and Andreae, 2000)
- There are significant differences between prescribed emission and data based estimates
- The DMS production should be dynamically simulated!

Importance of marine ecosystems



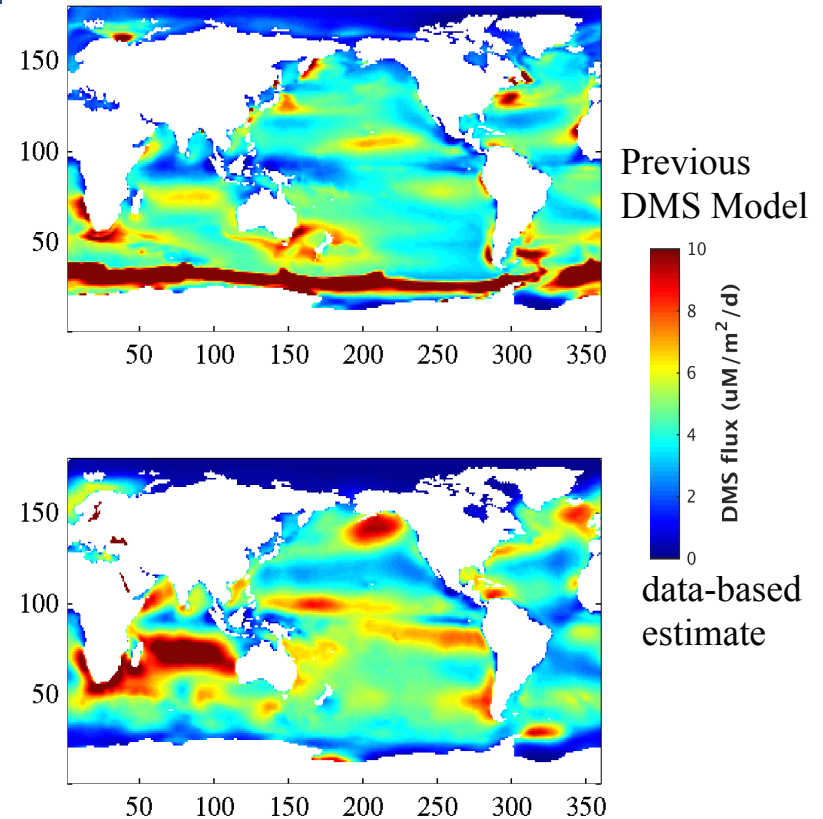
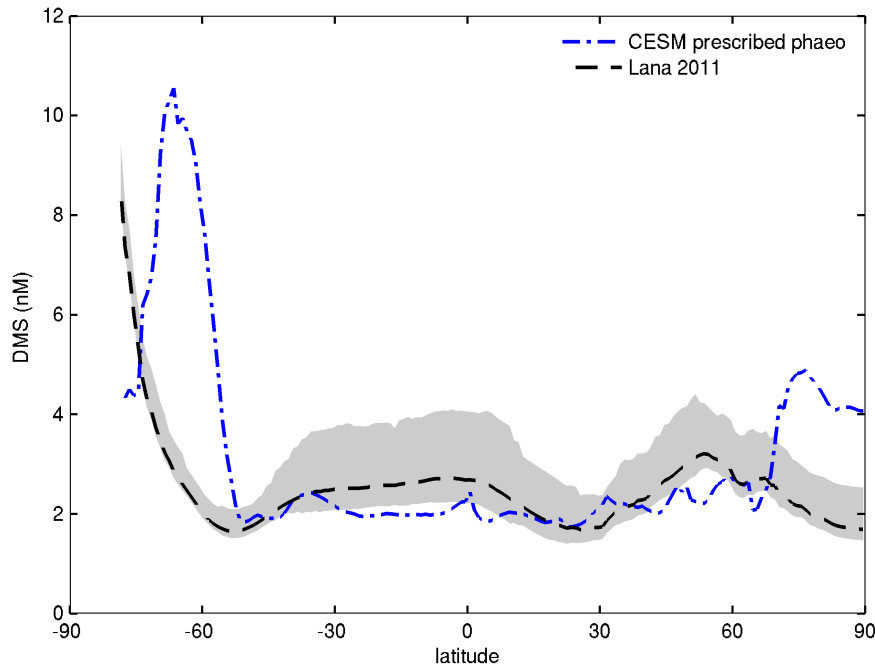
(a) DMS Kettle (nM)

(Vogt et al., 2010)

(b) DMS PlankTOM5 (nM)

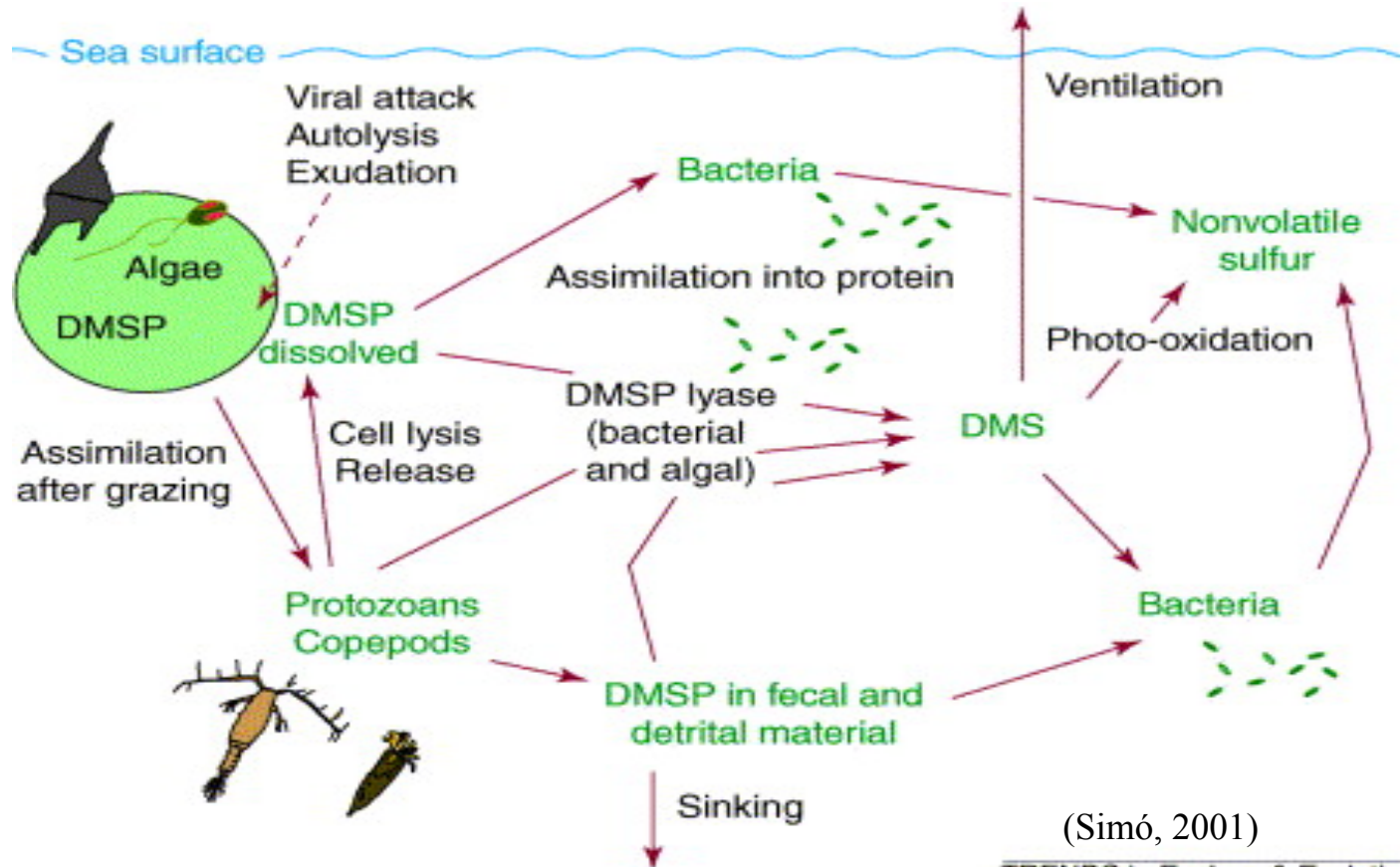
- The underestimated DMS concentrations at high latitudes are potentially due to the absence of *Phaeocystis* group (Vogt et al., 2010)
- *Phaeocystis* is important for high latitude DMS distribution

A previous DMS model



- In the earlier DMS module, *Phaeocystis* was simulated implicitly as a fraction of the global small phytoplankton. The simulated time and location of DMS peaks are biased.
- These results suggest the need of improving the representation of ecosystem structure

To simulate DMS...

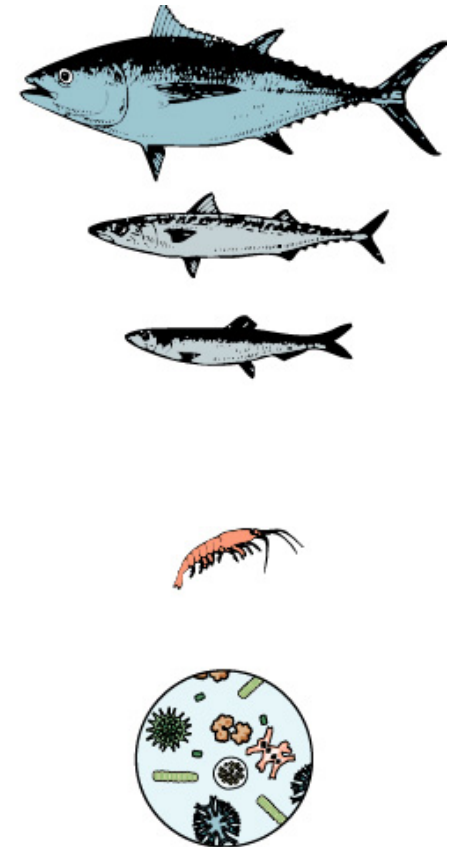
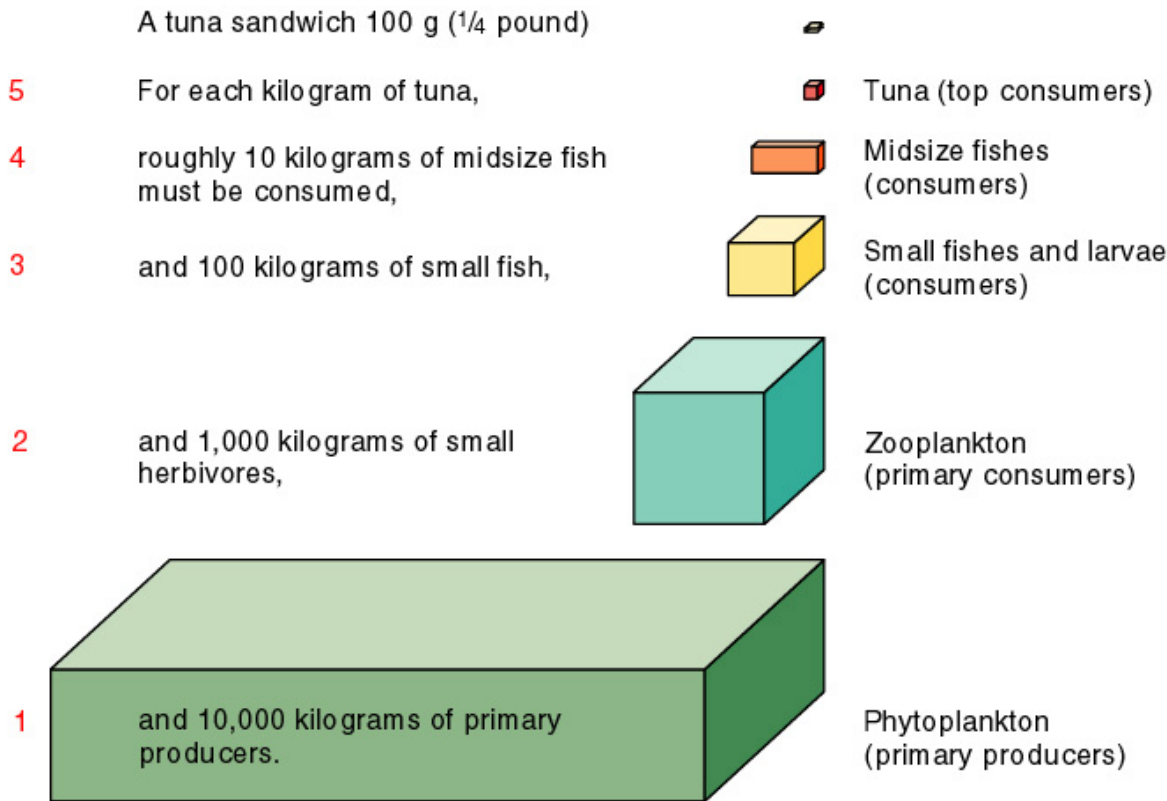


TRENDS in Ecology & Evolution

- Marine microorganisms are controlling various processes in the sulfur cycle!
- We first need to have the ecological/biogeochemical models

Ocean Biology

Trophic Level



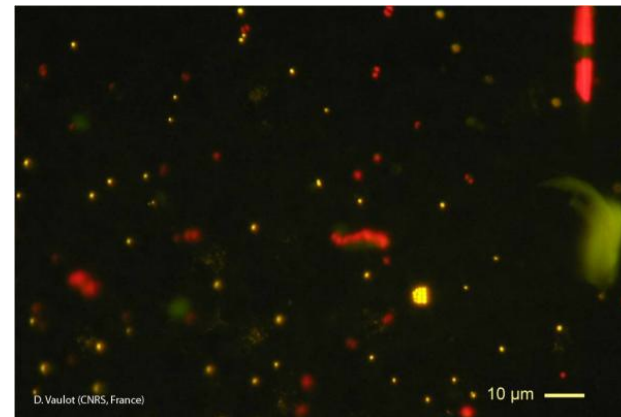
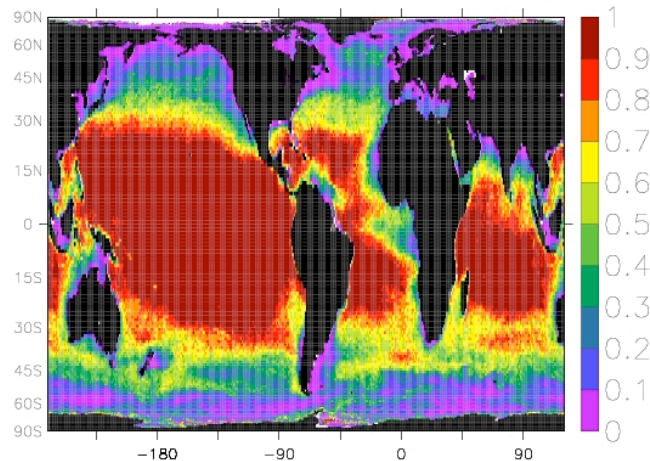
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- Higher trophic levels - Important for ecology, marine resources for humans
- Lower trophic levels - Important impact on global biogeochemical cycles
- Many BGC models adapt the concept of “**functional group**”. A function group is a collection of organisms that express a similar biogeochemical role.

Major Phytoplankton Functional Groups

Nano- and Pico-phytoplankton (nano 2-20 μm , pico <2 μm)

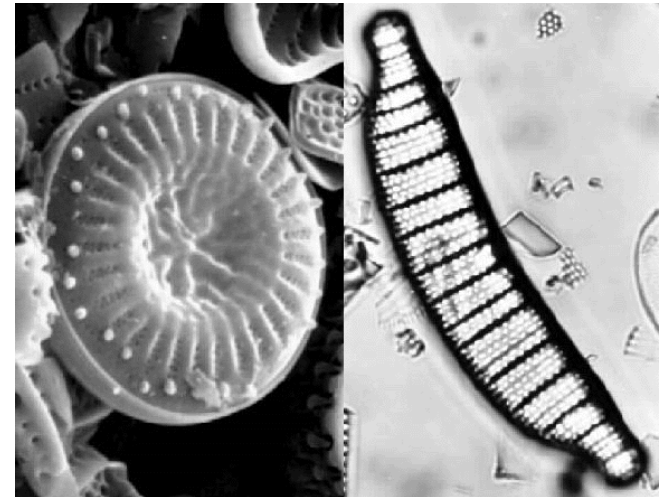
- Diverse classes of smaller species
- High surface/volume ratios
- Thrive under low nutrient conditions
- Dominate community in mid-ocean gyres
- Under strong grazing pressure (especially < 5 μM)
- Smaller contribution to sinking C export
- Usually considered as the background community.



Major Phytoplankton Functional Groups

Diatoms (20-500 μM)

- Outer frustule composed of biogenic silica
- Large phytoplankton, low grazing losses
- Low surface/volume ratio, thrive under high nutrient concentrations
- Often dominate phytoplankton blooms
- Often dominate sinking C export out of surface waters.
- Weak DMS producers



Major Phytoplankton Functional Groups

Coccolithophores (10-30 μM)

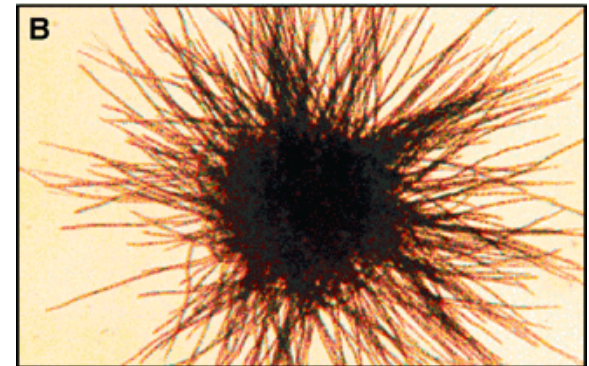
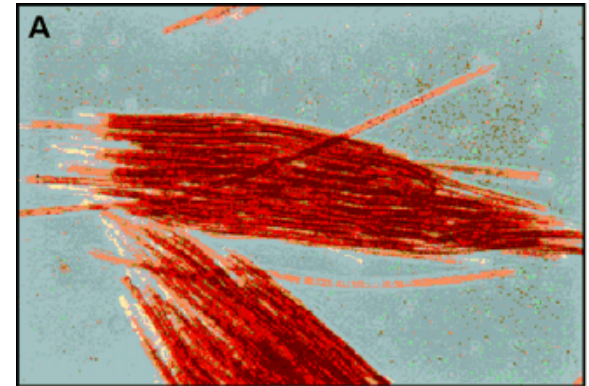
- Outer platelets, "coccoliths" of CaCO_3
- Carbonate requirement alters surface pCO_2
- Important component of sinking C export.
- Medium size, moderate grazing losses
- Moderate surface/volume, thrive under medium-high nutrient levels
- Bloom mainly at low to mid-latitudes, typically not found in coldest polar waters.



Major Phytoplankton Functional Groups

Diazotrophs (Cells 1 μM , colonies can be mm to cm)

- Capable of fixing dissolved N_2 gas into NH_4^+ .
- Source of new nitrogen to surface waters in tropical/subtropical regions. May influence total ocean nitrate and productivity over long timescales.
- Slow growth rates, low grazing losses.
- High light and Fe requirements.
- Thrive in warm, tropical and subtropical regions, not found at high latitudes.
- Do not produce DMS



Major Phytoplankton Functional Groups

Phaeocystis (solitary 3-9 μM , colonies to 3 cm)

- Thrive in colder, polar waters
- Exist in single cell and colonial forms
- Often dominate non-diatom blooms polar waters
- Colonies held together by carbohydrate rich, extra-cellular mucus (~50% of carbon). This can gives colonies higher C/N and C/P ratios, leading to efficient export of carbon.
- Strong producers of DMSP and DMS.

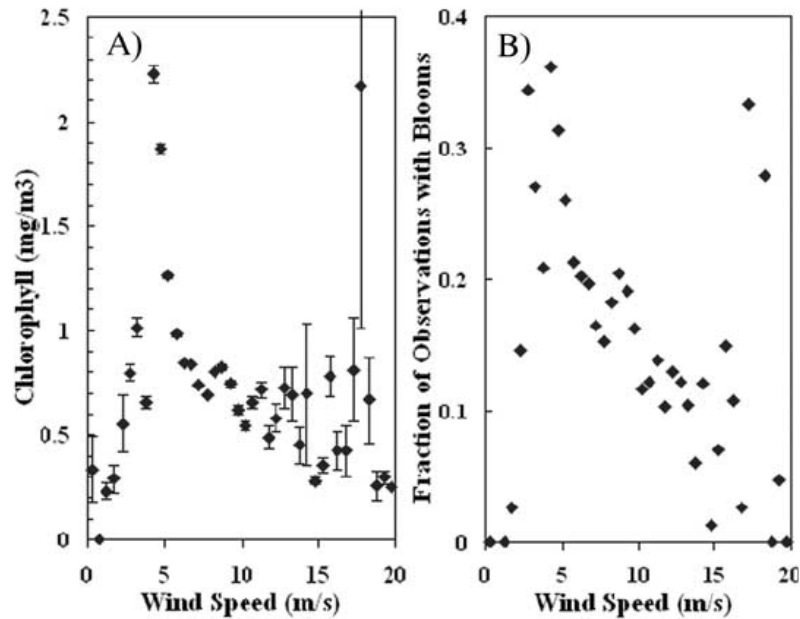


<http://www.vims.edu/phae/>

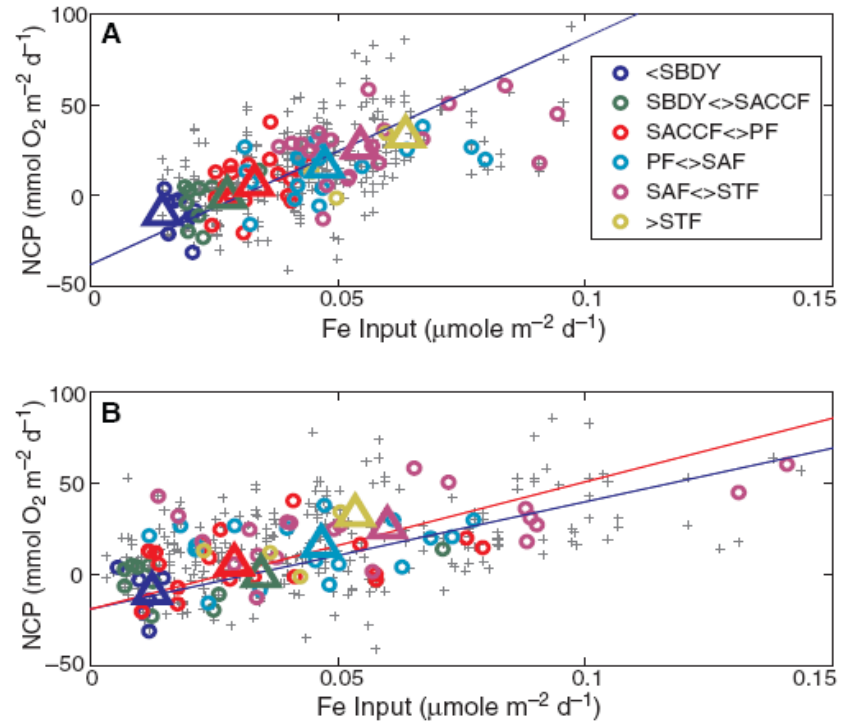


(<http://hazards.umwblogs.org>)

Controls of phytoplankton growth and competition



Chlorophyll concentrations and wind speed in Marginal Ice Zone (Fitch and Moore, 2007)

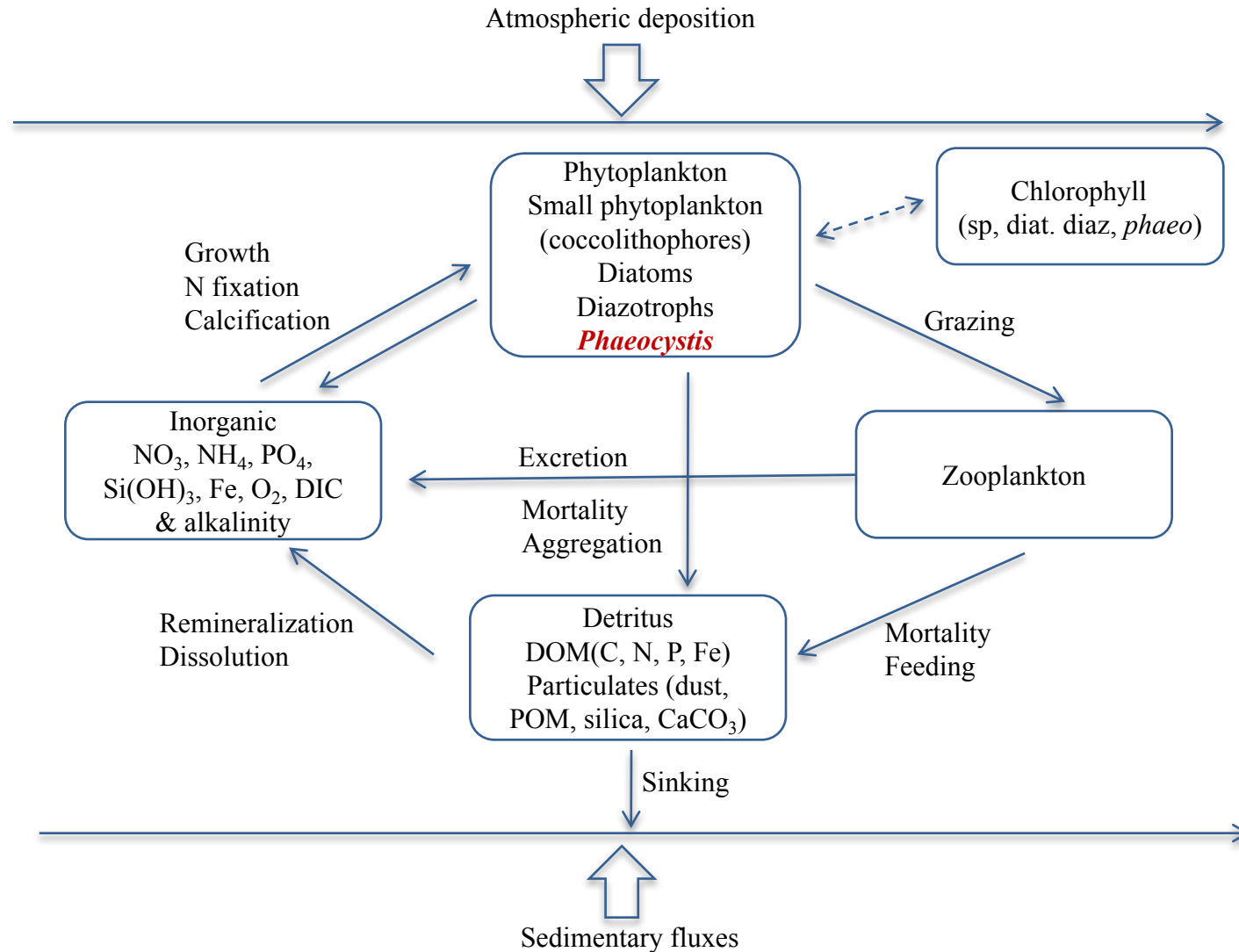


Spring (16%) and summer (84%) NCP measurements versus corresponding modeled annual (A) and 2-week (B) Fe deposition rates. (Cassar et al., 2007)

- The distribution of phytoplankton biomass and NPP is determined by the availability of light and nutrients (nitrogen, phosphate, iron, silicate).
- Different phytoplankton groups have different requirements for these factors

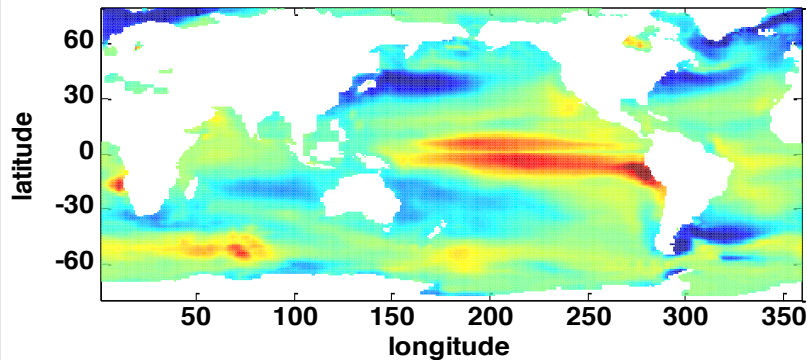
Model Description

Biogeochemical Elemental Cycling (BEC) model

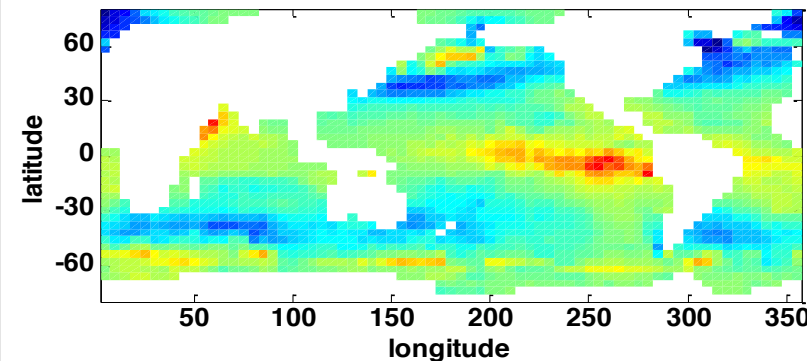


Model performance

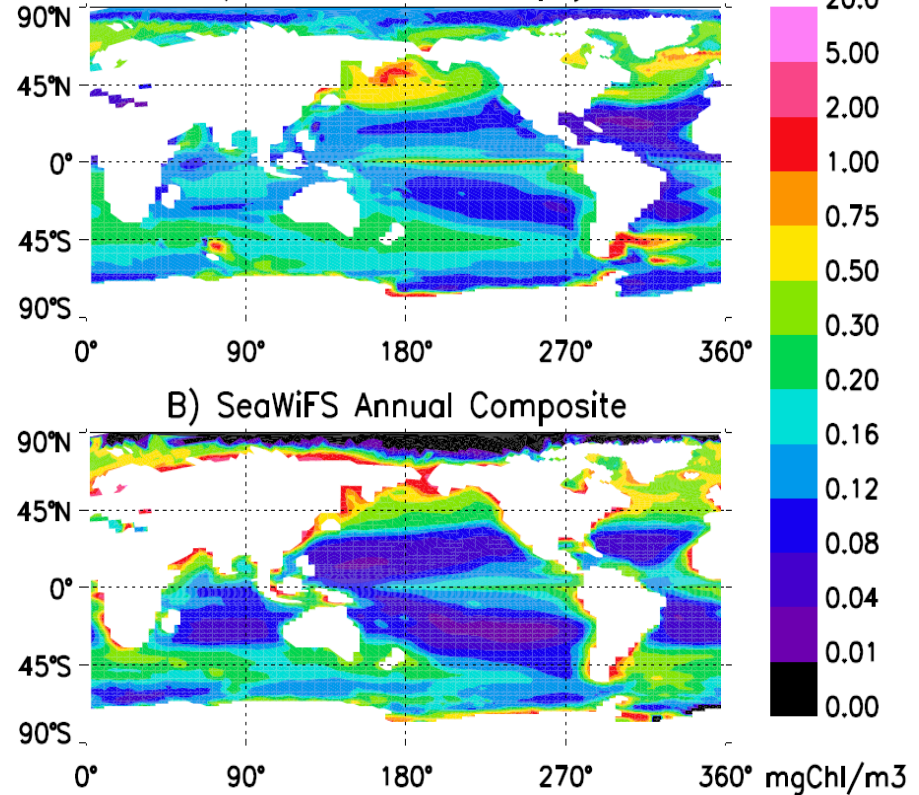
CO₂ flux in 2000 estimated using the CCSM (molC/m²/yr)



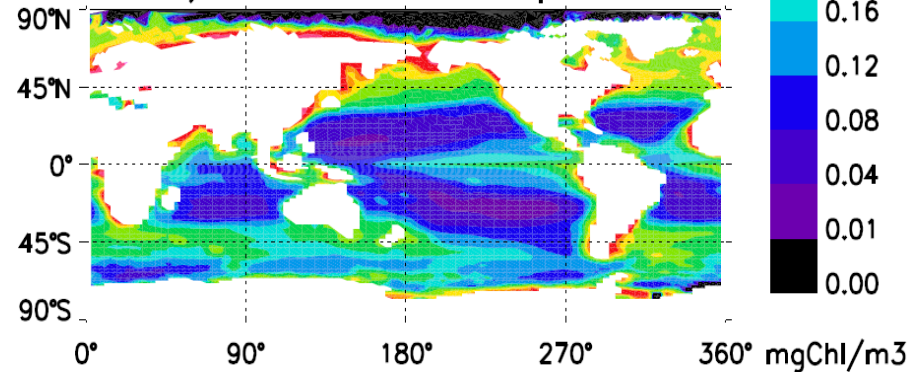
Observed CO₂ flux by Takahashi et al., 2009 (molC/m²/yr)



A) BEC Annual Chlorophyll



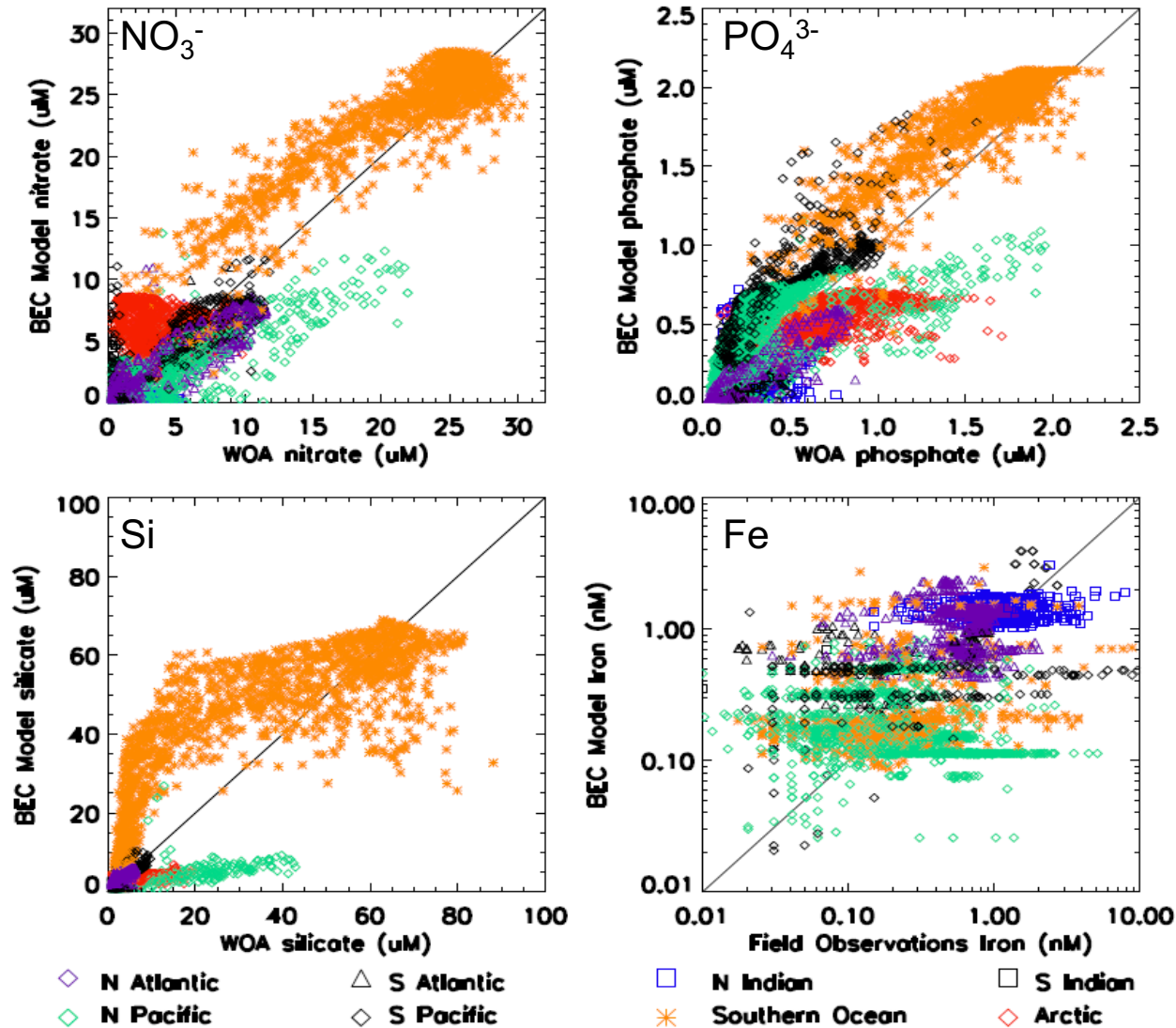
B) SeaWiFS Annual Composite



(Wang and Moore, 2012)

- The modeled global sea-air CO₂ flux is comparable to the observation-based estimate by Takahashi et al. (2009)
- The BEC model reasonably reproduces the patterns of surface chlorophyll concentrations.

Model performance



- Correlations between modeled and observed major nutrient concentrations are all higher than 0.85
- Regional biases remain

(Wang and Moore, 2012)

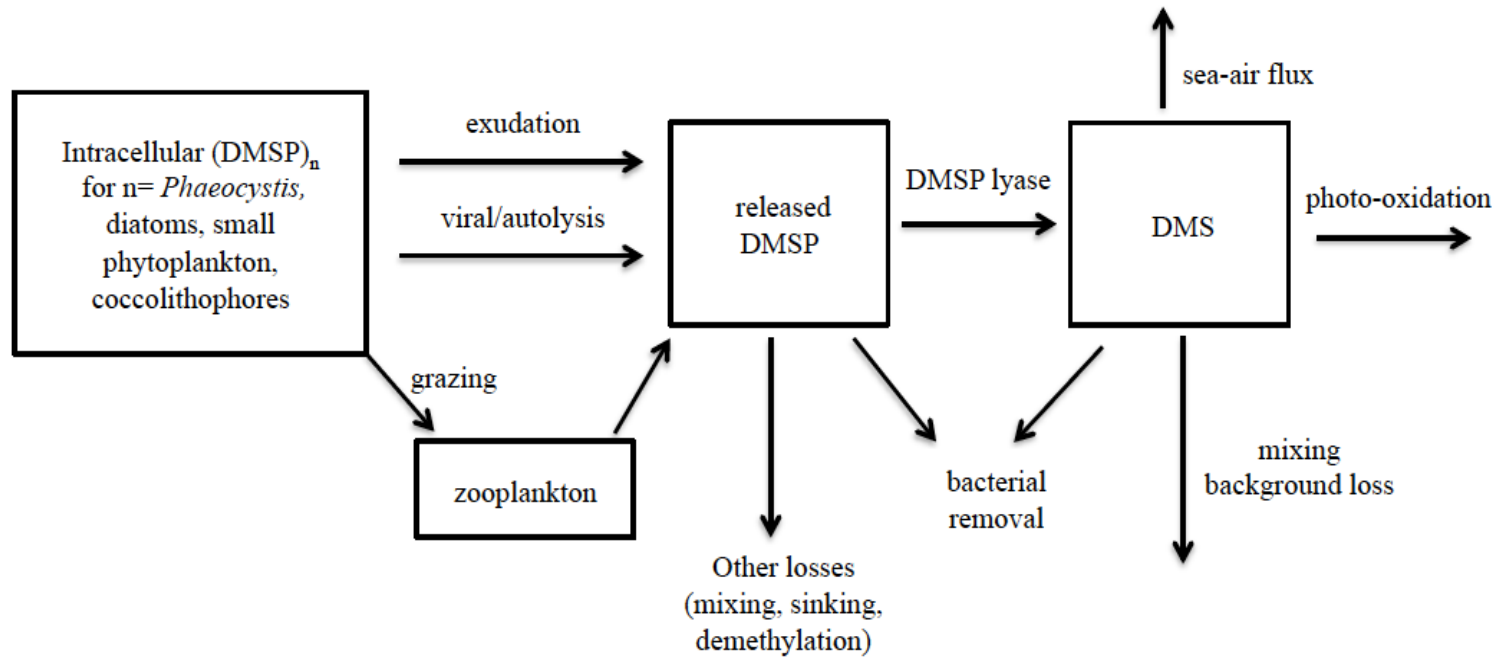
DMS module

$$Prod^{DMSP} = scale \times \left(\sum_i Prod_i^{DMSP} + Prod_Z^{DMSP} + Prod_{phaeo-dir}^{DMSP} \right)$$

$$Prod^{DMS} = yield \times k \times Prod^{DMSP}$$

$$Loss^{DMS} = Loss_B^{DMS} + Loss_{bkg}^{DMS} + Loss_{photo}^{DMS}$$

$$Loss^{DMSP} = Loss_{DMS}^{DMSP} + Loss_{bkg}^{DMSP}$$



Model Description

Biogeochemical elemental cycling (BEC) model

Ecosystem Module

5 phytoplankton groups
One zooplankton group
DOM, sinking particulates
Multiple growth limiting
nutrients

DMS module

Prognostic calculations of
the DMS production and
removal

Ocean circulation component in Community Climate System
Model / Community Earth System Model

Previous studies

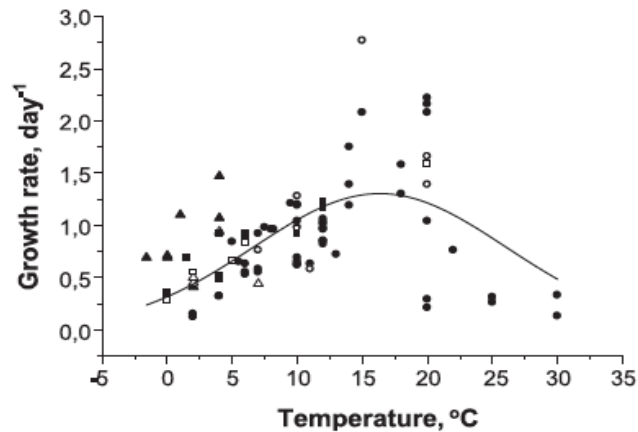
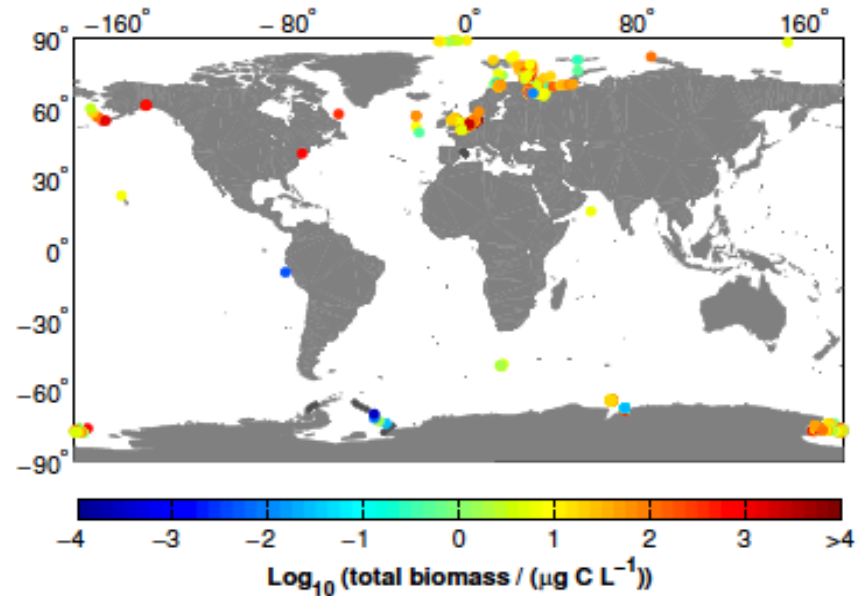
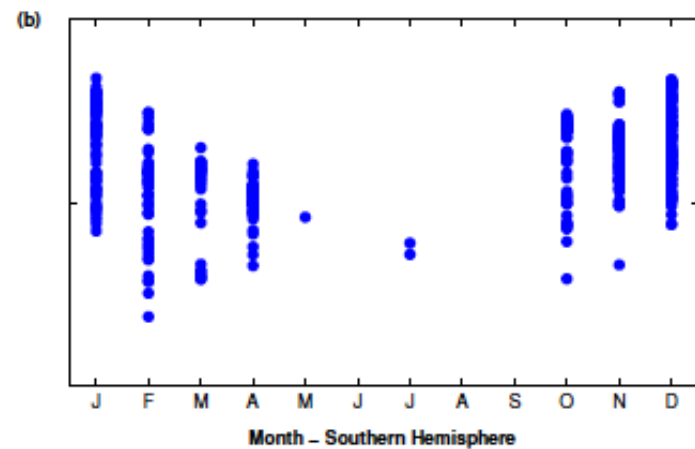
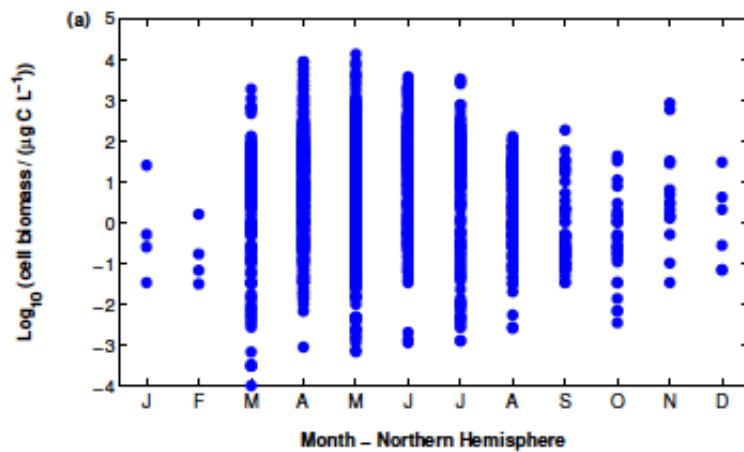


Fig. 3. Relationship between *Phaeocystis* specific growth rate μ and temperature ($\mu = \ln 2 \cdot v$, with v = division rate). The closed and the open symbols represent colony cells and solitary cells data, respectively. Squares, circles and triangles represent data for *P. pouchetii*, *P. globosa* and *P. antarctica*, respectively.

(Schoemann et al., 2005)



(Vogt et al., 2012)



(Vogt et al., 2012)

Incorporating *Phaeocystis*

- Two groups of *Phaeocystis*: One group for northern species, combining *P. globosa* and *P. pouchetii*; One group of *P. antarctica*
- Only the blooming colonial form considered

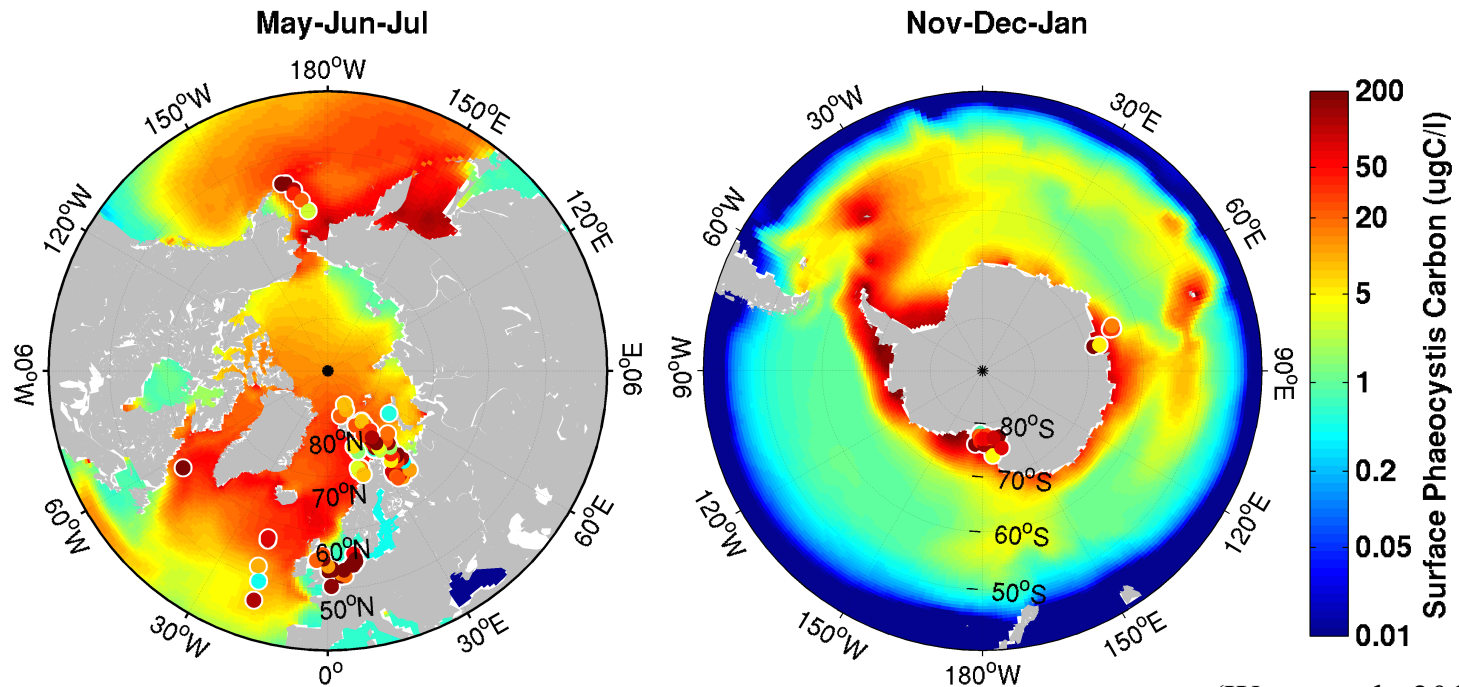
$$PCmax = PCmax \times \min\left(1, \frac{(thres - TEMP)}{(thres - Tpeak)}\right)$$

$$PCmax = \max(PCmax, 0)$$

$$graze = Z_{max} \times zooC \times \frac{phaeoC}{(phaeoC + diatC + Zgrz)}$$

- Key parameters: k_{NO_3} , k_{Fe} , grazing, T_{opt} , T_{thres} , and alpha
- The values of T_{peak} are 16.3 °C and 5 °C in NH and SH, respectively
- Grazing on *Phaeocystis* or diatoms is influenced by the total biomass of both groups
- In general, our preferred physiological parameters suggest that *Phaeocystis* is well adapted to low-light and nutrient-rich conditions

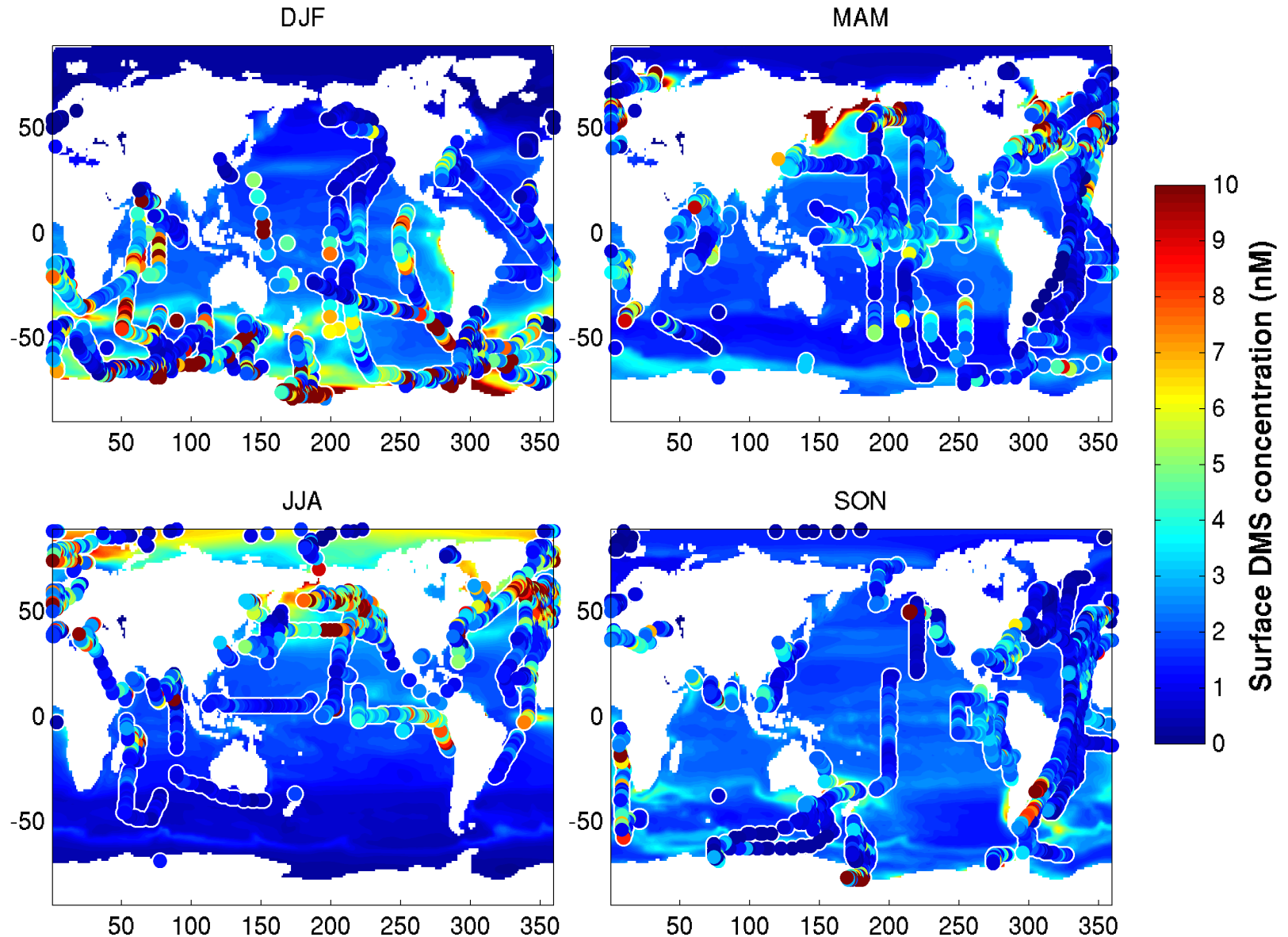
Simulated *Phaeocystis*



(Wang et al., 2015)

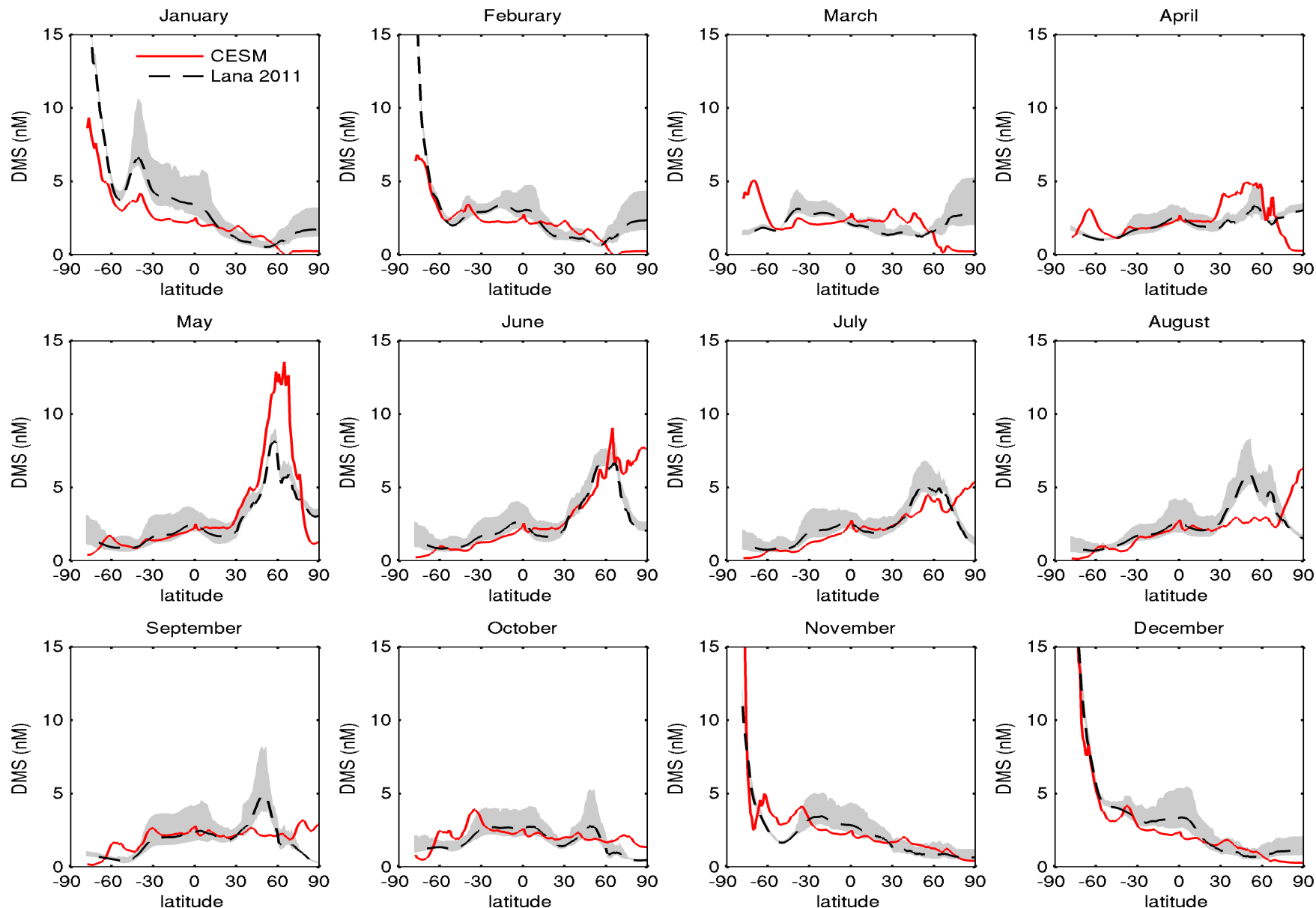
- In general, the model does a reasonable job reproducing the observed seasonality and spatial distributions of *Phaeocystis*
- *Phaeocystis* contributed ~13% of annual primary production and ~19% of sinking carbon export in the Southern Ocean (>40°S).
- Simulated *Phaeocystis* biomass cannot reproduce the extreme high concentrations sometimes reported in field studies due to averaging in large grid and underestimated surface nutrients

Simulated DMS



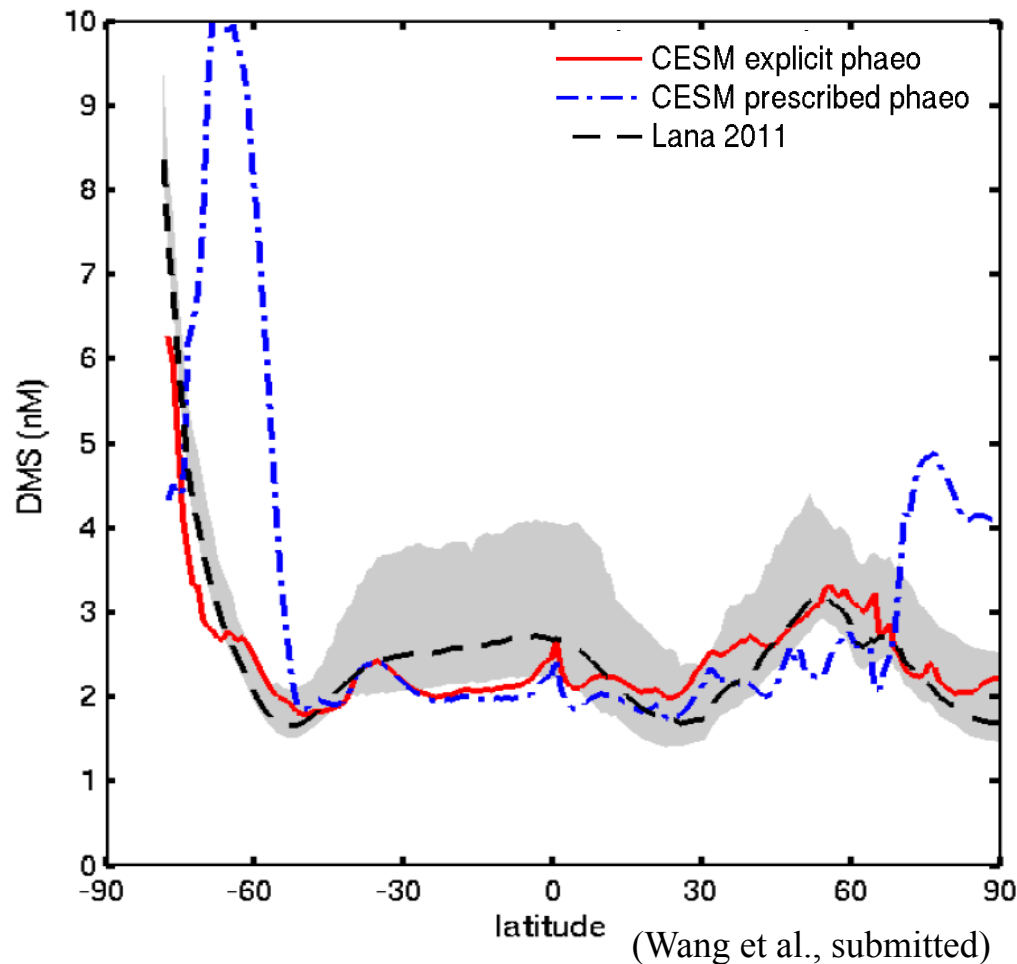
DMS maximum in the North Pacific in May and along the Antarctic coast in December, in agreement with *Lana et al.* (2011)

Simulated DMS



Simulated DMS shows strong bloom behavior at high latitudes, but weak seasonal cycling at low latitudes. Seasonality and distributions are in good agreement with observations.

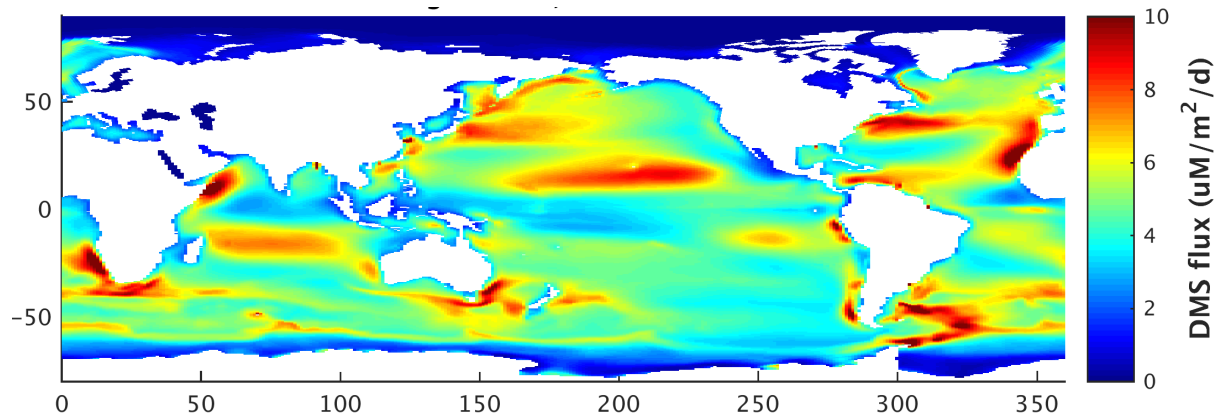
Zonal DMS distributions



- The simulated average DMS concentration for the surface ocean is 2.26 nM, comparable to data-based estimate of 2.34 nM
- Zonal mean DMS is clearly improved, and matches the observation-based estimate closely, with observed DMS peaks between 50° – 60° N and south of 60° S well reproduced

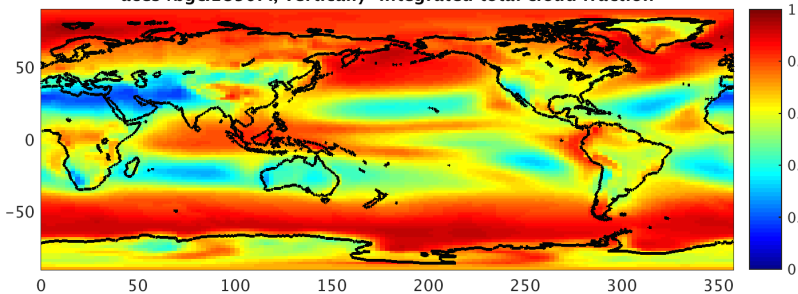
DMS-climate feedback

DMS flux (pre-industrial)



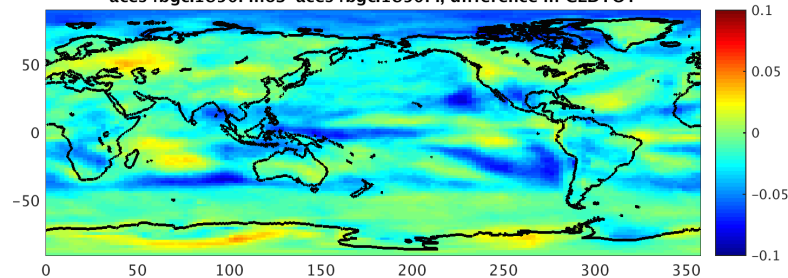
Pre-industrial

aces4bgc.1850PI, vertically-integrated total cloud fraction

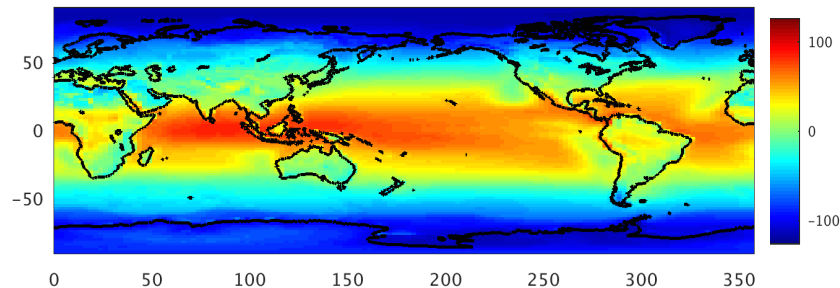


Difference (without DMS – with DMS)

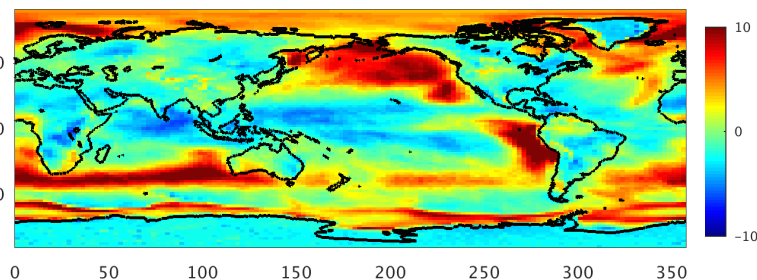
aces4bgc.1850PInoS - aces4bgc.1850PI, difference in CLDTOT



aces4bgc.1850PI, FSNT - FLNT (W/m^2)

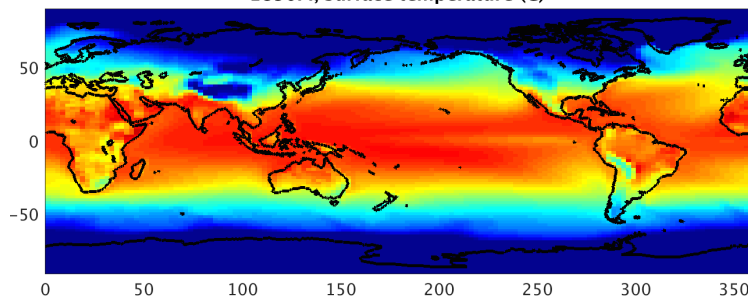


aces4bgc.1850PInoS - aces4bgc.1850PI, FSNT - FLNT (W/m^2)



Pre-industrial

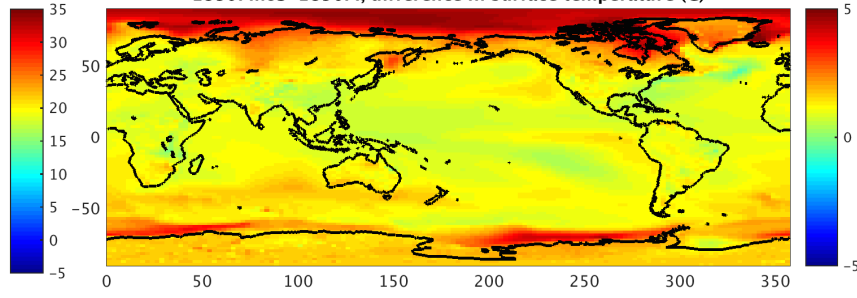
1850PI, surface temperature (C)



Surf Temp

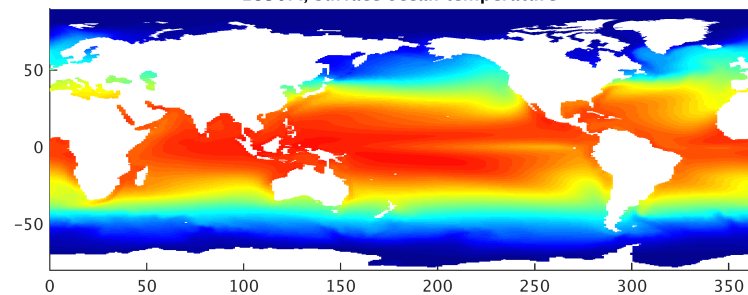
Difference (without DMS – with DMS)

1850PI_{noS}-1850PI, difference in surface temperature (C)

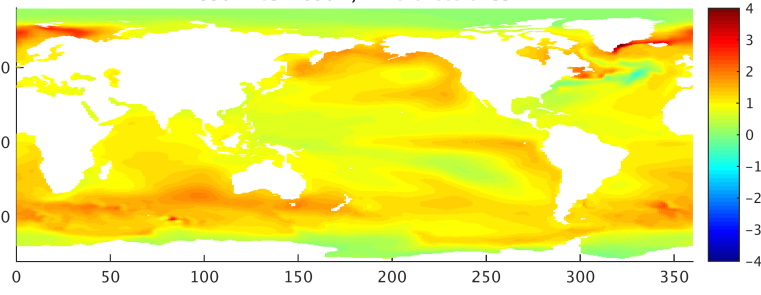


SST

1850PI, surface ocean temperature

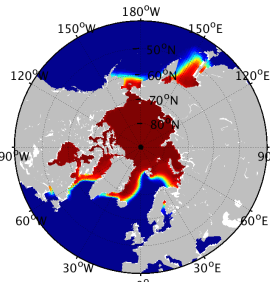


1850PI_{noS}-1850PI, Differences of SST

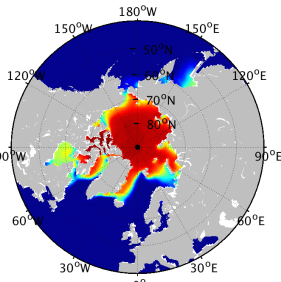


IFRAC

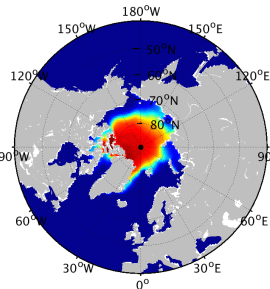
Feb-Mar-Apr



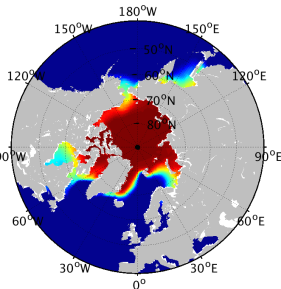
May-Jun-Jul



Aug-Sep-Oct

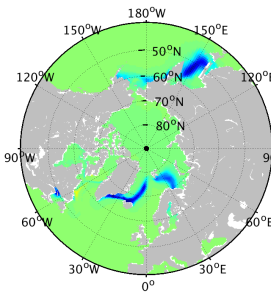


Nov-Dec-Jan

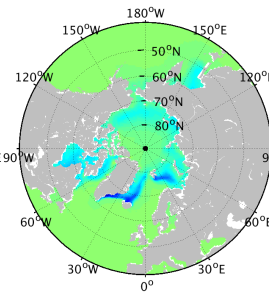


sea ice fraction

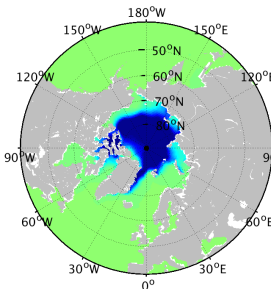
Feb-Mar-Apr



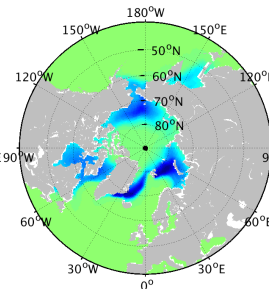
May-Jun-Jul



Aug-Sep-Oct



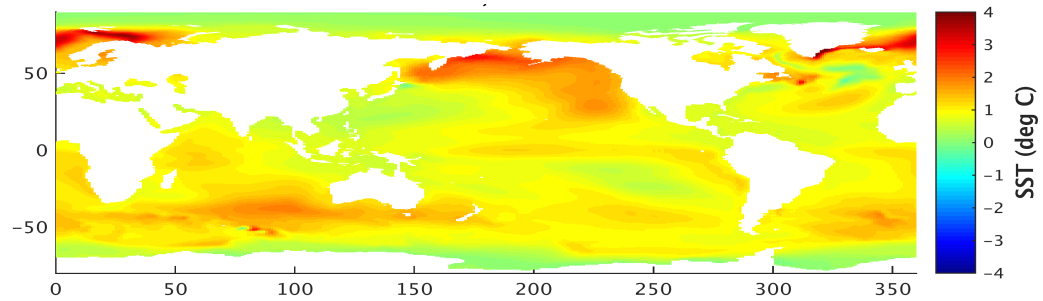
Nov-Dec-Jan



0.5
-0.5

Surface nutrients

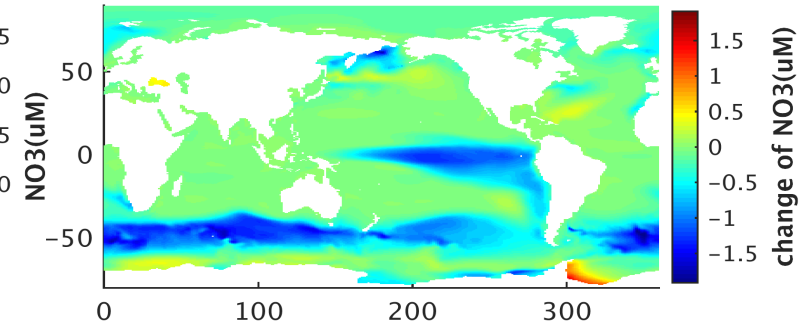
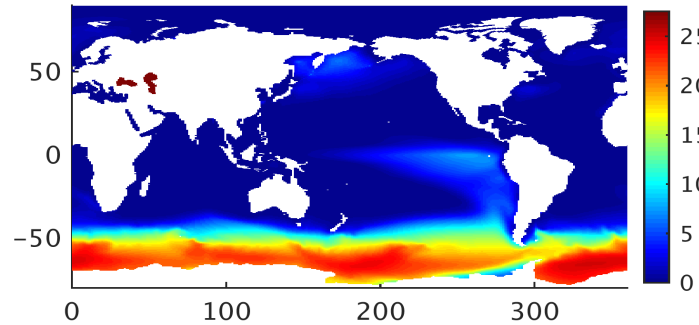
Warming SST in simulation without DMS flux into the atmosphere



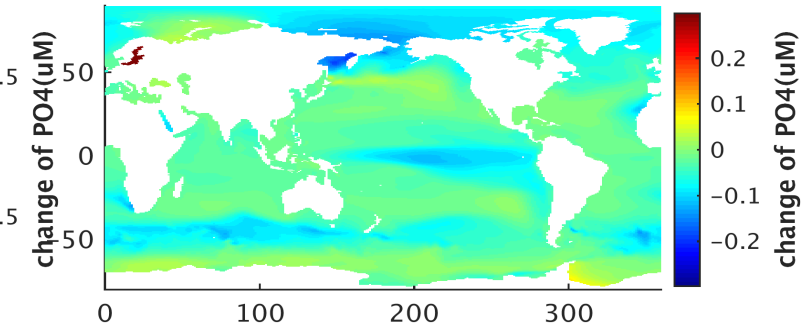
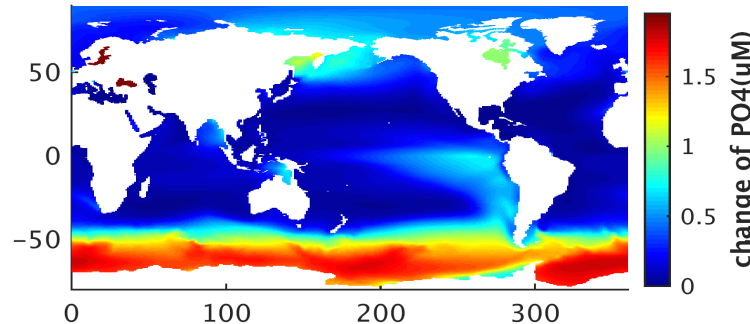
Pre-industrial

Difference (without DMS – with DMS)

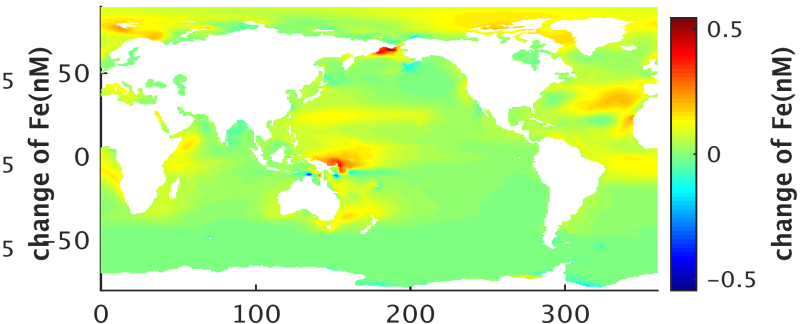
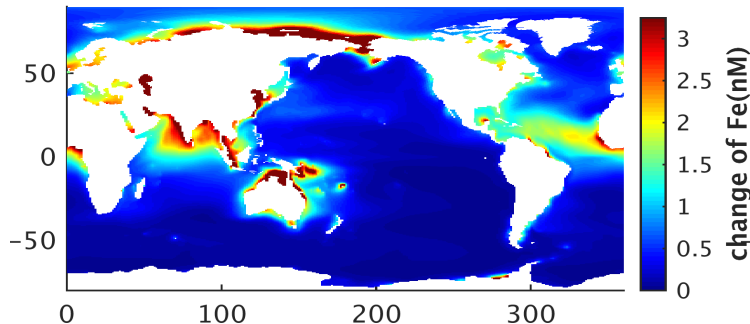
Lower surface NO_3



Lower surface PO_4



Higher surface Fe

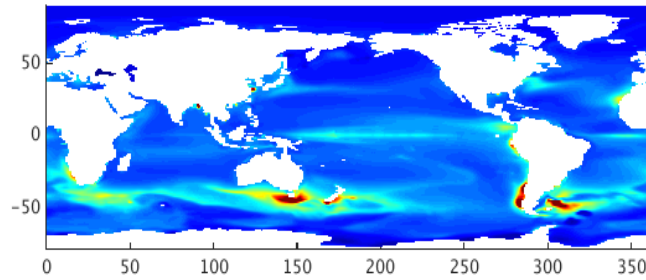


Phytoplankton and DMS

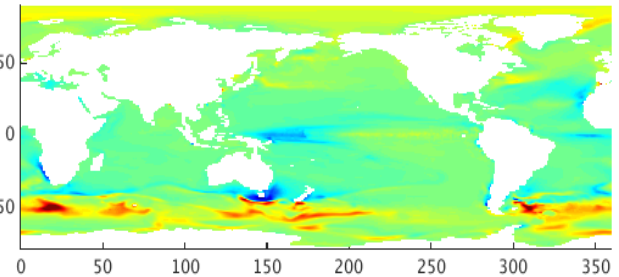
Pre-industrial

Difference (without DMS – with DMS)

Small phytoplankton
Increase 40 - 60 °S

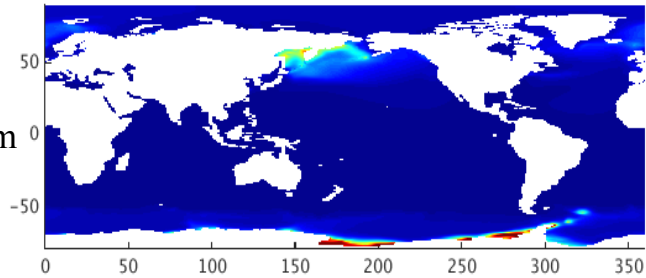


spC (mmol/m³)

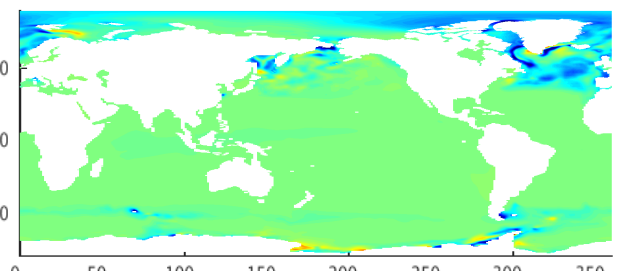


difference (mmol/m³)

Phaeocystis decrease
Similar change to diatom

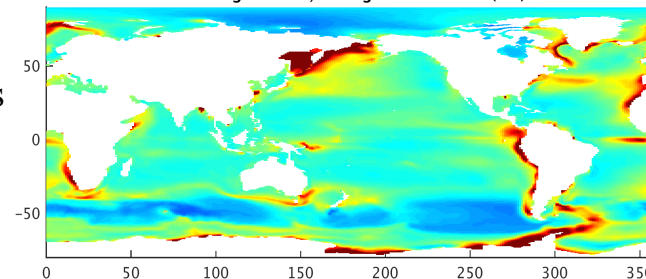


phaeoC (mmol/m³)

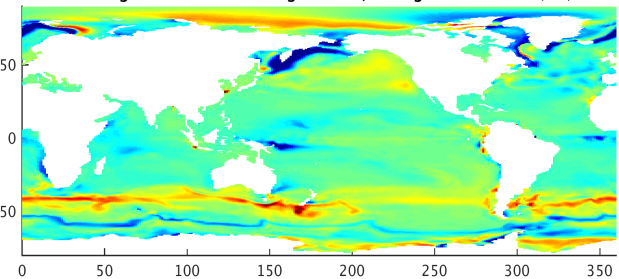


difference (mmol/m³)

DMS difference follows
marine ecosystems

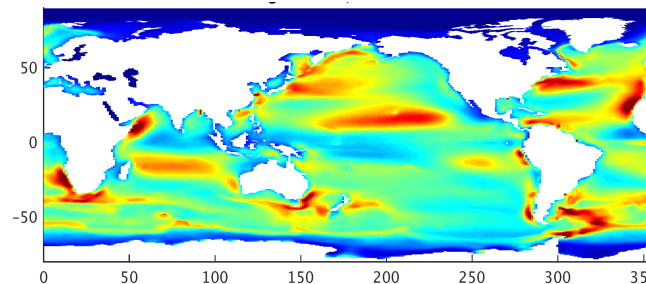


surface DMS (nM)

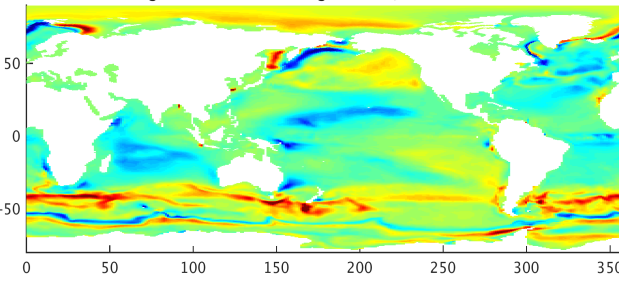


change of DMS (nM)

DMS flux also affected
by wind and sea ice



DMS flux ($\mu\text{M}/\text{m}^2/\text{d}$)

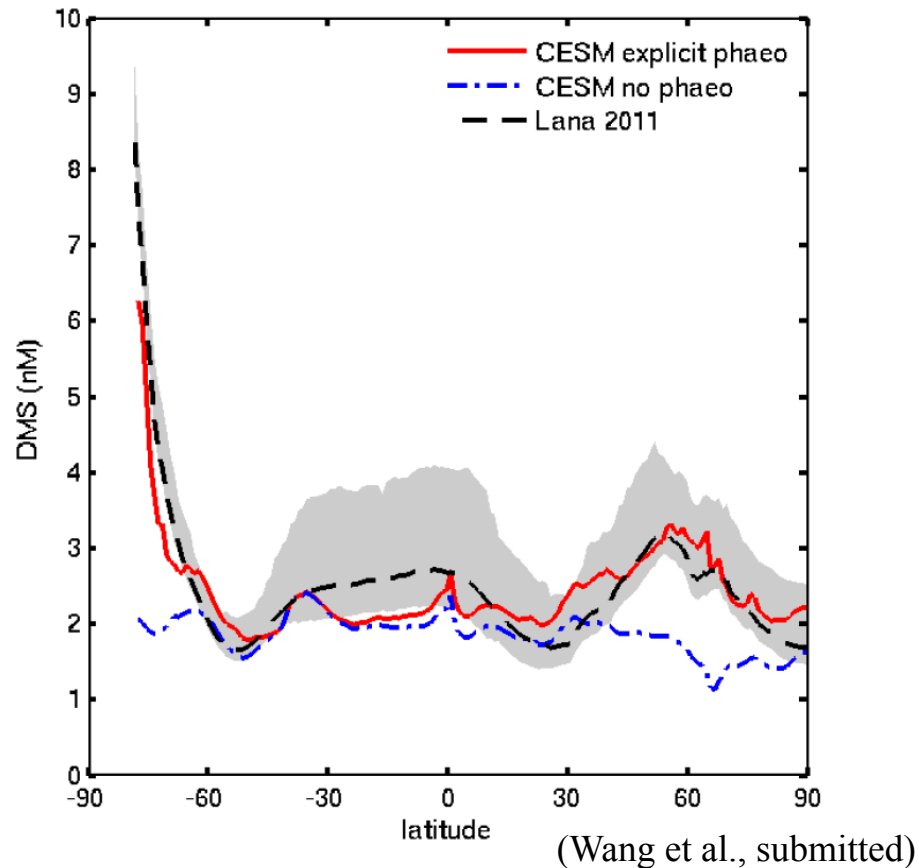


change of flux ($\mu\text{M}/\text{m}^2/\text{d}$)

There is no uniform feedback patterns.

Simulation	PI	PI _{noS}	PD	PD _{noS}	FUT	FUT _{noS}
Model top radiation (W/m ²)	0.44	1.06	0.61	0.93	2.6	3.18
Sea ice area (September, millions of km ²)	<p>The feedback seems weak, then is phytoplankton important?</p>					6.34
Sea ice area (February, millions of km ²)						1.95
SST (°C)						23.34
Primary Production						52.36
Sinking POC at 100m	7.63	7.51	7.62	7.5	6.65	6.37
Surface DMS (nM)	2.30	2.23	2.30	2.23	1.96	1.89
DMS flux (Tg/yr)	19.74	19.61	19.62	19.39	18.14	17.81

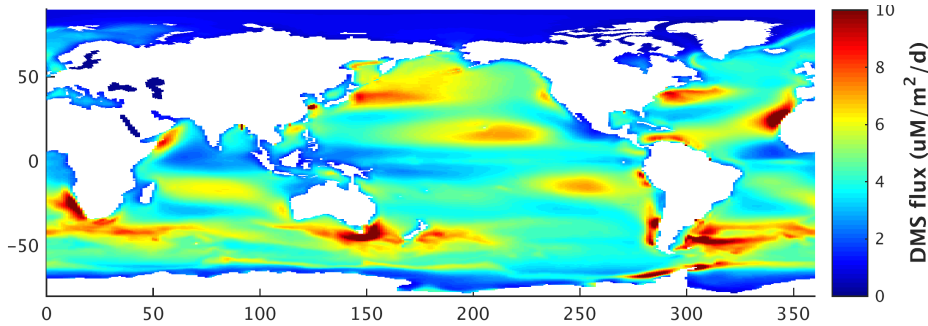
Phytoplankton cannot be ignored -*Phaeocystis*



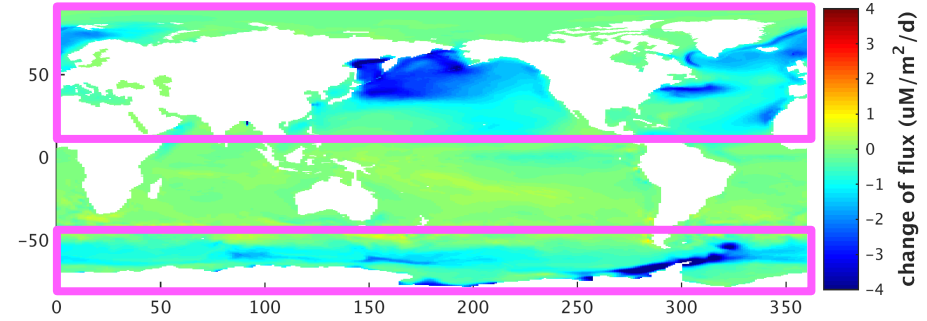
- *Phaeocystis* contributes up to 23% of upper ocean DMS during the growing seasons, and 8.6% of the annual DMS flux.
- *Phaeocystis* dominates DMS distribution at middle to high latitudes.
- None previous global DMS simulations included dynamically simulated *Phaeocystis*.

Without *Phaeocystis*

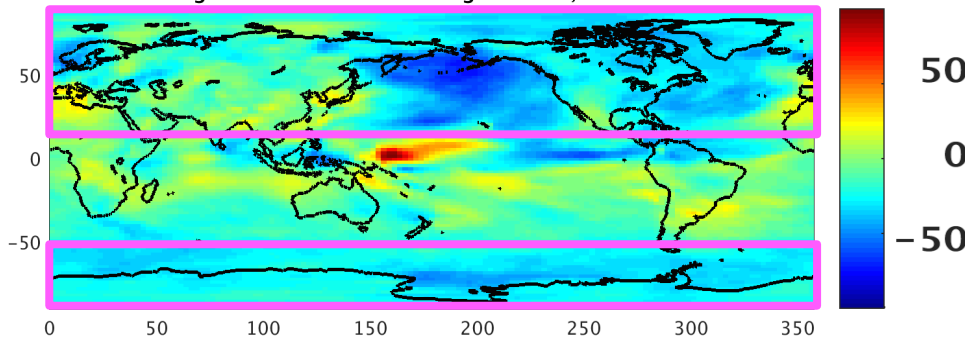
Sea-air DMS flux (2100)



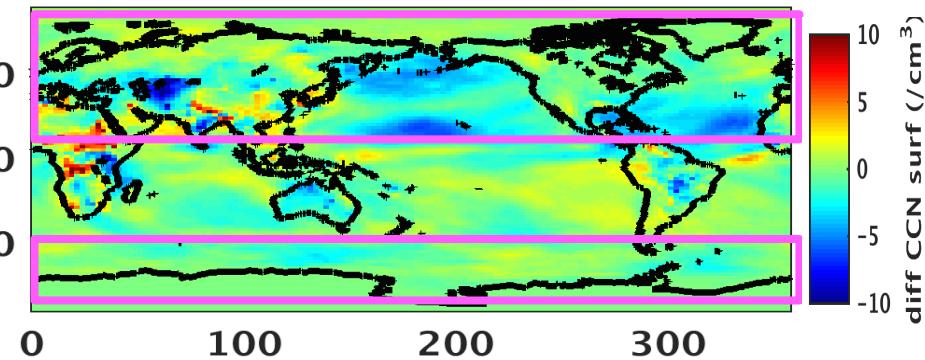
Change in DMS flux



Change in sulfate aerosol burden (ratio)



Change in surface CCN

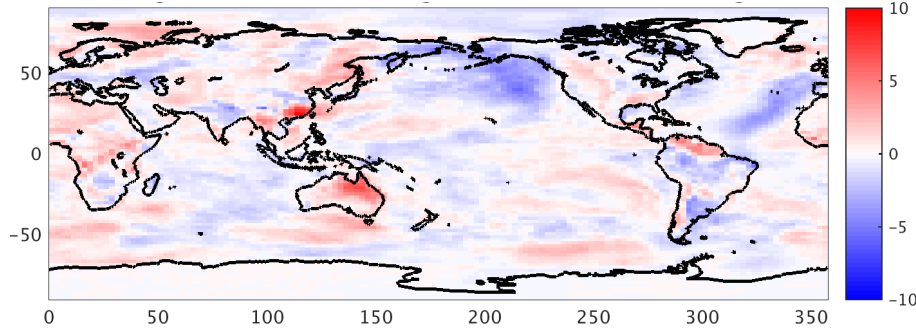


➤ *Phaeocystis* contributes 8.6% of the DMS flux

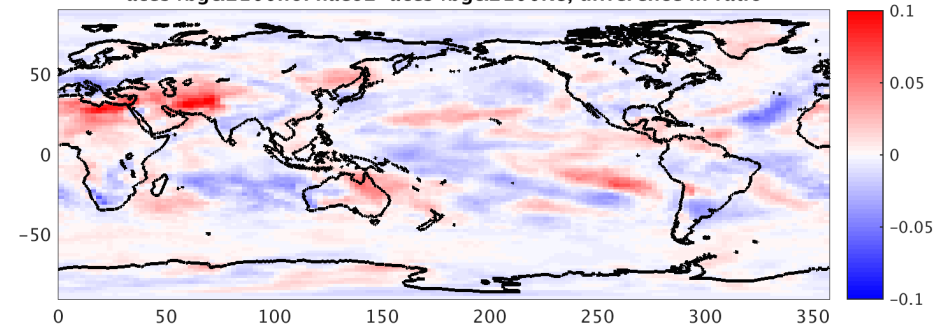


Without *Phaeocystis*

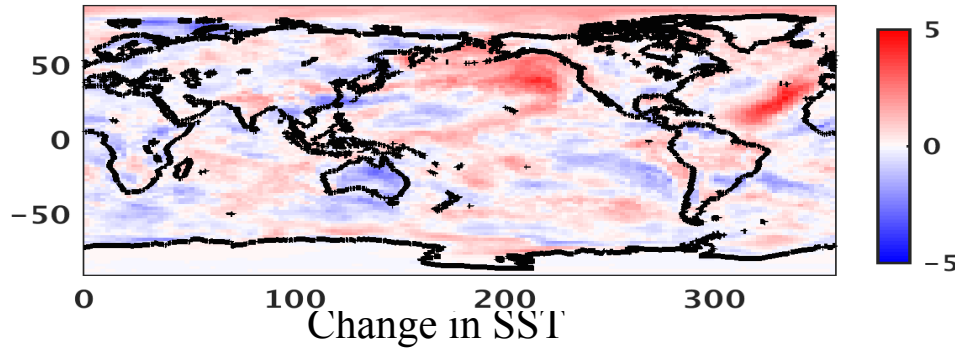
Change in LWP (g/m^2)



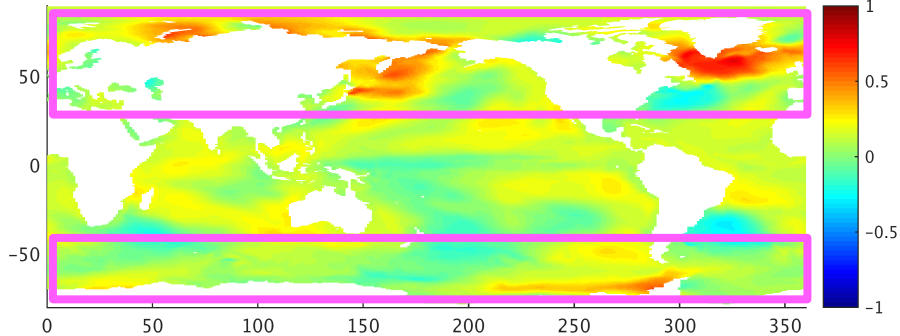
Change in cloud fraction (ratio)



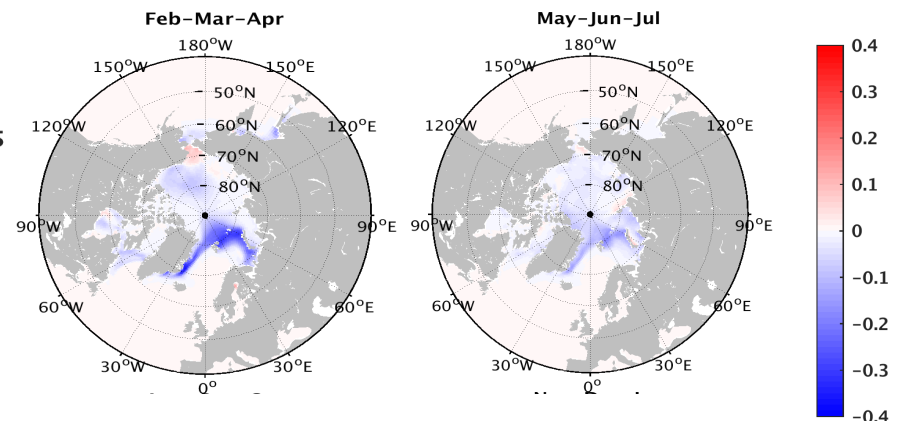
Change in total CRF (W/m^2)



Change in SST

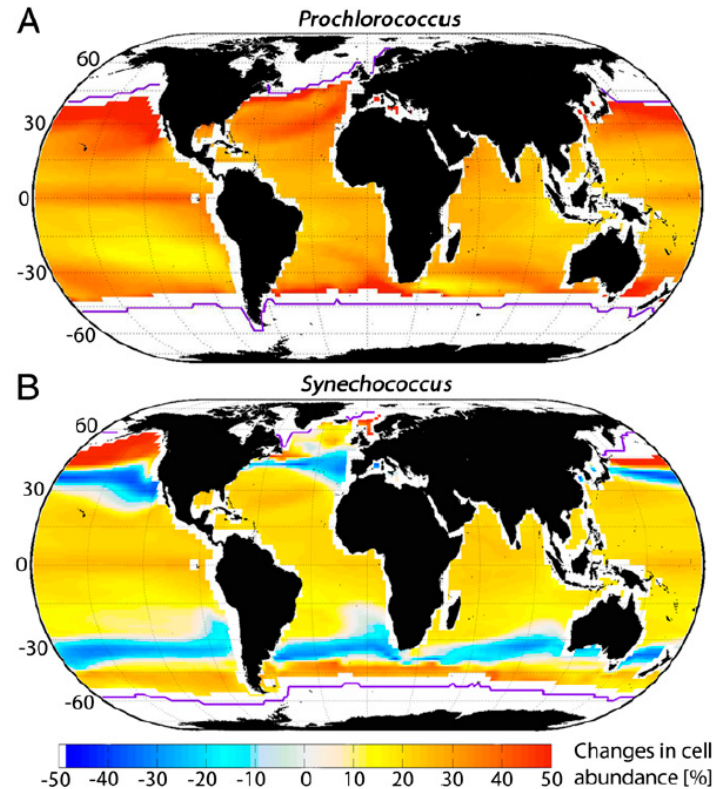


Change in IFRAC



- Without DMS produced by *Phaeocystis*, there are significant changes in the shortwave cloud radiative forcing and warms the surface ocean at high latitudes

Phytoplankton cannot be ignored - Cyanobacteria

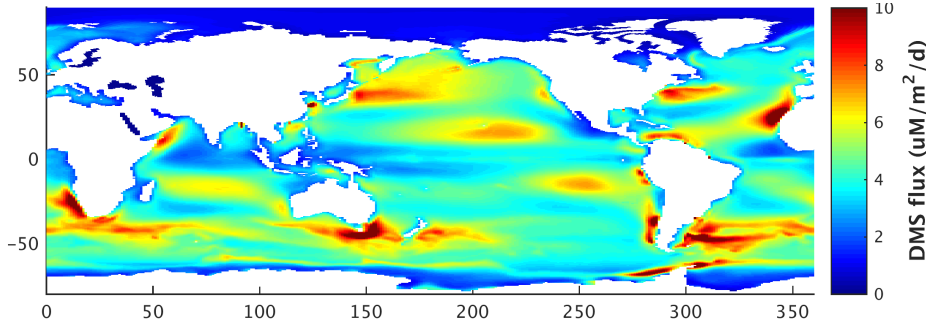


(Flombaum et al., 2013)

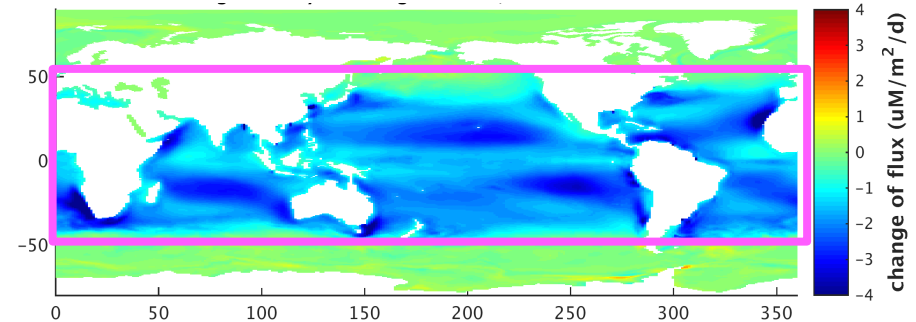
- Cyanobacteria group does not produce DMS.
- Fraction and total biomass of this group will increase in a warmer climate

Phytoplankton cannot be ignored

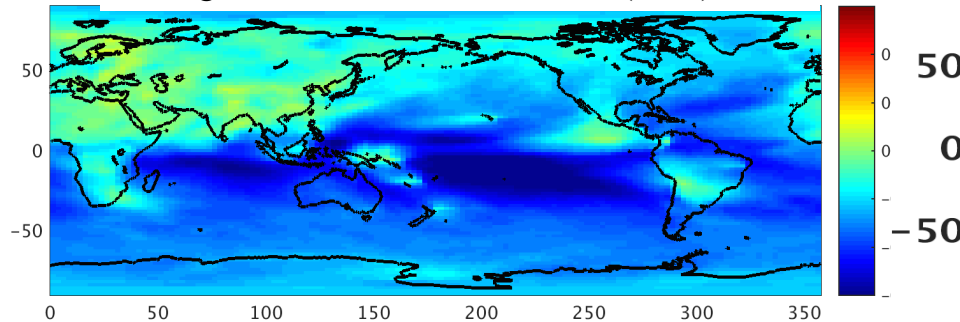
Sea-air DMS flux (2100)



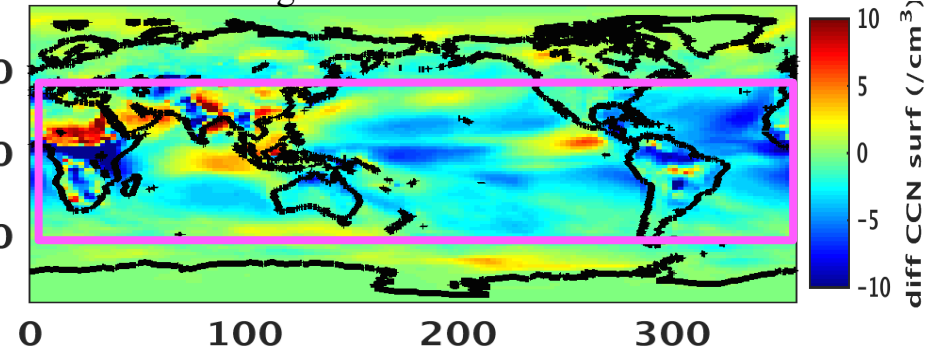
Change in DMS flux



Change in sulfate aerosol burden (ratio)



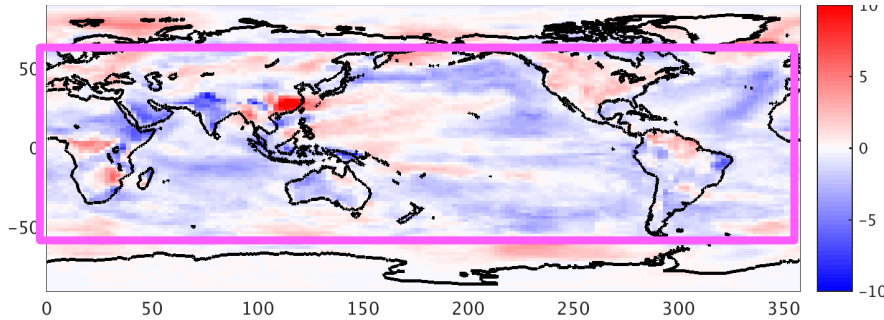
Change in surface CCN



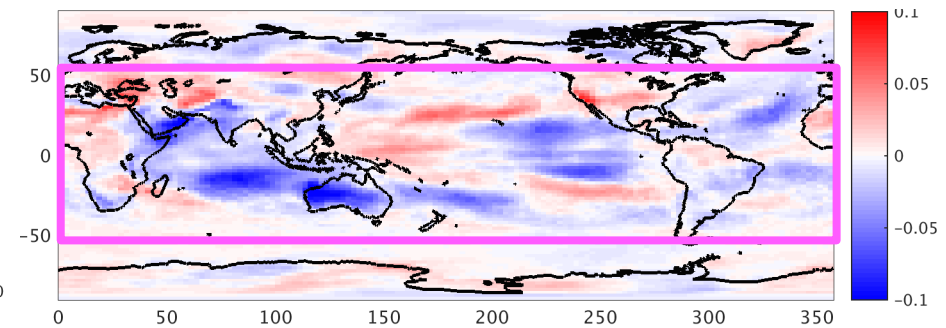
- Increasing cyanobacteria fraction leads to declines in DMS production at low-middle latitudes.
- If the maximum fraction of this group increases by 50%, DMS flux decreases $\sim 39\%$ at middle – low latitudes

Phytoplankton cannot be ignored

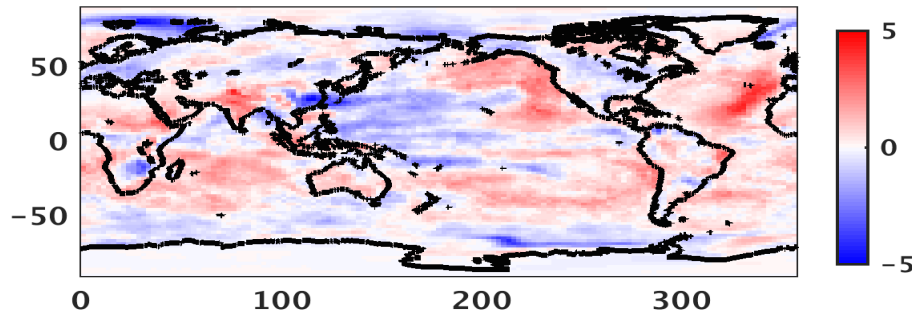
Change in LWP (g/m²)



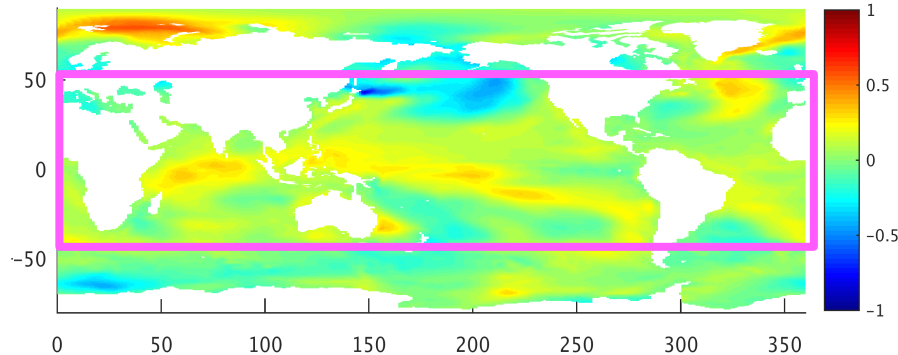
Change in cloud fraction (ratio)



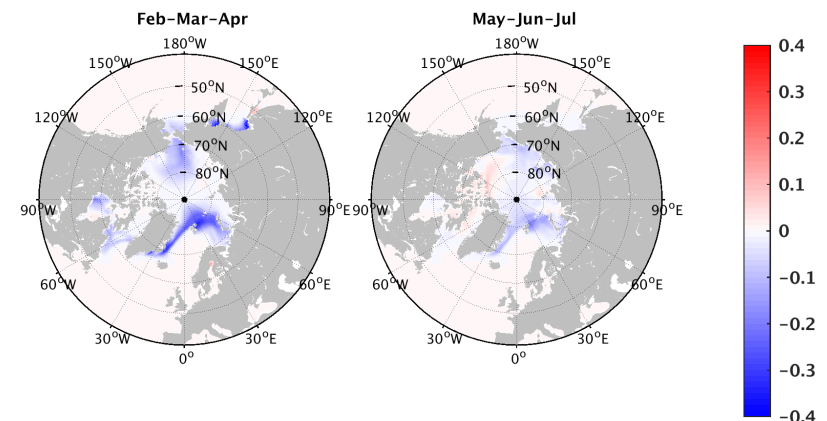
Change in total CRF (W/m²)



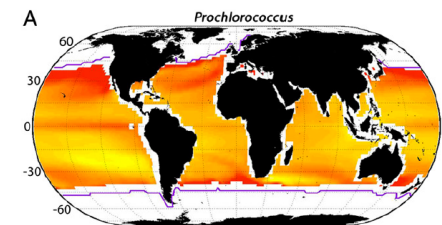
Change in SST (deg C)



Change in IFRAC



- Significant change in sw cloud radiative forcing at low to middle latitudes, $> 4 \text{ W/m}^2$ in some regions.



Summary

- Given the new explicit *Phaeocystis* representation, the DMS distribution shows significant improvements, especially regarding the amplitude and location of high latitude peaks
- The simulated mean surface DMS value is 2.26 nM. The total oceanic DMS source to the atmosphere is 20.3 Tg S/yr. Both are comparable to data-based estimates.
- Different phytoplankton groups play various roles in the DMS production. Shifts in phytoplankton community composition will feedback to climate
- DMS plays an important role in regulating climate
- Production of DMS varies with climate. It is therefore necessary to couple the dynamic DMS module in climate projections.

Thanks!

