



THETA-DATA: Introduction to rotational seismology and its potential uses

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**Rotational Seismic:
measure & analyze
rotational motion at
geophones,
*in addition to linear motion.***

**Potential applications:
spatial sampling, micro-
seismic, elastic wave, . . .**

Acknowledgements

- **Geokinetics**
- **Sandia National Laboratories**
- **Schlumberger Cambridge Research**

- **Applied Technology Associates**
- **Eentec**
- **MetTech / MetSensing**

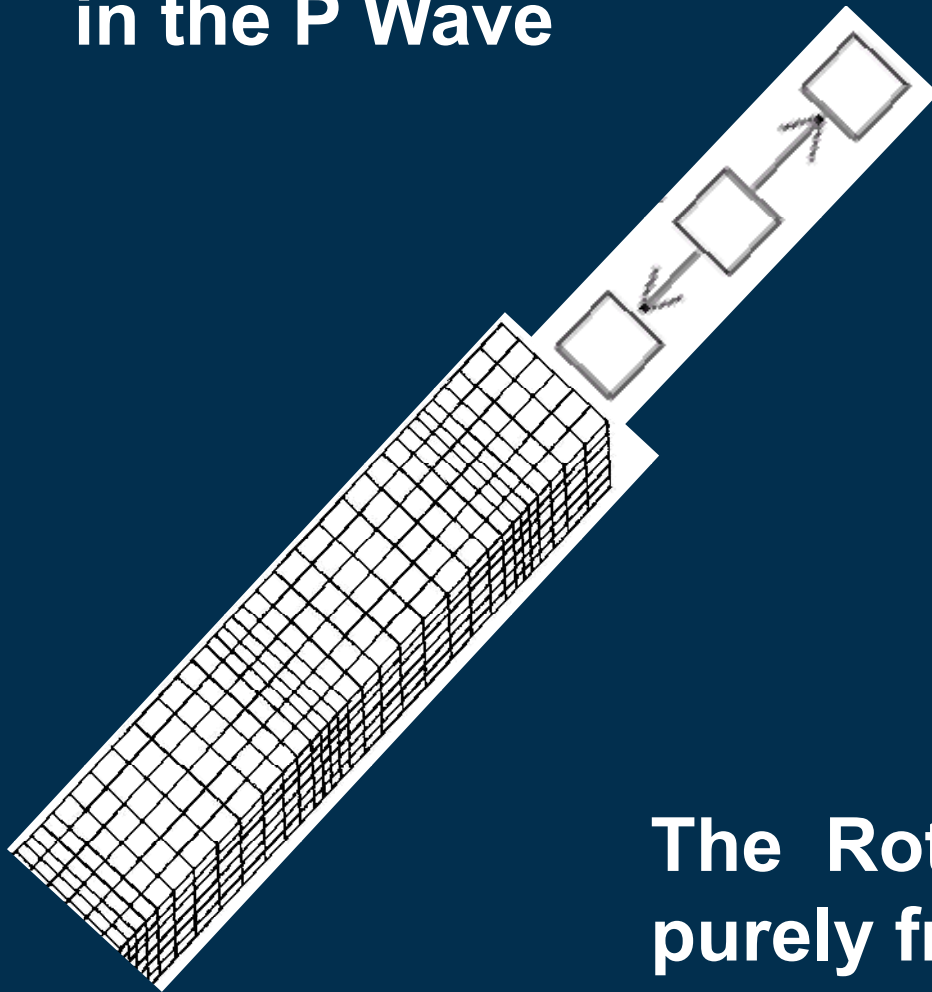
Discussions with many others . . .

Synopsis – Theta-Data Rotational Seismic

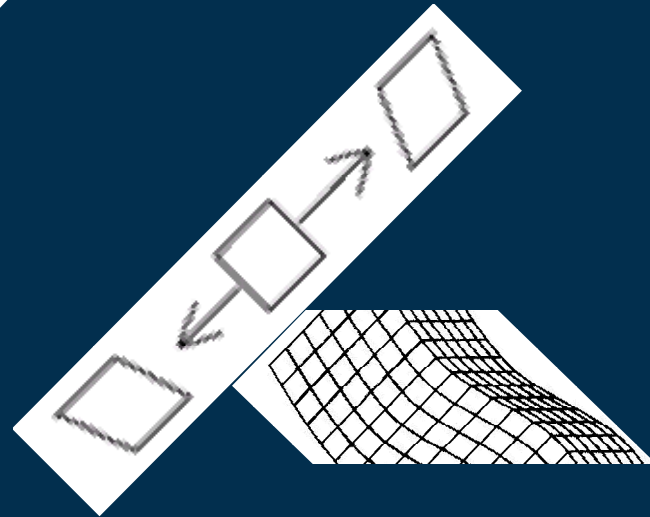
- **What is it and where did it come from?**
- **Sensors - How do we measure it?**
- **What does Theta-Data look like?**
- **What are possible Applications?**
 - **Spatial sampling enhancement**
 - **Shear Wave selectivity – microseismic**
 - **‘other’ uses**
- **Status and ‘Road Ahead’**

Rotational Seismic – Shear Selectivity

There is **NO ROTATION**
in the P Wave



There **IS ROTATION**
in the S wave



The Rotation Signal is
purely from the S wave

3-Theta Data – What is it?

Measure **Rotation**
around 3 axes - 3θ

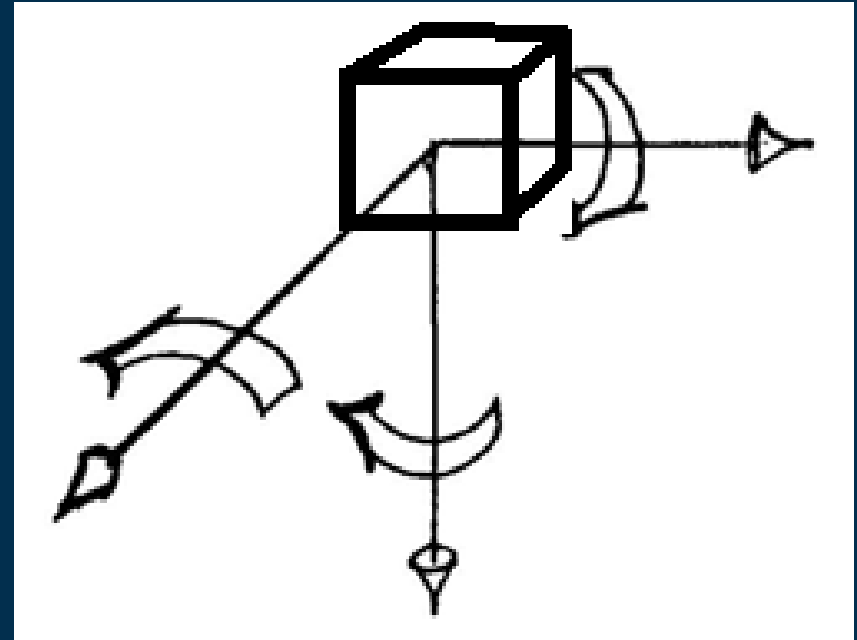
θ -x

θ -y

θ -z

Vector Curl
of displacement

Record **3- θ** Rotation in
addition to 3-C or 4-C



“Theta Data”
is comprised of
“Curl Traces”

Six Degrees-Of-Freedom (6-DOF) Seismic

Classical Mechanics:

Need **Roll, Pitch, Yaw**, as well as **x, y, z**

To **Fully Specify Motion**

for a finite body (“elemental cube”)

Y rotation =

Z gradient of **X** linear

▪ **X** gradient of **Z** linear

$$\underline{\theta} = \begin{bmatrix} \theta_x \\ \theta_y \\ \theta_z \end{bmatrix} \equiv \frac{1}{2} \nabla \times \underline{u}$$

$$\theta_y = \frac{1}{2} \left(\frac{\partial u}{\partial z} - \frac{\partial w}{\partial x} \right)$$

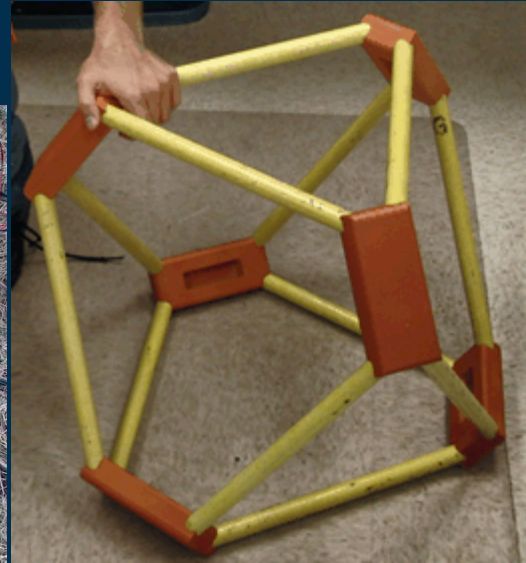
Origins – Where Did Rotational Seismic ‘Come From’?

- **Earthquake Seismology:**
 - **William H. K. Lee (USGS) & others**
 - **Int’l Working Group on Rotational Seismology (IWGRS)**
 - **BSSA (May, 2009)**
- **Sensor technology development - *decades*:**
 - **Defense (US and USSR)**
 - **inertial guidance, etc.**

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Rotational Sensors

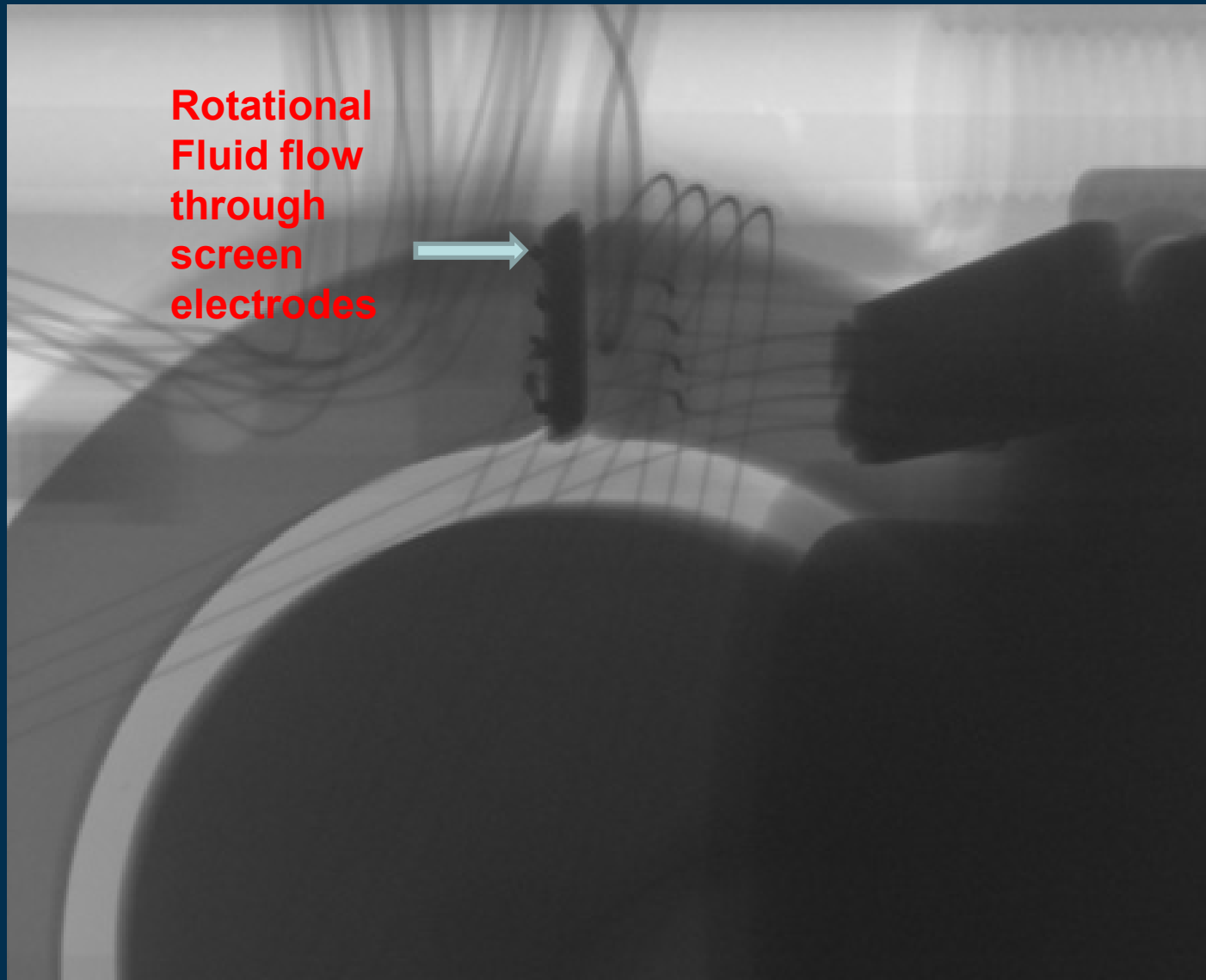


Typically flat spectral response for angular velocity.

Electrokinetic Sensors

- **Simple** Concept:
 - Ionic fluid in container is an inertial mass.
 - The ionic fluid flows through screen electrodes when there is motion.
 - Electrical current due to ionic fluid flow is a measure of **velocity** (linear or angular).
- Fluid in cylinder for **linear** motion
- Fluid in toroidal 'donut' for **rotational** motion
- Electrokinetic Sensor characteristics:
 - **Low noise** (order of 0.1 micro-radians sec)
 - **Low frequencies** possible

Electrokinetic Rotational Sensor

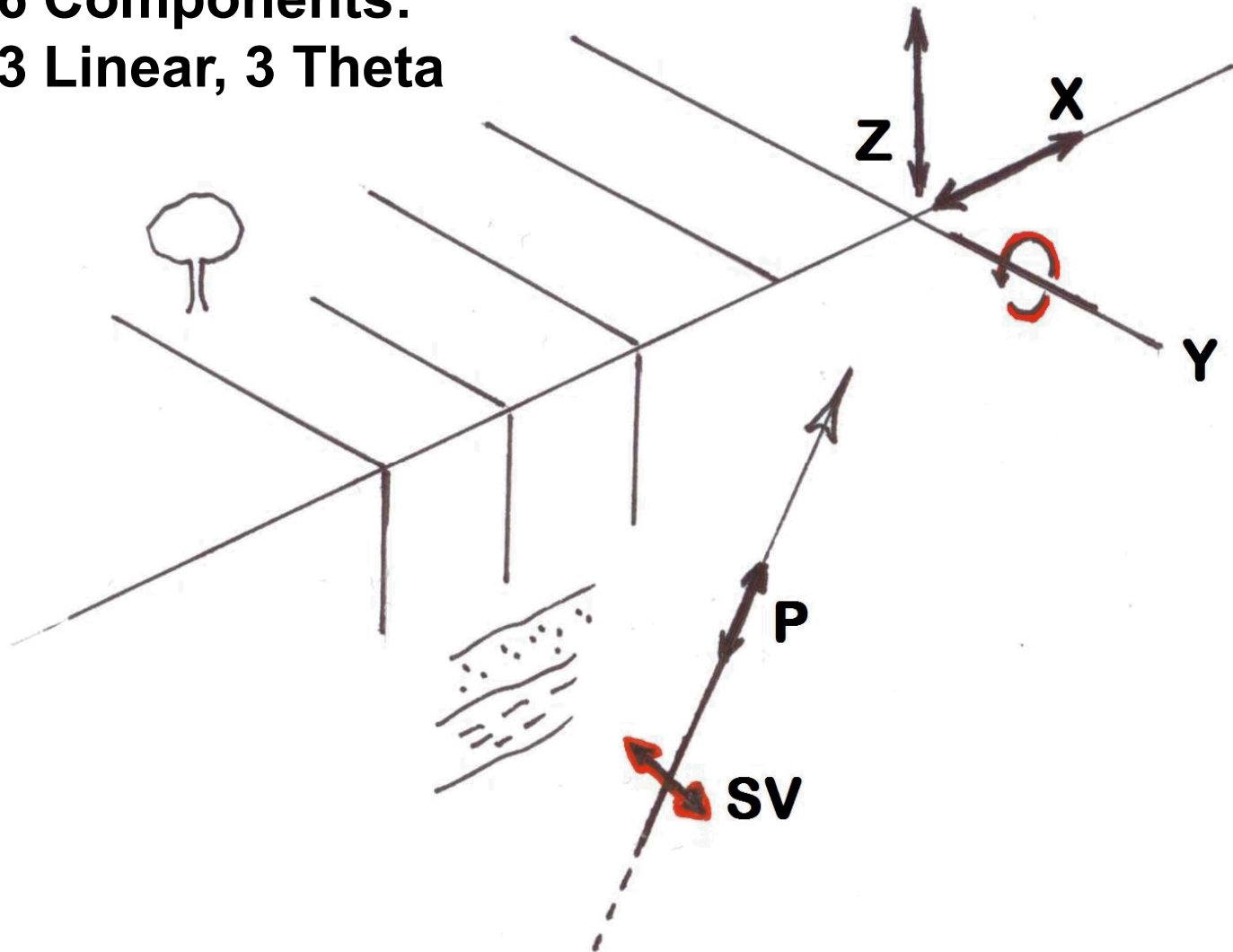


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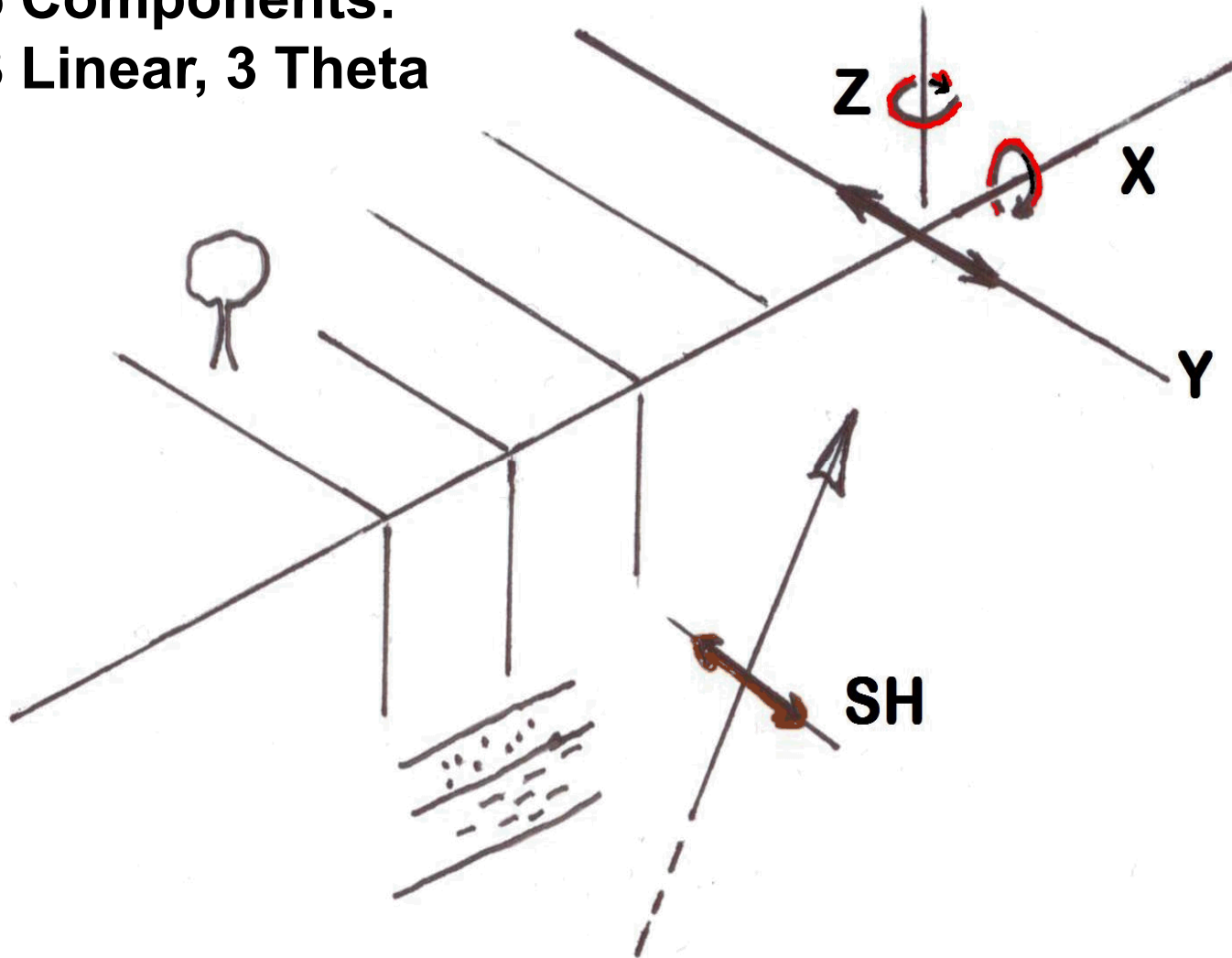
Transverse Rotational – TR Mode

**6 Components:
3 Linear, 3 Theta**

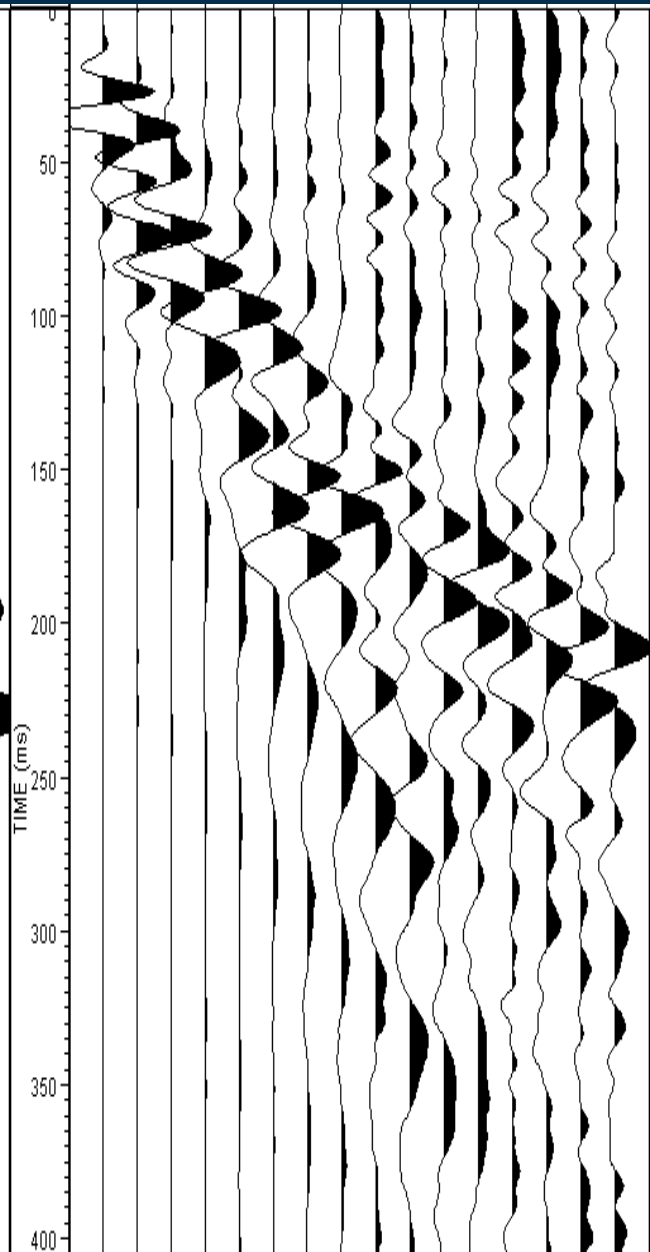
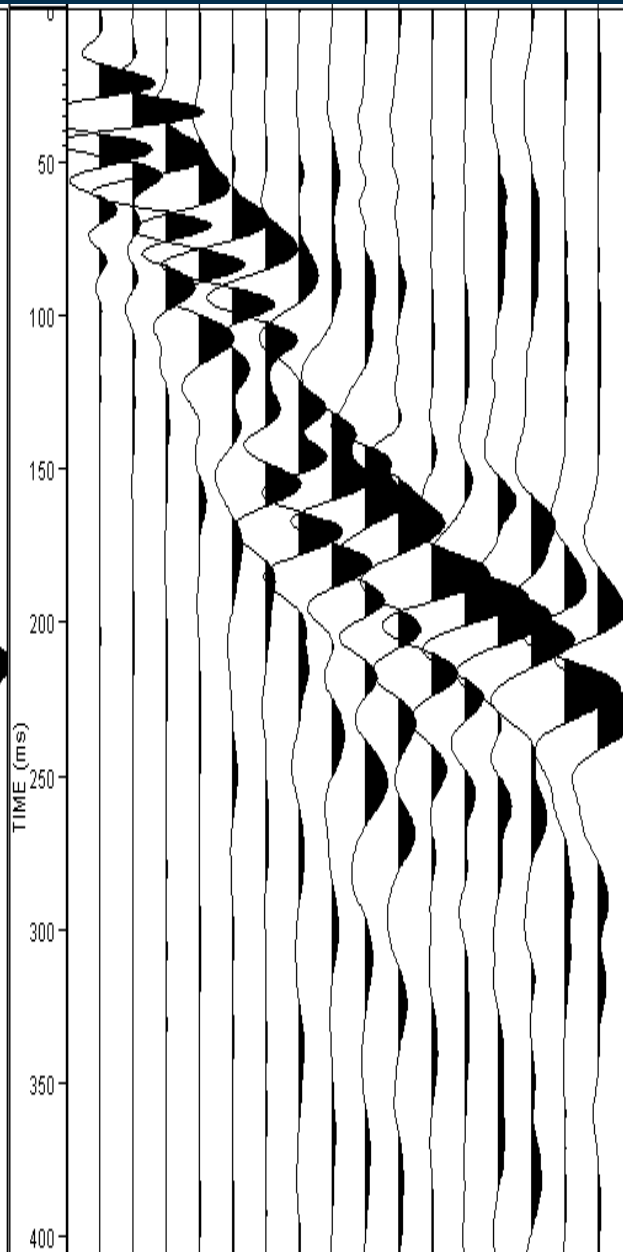
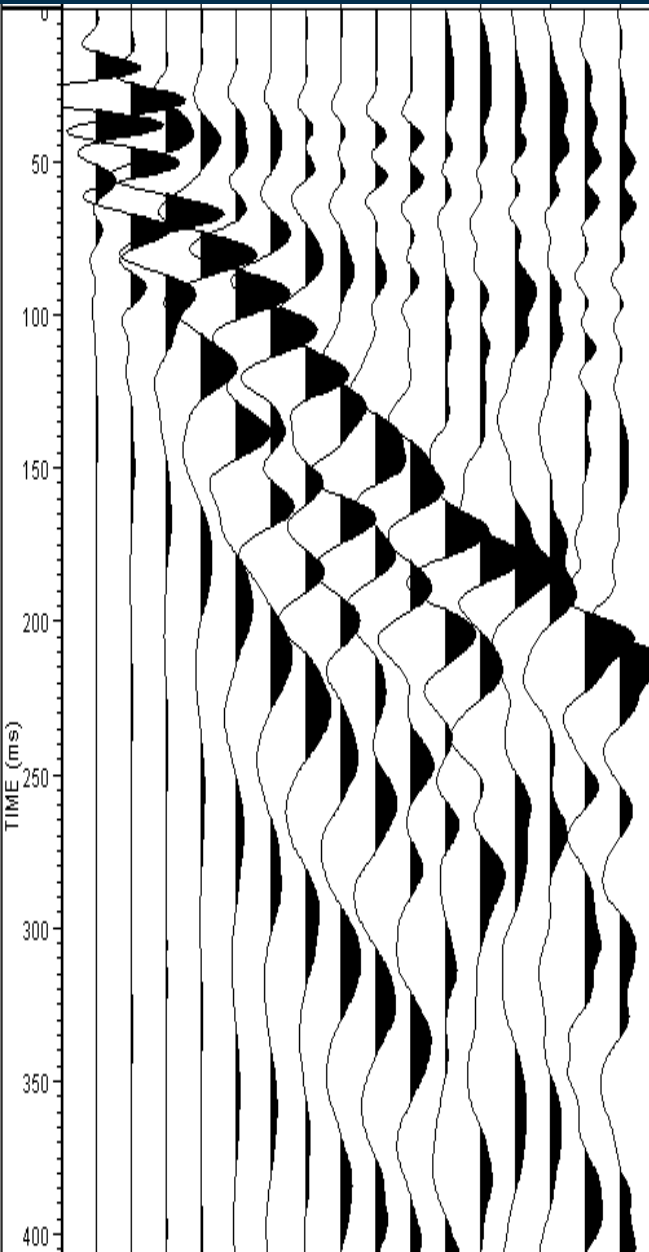


Transverse Linear – TL Mode

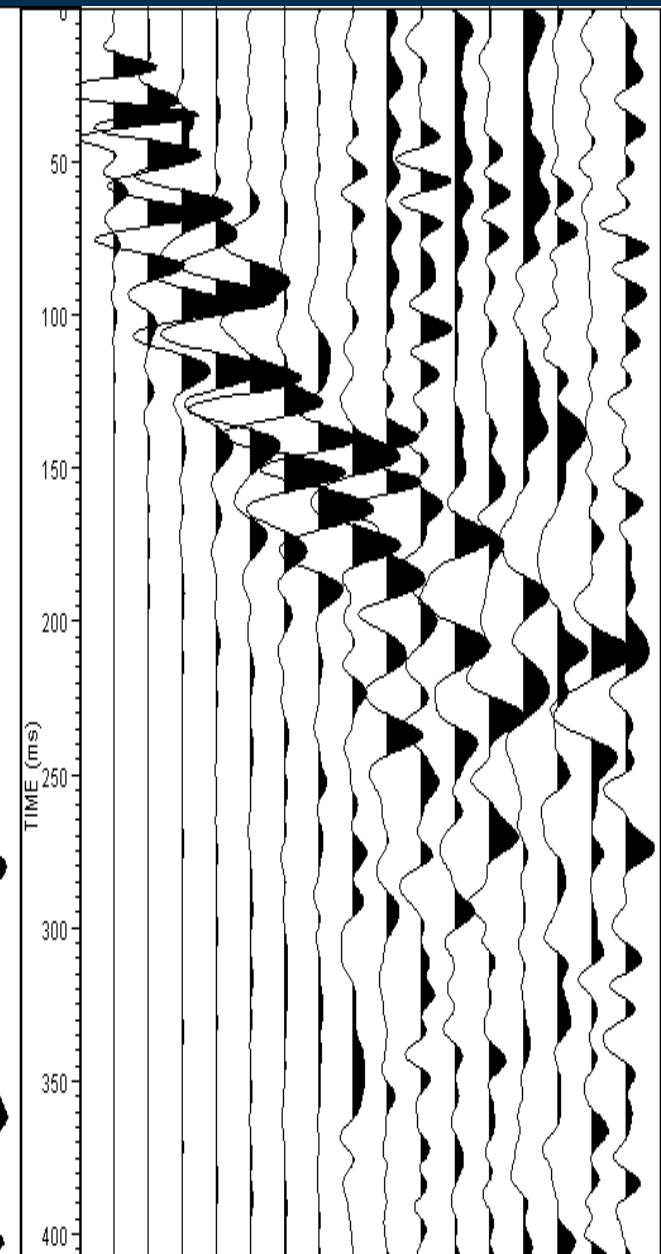
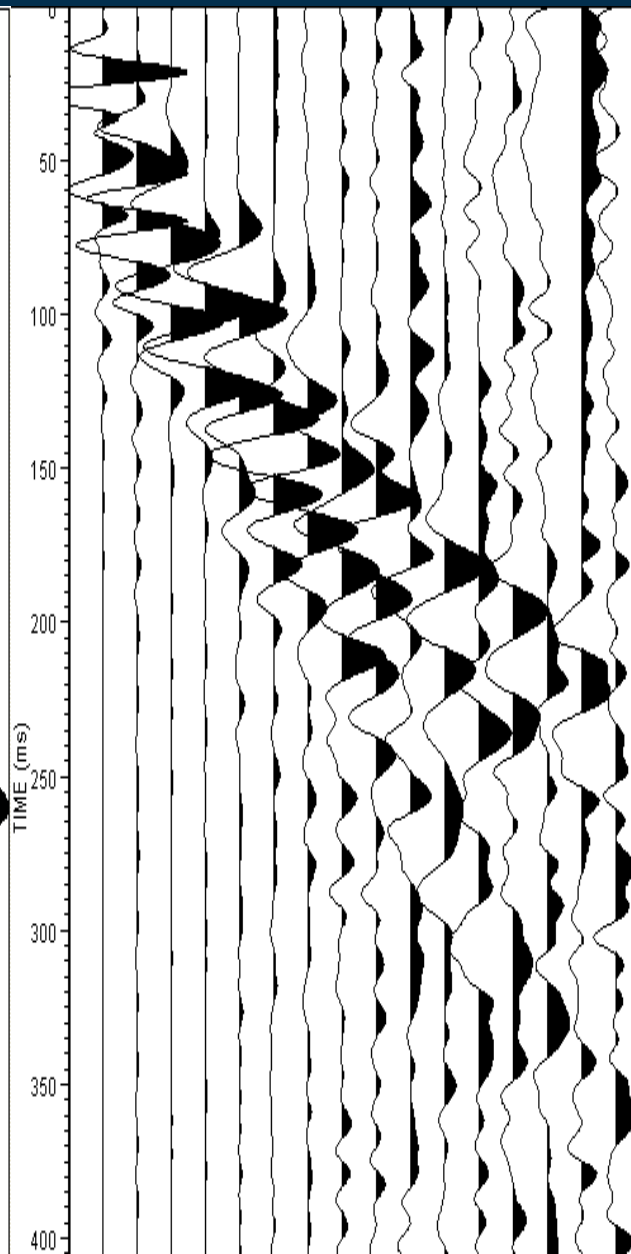
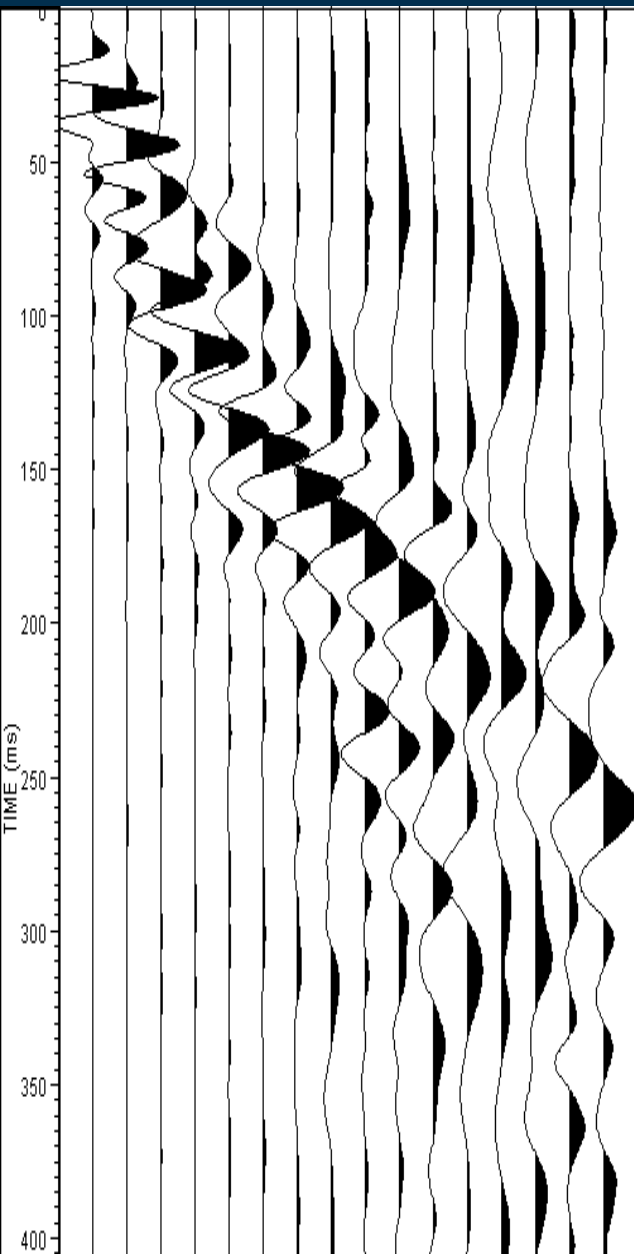
**6 Components:
3 Linear, 3 Theta**



PSV - 3C-Vert., 3C-Radial, 3 θ -Trans.



SH - 3C-Trans., 30-Vert., 30-Radial

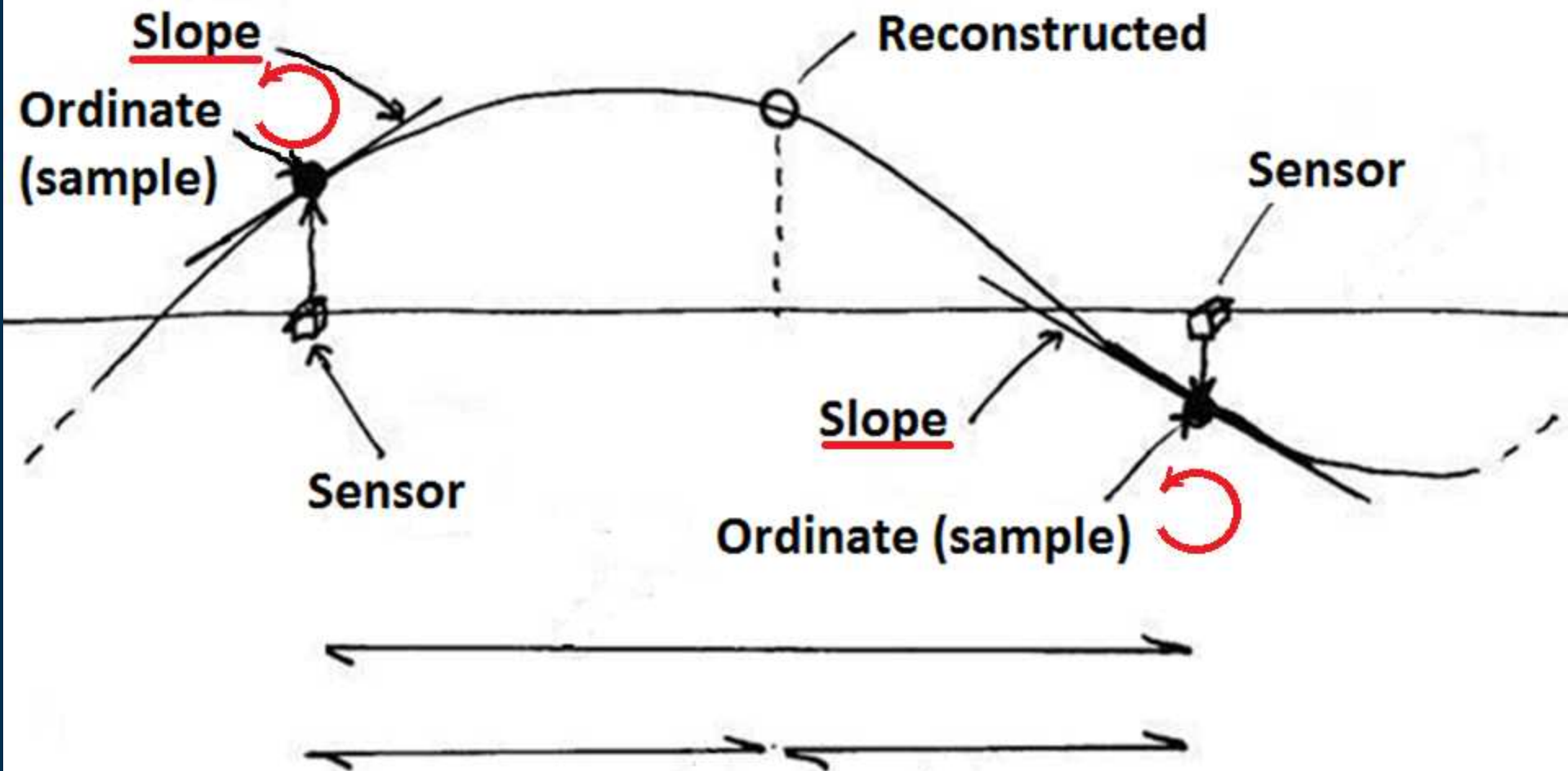


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Ordinate and Slope Sampling

DOUBLE the Effective Spatial Nyquist



For Z linear component:

Rotational yields the Spatial Slopes

At **Free Surface**, or **Water Bottom**, have zero traction/stress component, and thus zero strain component:

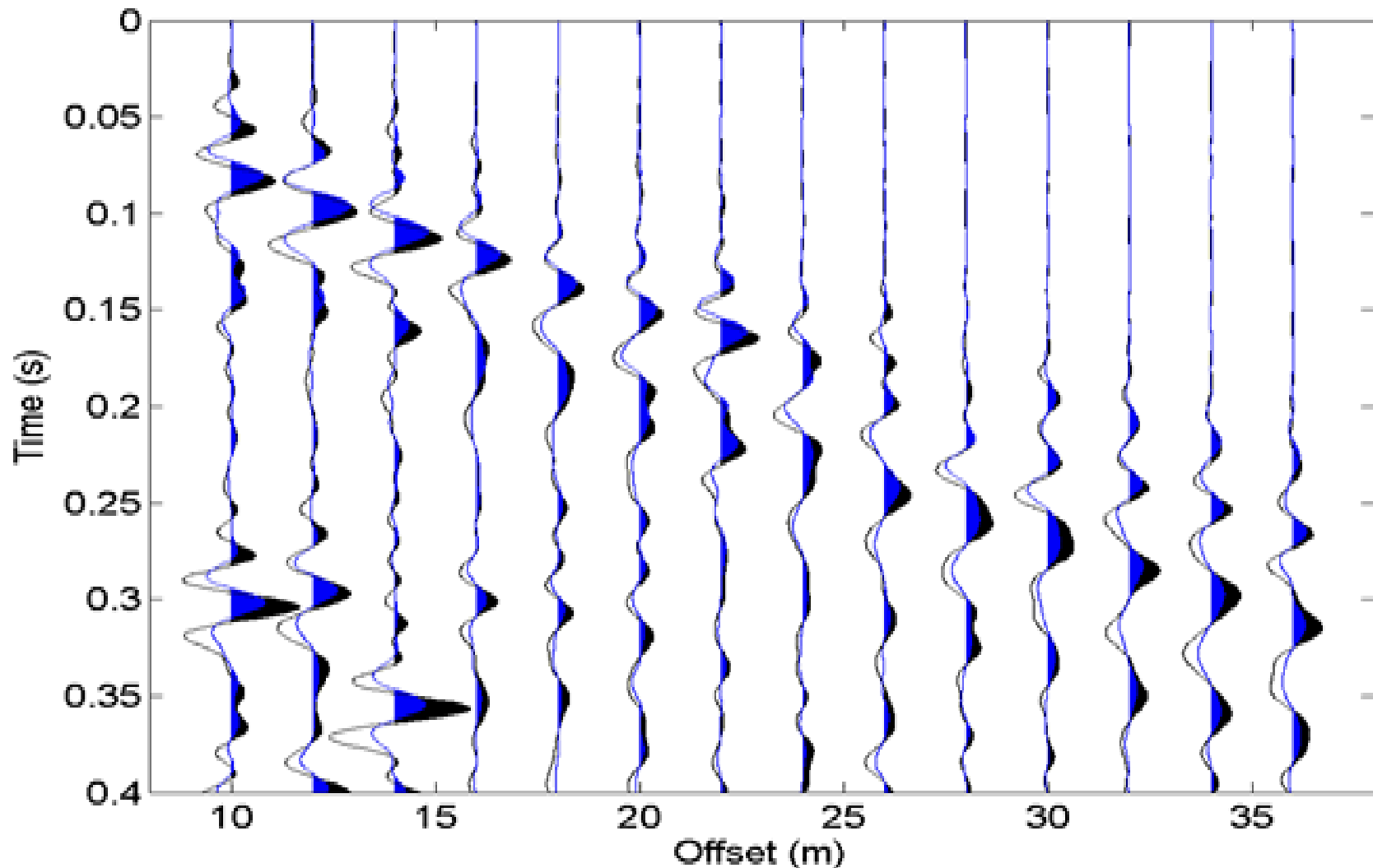
$$e_{yz} = \frac{1}{2} \left(\frac{\partial v}{\partial z} + \frac{\partial w}{\partial y} \right) = 0$$

Thus we have the **y spatial gradient** of the vertical linear component of displacement, w, given by the **x component of rotational** Theta Data:

$$\theta_x = \frac{\partial w}{\partial y}$$

Rotational Spatial Sampling – Horizontal Gradients of Linear Z

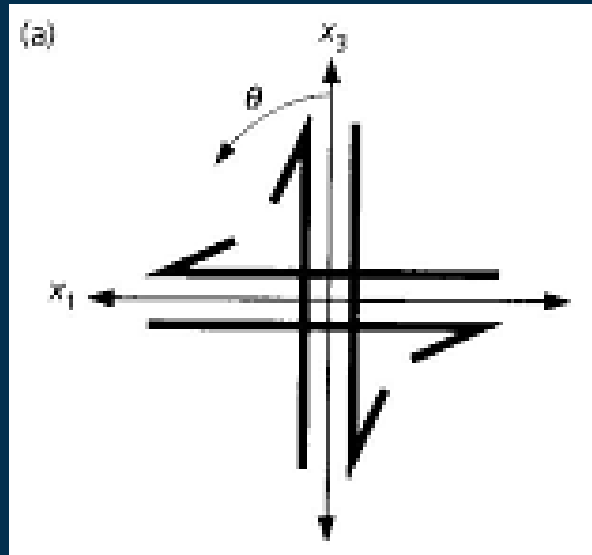
Everhard Muyzert, Schlumberger, EAGE, 2012



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Frac'ing / Micro-Seismic – Double Couple



P and S body waves

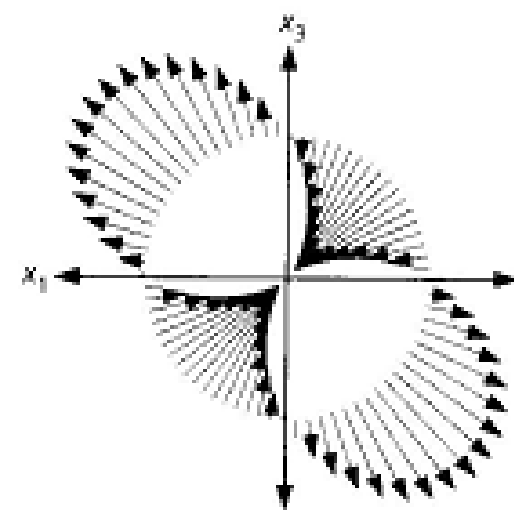
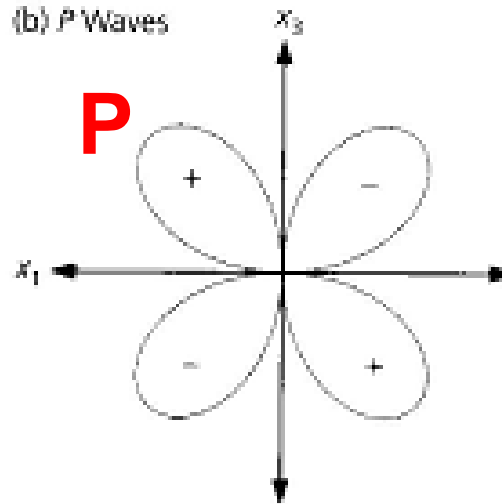
Motion **Amplitude**

And

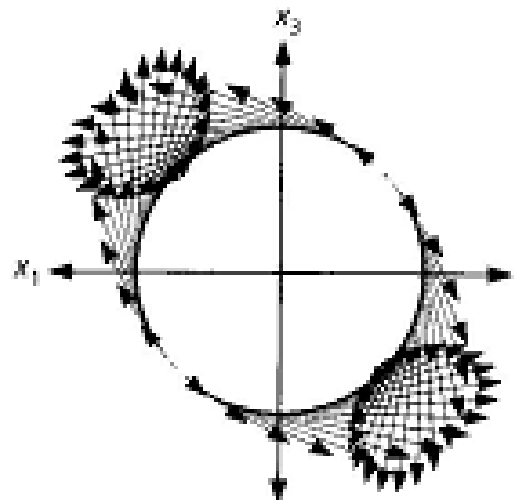
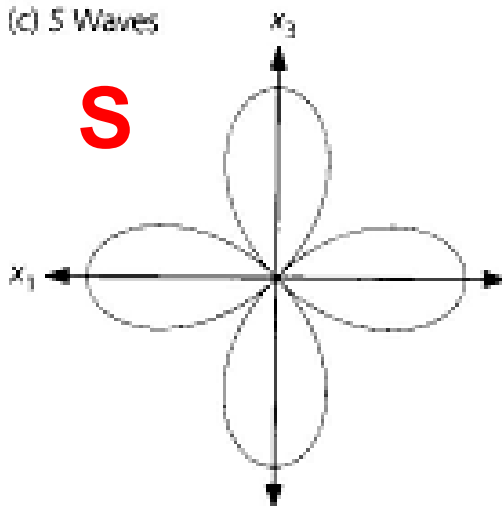
Motion **Direction**

S is STRONGER

(b) P Waves



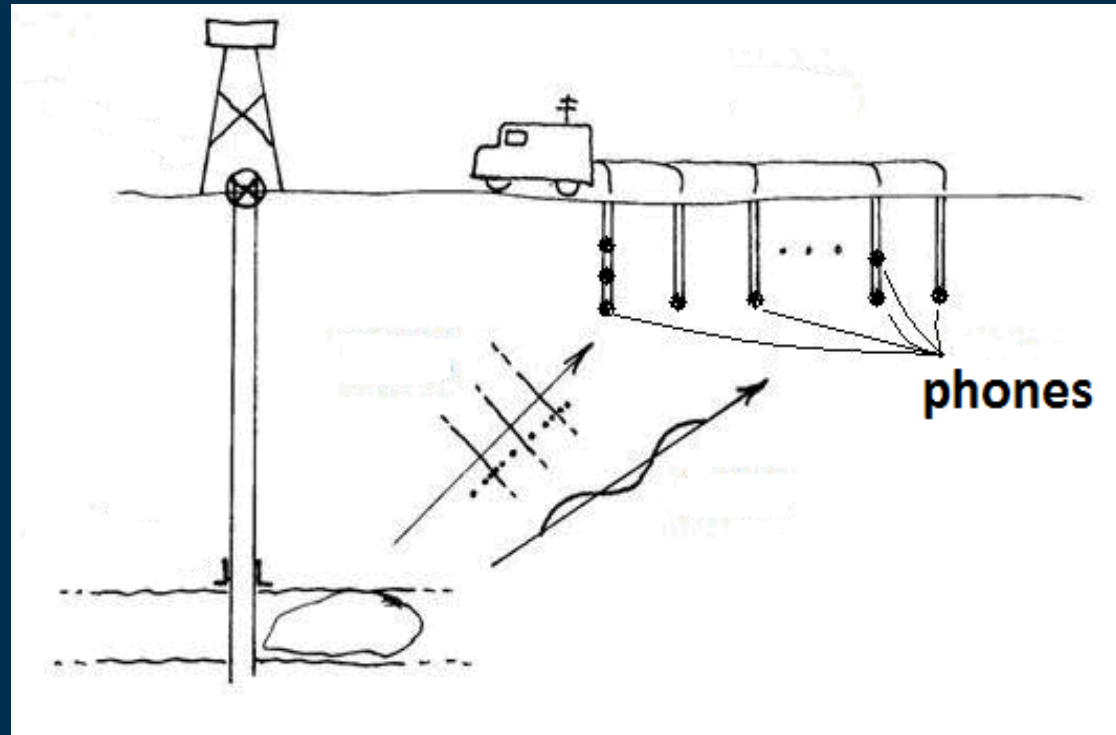
(c) S Waves



Micro-Seismic / Double Couple – Shear Detection

Rotational phones
below the free
surface, **away from
strong interfaces,**
and the associated P
& S **conversions.**

Shear Selectivity of
Theta Data possibly
may enhance
extraction of shear
body wave data



Induced Seismicity: Frac Rotational Monitoring ?

From earthquake seismology:

“Observations . . . showed that the amplitudes of **rotations** can be **one to two orders of magnitude greater than expected** from the classical elasticity theory”

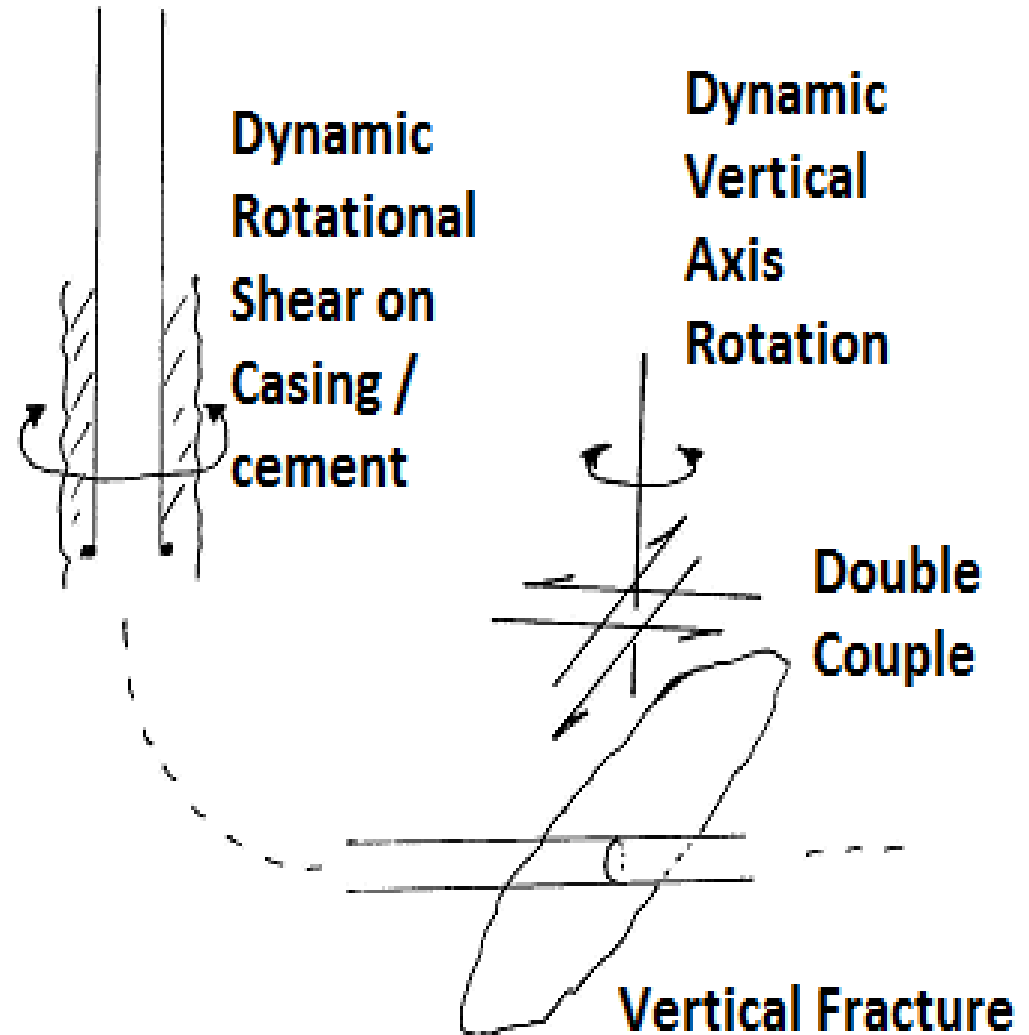
*Lee, Igel, & Trifunac, **2009**, “Recent Advances in Rotational Seismology”, Seis. Res. Ltrs., vol. 80,no. 3, May/June*

Frac Rotational Monitoring ?

**Vertical axis
dynamic rotation.**

**Release pre-
existing tectonic
stress.**

**Shear bond
strength of
cement-casing is
comparatively
weak.**

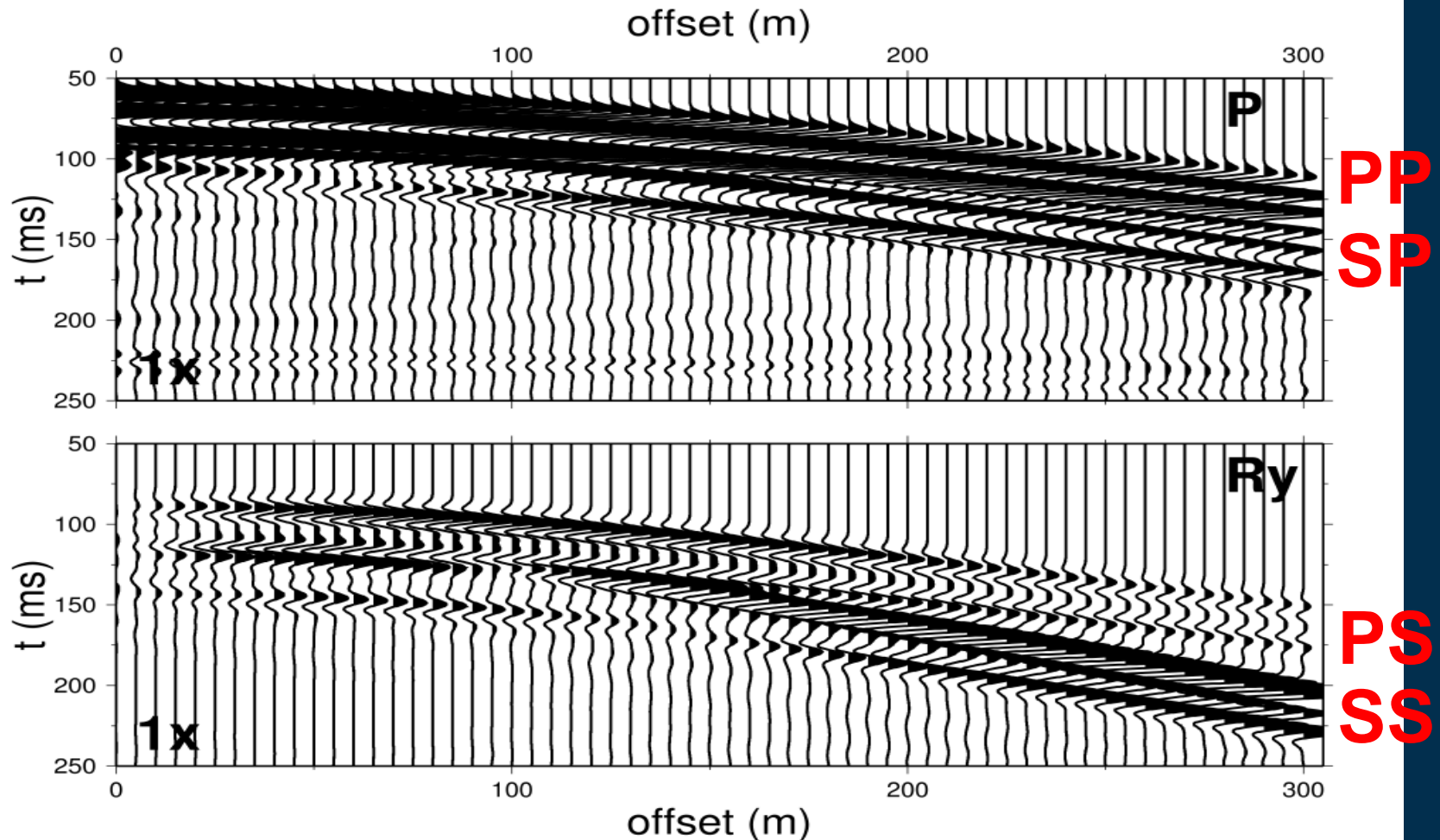


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Pressure (P wave only) + TR-Theta (S wave only)
*AVO for **gas** sandstone*
No Free Surface, buried phones

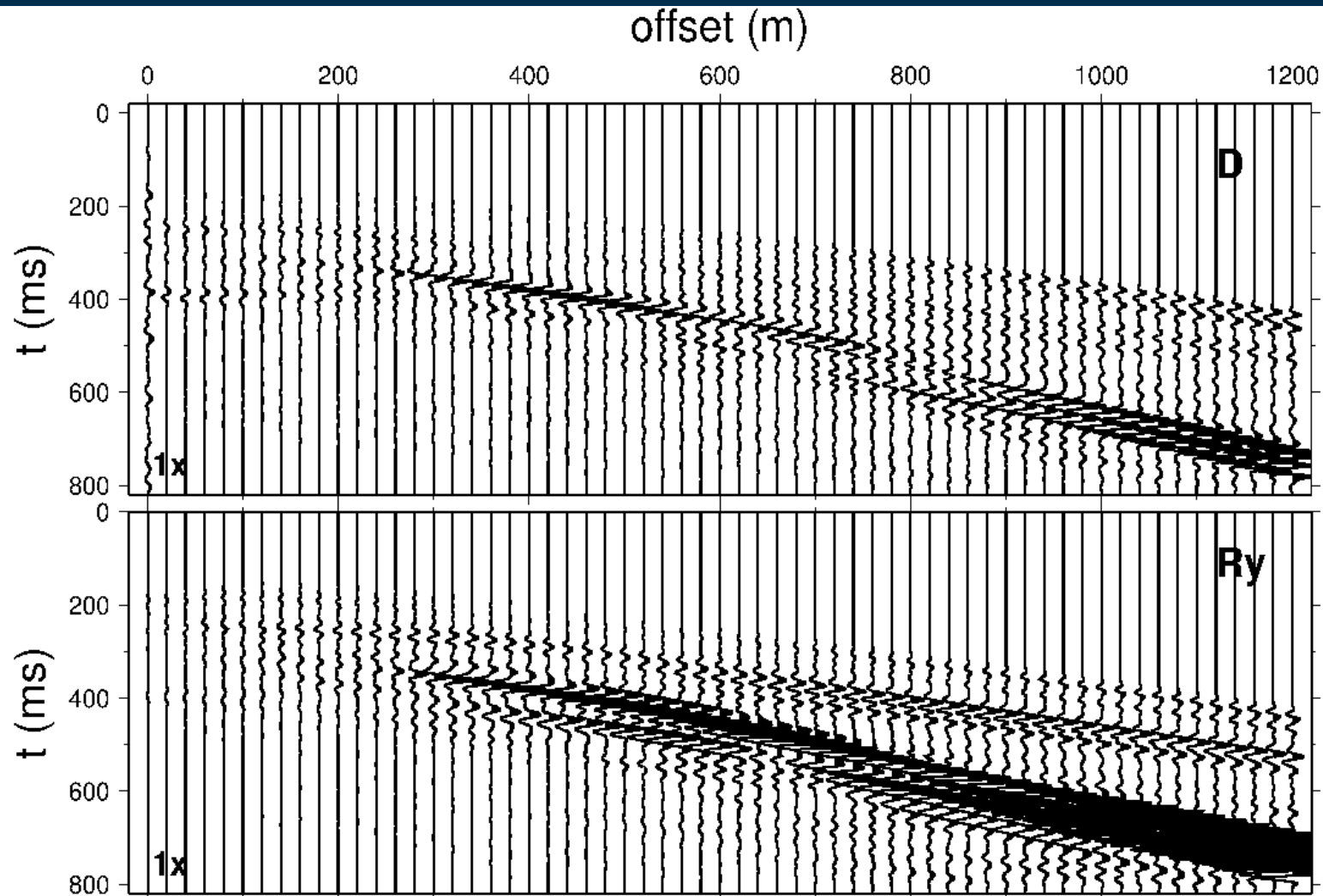
David Aldridge, et al., Sandia National Laboratories, 2007-2009



Pressure (P wave only) + TR-Theta (S wave only)

Free Surface, AVO for *brine* saturated sandstone

David Aldridge, et. al., Sandia National Lab, 2007-2009



Model: 25 m thick sandstone, $\phi = 0.25$, $\tau = 2.5$, $k_0 = 1D$

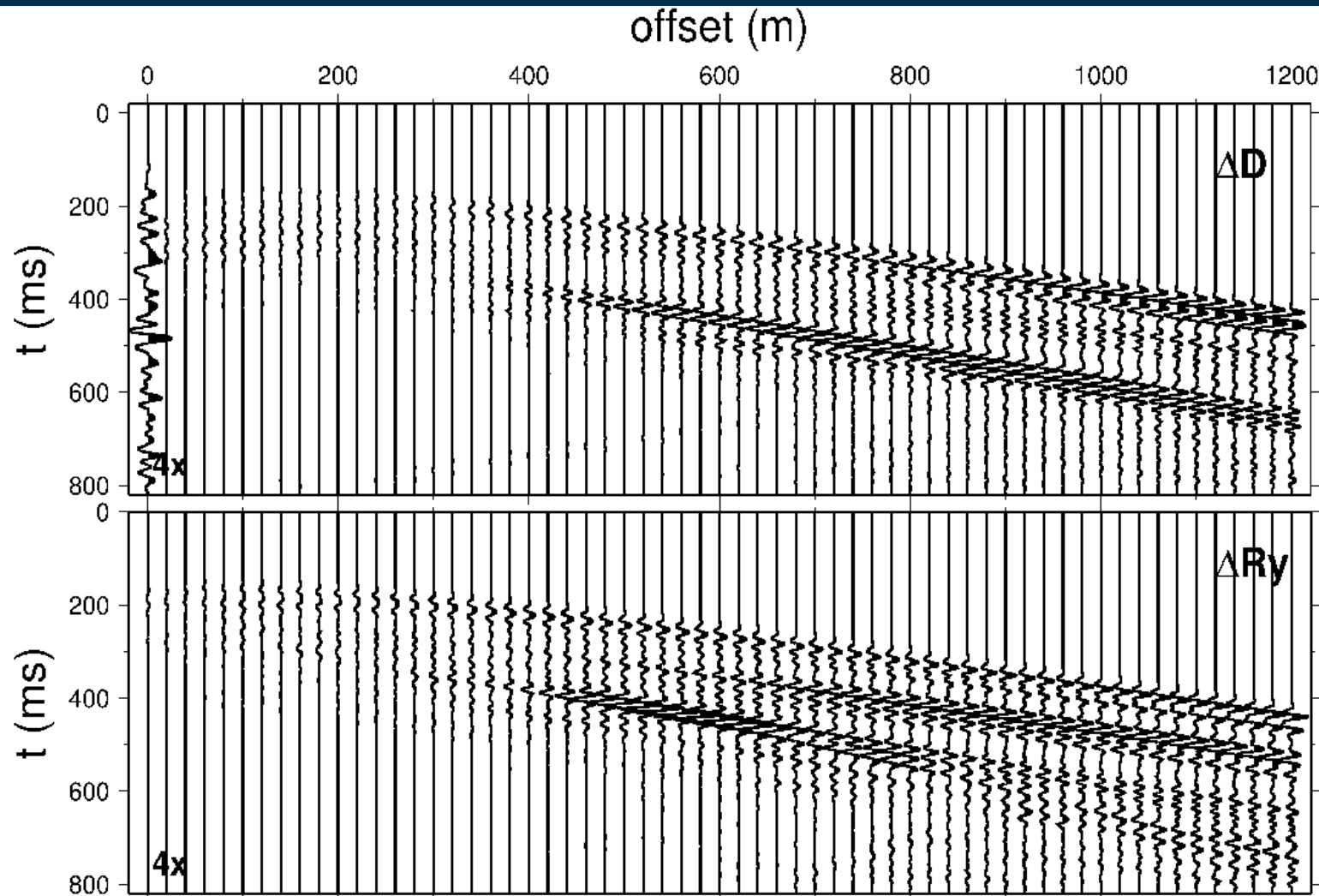
Source: Fz force, 30 Hz Ricker wavelet

Traces: 100%H2O; plot amp = 1.5×10^{-15}

Pressure (P wave only) + TR-Theta (S wave only)

Delta AVO for 25% CO₂ saturated sandstone

David Aldridge, et. al., Sandia National Lab, 2007-2009



Model: 25 m thick sandstone, $\phi = 0.25$, $\tau = 2.5$, $k_0 = 1D$

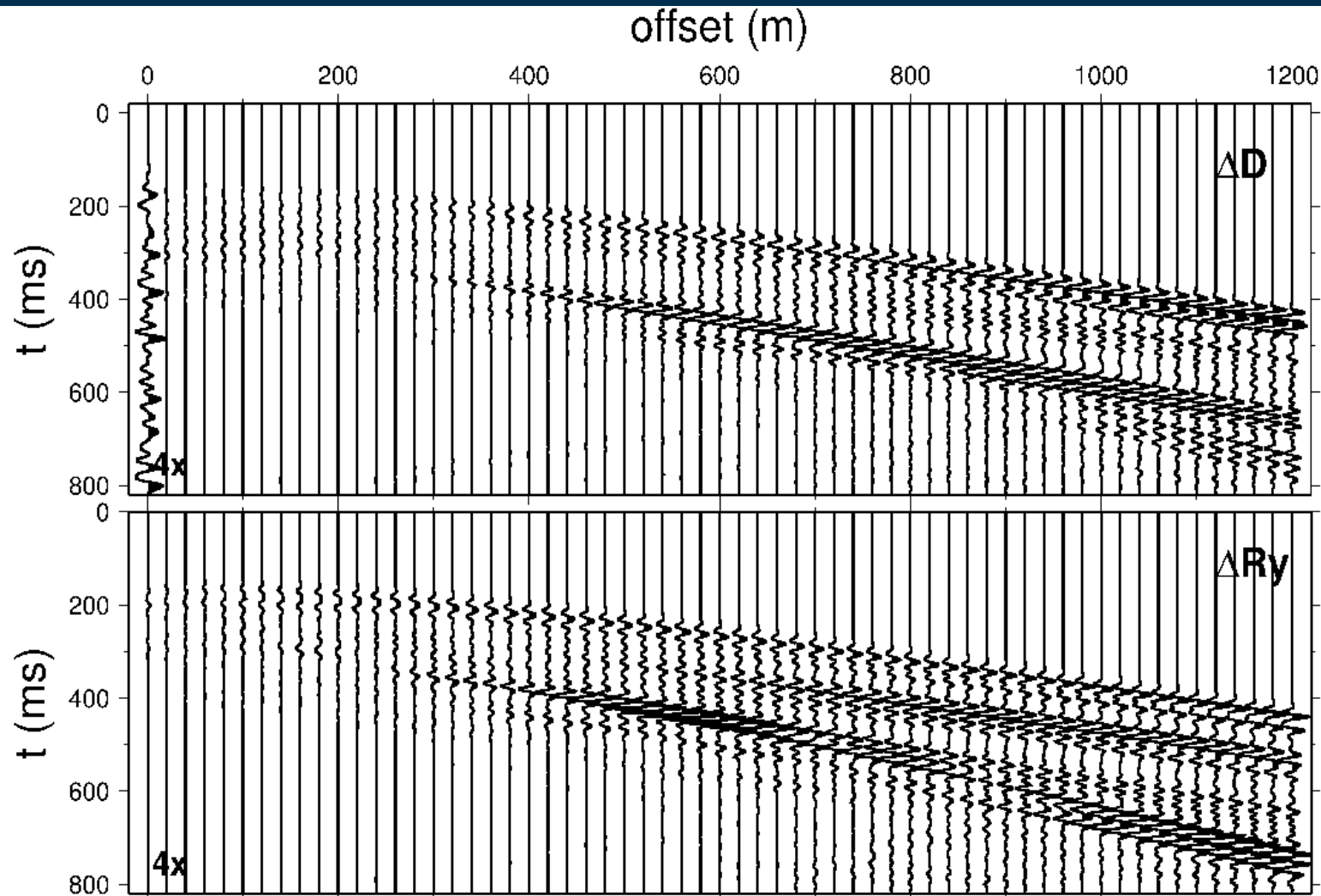
Source: Fz force, 30 Hz Ricker wavelet

Traces: 100%H₂O-25%CO₂; plot amp = 0.375×10^{-15}

Pressure (P wave only) + TR-Theta (S wave only)

Delta AVO for 50% CO₂ saturated sandstone

David Aldridge, et. al., Sandia National Lab, 2007-2009



Model: 25 m thick sandstone, $\phi = 0.25$, $\tau = 2.5$, $k_0 = 1D$

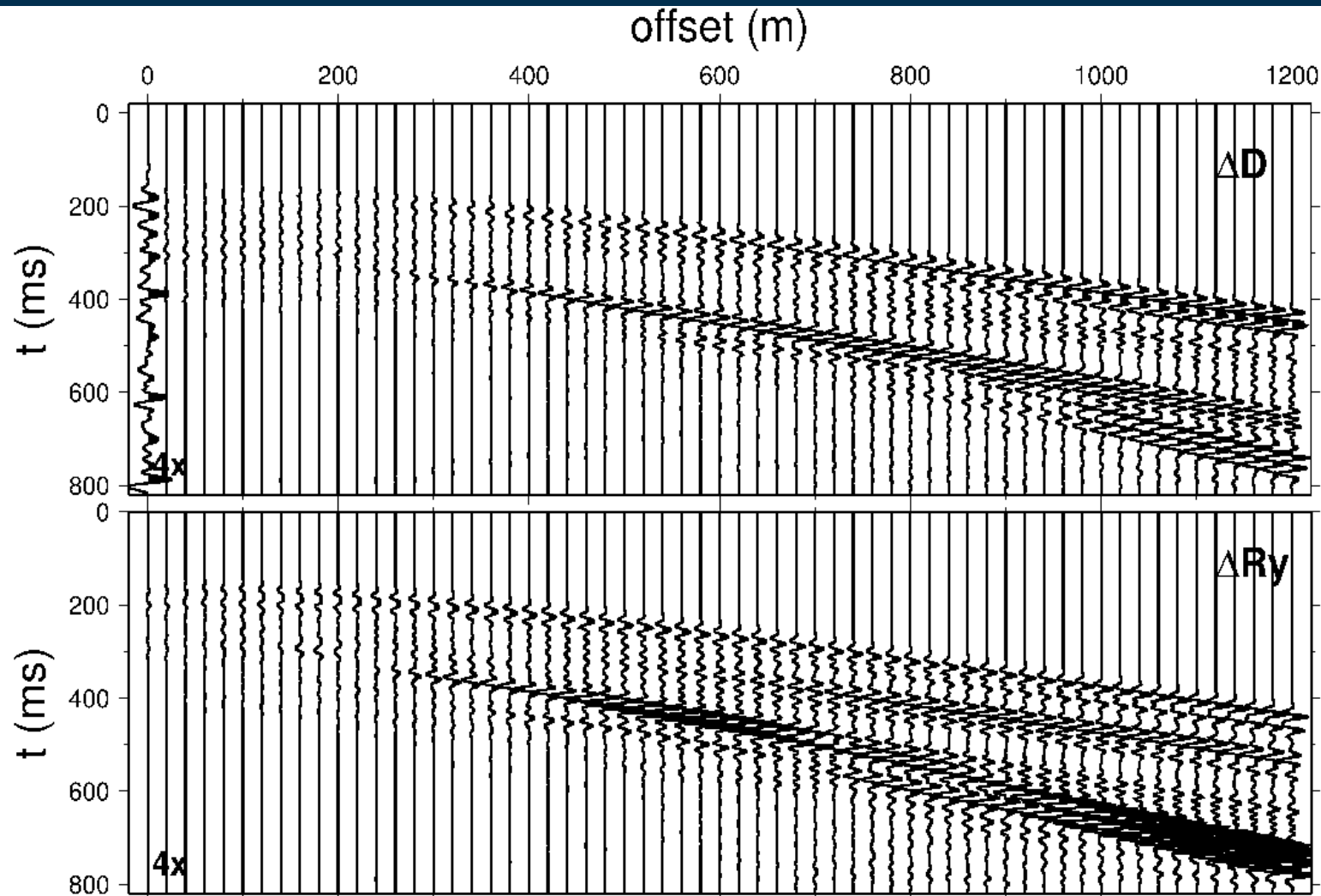
Source: Fz force, 30 Hz Ricker wavelet

Traces: 100%H₂O-50%CO₂; plot amp = 0.375×10^{-15}

Pressure (P wave only) + TR-Theta (S wave only)

Delta AVO for 75% CO₂ saturated sandstone

David Aldridge, et. al., Sandia National Lab, 2007-2009



Model: 25 m thick sandstone, $\phi = 0.25$, $\tau = 2.5$, $k_0 = 1D$

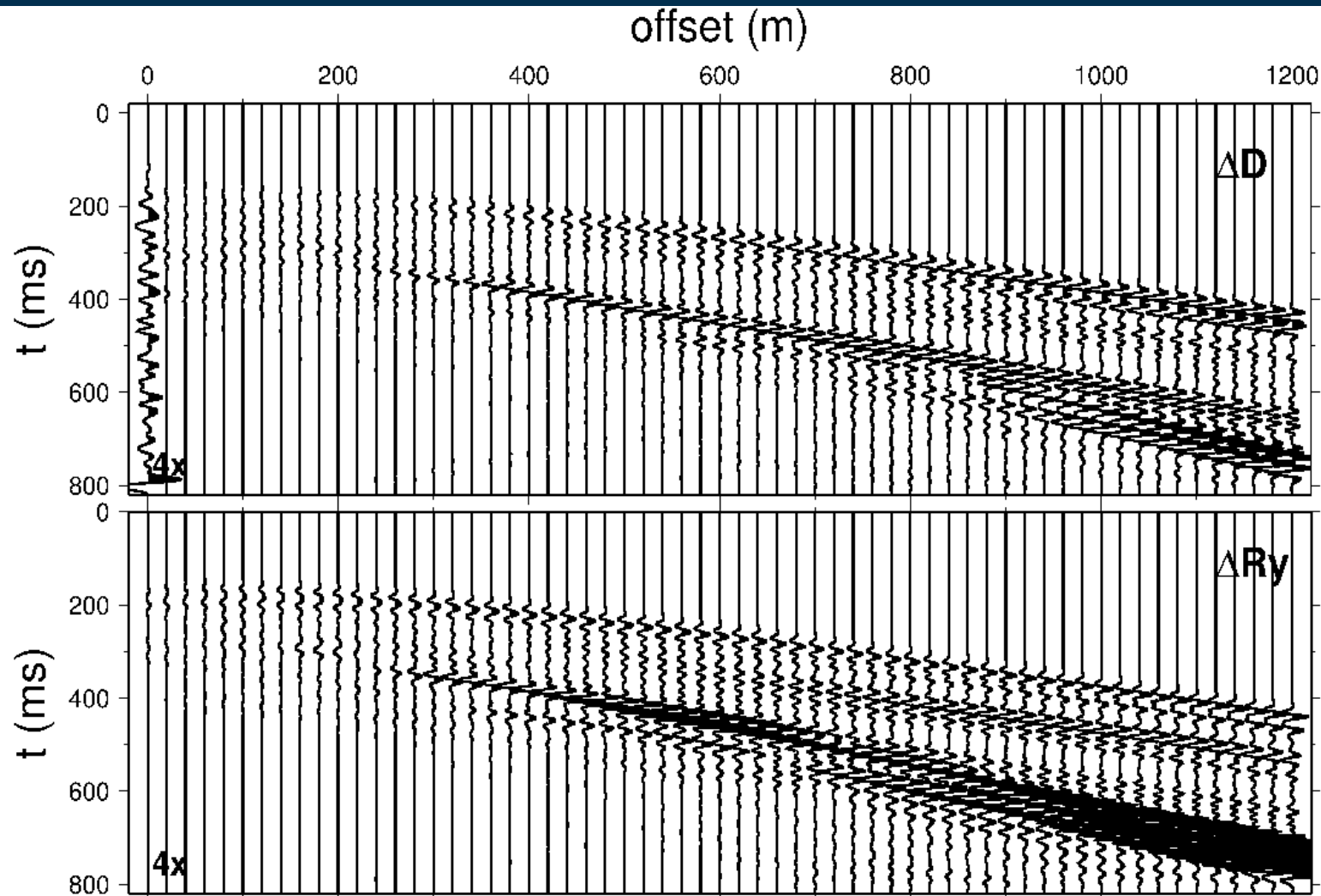
Source: Fz force, 30 Hz Ricker wavelet

Traces: 100%H₂O-75%CO₂; plot amp = 0.375×10^{-15}

Pressure (P wave only) + TR-Theta (S wave only)

Delta AVO for 100% CO₂ saturated sandstone

David Aldridge, et. al., Sandia National Lab, 2007-2009



Model: 25 m thick sandstone, $\phi = 0.25$, $\tau = 2.5$, $k_0 = 1D$

Source: Fz force, 30 Hz Ricker wavelet

Traces: 100%H₂O-100%CO₂; plot amp = 0.375×10^{-15}

Igel / Aldridge Point Array

At a **single** 3-C / 3- Θ sensor:

$$\text{S VELOCITY} = \frac{\text{transverse linear acceleration}}{\text{orthogonal transverse rotational velocity}}$$

Potential use for shear statics:

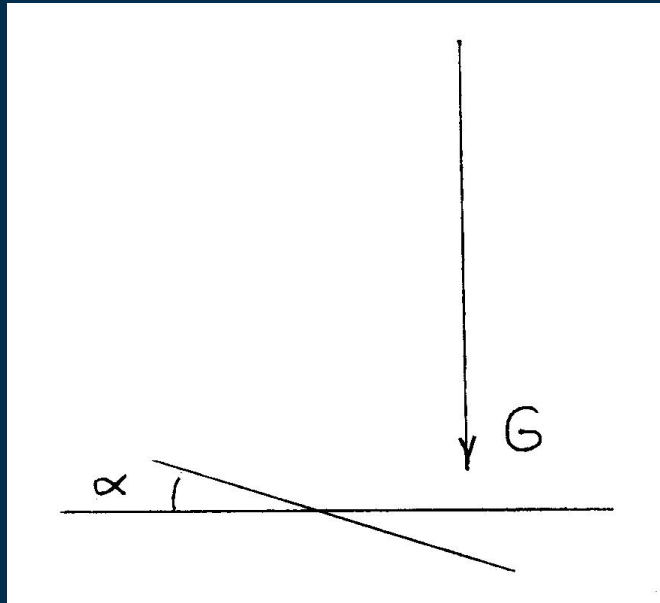
get shear velocity around each 6-C phone.

From 3-C *and* 3- Θ data can determine direction AND velocity of shear wave propagation.

Shear Waves / C-Waves

- Great effort in the 1980's pursuing shear sources, *but* always with 3-C linear phones that didn't differentiate P from S.
- Now we can use **Rotational Phones** to get pure shear **3- Θ** data as a 3-component vector.
- Related concept: use **Pressure Gradients** to get pure compressional **3- ∇P** data as a 3-component vector.
- **Processing** issues:
 - 3 components** with mode purity (gain, decon).
 - Direction and propagation velocity.

Tilt Correction of Linear using Rotation



If a **Horizontal Linear** Phone tilts by angle α , then it picks up error:

$$\text{Gravity} * \sin(\alpha)$$

Micro- or milli-radians of tilt give a potentially significant error compared to weak seismic reflections (C-waves).

Consider a worse case:
weak C-wave data on horizontal phone at
same time as strong ground roll in
uncorrelated vibroseis data

Possibly(?) Map Frac Density

Some precedent: Lg Scattering measurements in earthquake seismology & use of vertical rotation

Perhaps we can use rotation measurements to map scattering / fracture density:

- Use time, offset windows to target shales in Resource Plays. Compare Polarization modes.**
- Azimuthal Anisotropy of 6-C (3-C and 3- Θ):**
 - Amplitude Anisotropy**
 - Velocity Anisotropy**
 - Shear splitting**

Ground Roll – Rayleigh Waves

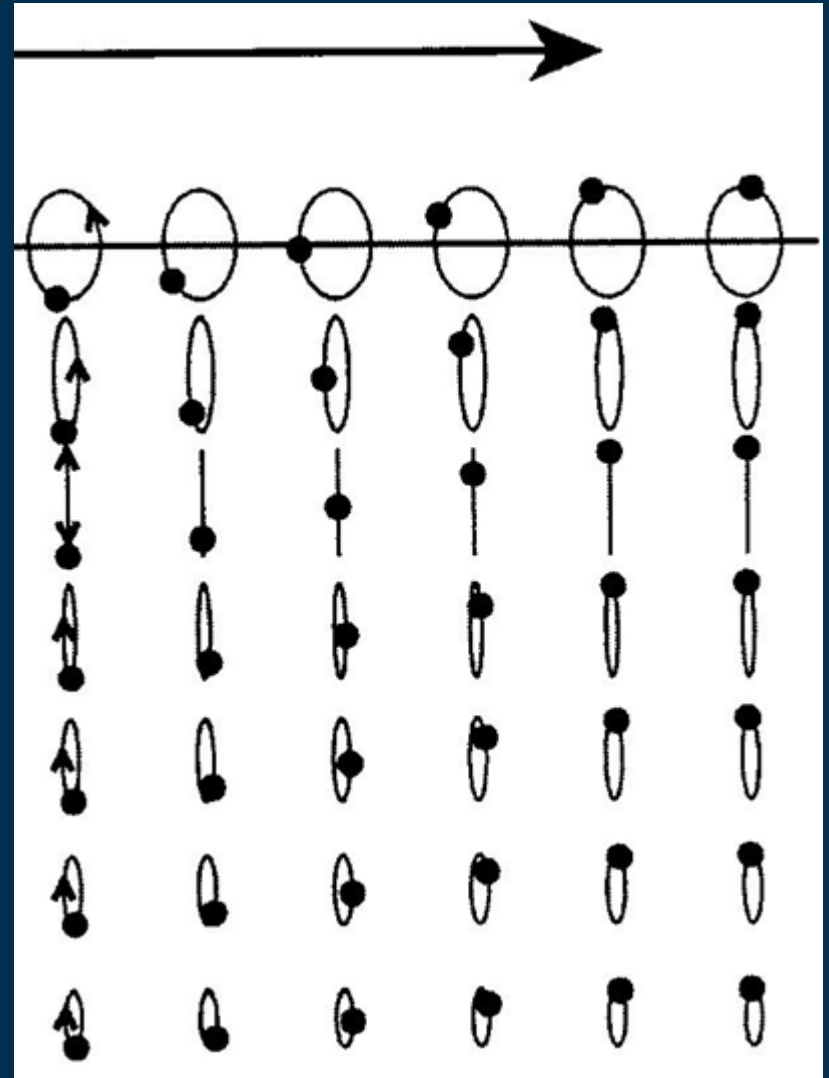
Classic Rayleigh Retrograde
Elliptical motion:

only for Poisson Solid (0.25)

For Poisson Ratio above
0.265, have **3 modes** of
polarization.

Hodograph of point is not
actually 'true' rotation.

Rayleigh wave / Rotational
Seismic: not simple.



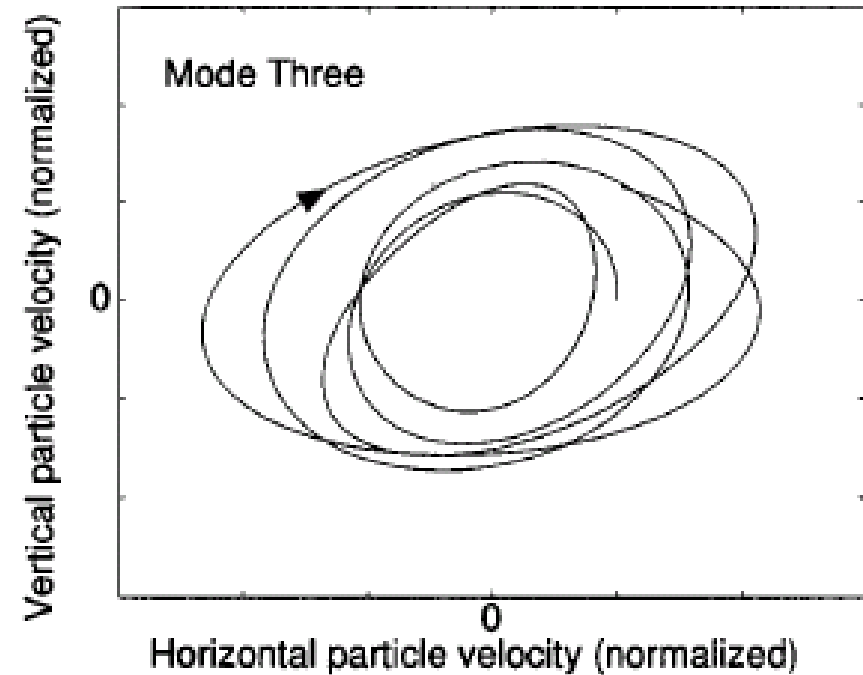
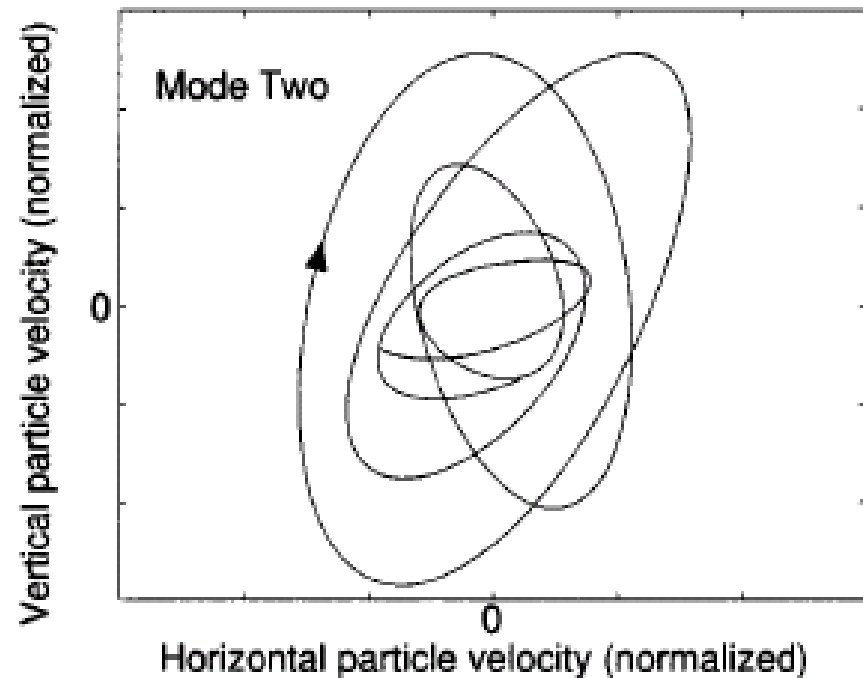
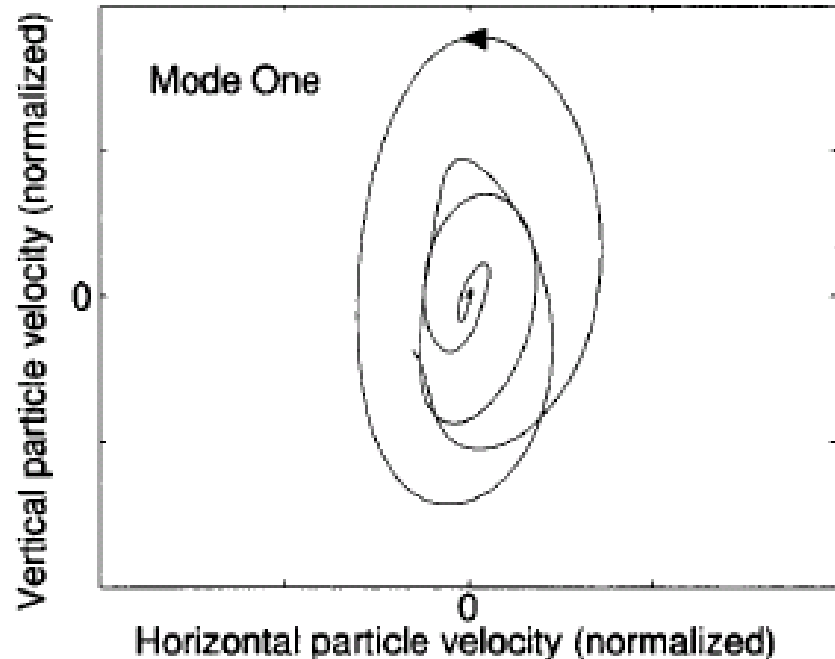
Rayleigh Modes: Hodographs

High Poisson Ratio
Near Surface

Mode 1 – Retrograde

Modes 2 & 3 – Prograde

Smith, et. al., JASA 1998



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Rotational Seismic – 'Status'

- E&P Operating Companies
 - Activity in only very few (~1) Companies.
 - No testing to date
- Geophysical (Equipment/Contractor) Cos.
 - **2 [3] contractors have done small scale tests**
- **3 +/- Rotational Phone manufacturers**
- Patent Status (8 companies)
 - Defensive
 - Perception that Rotational Seismic has potential to be a significant technology, *but*

Rotational Seismic – ‘Failures’

- **Failure in Market Research**
 - Confuse technical enthusiasm vs. willingness for expenditures / investment
- **Not Backward Compatible**
- No incremental path / no retrofit
- Not Market Driven – no user requests
- Not Proven –
 - **Wavefield** Scale in multiple locations
 - Meaningful **Geological** Scale

Thank you for your attention !

- Rotational **sensor prototypes**
- **Limited data** - no geological scale data
- **Many Applications** possible:

Spatial Sampling

Shear Selectivity – micro-seismic & C- wave

“Other” – AVO, Point Array, Tilt correction, Frac density, Rayleigh wave(?), etc.

- **Technically attractive for multiple purposes**
- **BUT, commercially ‘challenged’**

ROAD AHEAD ?

Acknowledgements

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