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EXPERIMENTAL MEASUREMENTS OF SHOCK PROPERTIES OF STISHOVITE

M. D. Furnish¹ and E. Ito²¹Experimental Impact Physics Dept. 1433, Sandia National Laboratories, Albuquerque NM 87185-0821 and,²Okayama University, Okayama, Japan

We have synthesized, characterized and performed Hugoniot measurements on monolithic samples of stishovite. Synthesis was accomplished in a multianvil press with pyrophyllite gaskets and carbon heaters. The samples had densities ranging from 3.80 to 4.07, corresponding to stishovite volume fractions of 0.7 to 0.87, a range confirmed by NMR analysis. They had no significant impurities except less than 1% carbon. Samples ~1 mm thick and 3 mm diameter were tested in reverse- and forward-ballistics modes on a two-stage light gas gun, using velocity interferometry diagnostics. Impact velocities ranged from 4.0 to 6.5 km/sec. Hugoniot stresses for the four successful tests ranged from 65 to 225 GPa. At higher stresses significant uncertainties arise due to impact tilt/nonplanarity issues. Results are consistent with earlier predictions of the stishovite Hugoniot based on quartz-centered Hugoniot data, static-compression (diamond-anvil cell) data and hydrostatic multianvil cell data. Release behavior appears to be frozen. These results are remarkable in view of the small size of the samples used.

BACKGROUND

Stishovite is a dense, high pressure polymorph of silica (SiO_2). It has an ambient density of 4.3 Mg/m^3 , is stable at stress above 7.5 GPa at room temperature, and is composed of a structure in which the silicon atoms are each surrounded by six oxygen atoms in an octahedral arrangement.

A high-pressure phase of quartz observed by Wackerle¹ in shock wave experiments was identified as stishovite by McQueen et al.² on the basis of its observed transition pressure (~14 GPa) and density. McQueen et al.² derived a Grüneisen ratio of $V(\partial P/\partial E)_V \equiv \gamma_0 = 0.9$, a heat of formation relative to quartz of $1.5 \times 10^{10} \text{ erg/gm}$, and a bulk sound speed of 10 km/sec for stishovite. This large heat of formation indicates why stishovite is so rarely found in nature under atmospheric conditions (found primarily at recent impact sites such as Meteor Crater, Arizona).

Grady et al.³ performed a set of experiments on shocked quartz using manganin gauges to measure stress histories, obtaining release information as well as Hugoniot states. Results from that study are shown in Figure 1, together with Hugoniots from other

sources as noted. Interestingly, that and subsequent studies (e.g. Chhabildas and Miller⁵) showed that releases are "frozen" down to about 8 GPa, then a transition to a lower-density phase occurs.

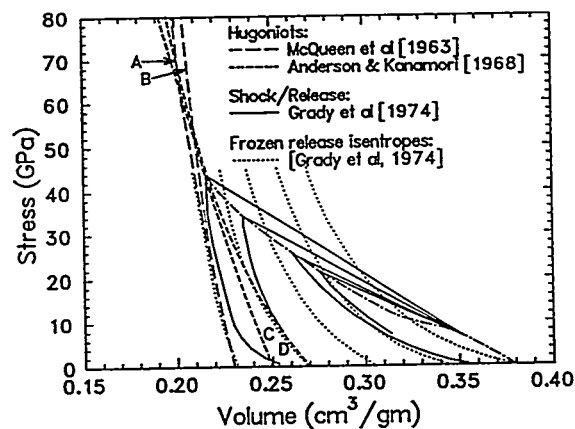


FIGURE 1. Summary of earlier experimental and theoretical studies. Curve "A" is centered at $p_0 = 4.3 \text{ gm/cm}^3$ ($V_0 = 0.233 \text{ cm}^3/\text{gm}$); "B" at 2.65 gm/cm^3 ($V_0 = 0.377 \text{ cm}^3/\text{gm}$) (McQueen et al²); "C" at 4.0 gm/cm^3 ($V_0 = 0.25 \text{ cm}^3/\text{gm}$); and "D" at 3.7 gm/cm^3 ($V_0 = 0.27 \text{ cm}^3/\text{gm}$) (Anderson and Kanamori⁴).

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Prior to the present study, no shock compression studies using stishovite samples had been performed. The goal of the present study is to measure the Hugoniot of silica beginning with stishovite samples (initial densities of 3.8 - 4.2 Mg/m³). This is interesting for two reasons. First, the thermal regimes exercised are quite different than for a Hugoniot centered on quartz. Second, the starting material is known to be stishovite in the present experiments.

SAMPLE PREPARATION AND CHARACTERIZATION

Samples of monolithic, polycrystalline stishovite (actually, a mixture of SiO₂ polymorphs) were prepared in the laboratory of Hideyuki Fujisawa of the University of Tokyo. A multianvil press, with tungsten carbide anvils, pyrophyllite gasket and graphite heaters, was used. The starting material was an activated silica (a gel). The sample was held at 1100 °C and 10 GPa for slightly over an hour, with a slow cool-down. Two samples were prepared this way; one was recovered whole, while the other was recovered in six disks (apparently fractured during the cooling/depressurization phase). The samples were about 3 mm diameter and 8 mm long. From this, samples about 1 mm thick and 3 mm diameter were obtained, and are shown in Figure 2.

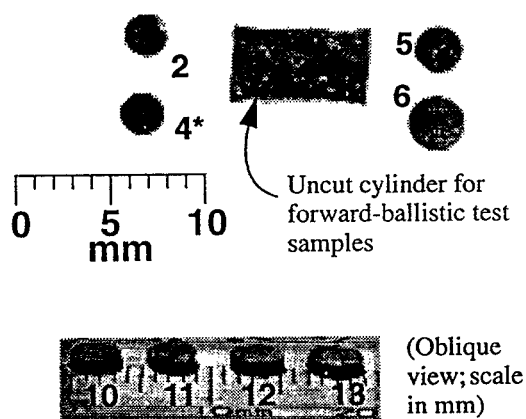


FIGURE 2. Stishovite samples used for present tests. Sample 4 was used for the successful reverse-ballistic test (ST-2); samples 11 - 13 were used for the successful forward-ballistic tests (ST-11, 12, 13, respectively).

The samples were characterized by various methods, including visual examination, X-ray fluorescence, density measurements, NMR, electron microprobe, x-ray imagery, ultrasonic measurements of sound velocities, and ultrasonic microscopy. The material studied appears to have a mass fraction of 80 - 91% stishovite (balance quartz/coesite). Electron microprobe and X-ray fluorescence characterizations showed minor carbon contamination (<1%), with no other significant impurities.

IMPACT EXPERIMENTS CONDUCTED

Although a total of 14 tests were performed (8 with stishovite samples), only 4 tests were found to provide useful data on stishovite. Three utilized a transmitted-save (forward-ballistic) configuration (Figure 3), while the fourth utilized a reverse-ballistic configuration (Figure 4). Selected parameters for these tests are shown in Table 1.

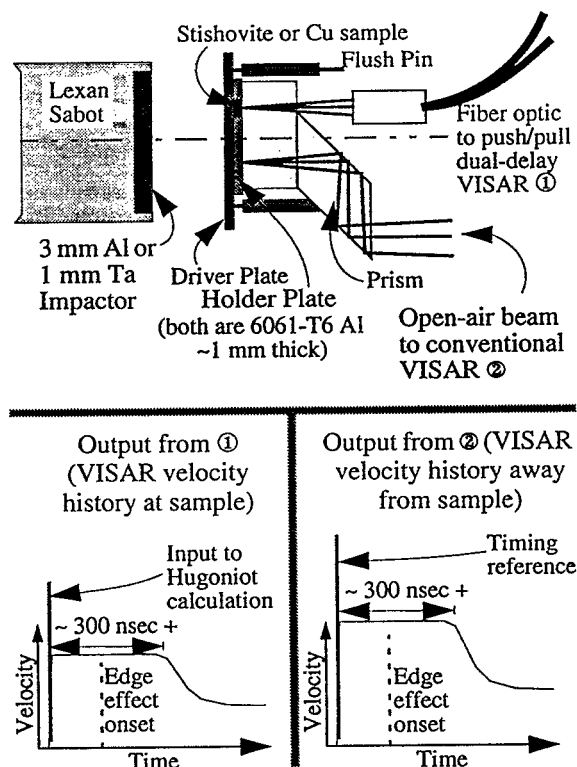


FIGURE 3. Forward-ballistic configuration.

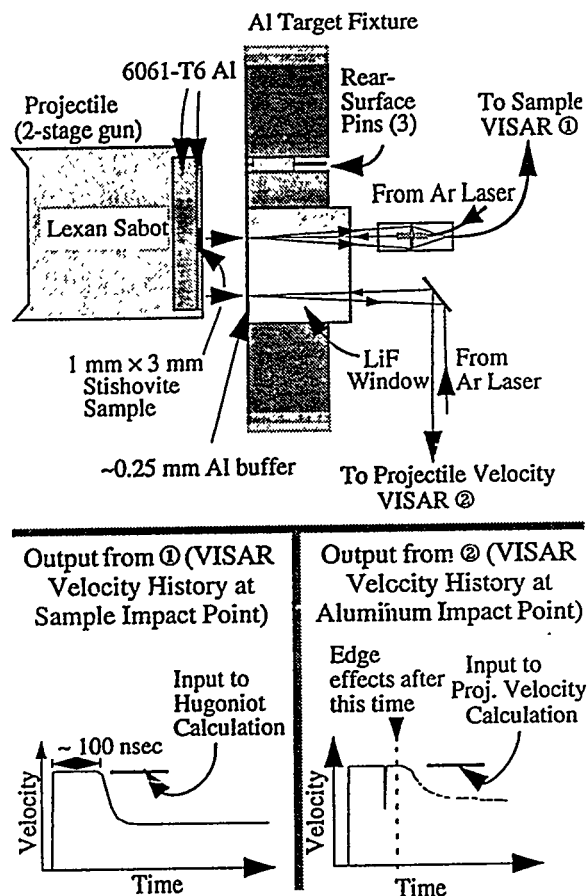


FIGURE 4. Configuration of reverse-ballistic stishovite tests

The primary difficulties for the forward-ballistic tests lay in obtaining a precise (2-3 ns) transit time, done by a careful time correlation of the two VISAR signals. The reverse-ballistic tests proved to be unreliable in delivering the sample to the proper point on the target, and in turn generating acceptable waveforms for both velocity interferometers.

TABLE 1. Test matrix

Test	ST2*	ST11	ST12	ST13
Projectile Vel. (km/s)	4.01	4.88	4.98	6.54
Impactor Material	—	Al	Ta	Ta
Impactor Thickness (mm)	—	2.995	1.030	0.779
Sample measurements -				
Thickness (mm)	0.859	0.931	0.978	0.977
Diameter (mm)	2.40	3.600	3.700	3.872
Density (Mg/m ³)	4.067	3.959	3.954	3.833

*Reverse-ballistic test; sample backed by 160 μ m epoxy.

RESULTS

Hugoniot results for the four successful tests are summarized in Table 2, and plotted in Figures 5 and 6. The results from the two methods seem to be mutually consistent, and are in substantial agreement with earlier quartz-centered Hugoniot data and with theoretical predictions. The primary source of uncertainty for the forward-ballistic tests is the tilt of the projectile.

The diamond-cell data⁶ probably slightly underestimate density because of the axial X-ray diffraction

TABLE 2. Stishovite Hugoniot states*

Test # $\rightarrow$	ST2	ST11	ST12	ST13
Strain	0.16(4)	0.24(3)	0.34(4)	0.33(10)
ρ_0 (Mg/m ³)	4.07(30)	3.96(10)	3.95(10)	3.83(10)
Particle Vel (k/s)	1.50(12)	2.06(9)	3.44(8)	4.36(25)
Stress (GPa)	59(4)	69(3)	136(9)	223(36)
ρ (Mg/m ³)	4.8(5)	5.2(3)	6.0(4)	5.7(1.1)
Shock Vel. (k/s)	9.6(1.4)	8.5(7)	10.0(9)	13.4(2.7)
ρ/ρ_0	1.19(5)	1.32(6)	1.52(10)	1.48(28)
V (m ³ /Mg)	0.207(19)	0.191(9)	0.166(11)	0.176(28)

*Figures in parentheses represent uncertainties

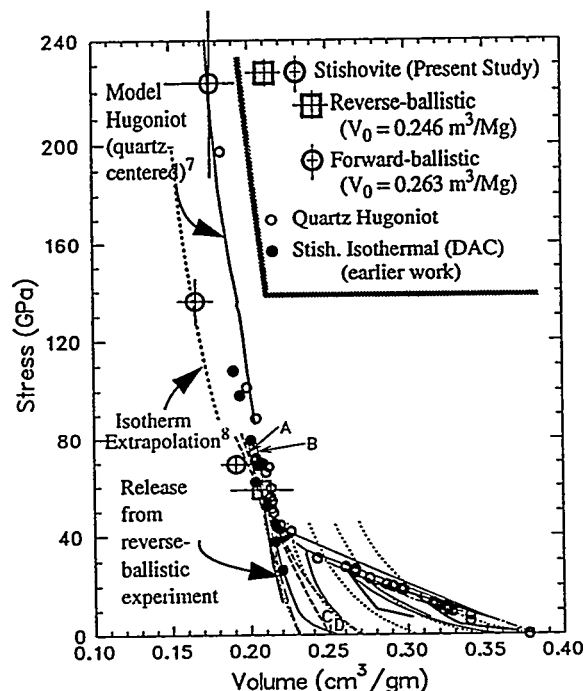


FIGURE 5. Present Hugoniot measurements compared with Fig. 1 curves. Small circles represent an aggregate of previous experiments centered on quartz; disks are diamond cell data⁶

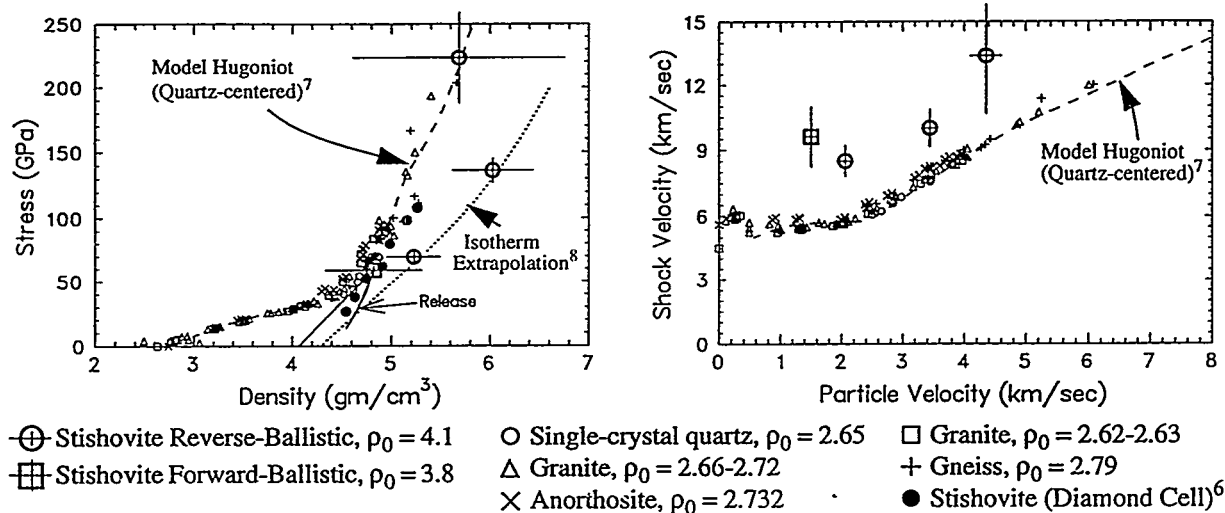


FIGURE 6. Hugoniot and release properties of stishovite, compared with representative rock Hugoniot data and stishovite diamond-cell data.

method used, while the isothermal cubic-anvil static data must be extrapolated from 11 GPa to the present region of interest.

Release paths may be derived from the reverse-ballistic test (Furnish⁹), and depend primarily on the time-of-arrival of the release in the velocity history. This release is plotted with the Hugoniot data and the theoretical Hugoniots and frozen-phase release isentropes in Figure 5. The corresponding waveform match is shown in Figure 7 (WONDY V with custom hysteresis model). Other calculated paths in Figure 7 show that the release inferred from this method is not seriously affected by edge effects.

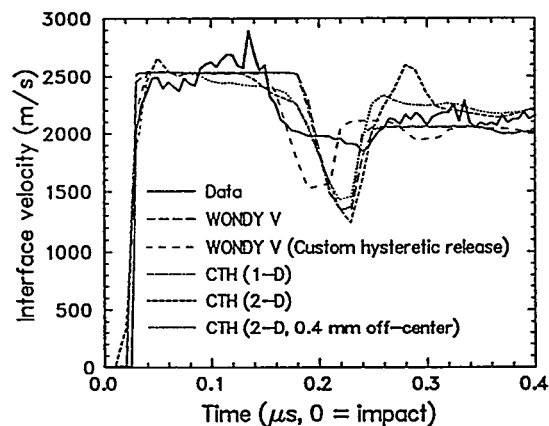


FIGURE 7. Waveform from ST2 reverse-ballistic test, compared with simulations from WONDY V and CTH codes. All simulations used Mie-Grüneisen EOS unless specified.

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