

## **Collaborative Research: The Importance of Organo-Iodine and Iodate in Iodine-127,129 Speciation, Mobility, and Microbial Activity in Groundwater at DOE Sites**

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Iodine occurs in multiple oxidation states in aquatic systems in the form of organic and inorganic species (iodide and iodate). This fact leads to complex biogeochemical cycling of I and its long-lived isotope,  $^{129}\text{I}$ , a major by-product of nuclear fission [1-3]. In order to assess the distribution of  $^{129}\text{I}$  and stable  $^{127}\text{I}$  in environmental systems, a sensitive and rapid method was developed which enabled us to determine isotopic ratios ( $^{129}\text{I}/^{127}\text{I}$ ) and speciate I via GC-MS [3] and AMS [5]. Results using this new method demonstrate that the mobility of  $^{129}\text{I}$  species greatly depends on the type of I species and its concentration [6,7], pH [8-10], and sediment redox state [8-10], with equilibration times taking up to 12 weeks [8-10]. At ambient concentrations ( $\sim 10^{-7}\text{ M}$ ),  $\text{I}^-$  and  $\text{IO}_3^-$  are significantly retarded by sorption to mineral surfaces and covalent binding to natural organic matter (NOM), while at concentrations traditionally examined in sorption studies (i.e.,  $10^{-4}\text{ M}$  or higher),  $\text{I}^-$  travels along with the water [7]. Iodate removal can also occur through incorporation into  $\text{CaCO}_3$  crystal lattice, e.g., at the Hanford Site [12-14]. Iodide and iodate interactions with NOM lead to covalent binding of I to a limited number of aromatic carbon moieties on the particle surface [15-17]. Iodine association with NOM is important in sediments, even when organic carbon concentrations are very low (e.g.,  $<0.2\%$  at Hanford Site) [12-14]. Removal of iodine from the groundwater through interaction with NOM is complicated by the release of mobile organo-I species [15-17]. A small fraction of NOM that is bound to iodine can behave as a mobile organo-I source [15-17], a process that we were able to numerically simulate using kinetic Michaelis-Menton-type redox-reactions and kinetic uptake reactions [3]. Field [1,6,12-14,18-19] and laboratory studies evaluating the cause for steady increases in  $^{129}\text{I}$  concentrations (up to  $1000\text{ pCi L}^{-1}$ , 3 orders of magnitude greater than drinking water limits of  $1\text{ pCi L}^{-1}$   $^{129}\text{I}$ ) emanating from radiological basins at SRS indicate that an increase of 0.7 pH units in groundwater over 17 years may explain the observed increased groundwater  $^{129}\text{I}$  concentrations [20]. Bacteria from a  $^{129}\text{I}$ -contaminated aerobic aquifer at the F-area of SRS can accumulate  $\text{I}^-$  at environmentally relevant concentrations ( $10^{-7}\text{ M}$ ), but account for only a minor fraction of total added iodide (0.2-2.0%) [21-23], indicating that bacterial  $\text{I}^-$  accumulation likely does not account for the high fraction (up to 25% of total I) of measured organo-I in groundwater [6,21-23]. However, enzymatic oxidation of  $\text{I}^-$  likely plays a greater role in iodination of NOM [20-23]. The contribution of bacteria to iodide oxidizing activity likely operates via superoxide [22], and organic acid [23] production.

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